

## 5. SITE 445, DAITO RIDGE, DEEP SEA DRILLING PROJECT LEG 58

The Shipboard Scientific Party<sup>1</sup>

### HOLE 445

**Date occupied:** January 11, 1978

**Date departed:** January 17, 1978

**Time on hole:** 6 days

**Position (latitude; longitude):** 25°31.36' N; 133°12.49' E

**Water depth (sea level; corrected m, echo sounding):** 3377

**Water depth (rig floor; corrected m, echo sounding):** 3387

**Bottom felt (m, drill pipe):** 3382

**Penetration (m):** 892.0

**Number of cores:** 94

**Total length of cored section (m):** 892.0

**Total core recovered (m):** 619.52

**Core recovery (%):** 69

#### Oldest Sediment Cored:

**Depth sub-bottom (m):** 892.0

**Nature:** conglomeratic sandstone

**Age:** early middle Eocene

**Measured velocity (km/s):** 3.6

#### Basement:

**Depth sub-bottom (m):** not reached

**Principal Results:** Site 445 is in a small basin in the Daito Ridge, northwest Philippine Sea. Sediments range in age from middle Eocene to Pleistocene, and the cored interval is 892 meters thick. A hiatus in sedimentation occurred during early Oligocene. Occurrence of *Nummulites boninensis* is limited to resedimented debris-flow conglomerate beds. The carbonate sediments, mudstones, sandstones, and con-

glomerates are dominantly resedimented by slumping, debris flows, and turbidity currents. Shipboard analysis of paleomagnetic inclination of samples indicates that Site 445 migrated from an equatorial latitude to its present position over the last 50 m.y.

### BACKGROUND AND OBJECTIVES

#### Background

The marginal basins of the western Pacific owe their origin to a variety of processes, and deep drilling during Leg 58 focused on three sites in the Shikoku Basin to test several models for the spreading origin of that basin. Previous Leg 58 results (Sites 442, 443, and 444) indicated that three modes of spreading were possible in the Shikoku Basin: symmetrical spreading, single-limb spreading, and asymmetrical spreading.

The second part of Leg 58 deals with the triangular northwest portion of the north Philippine Sea immediately around the remnant arcs of the Daito Ridge and Basin province. This region has been of considerable interest to several workers, and prior study included deep drilling in the adjoining Kyushu-Palau Ridge at Site 296 (Karig, Ingle, et al., 1975), structural analysis of the basin (Karig, 1975; Hilde et al., 1977; Watts et al., 1977; Mizuno et al., 1975, 1979), and dredging (Mizuno et al., 1975, 1979; Shiki et al., 1976). Magnetic lineations were identified south of the Daito Ridge and Basin area (Louden, 1976; Watts et al., 1977), but none have been identified in the area itself. From this, it has been inferred that the Daito Ridge and Basin region is very old, and that in fact this portion of the northwest Philippine Sea incorporates older crust trapped behind the remnant arc of the Daito Ridge and Oki-Daito Ridge (Karig, 1975; Hilde et al., 1977; Watts et al., 1977; Mizuno et al., 1975, 1979). The trapping mechanism should show similarities with the hypothetical origin of the Bering Sea (Cooper et al., 1976).

Dredge hauls from the Daito Ridge and Basin region, the Oki-Daito Ridge (immediately south), and the Amami Plateau (immediately to the north) indicate that the geology of the region is extremely variable. Greenschist, hornblende schist, and serpentinite have been collected from the Daito Ridge, indicating some regional metamorphism (Mizuno et al., 1975; Shiki et al., 1976). Igneous rocks recovered from the Daito Ridge include andesite and diorite of island-arc origin. Dredge hauls from the Oki-Daito Ridge recovered basalt; andesite, granodiorite, and basalt were recovered from the Amami Plateau.

Perhaps the most interesting discovery from dredging is limestone samples containing *Nummulites boninensis*,

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a larger foraminifer of shallow-water origin. These Eocene samples (Konda et al., 1977) were obtained from these ridges and the Amami Plateau from water depths ranging from 1160 to 2340 meters. This fossil suggests either that the entire area was uplifted during Eocene time and subsequently subsided (Mizuno et al., 1975, 1979), or that the fossils were resedimented from coastal areas into deeper environments. If uplift had occurred, followed by subsequent subsidence, then the Daito Ridge and Basin area represents a remnant arc whose uplift was tied to deep-seated igneous and metamorphic activity or to trench tectonics early in its history (Mizuno et al., 1975).

Site 445 was located in a small sediment pond on the southern side of the Daito Ridge, along a seismic profile surveyed by the R/V *Kaiyo-Maru* (IPOD-Japan, 1977), shown in Figure 1. The seismic-reflection profile obtained by the D/V *Glomar Challenger* is shown in Figure 2.

### Objectives

The primary objectives for drilling at Site 445 were fivefold. Of prime importance was to determine the age of the oldest sediment and of the basement to determine in turn whether this portion of the northern Philippine Sea incorporates old crust trapped behind a remnant arc. A second objective was to determine the nature of the basement and to elucidate its crustal history. A third objective at this site was to determine by sedimentology and paleoecology the nature of the subsidence history of the northwest Philippine Sea. Fourth, climatic changes at the site, presumably due to its supposed northward drift, were to be determined by paleontology and paleomagnetism.

The location of Site 445 near the Kuroshio Current provides a unique opportunity to understand its circulation history. That history can be elucidated from sedimentology and paleontology of samples from Site 445; this was the fifth major objective.

### OPERATIONS

Drilling operations at Site 444 terminated on 9 January 1978. At 0300 hours seismic gear was streamed, and at 0400 hours the *Challenger* was under way to Site 445 on a course of 236°.

Weather and sea conditions deteriorated late in the evening of 9 January and through the early hours of 10 January. Accordingly, at 1130 hours, 10 January, winds up to 40 mph and extremely high seas necessitated slowing the ship's speed to 140 rpm in order to better ride out the storm. Seismic gear was also retrieved. Storm conditions prevailed until 0400 hours, 11 January, when the ship's speed was increased to 190 rpm and seismic gear was streamed. The *Challenger* headed for Site 445 on a course of 202°.

At 1600 hours, 11 January, a 16-kHz beacon was dropped, marking Site 445 (Figure 3). Water depth at the site was later established as 3382 meters by drill-pipe measurement. The seismic gear was retrieved; at 1714 hours the *Challenger* established an auto-mode positioning over the beacon, and RIH was started. Spud-in

for Hole 445 was at 0217 hours, 12 January. Core 1, 8.5 meters of nannofossil ooze, was recovered at 0304 hours, 12 January.

A program of continuous coring was followed, the time interval (on deck) between cores being about one hour (Table 1). The sediment nature was consistent and permitted exceptionally high core recoveries (in some cases exceeding 100 per cent; Table 1).

On 16 January, weather predictions showed that a frontal system with accompanying high winds and seas was headed for the site area. Conditions were moderate through the 16th and into early hours of 17 January. However, at 0830 hours, conditions affecting positioning were severe enough to force abandonment of Hole 445. Accordingly, POOH began, and at 1530 hours, the *Challenger* was under way to Site 446 on a course of 206° (Figure 3).

In all, 94 cores were taken, with a total sub-bottom penetration of 892 meters. The length recovered was 619.52 meters, or 69 per cent (Table 1).

### SEDIMENT LITHOLOGY

A single-bit hole was drilled at Site 445 by continuous coring. It penetrated a total of 892 meters of sediments of middle-Eocene to Quaternary age and recovered 94 cores, 619.52 meters in total length.

Sediments are dominantly redeposited biogenic components for the late Eocene to the Quaternary, and dominantly redeposited terrigenous components for the middle Eocene. Five lithologic units were distinguished, based on the biogenic components (units I-IV) and terrigenous components (unit V). Some of the units are divided into sub-units on the basis of dominant color (unit I), abundance of siliceous biogenic components (unit II), and abundance of calcareous biogenic components and sediment texture (unit V). These units and sub-units are conformable, except for an early-Oligocene or late-Eocene hiatus between units II and III, discerned from micropaleontological data. Figure 4 shows the stratigraphy at Site 445, and Figure 5 shows the vertical distribution of some lithologic components.

#### Unit I

Unit I is predominantly soft to firm nannofossil ooze and foraminifer-nannofossil ooze, with some associated nannofossil-foraminifer ooze, volcanic ash layers, and scattered volcanic ash. Sediments are variously colored, predominantly from gray to brown. Vertical distribution of the dominant colors allows further subdivision into two sub-units.

##### Sub-Unit Ia

This sub-unit consists largely of soft to firm nannofossil ooze and interbedded foraminifer-nannofossil ooze, which are locally clayey or vitric and are associated with nannofossil-foraminifer ooze. Dominant colors of these sediments are light gray and pale brown.

The upper interval consists of interbedded nannofossil ooze and foraminifer-nannofossil ooze, whereas the remaining interval is predominantly nannofossil ooze

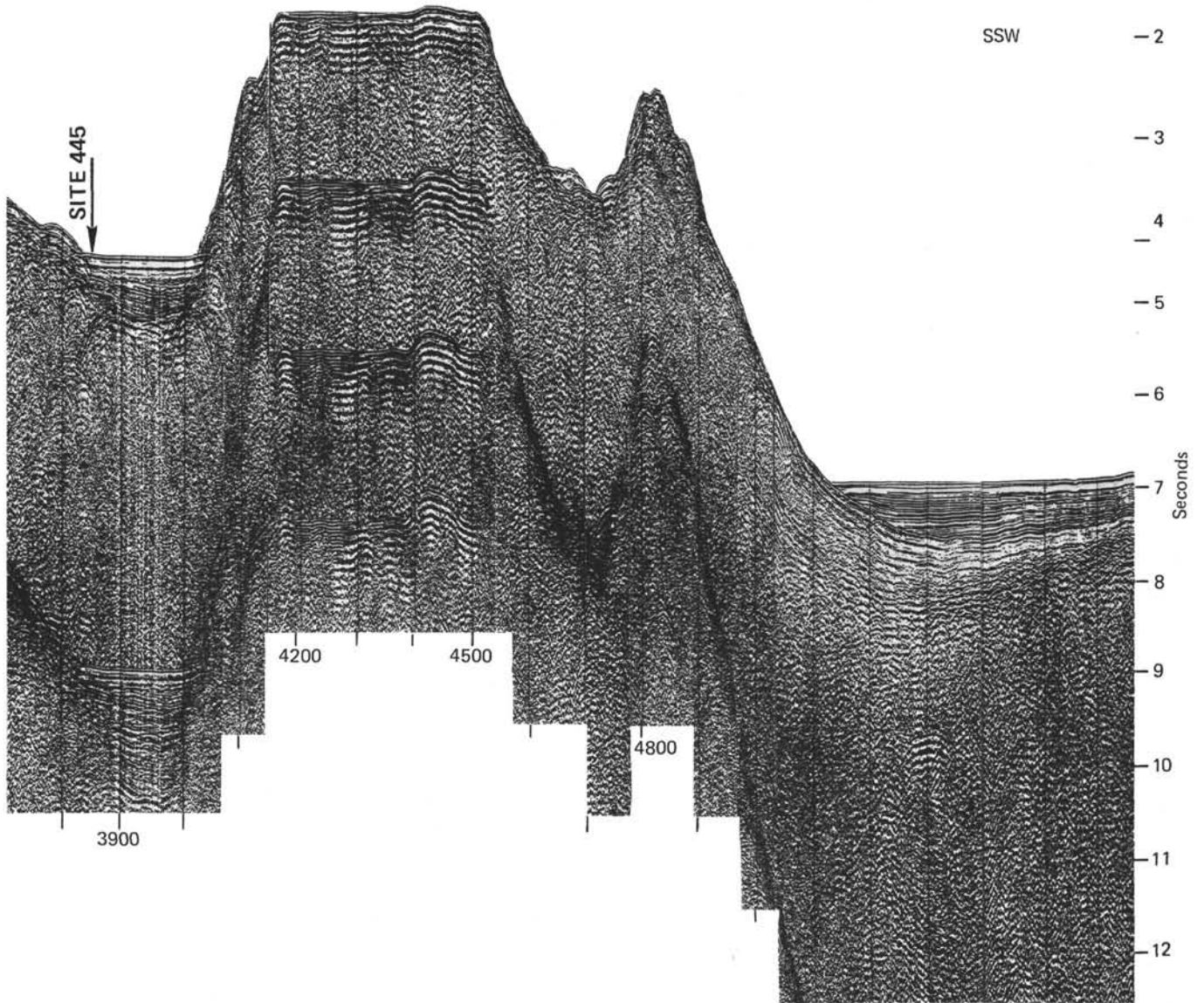


Figure 1. Seismic-reflection profile through Daito Ridge by R/V Kaiyo-Maru.

with subordinate amounts of foraminifer-nannofossil ooze. The sediments are mostly of silty clay texture (nannofossil ooze) to sandy mud texture (foraminifer-nannofossil ooze). A graded sequence identifiable to the partial Bouma sequence T(d,e) is observed in Core 3-3.

