The Shipboard Scientific Party1

SITE DATA

Date Occupied: 10 February 1973

Date Departed: 13 February 1973

Position: 74°32.29'S; 174°37.57'E

Water Depth: 491 corrected meters (echo sounding)

Water Depth (adopted): 495 meters (drill pipe from rig floor)

Total Penetration: 346.5 meters

Number of Cores: 38

Total Section Cored: 332.5 meters

Total Section Recovered: 83.4 meters

Percentage Core Recovery: 25%

Oldest Sediment Cored:

Depth below sea floor: 346.5 meters Lithology: Silty clay Age: Early Miocene

Principal Results: The stratigraphic sequence penetrated at Site 273 is similar to those at Sites 270-272, the other Ross Sea sites. The post-Miocene portion of the section here is extremely condensed or more likely missing in large part and was probably removed by the same glacial erosion event (~3-5 m.y.) that truncated dipping beds at Sites 270-272. Lithified silty clay constitutes the predominant lithology and, as elsewhere, it is mostly unstratified and contains marine microfossils. The environment of deposition for these strata is accordingly considered to be a seaway which received icebergs from the adjacent land areas and probably from shelf ice. Clasts in the Miocene clays differ from those in the eastern Ross Sea and those of Pliocene age in that diabase and basement rock types are present which are characteristic of the rocks exposed in the Transantarctic Mountains. The fact that the sequence dates back to at least the middle Miocene indicates that ice has been eroding the Transantarctic Mountains from mid-Miocene to Pliocene time. A pronounced reflecting horizon



lying at 0.3 sec subbottom appears to mark an earlier erosional or scoured surface and can be traced over many tens of miles and which must be pre middle Miocene in age. A sharp increase in the hardness and velocity of the silty clay was encountered at 276 meters subbottom and probably accounts for the seismic horizon. The significance of the geologic event giving rise to this erosional surface is not yet wholly known. Traces of methane and ethane were first detected at about 150 meters subbotton, but showed no systematic or significant increase downhole.

BACKGROUND

Site 273 is located in the west-central portion of the Ross Sea on the flank of an erosional valley which bounds the western edge of Pennell Bank (see Figure 1). The water depth at the site is 495 meters. The previous geophysical work of Houtz and Davey (1973) has been unsuccessful in determining the exact depth to basement at this site. However, extrapolation of nearby seismic sonobuoy data strongly indicates an excess of 2 km of low-velocity sediment in the vicinity of the site. Although the site was originally selected on the basis of an Eltanin 52 seismic profile, a higher resolution seismic profile was obtained from Glomar Challenger when it passed over the tentative site on a northwest-southeast traverse enroute to Site 270. This seismic profile is shown in Figure 2 along with the regional morphology. Gently dipping strata are truncated near the sea floor along the crest of the bank lying toward the northwest. Equivalent sediments appear to be missing in large part at Site 273, but similar flat-lying sediments appear to be truncated at the valley walls to the southeast of Site 273. A very pronounced reflecting horizon lies at a subbottom depth of about 0.3 sec (~250 m) and may represent a much older erosional valley nearly coincident with the

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Figure 1. Location of Site 273 and bathymetry. Contours in meters (corrected). Solid line shows location of Challenger seismic profile across the site shown in Figure 2.

present valley. A clear reversal in the dip of this horizon occurs near the central portion of the valley, and care was taken to ensure the site was not located near this structure.

The objectives at Site 273 were to sample down to 650 meters below the sea floor in order to contrast the history of sedimentation, glaciation, volcanism, and bio-stratigraphy, with that of the eastern Ross Sea.

OPERATIONS

Site 273 was approached on a heading of 002°. The beacon and a spar buoy were dropped while underway at 0508 on 10 February. After retrieving the towed geophysical gear, the ship returned to the buoy and positioned over beacon. Shallow water (491 m PDR corrected) and shifting winds and currents caused some difficulty in acquiring and maintaining a position directly over the beacon. Positioning in the automatic mode was not attained until 0857 hr.

While the drill string was being run and the hole spudded in, a sonobuoy record was obtained. Three clear reflecting horizons are found at 0.045, 0.16, and 0.30 sec (two-way travel time). These correlate well with horizons which are observed in the several profiler records and which can be followed for significant distances (Figure 2).

Hole 273 was spudded in at 1200 on 10 February and was cored continuously until 2030 hr when positioning problems caused by a failure of the ship's gyro compass required the bottom-hole assembly to be withdrawn and the hole abandoned. Total penetration was 76 meters. Nine cores were taken with an average recovery of 37% (see Table 1).

Following repairs to the ship's gyro compass, a second hole, 273A, was spudded in at 2300 hr. After delays of several hours, caused by unacceptably large (ca 100 ft) excursions of the ship off the beacon, Hole 273A was drilled to 80.5 meters subbottom and continuously cored (except for one drilled interval from 90-99.5 m) to a subbottom depth of 156.5 meters by early afternoon of 11 February (Table 1). Drilling and coring operations



Figure 2. Glomar Challenger acoustic reflection profile on approach to Site 273. Vertical scale is in seconds of two-way reflection time. Location of profile shown in Figure 1.

Core	Date (Feb. 1973)	Time	Depth From Drill Floor (m)	Depth Below Sea Floor (m)	Length Cored (m)	Length Recovered (m)	Recovery (%)
Hole 2	273						
1	10	1212	505.0-509.5	0.0-4.5	4.5	4.3	96
2	10	1304	509.5-519.0	4.5-14.0	9.5	5.1	54
3	10	1337	519.0-528.5	14.0-23.5	9.5	0.4	4
4	10	1412	528.5-538.0	23.5-33.0	9.5	2.2	23
5	10	1533	538.0-547.5	33.0-42.5	9.5	1.1	12
6	10	1703	547.5-557.0	42.5-52.0	9.5	2.0	21
7	10	1758	557.0-566.5	52.0-61.5	9.5	5.0	53
8	10	1937	566.5-576.0	61 5-71.0	9.5	5.0	53
9	10	2112	576.0-581.0	71.0-76.0	5.0	2.8	56
Subtot	al .	2112	0,010,00110	, 1.0 , 0.0	76.0	27.9	37
Hole 2	724				70.0	21.5	57
nole 2	/3A						
1	11	0635	585.5-595.0	80.5-90.0	9.5	1.2	12
2	11	0735	604.5-614.0	99.5-109.0	9.5	cc	-
3	11	0840	614.0-623.5	109.0-118.5	9.5	2.4	25
4	11	1021	623.5-633.0	118.5-128.0	9.5	2.0	21
5	11	1119	633.0-642.5	128.0-137.5	9.5	0.0	0
6	11	1212	642.5-652.0	137.5-147.0	9.5	2.1	22
7	11	1349	652.0-661.5	147.0-156.5	9.5	2.8	29
8	11	1826	661.5-671.0	156.5-166.0	9.5	6.4	67
9	11	1940	671.0-680.5	166.0-175.5	9.5	2.4	25
10	11	2115	680.5-683.5	175.5-178.5	3.0	2.8	93
11	12	0005	683.5-690.0	178.5-185.0	6.5	6.2	95
12	12	0110	690.0-699.5	185.0-194.5	9.5	0.6	6
13	12	0155	699.5-709.0	194.5-204.0	9.5	4.7	49
14	12	0335	709.0-718.5	204.0-213.5	9.5	2.4	25
15	12	0435	718.5-728.0	213.5-223.0	9.5	0.6	6
16	12	0520	728.0-737.5	223.0-232.5	9.5	1.7	18
17	12	0610	737.5-747.0	232.5-242.0	9.5	4.1	43
18	12	0700	747.0-756.5	242.0-251.5	9.5	0.4	4
19	12	0745	756.5-766.0	251.5-261.0	9.5	0.0	0
20	12	0910	766.0-775.5	261.0-270.5	9.5	0.0	0
21	12	1002	775.5-777.5	270.5-272.5	2.0	0.4	20
22	12	1125	777.5-785.0	272.5-280.0	7.5	4.0	53
23	12	1250	785.0-794.5	280.0-289.5	9.5	2.1	22
24	12	1421	794.5-804.0	289.5-299.0	9.5	0.6	6
25	12	1531	804.0-813.5	299.0-308.5	9.5	2.5	26
26	12	1648	813.5-823.0	308.5-318.0	9.5	0.6	6
27	12	1822	823.0-832.5	318.0-327.5	9.5	0.7	7
28	12	2007	832.5-842.0	327.5-337.0	9.5	cc	
29	13	0745	842.0-851.5	337.0-346.5	9.5	1.8	19
Subtot	al				256.5	55.5	22
Total					332.5	83.4	25

