26. CALCAREOUS NANNOFOSSILS: LEG 19 OF THE DEEP SEA DRILLING PROJECT

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INTRODUCTION

Calcareous nannofossils are the minute $(1-20\mu)$ skeletal remains of planktonic members of the golden brown algae. They are essentially tropical or subtropical organisms but cosmopolitan and high-latitude species are known (Edwards, 1968a, b; McIntyre, Be, and Roche, 1970; Worsley and Martini, 1970;). However, remains are not common in high-latitude sediments. Therefore, it is not surprising that the Neogene assemblages found in the North Pacific and Bering Sea have low population density and species diversity. Paleocene and Maestrichtian assemblages from the North Pacific are more diverse, suggesting more moderate global climate, and/or large translatitudinal plate motion.

In the Bering Sea, six sites were drilled (Figure 1) and the oldest sediments cored are Miocene. Most analyzed samples are completely devoid of calcareous nannofossils and little interpretation is offered here of the significance of the few fossiliferous samples, other than to note in the site reports where they occur. The reader is referred to the site reports and to the foraminifera chapter in this volume for detailed paleoclimatic interpretation.

In the North Pacific, five sites (Figure 1) were drilled, two of which contain more diversified nannofloras than any site in the Bering Sea. Meiji Guyot (Site 192) at the northwestern terminus of the Emperor seamount chain contains a somewhat impoverished Neogene assemblage underlain by progressively richer ones of Oligocene, Eocene, and Maestrichtian age. The Mesozoic-Cenozoic boundary was cored here, thereby demonstrating the complete absence of Paleocene strata at this site. The Gulf of Alaska (Site 183) is devoid of Neogene nannofossils and was apparently below the carbonate compensation depth (CCD) throughout its Neogene depositional history. Below, extremely sparse nannofloras occur in a thick turbidite sequence between a thin (1 to 20 meters) Oligocene nannofossil-rich chalk bed and a 1- to 2-meter-thick upper lower Eocene nannofossil-rich limestone bed.

The other North Pacific sites (Figure 1) contain no nannofossils.

CALCAREOUS NANNOFOSSIL BIOSTRATIGRAPHY

Biostratigraphy is ultimately based on accurately determining the relative abundance or absence-presence relationships of taxa in samples. If specimens are too rare in samples, it is not generally possible to accurately determine relative abundances or stratigraphic ranges. For this reason, detailed biostratigraphic subdivision of the impoverished Neogene assemblages recovered in the Bering Sea and North Pacific is not attempted. Most biostratigraphic control for this part of the column is based on silicoflagellates and diatoms. The reader is referred to those chapters for biostratigraphy. Unfortunately, siliceous fossils are absent below the Neogene so that age determinations for this part of the column are entirely based on planktonic foraminifers and calcareous nannofossils for Meiji Guyot (Site 192) and nannofossils alone for the Gulf of Alaska (Site 183).

No new biostratigraphic subdivisions are proposed for the North Pacific Paleogene. Instead, samples have been assigned to Martini's (1971) standard zones where feasible or assigned to an epoch or part thereof where a zonal assignment could not be made. Tables 1 and 2 are the nannofossil range charts of the North Pacific. In these range charts, A indicates more than 1 specimen per field of view; C, 1/10 to 1 specimen per field of view; R, less than 1/100specimen per field of view; and X, presence noted, abundance not determined. The field of view was at 640X magnification. Usually, as was done in this study, smear slides are used for age determinations. They frequently contain in excess of 30 specimens per field of view so that 20 minutes of observation at a rate of 5 fields of view per minute would total 100 fields or 3000 specimens. Unfortunately many of the samples studied here have far less than 1 specimen per field of view so that only a few specimens were found in 20 minutes of searching. Therefore, in the Paleogene portion of Site 183 where nannofossils are extremely scarce, with the exception of the two carbonate beds, and where no other fossil control is available, searching was continued long past twenty minutes and for as much as 4 to 5 hours for some critical samples. Even with these long searching times, the total number of specimens observed did not exceed a few hundred to a thousand, most of which were not age diagnostic. It is not improbable that still longer searching will reveal even more species for these samples.