#### Sub-Unit Ib

This sub-unit is distinguished from sub-unit Ia by sediment color; it consists dominantly of firm, light-gray to white nannofossil ooze, associated with interbeds of foraminifer-nannofossil ooze of various thicknesses.

Foraminifer-nannofossil ooze is dominant in the upper part (Core 10). Nannofossil ooze locally passes to darker-colored, clayey nannofossil ooze which contains clay (15-20% of total).

Coarse-tail graded bedding 5 to 10 cm thick is commonly found in color-band alternations in the middle part of the sub-unit. It starts from very fine sand- to silt-

sized nannofossil-foraminifer ooze or foraminifer ooze with parallel laminae just above a sharp contact, which fines upward to silty-clay-sized nannofossil ooze. It may be identified to the partial Bouma sequence T(d,e). In general, bioturbation is slight to moderate in the middle to lower part. Medium- to coarse-grained calcareous sand occurs only at the lowest part of this sub-unit. Containing about 60 per cent rock fragments, it rests just above the scoured bottom (the lowest boundary of sub-unit Ib) with a 10-cm-thick graded bed and passes upward to nannofossil ooze. This represents another type of graded sequence, equivalent to T(a,b,e).

#### Unit II

Unit II sediments are stiffer than those of Unit I. The dominant lithology is a gray, brown, or pink, very firm nannofossil chalk, with varying amounts of siliceous, clayey, and volcanic components. Mudstone occurs locally. It tends to be concentrated at the middle part of



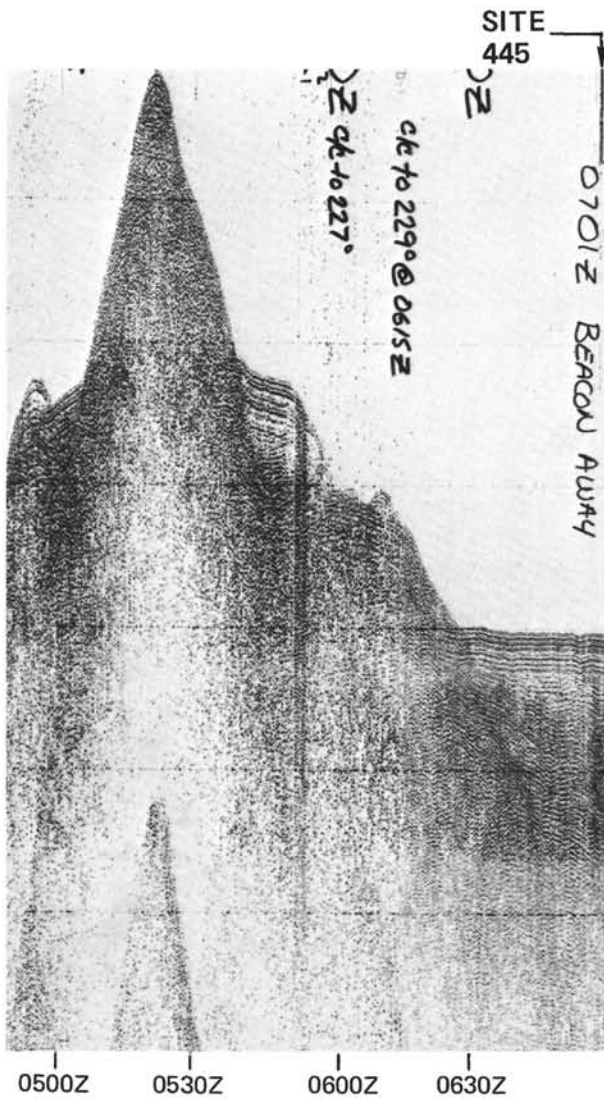


Figure 2. Glomar Challenger seismic-reflection profile approaching Site 445. See Figure 3 for location.

this unit (sub-unit IIb and the upper part of sub-unit IIc). Bioturbation is very common and intense. Calcareous nannofossils occur abundantly throughout the unit. The abundance of siliceous fossils, mostly radiolarians and sponge spicules, varies, and this makes it possible to subdivide the thick sequence of unit II into four sub-units. In sub-units IIa and IIc, siliceous fossils are nearly absent, whereas in sub-units IIb and IId they range from several per cent to around 10 per cent or more.

A distinctive aspect of unit II is the extensive occurrence of various sedimentary structures, such as graded sequences and bioturbation, locally associated with slump structures and microfaults.

The graded sequences are coarse-tail and mostly consist of resedimented calcareous biogenic components accompanied by accessory amounts of terrigenous coarse-grained materials at the basal interval, and in some cases by a thin layer of pelagic clay in the upper interval. They include the typical Bouma sequence T(a-e) and partial Bouma sequences T(c-e), T(d,e).

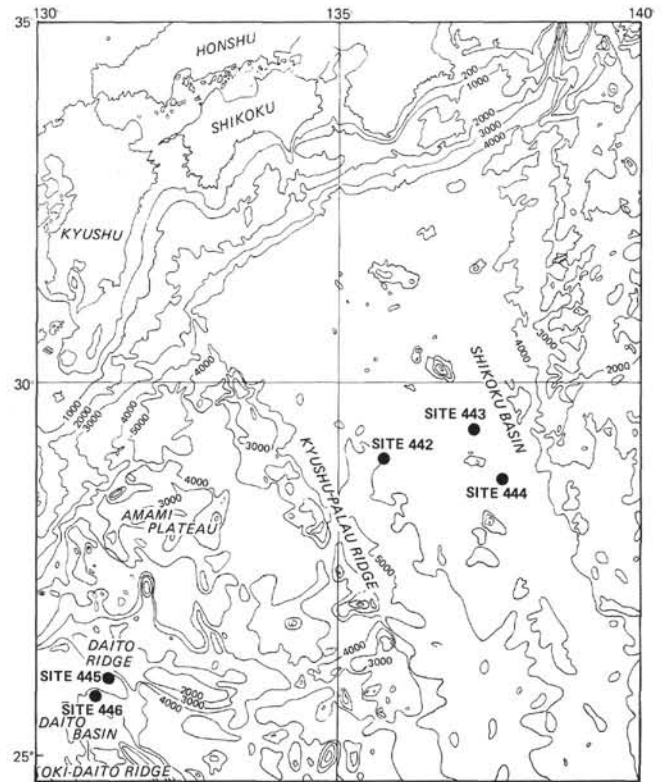


Figure 3. Site location map.

Occurrence of T(a-e) is restricted to the upper part of sub-unit IIa and the lower part of sub-unit IIc, and it mostly overlies, with sharp, scoured contact, the underlying sediments, which are associated with slump structures. It is usually 35 to 60 cm thick, with rare cases more than 5 meters thick.

Interval a consists of foraminifer chalk or calcareous chalk with lithic fragments and reworked fossils; it is coarse- to very coarse-sand size at the base and medium- to fine-sand texture at the top. In some cases it is partially replaced by calcareous sand with detrital grains of several tens per cent. This interval gradually passes upward to interval b, parallel laminated nannofossil-foraminifer chalk of fine-grained-sand size. This passes upward to interval c, foraminifer-nannofossil chalk (very fine-grained-sand to muddy-sand texture) with micro-cross-laminae and (or) convolute lamination. Above this, interval d is parallel laminated foraminifer-nannofossil chalk (sandy-mud to silty-clay texture). Interval e is a thick, faintly laminated or non-laminated nannofossil chalk and (or) clayey nannofossil chalk (silty-clay to clay size) with a clay content of clayey materials from a trace to 15 per cent of the total. This interval tends to be slightly bioturbated in the lower part, and moderately to intensely bioturbated in the upper part. Clayey nannofossil chalk is in some cases replaced by nannofossil mudstone.

T(d,e) is very common throughout unit II. The thickness of each bed is usually several tens of centimeters, but often decreases to around 20 cm or increases to more than 4 meters.