TABLE 1 Coring Summary, Site 273

were then suspended for about 3 hr to thaw out frozen air lines. After Cores 8, 9, and part of 10 had been cut, nearby small drifting icebergs forced another temporary suspension in operations. Because there was no assurance that the USCG icebreaker *Burton Island* could or would deflect these icebergs, the hole was filled with mud, and the drill string pulled until only 20 meters of the bottom-hole assembly remained buried with the power-sub close to the rig floor. Shortly after this precaution was taken, the *Burton Island* attempted to and succeeded in pushing a small approaching iceberg off its "collision course" with *Glomar Challenger*.

Drilling and coring resumed at about 2300 on 11 February, and the remainder of Hole 273A was continuously cored to a subbottom depth of 337 meters by about 2000 on 12 February. At that time the sand line parted just after retrieving Core 28, and sending the barrel down for Core 29. The overshot and about 5 meters of wire rope fell to the bottom of the string. Core 29 was cut and about 3 hr were spent fishing for the overshot with no success. Drilling and coring were terminated, the hole filled with 11.4 lb/gal drilling mud from 346 to 130 meters subbottom and a cement plug from 130 meters subbottom to the mud line. The mud line was cleared at 0400 on 13 February, the drill string, bottom-hole assembly, and Core 29 were recovered, and the ship got underway at 0814 hr.

A total of 29 cores was taken at Hole 273A penetrating 256.5 meters and recovering 55.5 meters for an average recovery of about 22%. Recovery on in-

dividual cores exceeded 50% only twice. Poorest recovery occurred between about 250 and 275 meters subbottom in what was apparently unconsolidated sand. Significant traces of methane and ethane were found in many of the cored sediments and were carefully monitored.

LITHOLOGY

At Site 273, 346.5 meters of a sedimentary sequence were cored in two holes, with a recovery of 83.4 meters. The sequence was divided into two units on the basis of the degree of lithification of the sediments penetrated, and each of these units was further divided into two subunits (Table 2). Texture of sediments from Site 273 is shown in Figure 3 and size distribution in Figure 4.

Unit 1

This unit encompasses sediments which are soft to stiff; further subdivision is made on the basis of color, biogenic components, and clast content.

Subunit 1A: This 80-cm thick unit consists of soft, dusky yellow-green diatom silty clay. Its boundary with the underlying Subunit 1B is sharp, being marked by an abrupt color change and a change in lithification from soft to stiff. This thin uppermost unit contains up to 40% diatom frustules, and a small amount of sand, but no larger clasts. In color, texture, and composition it resembles the thin diatom-rich sequences which occur in the top meter of Sites 270 and 272. On the basis of its contained diatoms and fine-grained character it appears to correlate with units of similar lithology at Sites 270 and 272.

Sedimentation appears to have occurred at a rate of approximately 0.5 m/m.y. through Subunit 1A.

Subunit 1B: This unit has a maximum thickness of 41.7 meters. It lies between the base of the diatom-clay unit above and the base of Core 5, below which there is a sharp increase in lithification. The unit consists of soft to stiff pebbly sand-silt-clay throughout. Colors are mainly dark greenish-grays, although locally there is a gradational darkening to olive-black. No bedding is discernible within the unit.

Generally poor sorting of sediments classified in this unit is shown in size frequency histograms, where grain distributions are spread fairly evenly over a wide range of size classes. Three samples plotted on the accompanying textural diagram show that there is a considerable uniformity in texture.

Diatoms are the chief biogenic component, but concentrations of frustules only locally exceed 10%. Micarb is present throughout, usually in concentrations of less than 1%. In Core 5, two beds of yellowish-green diatom silty clay, each about 20 cm thick, occur interbedded with gray silty clays. Paleontological evidence (from diatom studies) suggests that the age relationship of the diatom-rich beds to those above and below is anomalous; the possibility that these beds have been emplaced by slumping must therefore be considered. There is also some indication of grading within the beds; preliminary examination shows their bases to consist of spicule-rich silts, which pass upwards into diatom clays



Figure 3. Grain size analyses from lithologic Units 1 and 2 at Site 273. Numbers next to symbols refer to core and section.

	Lithologic Units, Site 273													
Unit	Lithology	Subbottom Depth (m)	Thickness (m)	Age										
1A	Soft, diatom silty clay	0.0-0.8	0.8											
1B	Soft to stiff, diatom- bearing pebbly sand-silt- clay, unbedded	0.8-42.5	41.7	Recent to Gauss or Younger										
2A	Semilithified, pabbly silty clay, diatom bearing sparsely bedded	42.5-166.0	143.5	Lower most upper Miocene to mid- Miocene										
2B	Semilithified pebbly silty clay, some diatom bearing, unbedded	166.0-346.5	180.5+	Mid to ? early Miocene										

	TA	BLE	2	
thala		Thite	Cito	272



Figure 4. Sediment size distribution at Site 273.

at the top. This suggests a sedimentary, rather than a deformation origin.

Granules and larger clasts are present throughout the subunit. Clast density averages less than 1% throughout the interval, but rises to an estimated maximum of 3%-5% in Sections 3 and 4 of Core 2. Clasts with a maximum diameter greater than 10 mm occur in densities of 2 to 8 per meter of core, without any systematic changes in concentration throughout. Metamorphic and igneous rock types dominate the clasts, with granites, gneisses, basalts and reddish-gray quartzites being the most common lithologies. Diabases are rare in the unit. The sedimentary rock suite includes lithified poorly sorted sandy silty claystone, dark gray argillites, and rare sandstones. Subsequent study of the pebbles from the working half of the core shows that schist is also an important lithology.

Sedimentation rates have been estimated at approximately 31 m/m.y. for Subunit 1B.

Unit 2

The top of Unit 2 marks a change from stiff to semilithified silty clays, a change which probably coincides with the upper of two prominent reflectors in the seismic profile for the site. Lithologically, the unit is relatively uniform throughout, and the subdivision into two subunits has been made, somewhat tentatively, on the presence of poorly developed bedding in the upper part.

The whole unit consists of predominantly greenishgray silty claystones which are mainly semilithified, but which contain thin (avg 20 cm thick) lithified intervals. Granules and pebbles occur scattered throughout, and sand is a consistent component, occurring in quantities ranging up to 18%.