NANNOFOSSIL PALEOECOLOGY OF THE NORTH PACIFIC

Background

Paleoecology, like biostratigraphy, is essentially dependent upon determining relative abundance or presenceabsence of taxa with samples. In biostratigraphy, the age of the sample is unknown and relative abundance or presenceabsence is assumed to have evolutionary and therefore age significance. The reverse is true in paleoecology where the age of a sample must be known before attaching environmental significance to relative abundance or presenceabsence of taxa. This is, at least in part, circular if the same fossil group is used for both age determination and

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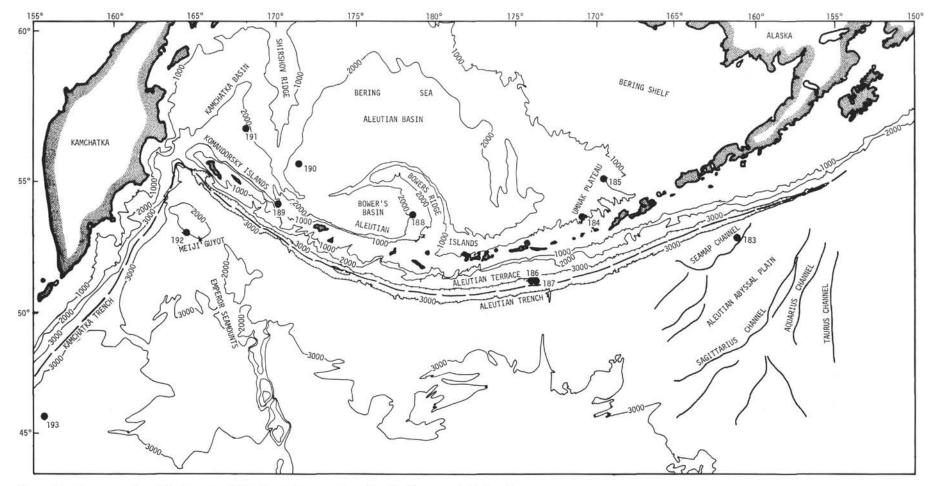


Figure 1. Base map, Leg 19, showing drill sites in the northern Pacific Ocean and Bering Sea.

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Sample	Age	Zone	Chiphragmalithus acanthodes	Markalius astroporus	Discoaster barbadiensis	Braarudosphaera bigelowi	Discoaster binodosus	Reticulofenestra dictyoda	Coccolithus chiastus	Discoaster deflandrei (cf.)	Reticulofenestra dictyoda (cf.)	Braarudosphaera discula (cf.)	Coccolithus eopelagicus	Discoaster falcatus (cf.)	Ericsonia fenestriata	Cyclococcolithus formosus	Cyclococcolithus luminis	Sphenolithus moriformis	Discolithina multipora	Chiasmolithus omaruensis	Reticulofenestra omaruensis	Coccolithus pelagicus	placoliths (indeterminate)	Discolithina pygmaea	Reticulofenestra scissura	Bramlettis serraculoides	Chiasmolithus solitus	Cyclococcolithus sp.	Fasciculithus sp.	Helicopontosphaera sp.	Reticulofenestra sp.	Chiasmolithus sp.	Ericsonia subdisthica	Marthasterites tribrachiatus	Reticulofenestra umbilica	Isthmolithus unipons
25(CC)	LM. Oligocene		-	.0905		100	23	x		12.00	x		x			1000	-		x	-	1	x	-	x			F	x			104			-		1
26-1 (90) 26-1 (140) 26(CC)	D. M. Ongovene					x		x				x							x x		x	x			x						x	x x				
27-3 (30) 27-3 (45) 27-3 (90) 27-3 (120) 27-4 (30) 27-4 (60) 27-4 (108)	Lower Oligocene																																			
28-1 (121) 28-2 (20) 28-2 (50) 28-2 (125) 28-3 (50) 29-1 (133)	(barren)																																			
29-1 (145) 30-1 (130) 30-1 (140) 31-1 (148) 31-1 (70) 31-1 (80)	(barren)							R			C																									
31-1 (100 31(CC) 32-1 (43)	Lower Oligocene							X X			x											X X	R		x								x			
32-1 (70) 32-1 (130) 32(CC)	(barren)																																			
33(CC) 34(CC)	Lower Oligocene							X X			x				x			x		x		x		x								x			X X	
35-1 (30) 35-1 (47)	(barren)																					_													T	
35-1 (53) 35-1 (55) 35-1 (60) 35-1 (112) 35(CC) 36-1 (90)								x x			x											x x			x x							x				
36(CC) 37-1 (105) 37(CC) 38-2 (20) 38-2 (40) 38-2 (60)	Middle to Upper Eocene							x														x			X R x			x			x					
38-2 (70) 38-2 (80) 38-2 (90) 38-2 (120) 3802 (140) 38-3 (68)				x	x			x c								x		x x				x			x x	x			x					1	R	x
38-3 (86) 38-3 (130) 38(CC) 39-1 (53) 39-1 (54) 39-1 (91)	Lower Eocene		x	x x	x	x	x		x	x			x	x		x x	x	x				x			x x x		x	x	x	x				x x	x	x
39-1 (113) 39-1 (121) 39-1 (130) 39-1 (145) 39(CC)								x			x											x													x x	