TABLE 1  
Site 445 Coring Summary

Cores	Date (Jan., 1978)	Time	Depth From	Depth Below	Length Cored (m)	Recovery (m)	Recovery (%)
			Drill Floor (m)	Sea Floor (m)			
445-1	12	0304	3382.0-3390.5	0.0-8.5	8.5	8.50	100
2	12	0408	3390.5-3400.0	8.5-18.0	9.5	6.10	64
3	12	0505	3400.0-3409.5	18.0-27.5	9.5	9.61	101
4	12	0610	3409.5-3419.0	27.5-37.0	9.5	9.50	100
5	12	0706	3419.0-3428.5	37.0-46.5	9.5	3.61	38
6	12	0805	3428.5-3438.0	46.5-56.0	9.5	8.62	91
7	12	0906	3438.0-3447.5	56.0-65.5	9.5	9.10	96
8	12	1000	3447.5-3457.0	65.5-75.0	9.5	8.27	87
9	12	1057	3457.0-3466.5	75.0-84.5	9.5	0.38	4
10	12	1155	3466.5-3476.0	84.5-94.0	9.5	7.90	83
11	12	1309	3476.0-3485.5	94.0-103.5	9.5	6.55	69
12	12	1412	3485.5-3495.0	103.5-113.0	9.5	6.00	63
13	12	1513	3495.0-3504.5	113.0-122.5	9.5	6.10	64
14	12	1615	3504.5-3514.0	122.5-132.0	9.5	1.95	21
15	12	1717	3514.0-3523.5	132.0-141.5	9.5	2.40	25
16	12	1817	3523.5-3533.0	141.5-151.0	9.5	5.34	56
17	12	1922	3533.0-3542.5	151.0-160.5	9.5	2.27	24
18	12	2020	3542.5-3552.0	160.5-170.0	9.5	6.66	70
19	12	2120	3552.0-3561.5	170.0-179.5	9.5	5.54	58
20	12	2230	3561.5-3571.0	179.5-189.0	9.5	3.46	36
21	12	2335	3571.0-3580.5	189.0-198.5	9.5	3.22	34
22	13	0047	3580.5-3590.0	198.5-208.0	9.5	1.76	19
23	13	0147	3590.0-3599.5	208.0-217.5	9.5	5.11	54
24	13	0247	3599.0-3609.0	217.5-227.0	9.5	7.11	75
25	13	0349	3609.0-3618.5	227.0-236.5	9.5	8.50	89
26	13	0446	3618.5-3628.0	236.5-246.0	9.5	8.61	91
27	13	0546	3628.0-3637.5	246.0-255.5	9.5	8.34	88
28	13	0644	3637.5-3647.0	255.5-265.0	9.5	7.60	80
29	13	0815	3647.0-3656.5	265.0-274.5	9.5	5.58	59
30	13	0915	3656.5-3666.0	274.5-284.0	9.5	6.98	73
31	13	1021	3666.0-3675.5	284.0-293.5	9.5	9.60	101
32	13	1116	3675.5-3685.0	293.5-303.0	9.5	3.77	40
33	13	1220	3685.0-3694.5	303.0-312.5	9.5	6.95	73
34	13	1318	3694.5-3704.0	312.5-322.0	9.5	7.45	78
35	13	1420	3704.0-3713.5	322.0-331.5	9.5	5.60	59
36	13	1521	3713.5-3723.0	331.5-341.0	9.5	7.52	79
37	13	1620	3723.0-3732.5	341.0-350.5	9.5	7.65	81
38	13	1735	3732.5-3742.0	350.5-360.0	9.5	8.15	86
39	13	1838	3742.0-3751.5	360.0-369.5	9.5	5.71	60
40	13	1953	3751.5-3761.0	369.5-379.0	9.5	4.37	46
41	13	2059	3761.0-3770.5	379.0-388.5	9.5	8.32	88
42	13	2215	3770.5-3780.0	388.5-398.0	9.5	8.25	87
43	13	2330	3780.0-3789.5	398.0-407.5	9.5	6.20	65
44	14	0048	3789.5-3799.0	407.5-417.0	9.5	6.17	65
45	14	0154	3799.0-3808.5	417.0-426.5	9.5	6.96	73
46	14	0305	3808.5-3818.0	426.5-436.0	9.5	9.06	95
47	14	0415	3818.0-3827.5	436.0-445.5	9.5	8.46	89
48	14	0529	3827.5-3837.0	445.5-455.0	9.5	8.51	90
49	14	0646	3837.0-3846.5	455.0-464.5	9.5	7.24	76
50	14	0802	3846.5-3856.0	464.5-474.0	9.5	3.15	33
51	14	0920	3856.0-3865.5	474.0-483.5	9.5	5.48	58
52	14	1117	3865.5-3875.0	483.5-493.0	9.5	7.02	74
53	14	1240	3875.0-3884.5	493.0-502.5	9.5	7.10	75
54	14	1358	3884.5-3894.0	502.5-512.0	9.5	9.74	103
55	14	1522	3894.0-3903.5	512.0-521.5	9.5	9.96	105
56	14	1643	3903.5-3913.0	521.5-531.0	9.5	7.27	74
57	14	1816	3913.0-3922.5	531.0-540.5	9.5	9.56	101
58	14	2013	3922.5-3932.0	540.5-550.0	9.5	8.01	84
59	14	2127	3932.0-3941.5	550.0-559.5	9.5	9.69	102
60	14	2238	3941.5-3951.0	559.5-569.0	9.5	5.14	54
61	14	2347	3951.0-3960.5	569.0-578.5	9.5	9.67	102
62	15	0059	3960.5-3970.0	578.5-588.0	9.5	9.69	102
63	15	0206	3970.0-3979.5	588.0-597.5	9.5	5.98	63
64	15	0314	3979.5-3989.0	597.5-607.0	9.5	9.59	101
65	15	0432	3989.0-3998.5	607.0-616.5	9.5	3.82	40
66	15	0541	3998.5-4008.0	616.5-626.0	9.5	6.42	68
67	15	0826	4008.0-4017.5	626.0-635.5	9.5	1.27	13
68	15	1046	4017.5-4027.0	635.5-645.0	9.5	2.25	24
69	15	1242	4027.0-4036.5	645.0-654.5	9.5	5.47	58
70	15	1417	4036.5-4046.0	654.5-664.0	9.5	1.78	19
71	15	1547	4046.0-4055.5	664.0-673.5	9.5	5.31	56
72	15	1720	4055.5-4065.0	673.5-683.0	9.5	2.05	22
73	15	1846	4065.0-4074.5	683.0-692.5	9.5	4.47	47
74	15	2021	4074.5-4084.0	692.5-702.0	9.5	7.46	79
75	15	2159	4084.0-4093.5	702.0-711.5	9.5	9.00	95
76	15	2333	4093.5-4103.0	711.5-721.0	9.5	7.71	81
77	16	0121	4103.0-4112.5	721.0-730.5	9.5	7.33	77
78	16	0311	4112.5-4122.0	730.5-740.0	9.5	8.47	89
79	16	0446	4122.0-4131.5	740.0-749.5	9.5	9.87	104
80	16	0606	4131.5-4141.0	749.5-759.0	9.5	8.35	88
81	16	0730	4141.0-4150.5	759.0-768.5	9.5	8.32	88
82	16	0903	4150.5-4160.0	768.5-778.0	9.5	6.63	70
83	16	1029	4160.0-4169.5	778.0-787.5	9.5	7.33	77
84	16	1206	4169.5-4179.0	787.5-797.0	9.5	4.50	47
85	16	1348	4179.0-4188.5	797.0-806.5	9.5	6.50	68
86	16	1517	4188.5-4198.0	806.5-816.0	9.5	5.63	59
87	16	1643	4198.0-4207.5	816.0-825.5	9.5	8.36	88
88	16	1816	4207.5-4217.0	825.5-835.0	9.5	8.07	85
89	16	2018	4217.0-4226.5	835.0-844.5	9.5	9.74	103
90	16	2153	4226.5-4236.0	844.5-854.0	9.5	4.42	47
91	17	0013	4235.0-4244.5	854.0-863.5	9.5	7.15	75
92	17	0308	4245.5-4255.0	863.5-873.0	9.5	4.55	48
93	17	0608	4255.0-4264.5	873.0-882.5	9.5	6.43	68
94	17	0807	4264.5-4274.0	882.5-892.0	9.5	7.20	76
			Totals		892.0	619.52	69

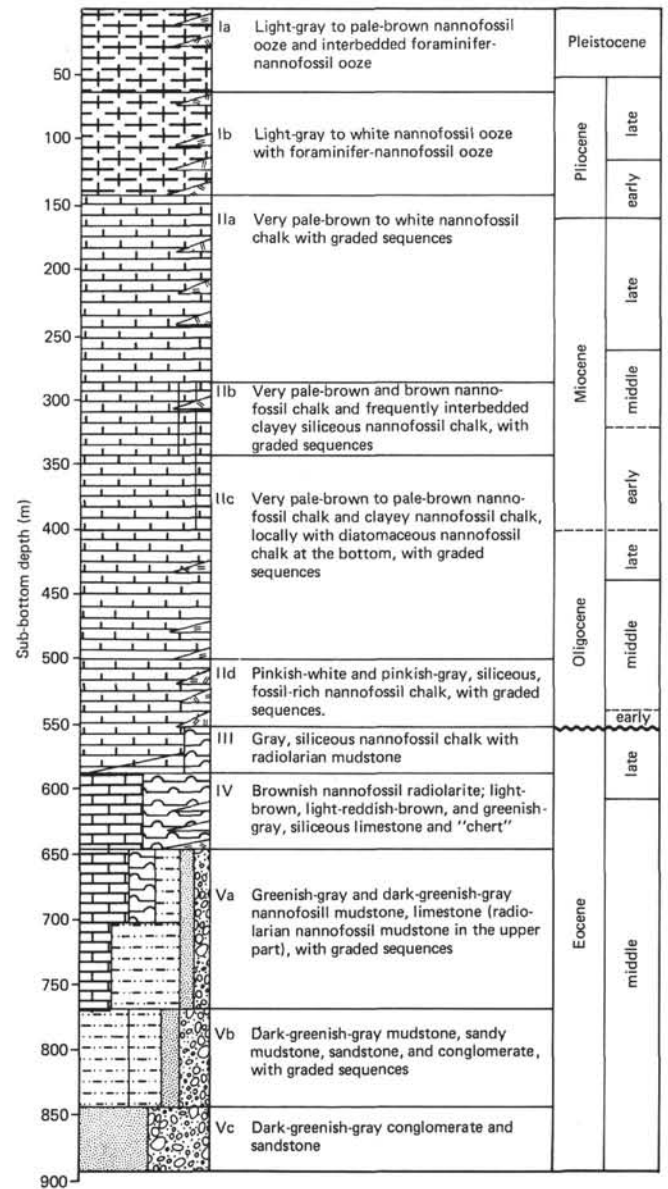


Figure 4. Stratigraphic summary, Site 445.

T(c-e) and T(a,b,e) are very scarce. The former is found in the lowest part of sub-unit IIb, sub-unit IIc, and sub-unit IId, and the latter is restricted to sub-unit IIa and sub-unit IIb.

T(b-e) is only observed in sub-unit IId.

Through all the types of sequences, the intervals a, b, and c generally include tests of shallow-water benthic foraminifers, and the a in the T(a-e) sequence contains many lithic fragments of very coarse-grained-sand to granule size, including clasts of muddy sediments formed earlier.

Slump structures occur in some intervals of unit II. They are concentrated in the middle to upper part of sub-unit IIa and in the upper and lower parts of sub-unit IIc. They are represented by slump folds extending vertically through an interval of 2 to 4 meters in most cases, and also by slump blocks of smaller magnitude. Some of them are accompanied by the typical Bouma