The sediments of Unit 2 are consistently slightly finer and more clay-rich than those of Unit 1. This situation parallels that at Site 270, where early Miocene and late Oligocene sediments, i.e., those below a suspected regional unconformity, are finer than those above, and fall into a silty claystone category (Barrett, this volume). At present age control is not by itself sufficient to establish that the lithification break at Site 273 represents the same unconformity as at Sites 270-272.

Diatoms are present throughout the unit, most commonly occurring in proportions estimated to be about 3%-4%, but reaching an estimated 35% in silty clays of Core 6A. A second zone of above-average diatom concentration (6%-8%) occurs through Cores 17-22A. Carbonate is present in trace amounts through most of the sequence, and three (thin) intervals, in Cores 7, 7A, and 24A are calcareously cemented through 20-cm-thick zones; in the latter, CaCo₃ content reaches 31%. Small broken shell fragments are very rare and have been recorded only from Cores 4A and 17A.

The distribution of pebbles and granules throughout Unit 2 does not show any readily discernible pattern. Clasts more than 4 mm maximum diameter range in estimated abundance from less than 1%, to a maximum of 2%-3% in Core 22A. Similarly, the density of larger clasts, i.e., those above 10 mm maximum diameter, is relatively constant throughout. These number 2 per meter on the average throughout Subunit 2A, and slightly under 3 per meter in 2B. In Core 22A the concentrations reach 12 per meter. Clasts above 4 mm are absent from the diatom silty clays of Core 6, but are present in most of the bedded intervals (except in the top of Core 9).

The most common clast lithologies include quartzdiorites, diabases, quartzites, and granites; clasts of sedimentary origin are fewer, but include argillites, rare silty limestones, poorly sorted sandstones, and a single clast of glauconitic sandstone.

Bedding occurs in thin intervals throughout Subunit 2A; no bedded interval exceeds 30 cm in thickness, however, and the sequences are widely separated by massive, uniform sections. Most commonly, bedding is indicated only by slight color differences, as lighter greenish-gray lenticular areas discernible within a darker gray sequence. Suggestions of grading occur in one 20-cm-thick bedded interval in Core 8, Section 4; here, darker gray laminae 2 mm thick are separated by interbeds 2 cm thick, which appear in some to be distinctly coarser near their bases. This feature is not consistently present, however, and cannot be seen in each interbed. Neither was it clear whether the coarse areas of the interbeds extended across the width of the core or whether the distribution of coarse material was in essentially random patches with gradational margins.

Histograms of grain-size frequency distributions from both the stratified and nonstratified intervals in Subunit 2A show that sorting within the subunit is generally poor. The analyses are too few to demonstrate any relationship between sorting and stratification, although one of the stratified intervals (4A-2, 34) shows a welldeveloped mode in the fine silt class.

In Cores 1A, 6A, and 17A of Unit 2, there are also structures which resemble bedding, but which are almost certainly due to drilling deformation. These consist of 2-3 mm thick, dark, clayey laminae 2-4 cm apart, which are down-bent at the core margins. These structures are confined to localized sections of sediment which have a stiff, rather than a semilithified consistency.

Bedding surfaces in Cores 9, 4A, and 9A show dips of 5°-10°. In Core 7, Section 2 there is clear evidence of bedding disturbance around a dropped pebble; bedding surfaces are distorted downwards below the clast, and those adjacent to it are sharply truncated against the clast—probably through loading deformation. The same core section, 50 cm below, displays possible load casts with attendant soft-sediment clasts identifiable by their color differences.

Bedding in Subunit 2B is extremely rare. There is, however, an interval of fine sand which was recovered in Core 21A, below a 2-core interval of fine sand which was recovered in Core 21A, below a 2-core interval of no recovery. The sand overlies a 30-cm-thick intervalof lithified bedded claystone in Core 22, which again, shows vague soft-sediment deformation. The position of the sand bed coincides roughly with the lower of two pronounced reflectors seen in the seismic profiles from this site. Below Core 22, Subunit 2B is generally uniform and unbedded, although there are intervals 10-20 cm thick which show pronounced fissility, a feature which reflects differing degrees of competence and may or may not be related to the development of bedding planes.

Sedimentation rates average approximately 28 m/m.y. throughout Unit 2.

Interpretation

Deposition of terrigenous sediment is interpreted as having occurred within a marine environment at this site since probably early Miocene times, although not necessarily continuously. The marine nature of the environment is inferred from the constant presence of marine diatoms in generally low, but fluctuating, abundances throughout. Marine foraminifera are also present throughout (see Biostratigraphic Summary), which further supports such an interpretation.

Coarse debris, consisting of sand, granules, pebbles, and cobbles, is present throughout the sequence, and was presumably introduced by ice rafting. The general uniformity and lack of sedimentary structures through most of Subunit 2B suggests little interruption to a process involving a steady rain of terrigenous material of diverse size and composition. There is little evidence of bottom current activity strong enough to affect significant sorting of the deposits. The sand bed in Core 21A is an exception and might have resulted from winnowing out of fine material, possibly on a local (or even regional) elevated area. The sparse bedding in Subunit 2A probably results from the sporadic generation of bottom currents, possibly associated with ice-shelf freezing and increased bottom-water production. The limited soft-sediment deformation observed has probably resulted from local slumping. The dips observed in some beds may be primary, but are more likely to result from small-scale slumps.

Subunit 1B shows no evidence of current activity. Examination of bottom topography and of seismic profiles for the site suggests that this unit was probably slumped in from high areas to the southeast. Paleontological evidence is in accord with such a suggestion; throughout Unit 1 Recent diatoms are mixed with those of late to middle Miocene age. The diatom-rich beds at the base of the unit, which show evidence of grading, may have been deposited from turbidity currents associated with slumping of local sediment piles.

Subunit 1A, the soft, apparently pebble-free thin diatomaceous unit at the top seems to indicate a lessening of ice rafting, with a correspondingly high contribution of pelagic detritus, although the regional extent and significance of this thin bed cannot be evaluated at present.

Preliminary examination of clasts suggests that most have been derived from the west, although precise matching of lithologies with those known to crop out in adjacent areas of the Antarctic continent has not yet been affected. One feature which is noteworthy is the abundance of diabase clasts in Unit 2 and their rarity in Unit 1, indicating a marked change in the sediment supply. The diabases probably represent those of the Ferrar Group of the Transantarctic Mountains, and if this is so, their rare occurrence in the younger units could be attributed either to erosion into basement by outlet glaciers, or to the termination of East Antarctic ice coming directly through the mountains near Site 273. Either event may be associated with increasing elevation of the Transantarctic Mountains.

PHYSICAL PROPERTIES

Sonic-velocity and wet-bulk density measurements were made on nearly all cores from this site. Representative data are plotted in Figure 5 adjacent to the lithologic description. The bulk density values have been determined from the analog GRAPE records assuming a linear interpolation between the "in liner" calibration traces for 2.6-in. diameter distilled water ($\rho = 1.00$ g/cc) and 2.6-in. diameter aluminum ($\rho = 2.60$ g/cc). No corrections for varying diameters have been applied to the bulk density data, so the values plotted should be considered minimal. Sonic-velocity measurements for Cores 1-5 were made on split and unsplit sections. The measurements for Cores 6-9 and 1A-27A were made on chunks of stiff, semilithified, and lithified sediment.

The observed downhole variation in bulk density is minimal with no clear overall trend. Densities generally lie between 1.70 and 1.90 g/cc (nominal). Notable exceptions are found in Cores 5, 7A, and 22A and are related to occurrences of cobbles (Core 5), a carbonate lithified zone (Core 7A), and a zone of clayey silt with a high pebble concentration (Core 22A).