TABLE 1 Nannofossils, Leg 19

TABLE 2A Nannofossils, Leg 19

	Hole 192		ula	Watznaueria barnesae	Gephyrocapsa caribbeanica	Reticulofenestra dictyoda (cf.)	Gephyrocapsa ericsoni	Cruciplacolithus helis (aff.)	Emiliania huxleyi (4μ)	Emiliania huxleyi	Pseudoemiliania lacunosa (cf.)	Cyclococcolithus leptoporus	Tetralithus murus (aff.)	Gephyrocapsa oceanica	Coccolithus pelagicus (bridge)	Coccolithus pelagicus	Placoliths (indeterminate)	Reticulofenestra pseudoumbilica	Reticulofenestra pseudoumbilica (cf.)	Tetralithus pyramidus	Discoaster sp.	Gephyrocapsa sp.	Syracosphaera sp.	Reticulofenestra bisecta	Syracosphaera clathrata (cf.)	Coccolithus eopelagious	Syracosphaera japonica	Sphenolithus moriformis	Cyclococcolithus neogammation	Chiasmolithus oamaruensis	Reticulofenestra scissura	Reticulofenestra sp.	Coccolithus tenuistriatus	Reticulofenestra umbilica (cf.)	Reticulofenestra umbilica	Placoliths (2-3µ)
Sample	Age	Zone	Micula	Wat	Gep	Reti	Gep	Cruc	Emi	Emi	Pseu	Cyc	Teti	Gep	Coc	Coc	Plac	Ret	Ret	Tet	Disc	Gep	Syr	Ret	Syr	Coc	Syr	Sph	Cyc	Ŀ.	Ret	Ret	Co	Ret	Ret	Pla
1-1(80) 1(CC) 2-1(114) 2-2(80)	Upper Pleistocene	NN21						R X R		R X R		X R		R R	C X	A X R C			R			x														
2-3(80) 2-4(80) 2(CC) 3-1(79) 3-2(70) 3-3(79)	(barren)																																			
3-4(75) 3(CC) 4-1(80) 4-2(80) 4-3(80) 4-4(80) 4-6(80) 4-6(80) 4-6(80) 4-(CC) 5-1(80) 5-1(84) 5-2(80) 5-3(80)	Middle Pleistocene	NN20 (?) NN19		x	C A X	c c				R C X	c c	C R	x	c c	0		x x		R C	x																
5-4(80) 5-5(80) 5-6(140)	(barren)																																			
5(CC)	L. Pleist.				-										x			-	-			-	-	-	x				-		-		-	-	1	
6-1(120) 6-2(80) 6(CC) 7-1(100) 7-2(40) 7-3(80) 7-4(80)	(barren)																																			
7(CC)	U. Plio.				_						Х	Х				1			х		X													_		
8-1(125) 8-2(94) 8-3(120) 8-4(68) 8-6(80) 8-6(80) 8(CC) 9-1(130) 9-2(80) 9-3(80) 9-5(80) 9-6(80) 9(CC) 10-1(140) 10-2(80) 10-3(80)	(barren)																x																			
10-4(80) 10(CC) 11-1(top) 11-2(80) 11-(CC) 12-1(top) 12-2(80) 12-3(80) 12-4(80) 12-4(80) 124(CC) 13-2(80) 13-4(80) 13-5(80) 13-5(80) 13-4(20) 14-1(110) 14-2(80) 14-3(80) 14-2(80)	Lower Pliocene				R R	C X X C R X			R			x x		с		C C C X C C C X R X X	x x c c		C C C C R			X R X R X C C	R													