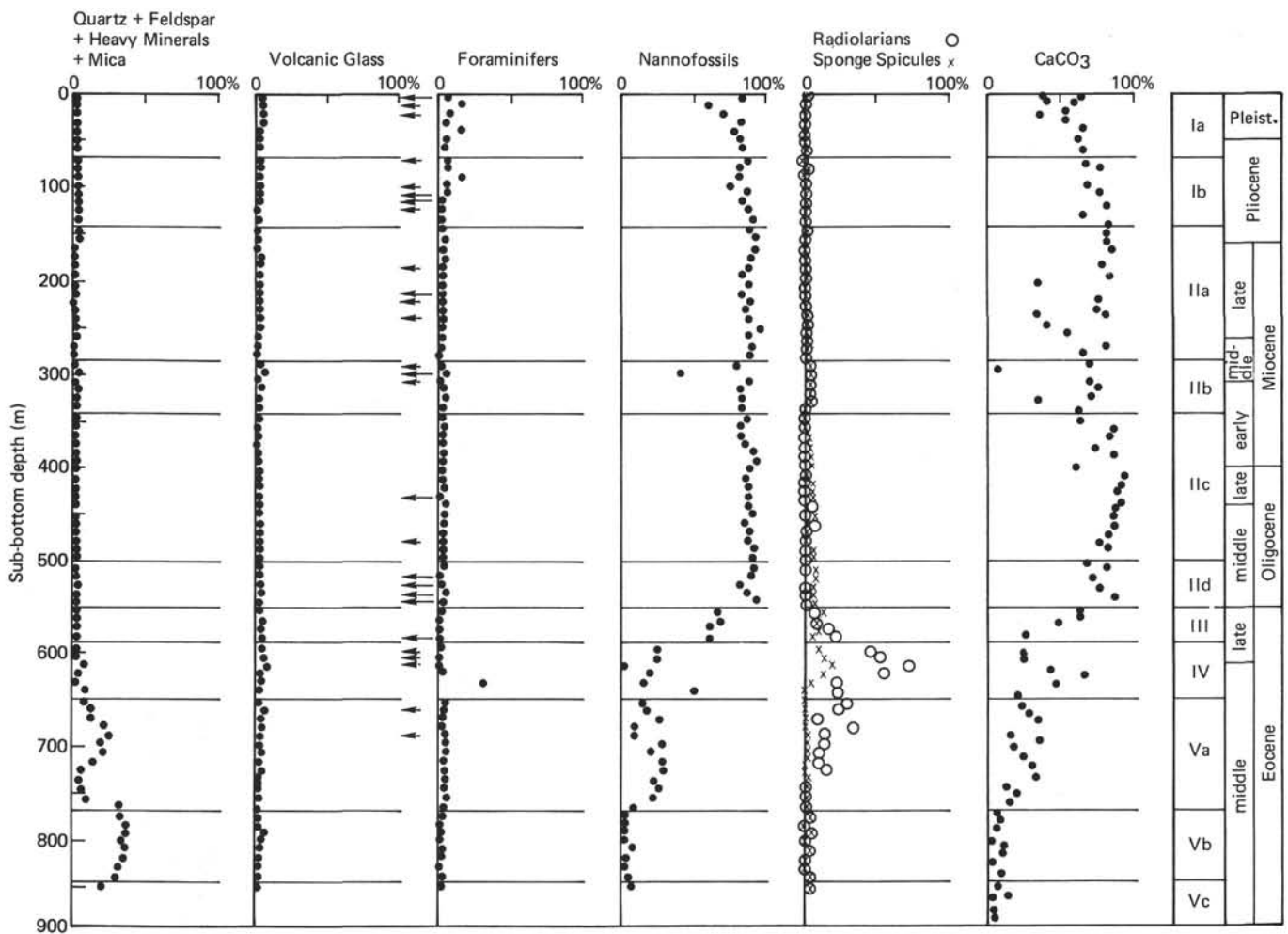


Figure 5. Composition of dominant lithology, Site 445, determined by smear-slide observation, with CaCO<sub>3</sub> data from shore-based analysis. Plotted are average values of the components of dominant lithology in each core. Arrows in the volcanic glass column indicate the interbedded ash layers as a minor lithology: long arrows, volcanic ash or tuff interbeds; short arrows, vitric interbeds.

sequence T(a-e) and are definitely eroded, as shown by scoured, sharp basal contacts.

These sedimentary structures suggest that the larger part of unit II was resedimented by turbidity currents.

**Sub-Unit IIa**

This sub-unit consists mostly of nannofossil chalk. Clayey nannofossil chalk, marly nannofossil chalk, nannofossil mudstone, vitric nannofossil chalk, and nannofossil tuff occur as thin interbeds. Nannofossil-foraminifer chalk and foraminifer-nannofossil chalk occur as very thin layers in the basal coarse-tail graded sequence. Dominant color of sediment is very pale brown to white.

Volcanic-glass shards are generally absent, but they range from 30 to 60 per cent. Siliceous fossils are very rare.

Graded sequences with bioturbation are common. Slump structures occur through the upper half of this sub-unit.

**Sub-Unit IIb**

This sub-unit consists dominantly of nannofossil chalk, with varying amounts of clayey, siliceous, and volcanic components. Nannofossil chalk is frequently interbedded with clayey nannofossil chalk and nannofossil mudstone. The dominant color of the sediments is the same as in sub-unit IIa.

Bioturbation is generally intense in the upper part of interval e; this tends to relate to increasing clay content, which produces color-bands.

**Sub-Unit IIc**

This sub-unit is similar to sub-unit IIb. It consists mainly of nannofossil chalk, with frequent interbeds of clayey nannofossil chalk, nannofossil mudstone, and foraminifer-nannofossil chalk. Generally, clayey components are abundant in the upper part of this sub-unit, and silty claystone only occurs in the uppermost part.

The sediments are dominantly very pale brown, pale brown, and brown. Color changes frequently form col-

or-band alternations, mostly varying with the amount of clayey component, particularly in the upper part of the sub-unit.

Graded sequences, with associated bioturbation, are extensively developed. The thickness of the beds ranges from 50 to 450 cm, mostly 80 to 100 cm.

The middle to upper part of the sub-unit often includes thin beds of another type. They are usually 30 to 40 cm thick and rest upon a sharp contact; they consist of very dark-colored nannofossil mudstone, passing upward to lighter-colored nannofossil chalk, either with or without a basal silty layer.

Slump structures occur in the upper part and the lower part.

#### Sub-Unit IIc

This sub-unit is dominantly nannofossil chalk, with varying proportions of clayey, siliceous, and volcanic components. The dominant color is pinkish white and pinkish gray to light gray and gray.

Volcanic ash is rather common throughout this sub-unit, both as interbeds and as an admixture.

Sedimentary structures are largely represented by graded bedding, intense bioturbation, and wavy lamination. In the upper half of the sub-unit, thick beds are intensely bioturbated. The intense bioturbation very likely obliterates the bedding which might have existed. The dark-colored (black to dark-gray), tuffaceous sediments, either clayey or non-clayey, have a unique sedimentary structure in the lower part of the sub-unit. They rest on nannofossil chalk with a sharp boundary and pass upward to nannofossil chalk, and the structure is very similar to that of the middle to upper part of sub-unit IIc. Dish structures are observed in Core 57-7.

#### Unit III

Unit III consists dominantly of siliceous nannofossil chalk with less than 10 per cent clay, radiolarians from 4 to 25 per cent, and sponge spicules 3 to 15 per cent (combined radiolarians and sponge spicules about 30 per cent maximum). The sediment is dominantly pinkish gray and light gray to gray.

Sedimentary structures are dominated by extensively developed parallel and wavy laminations (very thin-bedded siliceous nannofossil chalk) and by intense bioturbation. Micro-faults are found in Core 61-3.

#### Unit IV

Unit IV is essentially characterized by frequent occurrence of radiolarians and well-indurated rocks such as siliceous limestone and "chert."

The middle to upper part is largely composed of radiolarite, with interbeds of more-calcareous, clayey, or ashy layers. Dominant color is olive gray to light brown to light yellowish brown. Siliceous fossils form 60 to 80 per cent of the total in the radiolarite, of which two thirds to four fifths consists of radiolarians and the remainder of sponge spicules.

Bioturbation is moderate to intense and extensive. Graded sequences are very poor, represented rarely by some intervals similar to T(d,e). Instead, sandy to silty,

very thin layers (a few millimeters) are thin bedded, without definite upper and lower boundaries.

In the middle part of this unit (Cores 64 and 65), are interbedded thin layers (10 to 20 centimeters thick) of "chert", siliceous limestone, and siliceous marly limestone.

Core recovery for the lower part of Unit IV (Cores 67 and 68) is not good. The lower part is composed largely of interbedded limestone and chert, very hard and massive to parallel thin-laminated. Color is very variable: light reddish brown, to very dark grayish brown, olive gray, greenish gray, and dark greenish gray. Visual discrimination of siliceous limestone and chert is very difficult.

A thin interbed of greenish gray graded sandstone (about 10 centimeters thick) is present in Core 68-1. It begins as a coarse grained sandstone just above a sharp basal contact (Core 68-1, 95 cm), and passes upward to faintly laminated to massive sandy limestone and well-laminated nanno limestone, successively. The sandstone contains many reworked foraminiferal tests of *Asterocyclina* cf. *penuria*, *Operculinoides* sp., and smaller type foraminifers of benthic and planktonic forms.

#### Unit V

Unit V is essentially characterized by abundant terrigenous components, in contrast to the overlying units. Although siliceous and calcareous biogenic components are present in considerable amount in the upper part, they decrease downward; on the contrary, terrigenous components of various grain size gradually increase and become coarser downward, and the lower part consists only of conglomerate and sandstone. This makes it possible to subdivide this unit into three sub-units.

The dominant lithology of unit V is mudstone, sandy mudstone and limestone (sub-units Va and Vb) and conglomeratic sandstone (sub-units Va to Vc), greenish gray to dark greenish gray in color.

The lithological suite of mudstone, sandy mudstone, and limestone shows very thin-bedded or laminated sedimentary structure in general. Different-sized grains, very fine sand to silt and mud, repeatedly occur in laminae of 1 to 2 cm. Also, the extensive occurrence of thin-bedded T(d,e) sequences characterizes this lithologic suite. Small-scale slump folds are rarely developed, and microconvolutions and microfaults are common features.

Conglomeratic sandstone is developed throughout unit V. It occurs in sub-units Va and Vb as interbeds, thickness ranging from about 5 to 120 cm. On the other hand, it makes massive, very thick beds in sub-unit Vc, which has little finer sediment.

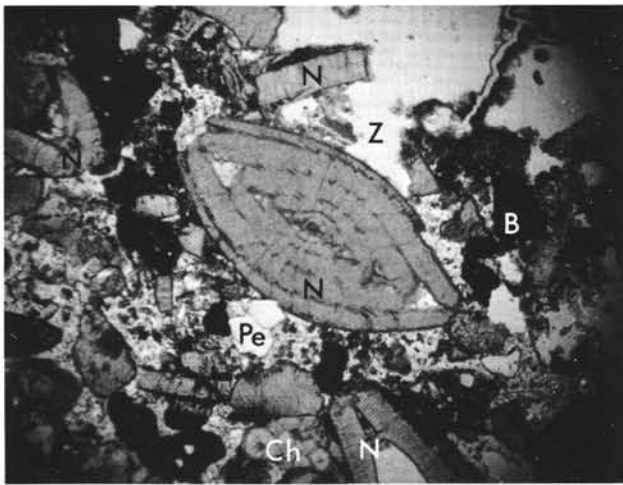
Conglomerate is of debris flow origin and consists of variously sized clasts and calcareous fossil tests, mainly of larger foraminifers, with sandy grains and small amounts of muddy matrix of nannofossil-rich mudstone (Figure 6).

Most clasts are chaotically arranged, but tend to fine upward from pebble and granule size to granule and very coarse-sand size. Clasts are relatively larger in sub-unit Vc, attaining diameters up to 15 cm. Clasts and

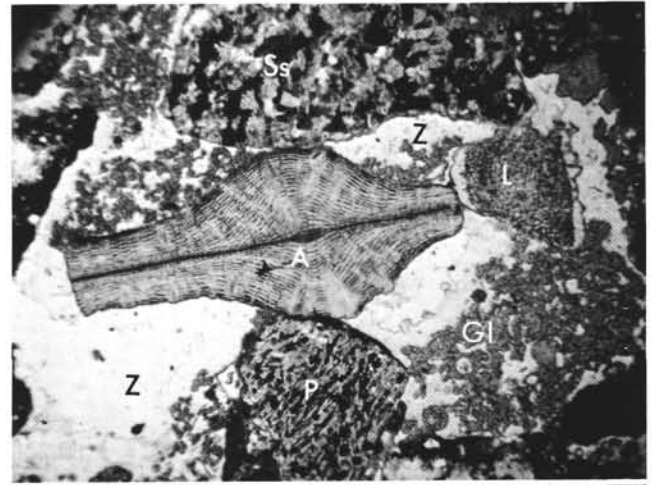




a



b



c

Figure 6. A. Photograph of slice sample of Nummulites-bearing conglomerate (Core 69-2, 64-68 cm; sub-unit Va). White spots are Nummulites tests, and dark spots are largely basalt and other igneous rocks. B. Nummulites-bearing conglomerate, showing tests of megalospheric form of Nummulites boninensis Hanzawa (Core 74-3, 59-61 cm, sub-unit Va). Besides Nummulites boninensis tests (N), fragments of basalt (B) and plagioclase (P), and highly altered, chloritized volcanic clasts with oolitic structure (Ch) are scattered in cementation materials of zeolite (Z), shown by white fills. Lower nicol only. C. Nummulites-bearing conglomerate, showing tests of Asterocyclina sp. cf. A. penuria Cole (Core 69-2, 8-12 cm; sub-unit Va). Asterocyclina sp. cf. A. penuria (A), clasts of sandstone (Ss), basalt (B), and limestone (L), and fragmental grains of glauconitic nannofossil clay or tuff (G1) are filled with zeolite (Z) shown by white color. Lower nicol only.

sand grains have various colors of green, brown, red, white, and black, and are variously shaped, angular to rounded. Under the microscope, the following rocks and minerals were distinguished: plagioclase phyric basalt (dominant), aphyric basalt (common), microdolerite (common), hornblende schist (rare), chert, sandstone, limestone, plagioclase (common), green hornblende (common), titaniferous augite (common), brown hornblende (rare), augite (common), olivine pseudomorphs (rare), chromian spinel (picotite) (common), epidote (rare). Heavy minerals (Sato, this volume) in conglomerate and sandstone include common hornblende and augite of volcanic origin, some minerals of schists, such as bluish-green amphibole and epidote, and chromite, which may have been derived from ultrabasic rock.