The velocity data suggest a slight general increase with depth of about 0.35 km/sec in the 350 meters of penetration. Although this gradient fits well with the gradient of 0.90 sec-1 determined by Houtz and Davey (1973) from sonobuoy data, it corresponds to only the uppermost 15% of the depth range fro which their data apply. This general trend is guite subdued in relation to the rather large variability found in seismic velocities from core to core and within individual sections of cores. Anisotropy in sonic velocities at this site is also quite variable and appears to show some changes with depth. In comparison with previous sites, the velocity anisotropy observed in Hole 273A is unique in that for more than half of the measurement pairs, velocities parallel to the core axis (normal to the bedding) are equal to or greater than those normal to the core axis (parallel to the bedding).

Two strong subbottom reflectors at 0.045 and 0.30 sec (two-way travel time) are observed on the seismic records. The shallowest of these clearly corresponds with the lithologic boundary between Units 1B and 2A at about 40 meters (bottom of Core 5 and/or top of Core 6). The deeper reflector must correspond closely to the base of the sand layer in Core 21A and the substantial density and velocity changes found in Core 22A. The two-way travel time calculated from the velocity measurements on the cored sediments is almost exactly 0.30 sec at the base of Core 21A. A third, rather weak reflector, is found at about 0.16 sec of two-way travel time and appears to correlate clearly with the significant increase in velocity found in Core 7A at about 150 meters subbottom. The calculated two-way travel time to the top of Core 7A is 0.165 sec.

BIOSTRATIGRAPHIC SUMMARY

Site 273 contains for aminifera and diatoms in good abundance with trace recordings of Radiolaria, but no nannofossils.

The stratigraphic sequence at this site will be divided into two paleontologic subdivisions. Each unit will be discussed separately with summary remarks to follow.

Cores 1-5

Foraminifera: Core catchers contain an abundant and well-preserved foram fauna consisting of planktonic and benthonic species; an age of Pleistocene/Pliocene is indicated.

Radiolaria: An age of (Brunhes?) Pleistocene is indicated by the radiolarian fauna in Core 1.

Diatoms: The diatom flora for this interval contains species previously recorded only from upper Miocene sediments with others of unknown stratigraphic range. A Recent to upper Miocene age is assigned.

Sample 5, CC through Core 29A

Foraminifera: A fauna is recorded consisting of 10-21 species contrasting markedly with those above in that they are small, thin walled, deformed, and are stained brown. The five dominant taxa show a close affinity to various Arctic faunas and also to some Ross Sea faunas. A determination of the depth of deposition is difficult based on the assemblage, but would indicate water depths of 100-300 meters with some exceptions, but probably none greater than 500 meters. The stratigraphic ranges for most of these taxa extend through this section with the exception of *Trochoelphidiella onyxi*, which occurs in the late Pliocene of the McMurdo Sound area.

Radiolaria: Only fragments were recorded.

Diatoms: Diatoms contained in this interval indicate a Miocene age with occurrences of some species which may prove useful in subdividing the sedimentary sequence. These possible subdivisions are based on land sections in California (Temblor), Maryland (Calvert), and Indonesia (Wonosari). Their abundances are indicative of the age of the formation. An exception is diatom 28, as yet a problematical taxon, which occurs in the mid-Miocene sediment of Site 269.

On the basis of diatoms, Cores 6-5A appear to be middle to upper Miocene; Cores 6A-18A are mid Miocene, and the cores below are (lower?) Miocene.

Concluding Remarks

An obvious discrepancy exists between the assigned ages based on foraminifera and diatoms. A portion of this problem derives from the inadequacy at the present time to discriminate Miocene from younger sediments. There also exists a distinct possibility that Antarctic diatoms cannot be directly correlated to distant land sections.

When considering the age assignments shown by foraminifera, it is important to recognize that a single species was used as the basis for the determination. It is hoped that further study will resolve this problem.

Foraminifera

Late Pliocene-Quaternary (1-4, CC)

Abundant and well-preserved foraminifera were extracted from the top four core catchers. The following taxa were noted in Sample 4, CC: Textularia sp., Nodosaria sp., Trifarina earlandi, Ehrenbergina glabra, Cassidulina neocarinata, Islandiella porrecta, Globocassidulina biora (dominant taxon), Cibicides sp., Globigerina pachyderma, and G. megastoma. Ostracodes, sponge spicules, and Radiolaria commonly accompany this fauna.

The size and preservation of foraminiferal tests in this interval contrasts markedly with that of tests in the underlying ?Miocene-Pliocene. The latter are generally small and thin walled and sometimes deformed and are often light brown to dark brown, whereas those of 1-4, CC are three to five times larger, are thick shelled, and have a bright white color with a high luster. The fauna in 1-4, CC has a late Pliocene-Quaternary character. If Globocassidulina biora (Crespin) is confined to Gauss sediments, as proposed by Fillon, these sediments may be late Pliocene rather than Quaternary. It is noteworthy that the genus Ehrenbergina is present in these very young sediments, but apparently not in the Oligocene, Miocene, and Pliocene sediments at Sites 270, 272, and 273. This is unexplained, for the genus is well represented in the New Zealand middle-upper Tertiary.

?Miocene-Pliocene (5, CC-9, CC; 1A-29A)

Ten samples from this thick and monotonous secession were examined, four from Samples 5, CC; 6, CC; 7, CC; and 8, CC and six from the lowermost 100 meters (17, CC; 18, CC; 22, CC; 25, CC; 28, CC; 29, CC). Foraminifera occur abundantly in most samples and preservation is generally good. Diversities are consistently in the range of 10 to 21 species.

The following taxa are present in the 10 samples listed above: Rhizamminidae, Lagena spp., Fissurina spp., Pseudonodosaria sp., Quinqueloculina sp., Bolivina sp., Trifarina sp., Uvigerina sp., Astrononion sp., Elphidium sp., Trochoelphidiella sp., Nonionella sp., Epistominella sp., Globocassidulina sp., Sphaeroidina sp., Pullenia spp., Rosalina sp., Buccella sp., Patellina sp., Eponides sp., Melonis sp., Cibicides sp., and Anomalinoides sp. In addition, a probable planktonic foraminifer of as yet undetermined genus and species was seen in Cores 17 and 18. It most closely resembles Candeina zeocenica, though this species is restricted to the Eocene of New Zealand, while at the present site it occurs in sediments otherwise dated as Miocene. It also closely resembles a form described by Krasheninnikov and Hoskins (1973).

Dominant taxa are *Elphidium*, *Trochoelphidiella*, *Uvigerina*, *Trifarina*, *Astrononion*, and *Globocassidulina*. The faunas show a close affinity to late Miocene/Pliocene faunas of various Arctic regions and also to some Ross Sea faunas. An interpretation of the depth of deposition is difficult. The abundance of members of the Elphidiidae suggests water depths in the region of 100 to 300 meters although the high count for *Uvigerina* and *Trifarina* points to slightly deeper water, but probably no greater than 500 meters. Most taxa range through the interval 5-29, CC and no useful biostratigraphic subdivisions appear possible from this initial work. These sediments are also difficult to date with any accuracy. A range of middle Miocene to Pliocene is proposed with the following possible refinement. The taxon *Trochoelphidiella onyxi* Webb occurs in the Pliocene of the McMurdo Sound area, and the occurrence of a closely comparable suite in several samples at Site 273 suggests that the 5-29, CC interval might have a Pliocene rather than Miocene age.