	1	-			1			1					
	Hole 192		Micula Watznaueria barnesae Gephyrocapsa caribbeanica	Reticulofenestra dictyoda (cf.) Gephyrocapsa ericsoni Cruciplacolithus helis (aff.)	Emiliania.huxleyi (4μ) Emiliania huxleyi Psuedoemiliania lacunosa (cf.)	Cyclococcolithus leptoporus Tetralithus murus (aff.) Gephyrocapsa oceanica	Coccolithus pelagicus (bridge) Coccolithus pelagicus Placoliths (indeterminate)	Reticulofenestra pseudoumbilica Reticulofenestra pseudoumbilica (cf.) Tetralithus pyramidus	Discoaster sp. Gephyrocapsa sp. <u>Syracosphaera sp.</u> Reticulofenestra bisecta Syracosphaera clathrata (cf.) Coccolithus eopelagicus	Syracosphaera japonica Sphenolithus moriformis Cyclococcolithus neogammation	Chiasmolithus oamaruensis Reticulofenestra scissura Reticulofenestra sp.	Coccolithus tenuistratus Reticulofenestra umbilica (cf.) Reticulofenestra umbilica	Placoliths (2-3µ)
			Micula Watznau Gephyro	eticulo ephyrc rucipla	miliani miliani suedoe	ycloco etralith ephyro	occolit occolit lacolit	teticulo teticulo 'etralitl	Discoaster sp. Gephyrocapsa Syracosphaera Reticulofenes Syracosphaera Coccolithus et	yracos pheno ycloco	Chiasm Reticul Reticul	Coccoli Reticul Reticul	Placolit
Sample 15-1(80) 15-2(80) 15-3(80) 15(CC) 16-1(120) 16-2(80) 16(CC)	Age (barren)	Zone	WAO	200	шша	0 1 0	x	A A F		2 2 2		0 1 1	
17-2(60) 17(CC) 18-1(80) 18-2(80) 18-3(80) 18-5(80)	Upper Miocene	NN12 (?)		C C X R	di la composita da la composit	C R R	C A C A C C		R X C	x			
18(CC) 19-1(80)	(barren)				2		1						
19-2(80) 19(CC) 20-1(80) 20-2(80) 20(CC)	(barren)						сх						
21-1(22) 21-2(20) 21(CC) 22-1(42) 22-2(65) 22-2(80) 22-3(100) 22(CC) 23-1(70)	Upper Miocene	NN12		A X X		с	CA CX CX X X X X	c x	R			x	x
23-2(90) 23(CC) 24-1(130) 24-2(80)	(barren)												
24(CC) 25-1(80) 25-2(80) 25(CC) 26-1(130) 26-2(109) 26(CC) 26(CC) 27-1(192)	Middle to Upper Miocene	NN11		x x x x			R X X X	R	X X A X X X	-	с	x	
27-1(135) 27-2(125) 27-3(20) 27-3(33) 27-3(76)	(barren)												
27-4(73) 27-5(80) 27(CC) 28-1(22) 28-1(22) 28-1(22) 28-1(22) 28-1(22) 28-1(22) 28-1(22) 29-1(112) 30-2(80) 30-1(102) 30-2(80) 31-2(80) 31-2(80) 31-2(80) 32-1(133) 32-2(52) 32-3(80) 32-2(52) 32-3(80) 32-2(52) 32-3(80) 32-2(52) 32-3(80)	Middle Miocene	NN5		x x x x			X X X X X X X R		A R R X X A X A	x - X		с	x
33-2(80) 33-3(80) 33(CC) 34-1(120) 34-2(72)				x			R R		A A X C R	х	x		