Calcareous fossil tests are megalospheric and microspheric forms of *Nummulites boninensis* (very abundant), *Asterocyclus* sp. cf. *A. penuria*, and *Operculinoides* sp., besides fragments of bryozoan and *Ostrea*. Microspheric forms of *N. boninensis* attain diameters of about 3 cm, but most of the tests are broken. *Nummulites* and *Asterocyclus* are identical with those reported from sea bottoms of the Oki-Daito Ridge, the Daito Ridge, and the Amami Plateau at depths of 1500 to 2300 meters by Mizuno and Konda (1977) and Mizuno et al. (1977).

Both lithologic suites locally develop the perfect Bouma sequence, T(a-e): in Core 71-1, conglomeratic sandstone forms the base of interval a and gradually passes upward to interval b, consisting of parallel-laminated sandstone, the interval c consisting of very fine-grained sandstone with microconvolution, and the interval d of nannofossil limestone with weak parallel lamination. This sequence ends in the uppermost interval e, composed of massive, muddy nannofossil limestone (71-1, 19-89 cm).

A primary sedimentary structure common to both lithologic suites in unit V is inclined bedding. In most cores and sections, beds dip at 5 to 10° or more. The direction of dip frequently changes from core to core and (or) from section to section. This structure may have resulted from frequent resedimentation by large-scale submarine slumping or sliding.

#### Sub-Unit Va

This sub-unit is dominantly thin-bedded and laminated, greenish-gray to dark-greenish-gray nannofossil mudstone, sandy mudstone, and nannofossil limestone, with interbeds of conglomerate and sandstone. In the upper part, mudstone, sandy mudstone, and limestone are often rich in partly dissolved radiolarians. Volcanic glass is usually present in amounts up to 3 per cent, but very locally reaches more than 10 per cent in the upper part.

Conglomerate and sandstone interbeds containing larger foraminifers are distributed throughout, but beds more than 50 cm thick occur in the middle to upper part (Cores 69-78).

#### Sub-Unit Vb

This sub-unit consists of dark-greenish-gray sandy mudstone, mudstone, and conglomerate. The general sedimentary features of sandy mudstone and mudstone are very similar to those of sub-unit Va, except for a general increase of sandy particles in mudstone. They are devoid of siliceous fossils and relatively poor in calcareous fossils. Conglomerate occurs as interbeds of 10 to 70 cm in the upper part and the lowest part.

#### Sub-Unit Vc

This sub-unit is exclusively dark-greenish-gray conglomerate, sandstone, and muddy sandstone. The upper part (Cores 90-92) consists of alternating conglomerate, sandstone, and muddy sandstone, whereas the lower part (Cores 93-94) consists of thick beds of conglomerate with larger clasts (up to 15 cm in diameter) and irregularly interbedded, ill-sorted sandstone. They show slump structures throughout.

### ORGANIC GEOCHEMISTRY

Organic-carbon and nitrogen contents were measured for 74 sediment samples. Results of the analyses are reported elsewhere (Waples and Sloan, this volume) and are plotted in Figure 7. In the upper part of the section (0-525 m sub-bottom depth), where virtually all the samples are pelagic (biogeneous carbonate represents more than 80% of the sediment), the organic-carbon and nitrogen profiles are very similar to those reported for Sites 442, 443, and 444 for hemipelagic sediments. Organic-carbon and nitrogen values decrease steadily from the sediment water interface (0.27 and 0.38%, respectively) to a depth of about 100 meters, below which they remain constant at about 0.06 and 0.015 per cent, respectively, throughout the pelagic-sediment interval.

Below 600 meters, the regularity of the organic-carbon profile is broken by anomalously high contents in many sediments. Most of the sediments from 600 to 885 meters are silty (~25% silt-size particles) or sandy (~15% sand-size particles), and represent reworked sediments deposited during episodes of slumping and (or) turbid flow (Figure 8). It is likely, therefore, that the organic debris contained in these sediments is also reworked.

Benthic foraminifers throughout this interval indicate that the original depositional environment was relatively shallow, conditions favoring the input of relatively greater amounts of terrigenous organic debris.

The C/N ratios for the silty material also support a terrestrial origin for much of the organic material. As grain size increases toward silt, there is a dramatic increase in the C/N ratio. Because terrestrial plants generally are much depleted in nitrogen compared to aquatic organisms, the organic material in the silt fraction probably is terrestrial. The C/N ratios of the sandy samples were about the same as for clays. This indicates

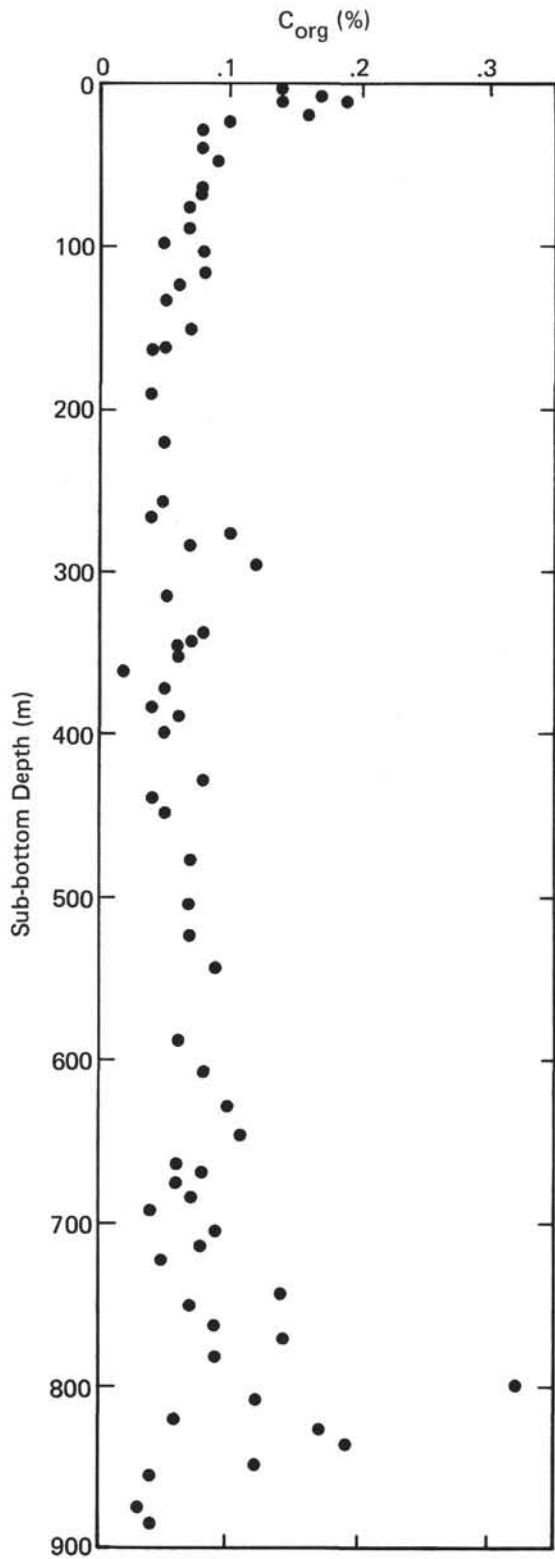


Figure 7. Per cent organic carbon versus depth (all sediments), Site 445.

that the sands did not bring much organic material with them; the small quantities of organic matter associated with these sediments are probably adsorbed on the clay-sized particles.

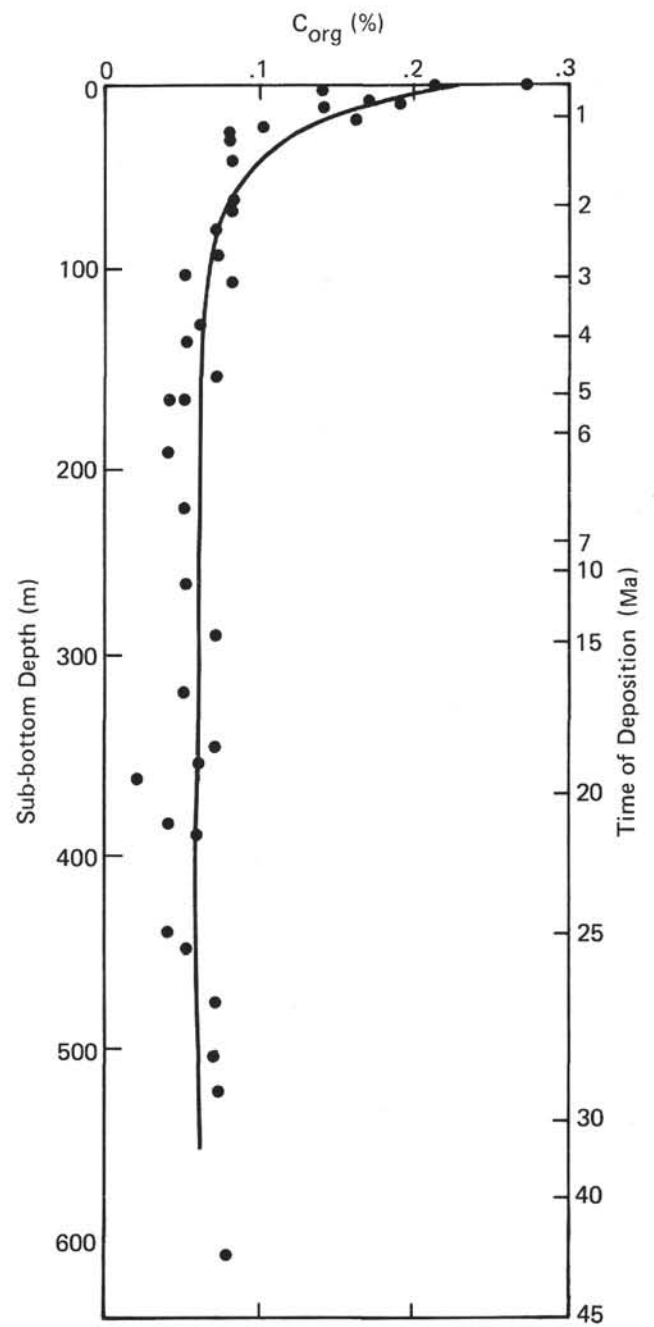


Figure 8. Per cent organic carbon versus depth (pelagic sediments), Site 445.

### INORGANIC GEOCHEMISTRY

From Hole 445, 14 samples were taken for interstitial-water studies. The data are presented in Table 2 and on Figure 9.