Nannofossils

Very rare coccoliths were found in sedimentological smear slides from Cores 6-1, 12-1, and 12, CC in Hole 273A, but the low numbers and poor preservation made identification and age assignment impossible.

Radiolaria

Sediments similar to the sediments at Sites 270 to 272 contain no or only traces of unidentifiable Radiolaria. However, some identifiable Radiolaria are present at the following horizons:

Core 1: common and well-preserved Radiolaria (Antarctissa denticulata, A. strelkovi, Theocalyptra bicornis, and Lithelius nautiloides): the Antarctissa denticulata Zone.

Sample 3-1, 120-124 cm: rare and well-preserved Radiolaria (Antarctissa denticulata, A. strelkovi, Theocalyptra bicornis, and Lithelius nautiloides); the Antarctissa denticulata Zone.

Sample 5-1, 78-81 cm: rare and well-preserved Radiolaria (*Antarctissa ewingi, Lithelius nautiloides*, and *Antarctissa denticulata*); Pliocene?

Diatoms

Diatom abundance and preservation is fair to good at this site. Cores 1 through 5 contain a mixture of diatoms comprising the *Coscinodiscus lentigenosus* Zone with the *Denticula antarctica* Zone. Core 6 through 21A-1-145 contains a portion of the *Denticula antarctica* Zone. Below this point it is unzoned, but shows a characteristic similar to Site 272 in that *Coscinodiscus apiculatus* is present.

Silicoflagellates

Sediments at this site generally contain low abundances and low diversity assemblages of silicoflagellates. The silicoflagellates from this site are well preserved and usually highly ornamented. A high degree of morphological variability makes a precise taxonomic placement of Site 273 silicoflagellates difficult. The taxonomic assignments made herein are therefore considered *sensu strictu*.

Core 1 contains many skeletal morphotypes of *Distephanus speculum* which have varying apical ring widths and spine lengths. No age assignment for the sediment of Core 1 is possible.

Samples 2, CC, 3, CC, and 6, CC are barren of silicoflagellates. Core 5 contains rare specimens of *Distephanus speculum s.l.* Silicoflagellates are fairly common in Sample 6A-1, 120-122 cm which contains the following silicoflagellate assemblage, suggesting a Miocene age: *Distephanus quinquangellus*, D. cf. crux, D. cf. *boliviensis*, *D. speculum*, and *D. cf. longispinus*. This assemblage also contains a variety of intraspecific morphotypes and a number of aberrant forms with incomplete skeletons.

Core 8A through Sample 28A, CC contain low abundances of *Distephanus speculum s.l.* No age assignment is possible for this interval of sediment.

SUMMARY AND CONCLUSIONS

The sedimentary sequence of the western Ross Shelf was sampled to a depth of 346.5 meters in two holes at Site 273. Marine sediments containing abundant clasts derived from icebergs characterize the sequence, which ranges in age from Quaternary to possibly early Miocene. Traces of methane and ethane occur from about 150 meters subbottom to the bottom of the hole.

Two lithologic units, each divided into two subunits, are distinguished at Site 273. Unit 1 is about 42 meters thick and at the top consists of 0.8 meters of distinctive yellow-green, diatom silty clay (Subunit 1A), closely similar to topmost sediments at Sites 270 and 272. The age of Unit 1 is uncertain, probably being in the range Recent to middle-Pliocene (Gauss). The presence of reworked mid-Miocene fossils makes definitive correlation with other sites difficult. Subunit 1B consists of soft to stiff pebbly silty clay with 20-cm-thick beds of yellowish-green diatom silty clay near the base. The latter show some indication of size grading. Clasts (pebbles and granules) include quartz diorites, diabases, granites, basalts, quartzites, argillites, and sandstones.

Unit 2 (324 m thick) marks a change to semilithified and some lithified silty clays with sand (up to 15%) and clasts. This change is apparent on the seismic profiles in the form of a prominent reflector at 0.05 sec. The geometry of the sequence above this reflector as well as the presence of reworked material indicate a sizable slump from the elevated area to the southeast. Another reflector, at 0.16 sec, apparently corresponds to an irregularly lithified sequence showing calcareous cementation in Core 7A.

Stratification is present in thin intervals throughout Subunit 2A (42.5-166 m subbottom) and includes color banding and possible graded bedding. Subunit 2B is generally unstratified, but contains a sand interval (Core 21A, 270.5-272.5 m) which overlies lithified bedded claystone (Core 22). The base of the sand corresponds approximately to a prominent seismic reflector at ~0.3 sec, which appears to truncate older beds further south. Clasts in Unit 2 include a plutonic type (approximating quartz diorite), diabases, quartzites, granites, argillites, silty limestones, and sandstones. The age of Unit 2 is not certainly established at present. Foraminifera suggest a maximum age of early Pliocene for the oldest sediments at Site 273. However, they have been assigned a mid to ?early Miocene age on the basis of a more extensive diatom flora.

Given the age uncertainties and probably slumping at Site 273, it is difficult to compare the stratigraphic section for the western Ross Sea with the eastern part, except that glacial marine sedimentation occurred in both areas in the Neogene. An advance of shelf ice very probably beveled sediments here, as well as at Sites 270, 271, and 272. The abrupt increase in lithification between Units 1 and 2 is similar to that observed at Sites 270 and 272. The regional configuration of the 0.3 sec and a deeper reflector suggests that the topographic valley in which Site 273 is located is quite an ancient feature, perhaps the locus along which earlier ice tongues or thick shelf ice advanced. One or more of these advances may have caused deep erosion of relatively hard sediment. It is probable that the valley floor near Site 273 was eroded in the interval 3-5 m.y. ago, as in the eastern Ross Sea.

The largely nonstratified lithofacies at this site can be attributed to deposition at most times when bottom circulation was restricted. The stratified lithofacies in Subunit 2A indicates active bottom currents and some slumping, but the apparent absence of burrowing suggests that bottom waters were not highly oxygenated.

Preliminary examination of clasts suggests that most have been derived from the west. The abundance of diabase clasts in Unit 2, and their relative absence from Unit 1, indicate a marked change in the sediment supply. It is probable that the diabases represent those of the Ferrar Group of the Transantarctic Mountains. The absence of these rock types from higher units within the glacial sequence could be associated with increasing elevation of the Transantarctic Mountains leading to erosion into basement by outlet glaciers, or to the damming of East Antarctic ice by the mountains.

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BIOSTRA	TIGRAPHY		AGE	(m) H	HOL	E	COLUMN	LITHOLOGIC	ACOUST. VEL.(kms-1) BULK GRAPE DENSITY SYRINGE
	RADS	DIATOMS	Huc	DEPT	273A	273	COLONIA	DESCRIPTION	POROSITYA 1.4 1.6 1.8 2.0
	POST- MIOCENE	MIXED 1 & 12	PLEISTOCENE	0	1 - 2 - 3 - 4 -			Soft, yellowish green <u>DIATOM SILTY CLAY.</u> Soft to stiff, dark greenish-gray, pebbly SILTY CLAY, some diatom- bearing.	60 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0
	FLIGENE			50	6 - - 7 - 8 - 9 -			No bedding discernible.	5,4-6.5 o cored o cored clasts 2.85 2.61 0 0 0
				100	1 - 2 - 3 - 2 - 3 - 2 - 2 			some diatom-bearing. Granules and larger clasts common. Bedding present in thin widely spaced intervals.	o 5.850 cored clast
			MIDDLE MIOCENE	150					
	BARREN		18	200	10 22 12 - 13 22 14 22 15 - 16 22			ified dark greenish-gray to olive black, pebbly SILTY CLAY and CLAY- STONE, some diatom- bearing. Granules and larger clasts common. Sand bed.	
				250	17 <u>7</u> 18 -19 20		•		U
		C. apiculatus	LOWER? MIOCENE	300	21 22 23 24 -25 26 27 28 28				0 2.4-4.00 0 0 0

Figure 5. Graphic hole summaries, Site 273.