,	fole 192			watznauerta barnesae Gephyrocapsa caribbeanica	Reticulofenestra dictyoda (cf.)	Gephyrocapsa erícsoni	Cruciplacolithus helis (aff.)	Emiliania huxleyi (4µ)	Emiliania huxleyi	Pseudoemiliania lacunosa (ct.)	Cyclococcolithus leptoporus	us murus (arr.)	Gephyrocapsa oceanica	nus pelagicus (ornuge)	Placoliths (indeterminate)	Dericulationestre neerdoumhilice	Reticulofenestra pseudoumbilica (cf.)	Tetralithus pyramidus	er sp.	Gephyrocapsa sp.	Syracosphaera sp.	Reticulofenestra bisecta	Syracosphaera clathrata (cf.)	Coccolithus eopelagicus	Syracosphaera japonica	Sphenolithus moriformis	Cyclococolithus neogammation	Chiasmolithus oamaruensis	Reticulofenestra scussura	Reticulofenestra sp.	Coccolithus tenuistriatus	Reticulofenestra umbilica (cf.)	Keticulorenestra umbilica	is (2-3µ)
Sample	Age	Zone	Micula	watznaueria b Gephyrocapsa	Reticulo	Gephyro	Cruciplac	Emulani	Emiliani	Pseudoer	Cyclocot	Tettallun	Coccoliti	Coccollin	Placolith	Dericula	Reticulo	Tetralith	Discoaster sp.	Gephyro	Syracosp	Reticulo	Syracosp	Coccolit	Syracosp	Sphenoli	Cycloco	Chiasmo	Reticulo	Reticulo	Coccolit	Reticulo	Keticulo	Placoliths (2-3µ)
34-2(90) 34-3(80) 34-4(77) 34-5(81) 34-6(92) 34(CC) 35-1(85) 35(CC) 36(CC)	Lower Pliocene	NN5			x x									;	x							X X X A X		x		x	x	C X	x					x

paleoecologic reconstruction. Therefore, without independent age control, paleoecological-biostratigraphic studies always contain an inherently large degree of subjectivity. This is especially true of the inferences about the Paleogene environment of the Gulf of Alaska (Site 183) where nannofossils are the sole fossil group recovered. The situation is better for Meiji Seamount where planktonic foraminifers are present also. For a detailed summary of the consequences of the paleoecological interpretations offered below, the reader is referred to the synthesis chapter in this volume.

Table 3 summarizes many of the significant physical, ecological, and taxonomic characteristics of four localities on the northern and eastern perimeter of the North Pacific (denoted by asterisks on Figure 1). As the paleontologic data of the table are mainly dependent on the absence or presence of taxa in the specified localities, it would be well to review the factors controlling nannofossil distribution in coeval deposits before interpreting the data. They are paleodepth of water, past levels of the carbonate compensation depth, past climate, and paleospastic position.

Information on the paleodepth of deposition for many oceanic (but not shelf) cores can be obtained if one assumes that all oceanic crust forms on ridge crests, subsequently founders as spreading proceeds, and is finally subducted into a trench. Sclater et al. (1971) have found a relationship between age of the oceanic crust and its depth below sea level. Most ridge crests average 2 to 3 km below sea level and ultimately founder to perhaps 5 km prior to subduction. When graphed, this relationship closely approximates a catenary curve of age versus depth. In this study, corrections have been applied to obtain paleodepths for the two oceanic sites. Only paleontologic criteria are available for determining paleodepth of the shelf sites but these were certainly much shallower than the oceanic ones.

Besides the depth to the sea floor, the carbonate compensation depth has also changed through time by as much as 1 km for the Atlantic (Hay, 1970; Ramsay, 1973). This means that even though "Sclater curves" have been applied to an oceanic locality, one cannot be sure of knowing the relation of the sediment-water interface to the carbonate compensation depth without additional information. Age versus CCD curves are available for the Atlantic (Ramsay, 1973), but not yet for the Pacific so that much of the information presented below with respect to this topic is subjective.