#### pH

pH averages 7.46, which is lower than the values for the IAPSO and surface-sea-water standards. pH varies somewhat down-hole, but does illustrate a trend to increase from the surface down to 296.5 meters, and from



TABLE 2  
Summary of Shipboard Geochemical Data for Hole 445

Sample (interval in cm)	Sample Number	Sub-Bottom Depth (m)	pH	Alkalinity (meq/kg)	Salinity (‰)	Ca <sup>++</sup> (mmol/l)	Mg <sup>++</sup> (mmol/l)	Cl <sup>-</sup> (‰)
	IAPSO		7.99	2.39	35.2	10.55	53.99	19.375
	SSW		8.32	2.41	35.2	10.48	53.29	19.444
445-1-4, 144-150	28	5.94-6.00	7.57	3.05	35.2	11.20	51.45	19.547
6-5, 144-150	29	53.94-54.00	7.36	3.33	35.5	15.83	43.40	19.822
11-4, 144-150	30	99.94-100.00	7.54	1.66	35.5	17.05	38.75	19.856
17-1, 140-150	31	152.40-152.50	7.21	0.87	35.2	20.58	33.09	19.925
22-1, 144-150	32	199.94-200.00	7.40	1.80	35.2	22.93	32.34	19.822
27-4, 140-150	33	251.90-252.00	7.38	1.76	36.3	25.46	30.22	19.959
32-2, 140-150	34	296.40-296.50	7.42	1.50	35.2	26.90	29.31	20.200
37-4, 140-150	35	346.90-347.00	7.18	1.70	34.6	27.86	29.59	19.512
42-5, 90-100	36	395.40-395.50	7.32	1.58	35.5	30.74	25.83	19.993
47-4, 90-100	37	441.40-441.50	7.37	1.31	36.3	35.53	21.51	20.062
52-4, 90-100	38	488.90-489.00	7.58	1.40	36.3	37.54	21.21	20.200
62-4, 140-150	39	584.40-584.50	7.63	0.68	36.3	39.56	21.91	20.097
65-2, 90-100	40	609.40-609.50	7.44	1.39	36.3	41.79	21.45	20.371
80-4, 90-100	41	754.90-755.00	8.08	0.36	36.3	45.84	26.14	20.406

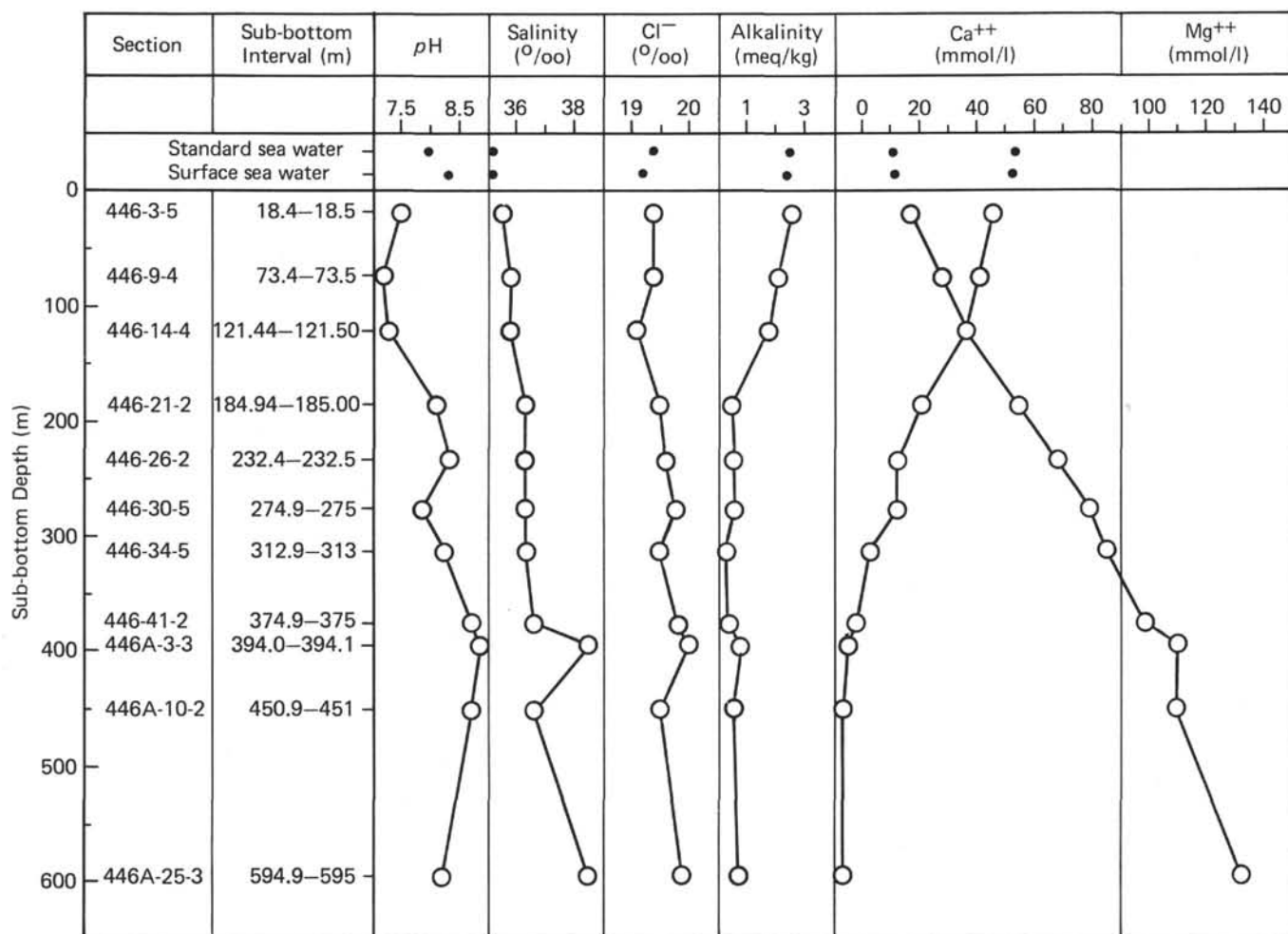


Figure 9. Interstitial-water geochemistry, Site 445.

346.9 to 755 meters. The separation of these two trends approximates the boundary between sub-units IIB and IIC.

#### Alkalinity

Alkalinity averages 1.6 meq/kg, lower than values for the two standards. Alkalinity shows little variation down-hole, except for a fairly steady decrease in values with increasing depth.

#### Salinity and Chlorinity

Salinity averages 35.7 per mill, and chlorinity 19.96 per mill, both averaging higher than the IAPSO and surface-sea-water standards. Both parameters increase with increasing depth, and, as expected, the trends of the two parameters correlate.

#### Ca<sup>++</sup> and Mg<sup>++</sup>

Ca<sup>++</sup> averages 24.5 mmol/l, higher than the standard values of 10.55 and 10.48. Mg<sup>++</sup> averages 30.44 mmol/l, which is lower than the standard values of 53.99 and 53.29.

Ca<sup>++</sup> illustrates a definite trend of increasing with depth, whereas Mg<sup>++</sup> generally decreases with depth. The crossover for the two trends occurs in Core 37 (346.9–347.0 m), which marks the sub-unit IIB/IIC boundary of very pale-brown and pale-brown, clayey nannofossil chalks; siliceous marly chalk; and white pinkish-gray and very pale-brown, clayey nannofossil chalks and nannofossil chalks.

The decrease of Mg<sup>++</sup> with depth is interrupted in Core 52, below which Mg<sup>++</sup> values show a slight tendency to increase.

### BIOSTRATIGRAPHY

Quaternary to middle-Eocene sediments were recovered at Site 445 (Table 3). This site, at a water depth of 3377 meters, is bordered by ridges in three directions with heights of 1000 to 2000 meters. This topographic feature created continuous and heavy reworking throughout the entire sequence.

Because the sedimentary basin is well above the CCD, calcareous microfossils are abundant. The preservation of foraminifers and calcareous nannofossils proves that this site has been well above the CCD from the middle Eocene to the present. Although the history of subsidence is not indicated by present paleontological data, further study of benthic foraminifers may provide this information.

All paleontological evidence indicates a strong influence of tropical water during the middle Miocene and afterward. The relatively rare tropical nannofossils of the early Miocene and Oligocene periods may indicate slight cooling from the early Oligocene to the early Miocene at this site; also, foraminifers are smaller, and microforaminifers dominate the fauna in the same sequence.

The Pliocene/Pleistocene boundary was identified in Core 6 by foraminifers and nannofossils. Although a well-preserved modern radiolarian fauna was observed

in Cores 1 and 2, radiolarians were not preserved in lower cores until the middle Miocene.

Foraminifers and nannofossils are abundant and well preserved in the Pliocene. Both the floral and faunal assemblages indicate Core 16 as the Miocene/Pliocene boundary.

Cores 17 to 28 represent the late Miocene and the middle and late early Miocene was recovered in Cores 29 to 35. The first overgrowth of nannofossils is in the upper Miocene. The sediment became too hard for quick examination of foraminifer assemblages in the lower cores, except for sporadic soft layers. Therefore, the age determination almost exclusively relied upon nannofossils in the early Miocene and older sediments.

The Oligocene/Miocene boundary was identified in Core 41. The middle and late Oligocene was represented by Cores 41 to 57, whereas the entire early Oligocene and the early middle Oligocene, representing at least 8 m.y., are represented by less than 15 meters of sediment. Therefore, a hiatus is suspected in the late Eocene to early middle Oligocene.

In the Oligocene and Eocene, which are mainly turbidite deposits, reworked fossils such as *Nummulites*, "larger" foraminifers, bryozoans, echinoid spines, and fragments of mollusks are common.

The abrupt appearance of a late-Eocene nannofossil assemblage was observed at Core 59-3. The top of the Eocene may be missing at this site. Radiolarians are abundant but not well preserved in Cores 63 to 65, and absent below Core 65.

Cores 66 through 89 represent the later middle Eocene. Both foraminifers and nannofossils confirm the age of the lower portion of this sequence as middle Eocene (47–48 m.y.).

#### Foraminifers

Continuous coring at Site 445 provided a foraminifer sequence from Neogene (N.23) through Paleogene (P.11/P.10; middle Eocene), suitable for biostratigraphic and environmental study.

In general, the foraminifers are abundant in the Pleistocene and Pliocene and decrease in number, size, and diversity from the Miocene to middle Eocene. They show no dissolution.

Although deposition is estimated to have been well above the CCD throughout the section encountered at this site, for some time spans foraminifers are rare or absent, and in some sections represented only by juveniles and microforaminifers. Therefore, fluctuating climate may have controlled this situation.

Only core-catcher samples were used for shipboard study, because rocks from the lower Miocene downward were so lithified (by foraminifer-study standards) that processing the material was time consuming.

Table 3 is a summary of the zonation of Hole 445.

In Cores 1 through 5, planktonic and benthic foraminifers are abundant and well preserved; they indicate a Pleistocene age and deposition well above the CCD.

Cores 6 through 15 are Pliocene, and again the condition of the entire foraminifer fauna indicates that

deposition was above the CCD. In Core 16, the Pliocene/Miocene (N.18; ~5 m.y.) boundary was detected; because this is defined by evolutionary appearances and concurrent ranges, precise definition must await study of the section samples. Cores 17 to 41 are Miocene. Throughout this interval, normal-sized foraminifers are very rare, and the index species used for the very short middle-Miocene zonation recognized in other parts of the world (i.e., Caribbean) were not seen.

Although the foraminifers are mostly in the microforaminifer size range, occasional normal-sized benthic forms indicate an open-sea ( $\geq 500$  meters) environment.