	2.70	F CH	OSSI	L TER	2			NOI	BLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SAM	LITHOLOGIC DESCRIPTION
					1	0.5	VOID	Т	GZ	59% clay Soft, dusky yellow green (56Y 5/2) 40% silt 55% DIATOM SILTY CLAY, sharp 1% sand color boundary at 82 cms. 40% diatoms
CUL						1.0	9 9 		cc	70% clay Soft, Slightly stiffer than above, dark greenish gray (SGY 4/1), 25% silt 76% 5% sand ULXTOW-BERNIMS SILTY CLAY 10% diatoms visible.
QUALERRART/TLAU					2	multanda			*	60% clay Density of pebbles ±1% throughout, max. size pebbles 58 mm, av. 33% silt 95% 7% sand 3% diatoms size (above 4 mm) = 15 mm.
A REAL PROPERTY AND A REAL					3	nutrantun.	D M		•	Soft, dark greenish-gray DIATOM- BEARING SILTY CLAY, with 26% silt 94% pebbles and granules. No 4% sand bedding visible. 4% diatoms 2% micarb
?					Ca	ore tcher			*	70% clay Soft, olive gray SILTY CLAY; some 25% silt 97% 5% sand pebbles and granules. 8mm 2% diatoms ?basalt. 1% micarb 1

Site	273	HOI	e		Lo	re 2	Cored In	terv	alt	4.5-14 m
Π		F CH/	OSSI	TER	N	5		NOL	(PLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITH0.SAM	LITHOLOGIC DESCRIPTION
					1	0.5	VOID			
						1.0	<u> </u>			60% Clay Soft to semi-stiff, dark 39% silt greenish-gray (SSY 4/1) 1% sand SILTY CLAY with pebbles. 2% diatoms No bedding visible.
IOCENE					2	other data	е — D — М —		•	Density of pebbles 1-2%, max. pebble diam. 40 mm, Av. pebble size (above 4 mm) = 1 mm. 37% silt 95% 3% sand 2% diatoms
QUATERNARY/PL					3	in the function of the second se	р_м_			55% clay 40% silt 97% 5% sclay 5% score 1% diatoms 1% micarb 5% score 5% score
					4	d readered	-D- M-		XM	66% clay Soft to semi-stiff, DIATOM- 30% silt 94% BEARING SILTY CLAY. No 4% sand bedding visible. 3% diatoms 2% micarb
					C Ca	ore	б		*	Bulk X-ray (9.9 m): Amorph -51.6% Ident - 48.4% Quar. - 38.2% K-Fe. - 16.0% Plag. - 20.6% Mica - 23.2% Chio. - 2.0%

Site	273	Hol	00001		Co	re 3	Cored In	terv	al:1	-23.5 m
		CHA	ARAC	TER	N	s		NOI	4PLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO.SA	LITHOLOGIC DESCRIPTION
NE					1	0.5	VOID			Soft, dark greenish-gray SILTY CLAY,
AKY/PLIOCE						1.0		T	*	60% clay Density of pebbles <1% max.
QUATERN					Car Car	ore tcher	M0		•	67% clay Soft, dark greenish-gray DIATOM 30% silt 95% BEARING SILTY CLAY with 3% sand pebbles, inc. 39 mm dark 3% diatoms gray ultramafic, with 2% micarb sulbrides.



2% diatoms 1% micarb

ite	273	Ho1	e		Co	re 6	Cored In	terv	a1:42	2.5-52 m
		F CH/	OSS1 ARAC	TER	N	s		NOI	APLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTION	METER	LITHOLOGY	DEFORMA'	LITHO.SA	LITHOLOGIC DESCRIPTION
					1	1.0	V01D		* XM CC GZ *	01ive black, stiff to semi- lithified SLLTY CLAY with pebbles. No bedding visible. 37% silt 98% 3% sand Pebble density <1% throughout, max. pebble diam. 36 mm, av. pebble diam. 15 mm. 5% clay Semi-lithified olive-black SILTY CLAY. 5% silt 97% 2% diatoms Semi-lithified olive-black SILTY CLAY. 2% diatoms 1% micarb
					C Cat	ore tcher	9 D			65% clay Dark gray (N3), stiff, slightly 32% silt 96% 3% sand calcareous SILTY CLAY. 2% diatoms Bulk X-ray (44.5 m): 1% micarb Bulk X-ray (44.5 m): Amorph. - 56.8% Ident. - 43.2% Quar. - 41.1% K-Fe. - 10.9% Plag. - 21.6% Mica - 19.1% Chlo. - 2.3% Pyri. - 1.1%

Explanatory notes in Chapter 1

Site 27	3	Ho1	е	(Core	7	Cored	Inte	rva	1: 5	2-61.5 m	Sit	te a	273	Hole	9	C	ore 8	Cored I	iter	/al:6	51.5-71 m	
AGE	ZONE	FOSSIL 2	OSSIL RACTE ONNBY	CECTION	action	METERS	LITHOLOG	şγ	DEFORMALION	LITHO. SAMPLE	LITHOLOGIC DESCRIPTION	AGE		ZONE	FOSSIL R	RACTER .	SECTION	METERS	LITHOLOGY	DEFORMATION	LITH0.SAMPLE	LITHOLOGIC	DESCRIPTION
MIDDLE/UPPER MIDGENE				1 2 3 4	0 1 2 2		VOID			* * *	60% clay Semi-lithified, olive black 37% silt 96% 37% silt 96% 38% sand Greenish-gray (56° 4/1) 1- 2% micarb 30 cms. 60% clay Pebble density <1%, max. pebble 40% silt 64% 35% micarb Giam. 3 mm, av. pebble 1% diatem 60% 60% clay 97% 37% silt 97% 37% silt 97% 37% silt 97% 37% silt 95% 65% clay 0live-black, semi-lithified 38 sand Silty CLAY. Very faint 95% bedding - 2 cm thick 38 silt 95% 38 diatoms differences. 18 micarb Max. pebble density <1%, max. pebble diam. 17 mm. 65% clay Dark greenish-gray, semi-lithified silty cal-careous SiltY CLAY. No 33% silt 93% 34% diatoms Dark greenish-gray (56Y 4/1), semi-lithified Silty CLA' 33% silt 93% 24% diatoms Dark greenish-gray (56Y 4/1), semi-lithified Silty CLA' 33% silt 93% 55% clay	AIJOOLEX NIOCENE					1 2 3 4	0.5	V010		* * & GZ *	60% clay 95% 3% diatoms 1-2% micarb 45% clay 7% 15% sand 30% 30% clay 95% 3% silt 7% 75% clay 95% 3% sand 95% 20% clay 95% 25% silt 97% 25% silt 99% 36% silt 99% 36% silt 99% 36% silt 96% 78 sand 3% diatoms 30% clay 96% 30% clay 96% 31% micarb 96%	 Olive black, semi-lithified SILTY CLAY, very faintly laminated above 103 cm., massive below. Pebbles <1%; rare granules along irregular erosion surface a 103 cm. Pebbles 5-7%, max. diam. 7 mm, av. diam. = 5 m Dark greenish-gray to olive black semi-lithified DIATOM BEARING SLTY CLAY. Mainly massive, but includes some zones of fissility. up to 20 cm thick. Pebbles and granules <1%, max. diam. 60 mm, av. pebble diam. = 14 mm. Dark gray (M3) to dark greenish- gray SILTY CLAY with pebbles Bedding developed below 125 cm, 2 mm thick darker gray, sharp-based fine laminae separated by 2 cm thick silt beds; suggestion of grading. Pebbles absent in bedded interval Olive black, passing downwards into dark greenish gray, also hair-like catice vering availed to bede