Besides being dependent on water depth and the CCD, the quantity of biogenic calcite in a sediment is also a function of surface water temperature at time of deposition. The colder the climate, the lower the specimen density and species diversity in surface waters and therefore in sediments. Climate at a site can change by either having climatic belts migrate across the site or having the site migrate across climatic belts. That both occur is demonstrated by the record of the ice ages for the first case or in the hypotheses of sea floor spreading and polar wandering for the second.

Meiji Guyot and the Gulf of Alaska have probably been strongly affected by all four of these processes. The question is, which dominated when? If a unique solution is attainable, information about paleoclimate, paleo-ocean chemistry, sea floor spreading, and absolute plate motion can be obtained. For this reason, interpretation is attempted regardless of the ambiguity in the small amount of data available in the hope it will stimulate further research into this complex and fascinating problem.

CONCLUSIONS

Available evidence seems to favor cool- to cold-water deposition for the Paleogene sediments of Meiji Guyot and the Gulf of Alaska. However, warm-water conditions are suggested for the Maestrichtian of Meiji Seamount. Species diversity (Table 3) in the North Pacific is only a fraction of that in Southern California. Only four species of *Discoaster* occur in the North Pacific compared to twentyone in coeval deposits in Southern California. If climates were similar for both regions at that time, discoasters should have been equally well represented in all four sites because they are known to occur abundantly in both shelf

CALCAREOUS NANNOFOSSILS

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	Hole 192A		1.00	Tetralithus aculeus (aft.) Reinhardtites anthophorus	Discoaster barbadiensis Watznaueria barnesae	Reticulofenestra bisecta	Lucianorhabdus cayeuxi Cretarhabdus conicus	Prediscosphaera cretacea	Arkhangelskiella cymbiformis	Reticulofenestra dictyoda (cf.)	Ciauxonuius uipiogrammus	Coccolithus eopelagicus	Cyclococcolithus formosus	Tetralithus gothicus	Chiasmolithus grandis Marthasterites inconsnictus	Ceratolithoides kamptneri	Cyclagelosphaera margereli	Lanternithus minutus	Sphenolithus moriformis	Tetralithus murus (cf.)	Chiasmolithus oamaruensis	Arkhangelskiella parca Poccolithus nelseicus	placoliths (indeterminate)	Tetralithus pyramidus	Sphenolithus radians	Isthmolithus recurvis	Kecticulorenestra scissura Cvlindralithus serratus	Chiasmolithus solitus	Zycodiscus spiralis	Cyclococcolithus sp.	Discolithina sp.	Rhabdosphaera sp. Fiffelithus turriseiffeli (cf.)	Reticulofenestra umbilica	Micrantholithus vesper	placolithus $(2-3\mu)$
Sample	Age	Zone	Micula	Tetrali Reinha	Discoa	Reticu	Lucian Cretarl	Predisc	Arkhai	Reticu	Criboe	Coccol	Cycloc	Tetrali	Chiasn	Cerato	Cyclag	Lanter	Sphen	Tetrali	Chiasn	Arkha	placoli	Tetrali	Sphen	Isthme	Cvlind	Chiasr	Zycod	Cycloo	Discol	Rhabd Fiffeli	Reticu	Micrar	placol
1-1(105) 1-2(77) 1-3(38) 1-4(65) 1-5(80) 1-6(53) 1-6(113) 1(CC)	Upper Upper Eocene- Lower Oligocene	NP19-22				A A A A A X X						x						R	x		R X	,	x			R X					x	x			
2-1(74) 2-2(54) 2-2(124) 2-3(62) 2-4(94) 2-5(103) 2-6(125) 2(CC) 3-2(89) 3-3(89) 3(CC)	Lower Upper Eocene	NP17-18	x			x x x x x x x x x x				x									x ²	¢	x	2	<				x			x			x	c c	
4-1(top) 4-2(75) 4-2(95)	Upper Middle Eocene	NP15-16			х	C C						Х	x		х						R R		1		x		с	>	¢					X X C R	C C
4-4(top) 4-4(72) 4(CC)[top]	Lower Eocene	NP12-14	R		x	A X						С	x		x				,	хx		,	K		x		С	,	ç				Х	C	С
4(CC)[bot] 5-0 5-0 5-1(70) 5-1(75) 5-1(85)	Middle Maestrichtian		X X A X A A	x	X A A A				x	R	x			C X C C	F	R F F	ł		5	X X 4 X 4 4		X R R R		Х			3	x	R R R						
5-3(55) 5(CC)	Lower Maestrichtian		x	x	x		хх	x	x	,	k 3	ĸ					x			x		x		x			3	x	х	ι.		3	x		