Fossils extracted from the indurated Oligocene are tiny and moderately well preserved. Based on the nanofossils, the Oligocene is estimated to range from within Core 41 to Core 59, Section 4. The Oligocene identified foraminifer zones are within Core 49 through Core 57, approximately 26.5 to 36+ m.y. In Cores 52 and 53-4, 68-70 cm, *Nummulites*, one of the "larger" foraminifers, was recovered from a greenish fragmented sandstone. Core 53 also contained other shallow-water foraminifers, such as *Baculogypsina*, *Gypsina*, *Sphaerogypsina*, and *Rupertina*. In the rubble, Eocene planktonic forms were also found, which indicate reworking in this section. The washed residue of 57,CC is ash, but it contained many planktonic foraminifers of Eocene (?) age, encrusted with ash.

In Core 76, from 711.5 to 721 meters, the middle-Eocene species *Globorotalia spinulosa* was encountered. This sample is dated approximately 44.5 to 45 m.y. Cores 79, 80, and 85 are also middle Eocene. The tentative foraminifer zonation shown in Table 3 indicates that sediments as old as 48 to 49 m.y. were penetrated in this hole.

Both the megalospheric and microspheric generations of *Nummulites* occur in abundance in many of the Eocene cores from Core 60 downward. These may be related to the species *Nummulites boninensis*, described by Hanzawa in 1947 and recognized by Mizuno and Konda (1977) in dredge hauls near Daito and Oki-Daito Islands. The rocks in which these *Nummulites* occur also contain other shallow-water elements such as the benthic genus *Amphistegina* s.l. and fragments of bryozoans, pelecypods, echinoid spines and *Ostrea*. The fauna is undoubtedly reworked. However, the unbroken condition of the tests of the foraminifers may indicate a short distance of transport, therefore suggesting that this site is close to the source.

#### Nannofossils

Upper-Pleistocene to middle-Eocene nannofossils occur at this site. Nannofossils are abundant in all cores except the middle-Eocene sequence, in which they are sporadic. Because of heavy and continuous reworking, however, dating of cores was difficult in some intervals. All reworked specimens observed at this site represent a few zones prior to the time of redeposition, with a maximum time difference of several million years. Nannofossils are well preserved in the upper sequence (upper Miocene and above); only slight etching, without

any sign of overgrowth, is recognized. Slight etching and moderate to heavy overgrowth, on the other hand, prevail in the lower cores. Recrystallization of nanofossils is strongest in the Oligocene and becomes relatively weak in the middle-Eocene turbidites. The age identification of cores is summarized in Table 3.

#### Pleistocene

Sections 1 to 5 of Core 1 belong to the *Emiliania huxleyi* Zone, whereas Section 1,CC contains an assemblage of the *Ceratolithus cristatus* Subzone. Reworked fossils are mostly Pliocene to early Pleistocene forms, but much older species, such as *Cyclicargolithus floridanus*, are also observed. Sections 2-2 to 3-5 belong to the *Pseudoemiliania lacunosa* Subzone, and a subtropical assemblage of the *Crenolithus doronicoides* Zone occurs in Section 4-1 to 6-3. The Pliocene/Pleistocene boundary was identified in Section 6-4.

#### Pliocene

Section 6-5 to 12-4 belong to the late Pliocene (*Discoaster brouweri* Zone). The disappearance of *Discoaster surculus* identifies the base of the *D. surculus* Subzone at Section 9,CC. Heavy reworking prevents subdivision of this sequence. The early-Pliocene cores, however, are recognized at subzone levels. Core 13 belongs to the *Discoaster asymmetricus* Subzone, whereas Cores 14 and 15 represent the *Sphenolithus neoabies* Subzone. The *Coratolithus rugosus* Subzone is identified in Sections 16-2 to 16-4. The *C. acutus* Zone occurs in Cores 16,CC to 18-2, and the Miocene/Pliocene boundary lies within this interval.

#### Miocene

Miocene assemblages of abundant nanofossils occur in Cores 18-4 through 41-1. Section 18-4 belongs to the late Miocene (*Triquetrorhabdulus rugosus* Subzone). Cores 18,CC to 25-3 belong to the *Discoaster quinqueramus* Zone, and the first occurrence of *Amaurolithus primus* identifies the base of the *A. primus* Subzone in Section 22-2. The preservation of nanofossils is good until Core 23; it becomes mostly moderately good in the Miocene sequence. The *Discoaster neohamatus* Zone occurs in Cores 25-5 to 27-5, and Cores 27,CC to 28-4 belong to the *Catinaster calyculus* Subzone. Because *Discoaster hamatus* occurs only sporadically, the base of the *Helicosphaera carter* Subzone is not clear. The base of the *Catynaster coalitus* Zone is recognized in Section 29-4 by only the occurrence of *Discoaster kugleri* in 29,CC. Cores 30-2 to 31-1 belong to the *Coccolithus miopelagicus* Subzone of early middle Miocene. Cores 31-3 to 35,CC represent the *Helicosphaera ampliaperata* Zone or the *Sphenolithus heteromorphus* Zone. The sporadic and rare occurrence of *H. ampliaperata* and reworking hampers identification of the boundary between these two zones. Similarly, the absence of *Sphenolithus belemnoides* prevents division of the early-Miocene cores between sections 36-2 and 41-1. These cores belong to the *Discoaster druggi* Subzone or *S. belemnoides* Zone.

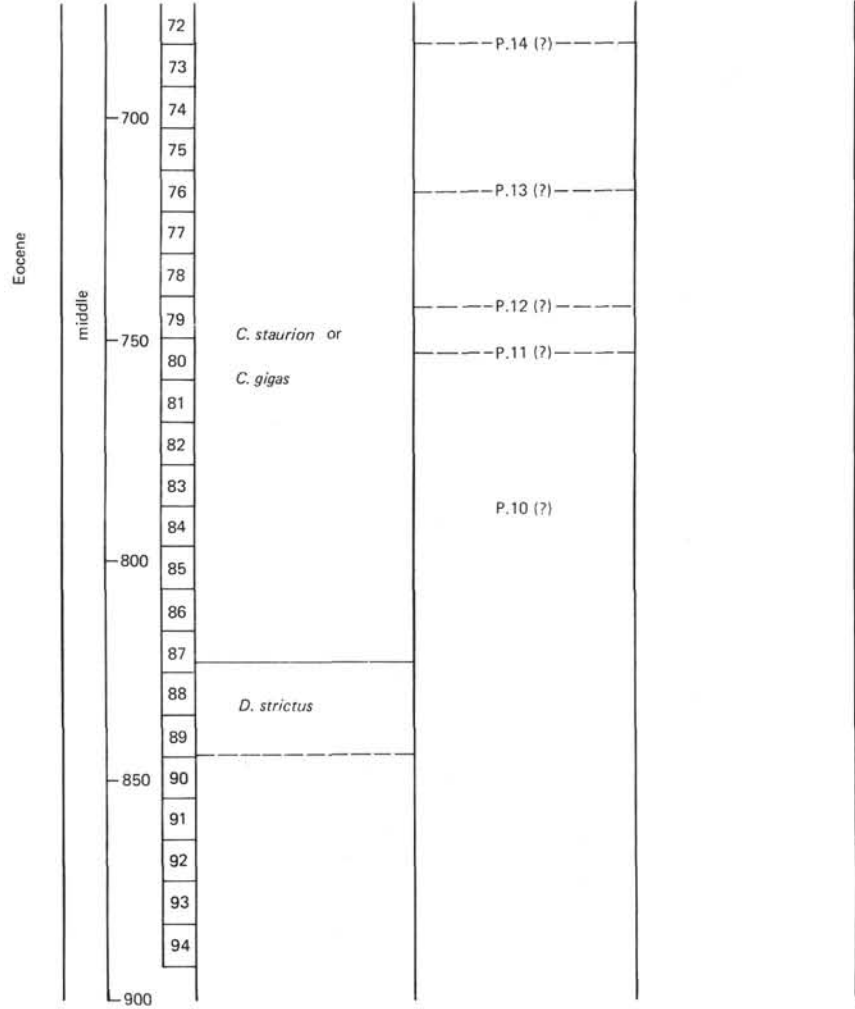
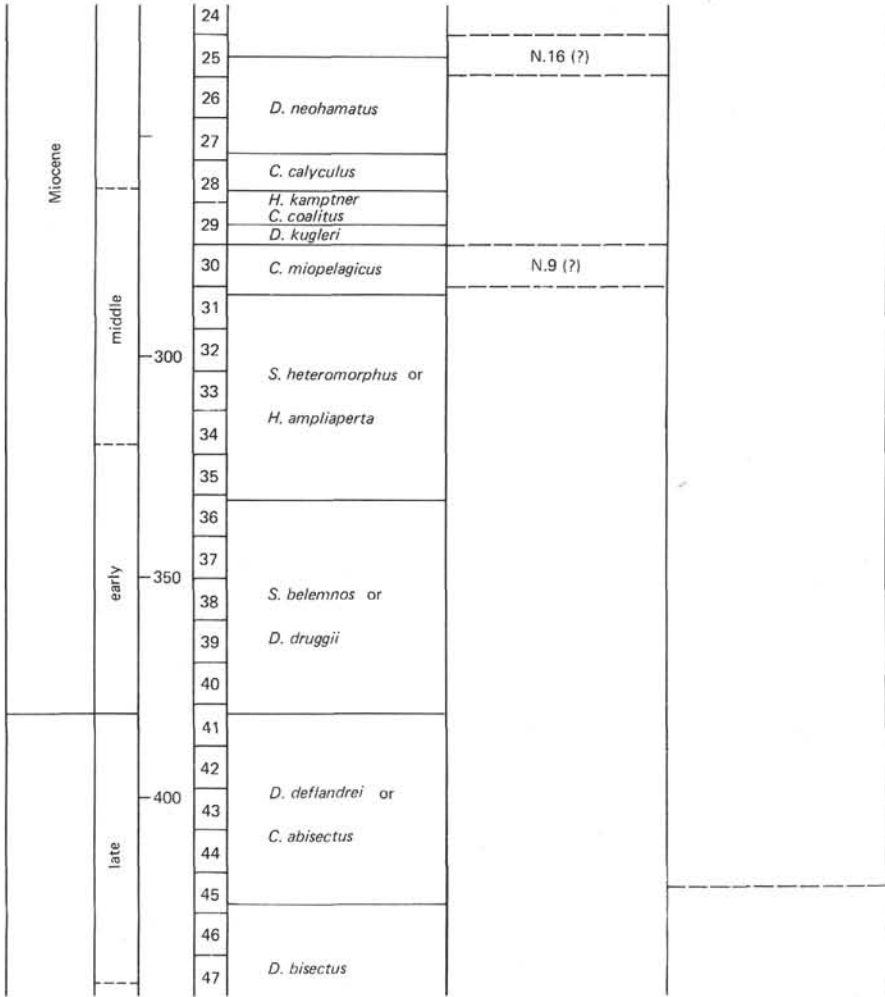


TABLE 3  
Biostratigraphic Zones, Site 445

Age	Depth (m) and Core No. 445	Nannofossil Zones and Subzones	Foraminifer Zones	Radiolarian Zones
Pleistocene	1	<i>E. huxleyi</i>	N.23	<i>L. haysi</i>
	2	<i>C. cristatus</i>		
	3	<i>P. lacunosa</i>		
	4	<i>C. doronicoides</i>	N.22	
	5			
	6	<i>C. macintyeri</i>	N.21	
7				
8				
9				
10	<i>D. surculus</i> or <i>D. tamalis</i>	N.19 (?)		
11				
12	<i>D. asymmetricus</i>	N.18 (?)		
13	<i>S. neoabies</i>			
14	<i>C. rugosus</i>			
15	<i>C. acutus</i>	N.17 (?)		
16	<i>T. rugosus</i>			
17	<i>A. primus</i>	N.17 (?)		
18				
19				
20	<i>D. berggrenii</i> (?)	N.17 (?)		
21				
22				
23				

TABLE 3 - Continued

Age	Depth (m) and Core No. 445	Nannofossil Zones and Subzones	Foraminifer Zones	Radiolarian Zones
Oligocene	middle	48	N.3/P.22 (?)	
		49		
		50		
		51		
		52		
		53		
	early	54	N.2/P.21 (?)	
		55		
		56		
		57		
		58		
late	550	58	P.17/P.18 (?)	
		59		
	600	59	D. saipanensis	
		60		
		61		
		62		
		63		
		64		
		65		
		66		
		67		
		68		
		69		
70				
71				



### Oligocene

Cores 41-3 to 45-3 are assignable to the latest Oligocene or the earliest Miocene (*Cyclicargolithus abisectus* Subzone or *Discoaster deflandrei* Subzone). Cores 45,CC to 52-1 belong to the *Sphenolithus ciperoensis* Zone, and Section 47-3 represents the base of the *Dityococcites bisectus* Subzone. Although rare, the consistent occurrence of *Sphenolithus distentus* identifies Cores 52-3 to 57,CC with the *S. distentus* Zone.