		F CH/	OSSI	IL TER	N	s		ION	PLE	
AGE	ZONE	FOSSIL	ABUND.	PRES.	SECTIO	METER	LITHOLOGY	DEFORMAT	LITHO. SAM	LITHOLOGIC DESCRIPTION
JPPER MIULENE					1	0.5	V01D		*	50% clay Bedding distinct in top 30 cm. 40% silt 80% Dark gray, fine beds 2 cm TR sand thick, alternate with 15% opaques coarser gray(sh-green 5% diatoms intervals. Beds dip 8-10° 60% clay to core axis. 38% silt 98 2% sand bedded interval.
UTUNIE/					2 Cat	ore	-0-			74% clay Massive, semi-lithified dark greenish gray (567 4/1) 25% silt 96% Siltr CLAY. 1% sand Pebbles <1%, max. pebble size 17 mm, av. pebble diam. = 58% clay 7 mm. Granules 1-2%. 40% silt Below 125 cm dark gray laminae present. 1-2 mm thick, 2-6 cm apart, downbent at core margin.

SITE 273

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Site	273	Hole	ole A Core 7 Cored Interval: 148-157.5 m									511	te a	273	Hole		Cor	e 8	Cored In	iterv	a1:1	56.5-166 m	
AGE	ZONE	FOSSIL P	ABUND. ABUND.	SECTION	METERS	LITHOL	OGY	DEFORMATION	LITHO.SAMPLE	LITHOLOGIC	DESCRIPTION	AGE	-	ZONE	FOSSIC CHARA VIND	DRES.	SECTION	METERS	LITHOLOGY	DEFORMATION	L1THO. SAMPLE	LITHOLOGIC	DESCRIPTION
MIDDLE MIDCENE				1 2 Cat	0.5				* (M (M *	53% clay 45% silt 90% 2% sand 8% diatoms 1% micarb 55% clay 44% silt 95% 1% sand 4% diatoms 1% micarb 58% clay 40% silt 95% 1% sand 4% diatoms 78% clay 40% silt 96% 2% sand 3% diatoms TR micarb Bulk X-ray (149.0 m): Amorph 53.7% 1dent 46.3% Caic 4.4% Doilo 30.3% Quar 28.2% K-Fe 10.5% Plag 18.8% Mica - 6.9% Chio 0.9%	 Medium-dark gray (N3) to dark greenish-gray DIATOM-BEARING SLITY CLAY. Seni-1+thified above 56 cm, stiff below. Dark laminae 1-2 mm thick, 1-3 cm apart, occur between 56 cm in Sec. 1 and 35 cm in Sec. 2. Bedding deforma- tion distinct around dropped pebble at 35 cm. Pebbles and granules -1%, max. pebble diam. >25 mm, av. pebble diam. >25 mm, av. pebble diam. > 7 mm. Pebbles and granules 3-5%, max. pebble diam. = 7 mm. Pebbles diam. = 7 mm. Lithified, CALCAREDUS SILTY CLAYSTONE, with pebbles. Wisps and irregular darker gray Clay lenses present. Possible loat sats and soft sediment deformation. Dark greenish-gray, stiff DIATOM- BEARNOS SILTY CLAY with pebbles. Max. pebble diam. 54 mm. Bulk X-ray (149.5 m): Amorph. = 55.15 Ident. = 44.95 Quar. = 40.45 K-Fee. = 14.05 Plag. = 26.75 Mica = 14.65 Chlo. = 1.65 Pyri. = 1.75 Amph. = 1.05 	MIDDLE MIOCENE					1 2 3 4		VOID	*0000000000000000000000000000000000000	* *** * * * *	56% clay 40% silt 96% 4% sand 3% diatoms 1% micarb 52% clay 45% silt 97% 3% sand 3% diatoms 70% silt 28% clay 2% sand 15% diatoms 50% silt 47% clay 93% 3% sand 5% diatoms 1% micarb	 Dark gray (N3) - dark greenish- gray semi-lithified SLLTY CLAY. Bedding traces clear above 35 cm. Dark gray clay-rich beds 2 cm thick alternate with silty, paler gray beds. Contacts sharp but irregular. Rare shell fragments. Pebbles and granules <1%, max. pebble diam. 45 mm, av. pebble diam. 45 mm, av. pebble diam. 45 mm. Olive-black (5Y 2/1) semi- lithified SLTY CLAY with pebbles. Gritty throughout. Irregular bedding developed at 110-130 cm. Beds include DIATOM-BEARING CLAYEY SLLT. Faint bedding at 85 cm. Pebbles and granules <1%, max. pebble diam. = 5 mm. Semi-lithified, olive black, gritty SLLTY CLAY with DIATOM-BEARING pebbles.

Core Catcher 60% clay 39% silt 97% 1% sand 1% micarb Very dark greenish-gray (5GY 3/1) pebbly SLLTY CLAY - semilithified. Pebbles and granules sparse.