TABLE 2B Nannofossils, Leg 19

	Dereeter	i biostratigraphic Da		
	Site 192	Site 183	Washington ^a	California ^b
Location	53°N,164°E	52°N,161°W	48°N,124°W	37°N,120°W
Water Depth (m)	3024	4718	subaerial	subaerial
Deposition Type	pelagic	pelagic and dis- placed shelf	shelf	shelf
Age Control				
Foraminifera	yes	no	yes	yes
Radiolaria	no	no	no	partial
Diatoms	no	no	no	partial
Silicoflagellates	no	no	no	partial
Magnetic	no	yes ^c	yes,	yes
Isotopic	no	no	yesd	yes ^e
Nannofossil Species	12-15	23	~50	~100
Discoaster Spp.	1	4	10	21
Sphenolith Spp.	2	1	3	3
Pontosphaerid Spp.	2 1	2	3 7 3	15
Helicopontosphaerid Spp.	-	1(?)	3	3
Cool-Water Species				
Isthmolithus recuvis	yes	no	yes	no
Chiasmolithus oamaruensis	yes	yes	yes	(?)
Lanternithis minutus	yes	no	yes	(?)
Water Paleotemperature	cool-cold	cold	cold	warm
Paleodepth (m)	<3000	~3000	bathyl(?)	shallow bathyl(?)
Zone Range	NP12-14	NP12-13	NP14-16	NP12-15

TABLE 3 Selected Biostratigraphic Data

^aWorsley and Crecelius, 1972. ^bBramlette and Sullivan, 1961. ^cSite survey. ^dTabor, 1972. ^eEvernden and Evernden, 1970.

and deep ocean sediments and also because they are known to have the deepest compensation depth of all Paleogene nannofossils. A latitudinal trend similar to that of the discoasters is observed for sphenoliths, helicopontosphaerids, and pontosphaerids, all of which, like discoasters, are thermophilic. However, unlike discoasters, these three groups have relatively shallow compensation depths and prefer shelf conditions (Edwards, 1968; Bukry 1970). Therefore, proximity to shore and water depth, as well as paleotemperature, probably contribute to their observed distribution pattern in the North Pacific margin.

Further evidence for upper Eocene cool-water deposition lies in the occurrences of *Isthmolithus recurvis*, *Lanthernithus minutus*, and *Chiasmolithus oamaruensis* without other upper Eocene "guide fossils" (Edwards, 1968; Worsley and Jorgens, in press). The nonoccurrence of middle Eocene chert in the North Pacific sites as opposed to ubiquitous occurrences at lower latitudes suggests deposition in the barren central water mass of the North Pacific during Eocene time. This could also account for the puzzling absence of siliceous microfossils at these sites despite their widespread occurrence and abundance at this time.

Maestrichtian assemblages recovered from Meiji Guyot have a generally tropical aspect. Members of the *Tetrali*thus complex similar to the tropical *T. murus* occur, but *Nephrolithus*, a coeval genus having cooler water preference, is absent. The occurrence of the tropical Ceratolithoides kamptneri and the extinction of Lucianorhabdus cayeuxi below that of Cylandralithus serratus are supporting evidence for generally warm surface waters. However, the absence of siliceous microfossils suggests deposition north of the equatorial belt of high productivity.

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APPENDIX

TAXONOMIC NOTES

The generally impoverished nannofloral assemblages from the North Pacific and Bering Sea differ slightly from lower latitude ones on the subspecific level. However, because their areal and stratigraphic distribution is so imperfectly known at this time and because only light microscopy has been used in identifying them, no new taxa have been erected and these variants have been included within established species. The following paragraphs describe fifteen forms which either differ slightly from original descriptions or exhibit wider morphologic variation than previously noted. All other species recorded in the range charts conform closely to previous descriptions. Reference to illustrations and synonomy of the taxa mentioned here may be found by consulting Loeblich and Tappan (1966, 1968, 1969, 1970a, b).

Emiliania huxleyi

A large size range is noted for specimens assigned to this species, with many approaching 4μ . The large specimens appear to be high-latitude variants. A differentiation of ranges among the sizes proved unfeasible using light microscopy.

Coccolithus pelagicus

Included is *Ericsonia muiri* of the middle and lower Cenozoic. Specimens possessing a bridge across the central area seem restricted to the Pleistocene.

Gephyrocapsa oceanica

Occurs rarely, probably in part because solution differentially attacks the bridge, rendering specific identification by light microscopy difficult to impossible.

Reticulofenestra aff. dictyocha

Includes generalized placoliths of many sizes having swasticoidal interference crosses which break into infinity signs upon rotation of the stage. This catchall probably includes several distinct species which are not easily separated in the light microscope. This type of placolith is the dominant nannofossil of the Bering Sea Miocene.

Cruciplacolithus aff. neohelis

This form coincides fairly closely with McIntyre's description of *C. neohelis*, but may be a variant of the bridge-bearing form of *Coccolithus pelagicus*. The central area is infilled with calcite and appears to be a cross aligned with the axes of the elliptical placolith.

Gephyrocapsa ericsoni

Identification of this small (2-3 micron) species is not certain in the light microscope. Very small *Gephyrocapsa* specimens are referred to it.

Reticulofenestra cf. pseudoumbilica

These specimens resemble R. pseudoumbilica except for their small (4-5 micron) size. They occur as low as the upper Miocene.

"Cyclococcolithus leptoporus"

Refers to all species within the "Cyclococcolithus" complex having a weak polarization figure and occurring in Miocene and younger sediments.

Pseudoemiliania cf. lacunosa

Refers to small species (about 4-5 microns) resembling this distinctive species. It is difficult to differentiate *P*. cf. *lacunosa* from the larger specimens of *Emiliania huxleyi*.

Geophyrocapsa caribbeanica

Refers to specimens within the *Gephyrocapsa* complex having crossbars more nearly aligned with the major axis of the elliptical placolith. They are thought to be the cold water form of G. oceanica.

Tetralithus aff. murus

Includes concave-convex tetraliths strongly resembling T. murus, but not having ray-tips which curve sharply clockwise when the convex side faces the observer. It is not certain whether they are true specimens of T. murus whose ray-tip curvature has been selectively removed by solution or if they are the evolutionary forerunners of T. murus. In any event, such forms seem to range throughout most of the Maestrichtian in many parts of the ocean, rather than being restricted to its upper part.

Reticulofenestra cf. umbilica

Refers to large placoliths which, like R. *umbilica*, have very large open central areas with respect to their diameters, but have an interference figure different from that of R. *umbilica*. They may be partially dissolved species of R. *bisecta*.

R. bisecta

Includes specimens larger than about 8μ conforming to the definition of Roth (1970). *R. bisecta* ranges from the Eocene to the base of the Pliocene in the North Pacific and Bering Sea, but becomes extinct near the base of the Miocene in lower latitudes.

Smaller specimens somewhat resembling it are referred to R. dictyocha. Those smaller specimens exactly resembling it are referred to R. scissura = Ericsonia ovalis.

Tetralithus aff. aculeus

Refers to specimens intermediate between Ceratolithoides kamptneri and T. aculeus.

Tetralithus cf. murus

Includes quadrate forms resembling T. murus, but which appear not to be concave-convex.

"Cyclococcolithus neogammation"

Refers to all cyclococcoliths above the Oligocene having a strong polarization cross and small central pore.