Image: Note: The second seco	Site 273	Hole	A	Cor	e 9	Col	red In	terval	: 16	6-175.5 m		Sit	e 273	Но	le A		Core	11	Cored In	nterv	val:	178.5-185 m		
Note 120	AGE ZONE	FOS CHAR TISSOJ	ACTER	SECTION	METERS	LITHO	ILOGY	DEFORMATION	LI I RU. SWIFLE	LITHOLOG	IC DESCRIPTION	AGE	ZONE	FOSSIL C	FOSSI ARAC	LL TER .SBW	SECTION	METERS	LITHOLOGY	DEFORMATION	LITHO.SAMPLE		LITHOLOGIC	DESCRIPTION
Site 273 Hole A Core 10 Cored Interval: 175.5-178.5 m	MIDDLE MIDCENE			1 2 Co Cato	re cher	VO			M c	55% clay 4% silt 97% 4% sand 3% diatoms 5% clay 40% silt 95% 5% clay 40% silt 97% 2% sand 2% diatoms 1% micarb	 Olive black, semi-lithified gritty-looking DIATOM- BEARING SILTY CLAY. No bedding visible, but intervals of greater fissility are developed. Pebbles and granules <1%, max. pebble diam. 29 mm, av. pebble diam. 29 mm, av. pebble diam. = 10 mm. Olive-black semi-lithified, gritty DIATOM-BEARING SILTY CLAY. No bedding discernible. Dark greenish-gray to olive black, stiff to semi- lithified SILTY CLAY. Max. pebble diam. 78 mm. Bulk X-ray (166.9 m); Amorph 53.7% Ident 46.3% Quar 38.1% K-Fe 18.6% PIG 25.6% Kaol 1.0% Mica - 13.9% Chlo 1.6% Mont 1.2% 	MIDDLE MIDCENE					1 1 2 3 4		V01D		* * *	58% clay 40% silt 2% sand 2% diator 55% silt 1% sand 2% diator 7R micard 55% clay 34% silt 1% sand 2% diator 7R micard 55% clay 33% silt 2% sand 2% diator 7% diator 7	98% is 97% 97% 97% 97%	 Dark greenish-gray (56Y 3/1) semi- lithified SILTY CLAY. Olive-black, semi-lithified, gritty-looking SILTY CLAY; angular granules common. Rare pebbles, No bedding visible. Slight fissility around 110 cm. Uniform lithology. Pebbles and granules <1% through- out, max. pebble diam. 29 mm, av. pebble diam. = 9 mm.
Understand Understand <td>Site 273 39V</td> <td>Hole FOS CHAR TISSOJ</td> <td>A ACTER ONDER</td> <td>SECTION</td> <td>NETERS</td> <td>LITH</td> <td>DLOGY</td> <td>DEFORMATION</td> <td>LI INO. SAMPLE</td> <td>5.5-178.5 m LITHOLOG</td> <td>IC DESCRIPTION</td> <td></td> <td></td> <td></td> <td></td> <td></td> <td>5</td> <td></td> <td>-D-</td> <td>10°0 DDoc</td> <td>*</td> <td>55% clay 43% silt 2% sand 1% diator TR micar</td> <td>98% 98%</td> <td>Dark greenish-gray to olive-black, semi-lithified, gritty- looking SLIT (LAY with pebbles. Intense fracturing (?fissility) through 10 cm thick intervals.</td>	Site 273 39V	Hole FOS CHAR TISSOJ	A ACTER ONDER	SECTION	NETERS	LITH	DLOGY	DEFORMATION	LI INO. SAMPLE	5.5-178.5 m LITHOLOG	IC DESCRIPTION						5		-D-	10°0 DDoc	*	55% clay 43% silt 2% sand 1% diator TR micar	98% 98%	Dark greenish-gray to olive-black, semi-lithified, gritty- looking SLIT (LAY with pebbles. Intense fracturing (?fissility) through 10 cm thick intervals.
Understand Diffiling sludge - dark greenish- gray (567 3/1) SiLTY CLAY, abundant pebbles. Diffiling sludge - dark greenish- gray (567 3/1) SiLTY CLAY, abundant pebbles. Site 2/3 Hole A Core iz Core iz Core il Z Core il Z <td></td> <td></td> <td></td> <td></td> <td>0.5</td> <td></td> <td></td> <td>0</td> <td>1</td> <td></td> <td></td> <td></td> <td></td> <td>_</td> <td></td> <td></td> <td></td> <td>F</td> <td></td> <td>2</td> <td></td> <td></td> <td></td> <td></td>					0.5			0	1					_				F		2				
Image: Second	DDLE MIOCENE			1	1.0			0000		55% clay 44% silt 98% 1% sand 2% diatoms	Drilling sludge - dark greenish- gray (5GY 3/1) SILTY CLAY, abundant pebbles. Pebble density '10%, max. pebble diam. = 14 mm.	GE	Se 273 BNOZ	EDSSTI 12	FOSS ARAC	IL TER Saya	SECTION	METERS	Cored I	DEFORMATION	LITHO.SAMPLE	185-194.5 m	LITHOLOGIC	DESCRIPTION
				2		•		00000	*	IR micarb 50% clay 46% silt 98% 4% sand 2% diatoms TR micarb	Pebbles and granules <1%, max. pebble diam. 12 mm, av. pebble diam. = 7 mm. Grayish-green to olive black semi-1ithifed, gritty- looking SILIY CLAY. No bedding visible. Fissile interval 105-120 cm.	MIDDLE MIOCFNE					1 Cor Catc	0.5	VOID	UQ o		60% clay 37% silt 3% sand 4% diato TR micar 59% clay 40% silt 1% sand 3% diato 1% micar	95% b 95% s	Olive-black, semi-lithified SILTY CLAY. No bedding visible. Pebbles and granules <1%. Very dark greenish-gray, stiff to semi-lithified SILTY CLAY. Granules and small pebbles rare.

Site 273		ł	tole	A	Core 13		Con	Cored Intern			194.5-204 m			e	273	Hol	2 A	C	ore 14	Core	d Inte	erval	; 204-	213.5 m		
ACC	ZONE		FOS CHAR/ TISSOJ	SIL ACTER	SECTION	METERS	LITHO	.0GY	DEFORMATION	LITHO.SAMPLE	LITHOLOGI	C DESCRIPTION	AGE		ZONE	FOSSIL 3	SSIL RACTE	PRES. 33	METERS	LITHOL	OGY	LEFUKMALION			LITHOLOGI	C DESCRIPTION
MIDDLE MIDGENE			1 2 3	0.5	V01	V01D	امل ال ال ال ال ال ال مؤلى م	* * *	60% clay 38% silt 96% 2% sand 3- 4% diatoms TR micarb 57% clay 42% silt 98% 1% sand 1% diatoms TR micarb 55% clay 43% silt 97% 2% sand 1- 2% diatoms TR micarb	<pre>Very dark greenish-gray (56Y 3/1), semi-lithified DIATOM- BEARING SILIY CLAY with pebbles. Bivalve mold. No bedding. Pebbles and granules <1% through- out, max. pebble diam. 60 mm, av. pebble diam. = 10 mm. Very dark greenish-gray (56Y 3/1) gritty-looking, semi- lithified SILIY CLAY with pebbles. Uniform lithology.</pre>	WIDDLE MIDCENE Sit	te	273	Hol	e A	1 2 ca	0.5- 1.0-	VOI	D D C C C C C C C C C C C C C C C C C C		: 213.	56% clay 40% silt 41% sand 3% diatoms TR micarb 60% clay 40% silt TR sand 2% diatoms TR micarb 60% clay 39% silt 1% sand 3% diatoms 1% micarb	96% 98% 95%	<pre>Very dark greenish-gray (56Y 3/1) semi-lithified, gritty- looking, pebbly SLLY CLAY. Rare shell fragments. No bedding discernible. Pebbles and granules ~1% through- out. Max. pebble diam. 50 mm, av. pebble diam. = 10 mm. Very dark greenish-gray (55Y 3/1) semi-lithified, pebbly SLLY CLAY. Rare shell frag- ments. Max. pebble diam. 23 mm.</pre>		
					4	to the sector of		$O \circ \circ O O$		55% clay 43% silt 97% 2% sand 2% diatoms TR micarb	Very dark greenish-gray, gritty- looking SILTY CLAY. Bedding absent.	AGE		ZONE	FOSSIL E.	RACTI 'ONNBY	PRES. B	METERS	LITHOL	OGY	DEFORMATION	L1100-2001 LC		LITHOLOGI	C DESCRIPTION	
					Ca	ore tcher			00	*	65% clay 32% silt 96% 3% sand 3% diatoms 1% mfcarb	Very dark greenish-gray, semi- lithified SILTY CLAY with pebbles. Bulk X-ray (198.8 m): Amorph 53.1% Ident 46.9% Quar 38.5% K-Fe 14.2% Plag 24.4% Mica - 15.2% Chio 2.2% Mont 3.6% Gyps 1.9%	MIDDLE MIDCENE	olar	natory	note	es in	1 C	0.5 1.0 Core atcher	V01	D	*		69% clay 30% silt 1% sand 3% diatoms TR micarb	96%	 Mainly an aggregate of apparently well=sorted pebbles. Some greenish-gray mud. Pebble sample shows size range 4-16 mm (Mean 10 mm). Sub- angular to subrounded. Greenish-gray SILTY CLAY with abundant pebbles. Max. pebble diam. 54 mm.

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