An assemblage of the early early to early middle Oligocene (*Helicosphaera reticulata* and *Sphenolithus predistentus* Zones) occurs in Cores 58-2 to 59-3. Because of heavy reworking, detailed age identification is impossible for this interval. Considering the length of time (8 m.y.) represented by this short sequence of sediment, a hiatus is suspected. The sudden increase of *Reticulofenestra umbilica* and *Discoaster saipanensis* below Sections 58-2 and 59-4 respectively suggests hiatuses at the top and at the base of this sequence. Although rare, *Isthmolithus recurvus* occurs consistently through this sequence and in the uppermost Eocene cores. This species is seldom preserved in low-latitude area, whereas it is common in high-latitude areas (Bukry, 1975). At nearby Sites 290, 291 and 292, this species was not observed. The cause of the sporadic occurrence of this species at this site is not clear at present.

### Eocene

The first occurrence of *I. recurvus* in Section 60,CC identifies the base of the upper-Eocene *I. recurvus* Subzone. Although *Chiasmolithus oamaruensis* does not occur at this site, extinction of *C. grandis* indicates the base of the *C. oamaruensis* Subzone in Section 63-1. Cores 63-3 to 65-1 and 66-1 to 71-4 belong to the *D. saipanensis* and *Discoaster bifax* Subzones, respectively. Below Core 65, hitherto ubiquitous nannofossils are scarce and sporadic. The *Chiasmolithus gigas* and *Coccolithus staurion* Subzones of the middle middle Eocene are identified in Cores 72-1 to 87-2. The turbidites of Cores 88 and 89 contain a moderately well-preserved assemblage of the *Discoaster strictus* Subzone (47.0–48.0 m.y.). The oldest reworked fossils found in these cores are *Discoaster lodoensis* and *D. sublodoensis*. Both species are considered to have become extinct about 48.0 Ma. Cores 90 and 91 contain a few poorly preserved nannofossils; the assemblage does not justify age identification, and Cores 92 to 94 are barren of nannofossils.

### Radiolarians

Radiolarians are sporadic at Site 445. Only Cores 1 and 2 had well-preserved and abundant radiolarians. Radiolarians in any abundance are again encountered only in Cores 63 through 65, but they are not well preserved.

Cores 1 and 2 are Pleistocene to Holocene and contain species which are still living in today's seas (*Spongaster tetras*, *Euchitonia furcata*, and *Ommatarus tetrathalamus*). Cores 3 through 30 are barren, and Cores 31 through 34 have a few middle-Miocene species,

with slight lower-Miocene reworking. Many cores from 35 to 58 are barren, but some contain abundant unidentified fragments of radiolarian spines and lattices.

Below the chert, the sediments are coarse-grained (sandy mudstones grading into conglomerates), and radiolarians are not preserved in this type of sedimentary regime. The moderately well-preserved radiolarians in Cores 63 through 65 contain species of the Eocene and Oligocene genera *Eusyringium*, *Podocyrtes*, *Dorcadopyris*, and *Theocyrtes*. No radiolarian zone can be assigned to Core 63, but Cores 64 and 65 belong to the late Eocene (*Podocyrtes chalara* Zone). There are no radiolarians preserved below Core 65.

### SEDIMENTATION RATE

An age–depth plot is shown in Figure 10. The ages of the sediment were obtained using the time-scales of Berggren (1972), Berggren and Van Couvering (1974), and Bukry (1975), and the modified Miocene time scale of Saito (1977). Table 4 shows sediment accumulation rates calculated for each stratigraphic unit.

The sediment accumulation rates show a systemic change down-hole. They are moderate to moderately high for the Pleistocene and Pliocene foraminifer-nannofossil oozes of unit I, and moderate for the Oligocene and Miocene chalks of unit II. The rates are lowest 9.3 during the early Oligocene and late Eocene, during deposition of the radiolarites and chert of unit III; they increase drastically in sandstones, conglomerates, and mudstones of unit V (early late and middle Eocene). The sediment accumulation rate curve shows remarkable similarity to the trend of the accumulation rate curve for Site 286 in the Hebrides Basin (Andrews, Packham, et al., 1975), in particular, and both Sites 285 and 286 in general (Klein, 1975). There, the high rates of accumulation were characteristic of debris-flow conglomerates and turbidite sandstones; the lowest rates were characteristic of biogenic, pelagic carbonate oozes, and intermediate and moderate rates were characteristic of resedimented nannofossil oozes.

The explanation for the sediment accumulation rate pattern at Site 286 is the same for Site 445. The highest rates of sediment accumulation are for the turbidite sandstones and the pebble to granule gravel conglomerates of slump and debris-flow origin of unit V. The lowest rates are characteristic of the unit IV radiolarites and cherts. Intermediate values are characteristic of units I, II, and III, where evidence of resedimentation by subaqueous gravity processes is common. That evidence includes preserved slump blocks, folds and faults, and graded foraminifer sands organized into partial Bouma sequences [T(a,b,e)] produced by turbidity currents. The slight increase in sediment accumulation rates for the Pleistocene oozes may well reflect higher productivity of the Kuroshio Current, detected at Sites 296 and 297 by Karig, Ingle, et al. (1975).

### PALEOMAGNETISM

Site 445 is a single-bit hole with about 900-meter penetration into the sediment cover. The topmost 200

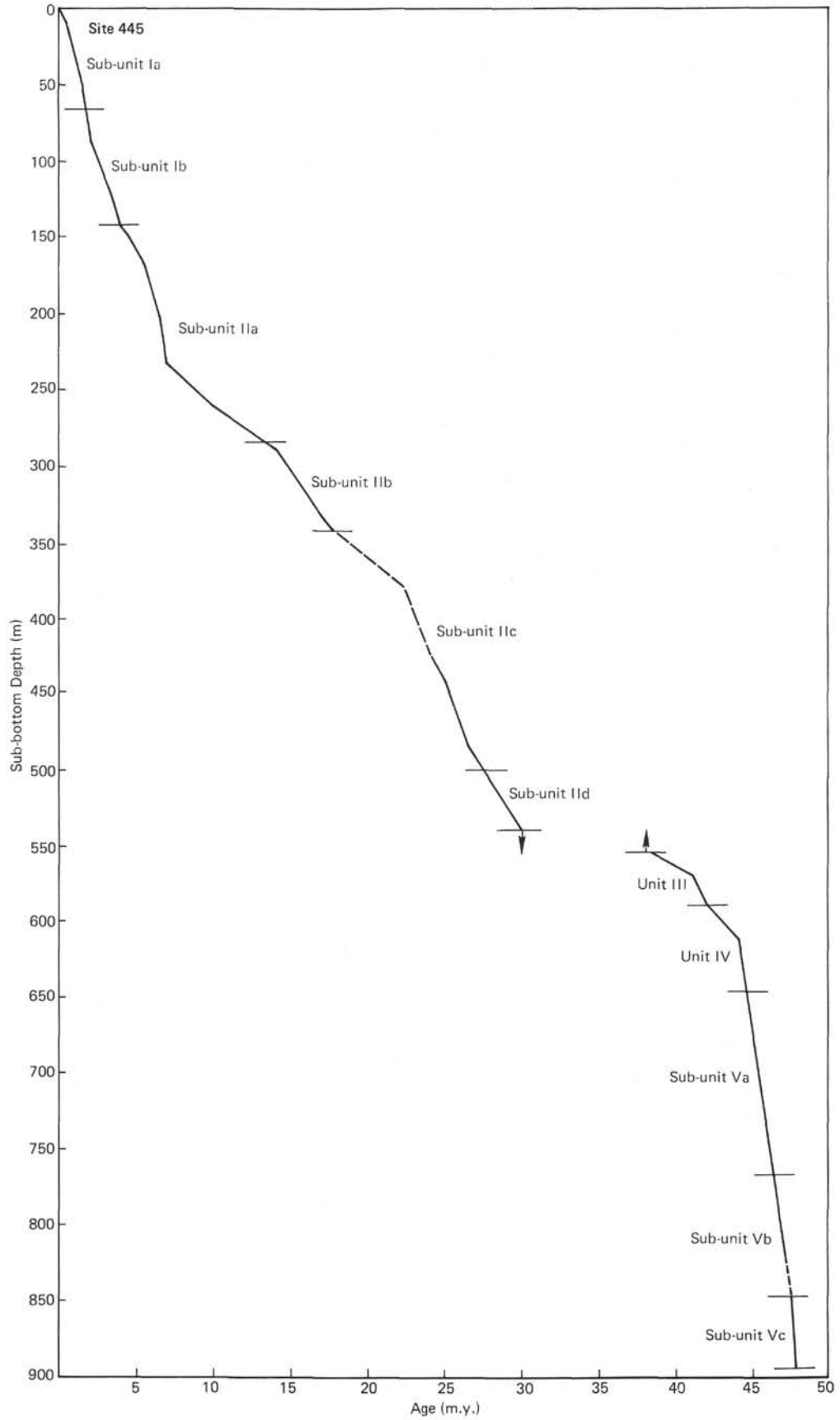


Figure 10. Sediment accumulation rate curve for Site 445, based on biostratigraphic age determinations.



TABLE 4  
Sedimentation Rates, Site 445

Unit	Depth (m)	Interval Thickness (m)	Sedimentation Rate (m/m.y.)
Ia	0.0-65.5	65.5	36.9
Ib	65.5-141.8	76.3	34.7
IIa	141.8-284.0	142.2	16.0
IIb	284.0-341.0	57.0	13.9
IIc	341.0-502.5	161.5	16.5
IId	502.5-551.7	49.2	18.9
III	551.7-588.9	37.2	9.3
IV	588.9-645.0	56.1	22.4
Va	645.0-768.5	123.5	72.6
Vb	768.5-844.5	76.0	95.0
Vc	844.5-892.0	47.5	237.5

meters consists of relatively soft sediments. In the other part, the sediments are lithified and mechanically stable. Paleomagnetism samples were taken on an average of 1.5 meters in the recovered cores. Three hundred and twelve of them were used for NRM and AF-demagnetized NRM measurements. Changes in NRM during AF treatment are significantly different from sample to sample, as shown in Figure 11. Because stepwise AF demagnetization for all of the samples was impossible within a limited time schedule, the AF demagnetization

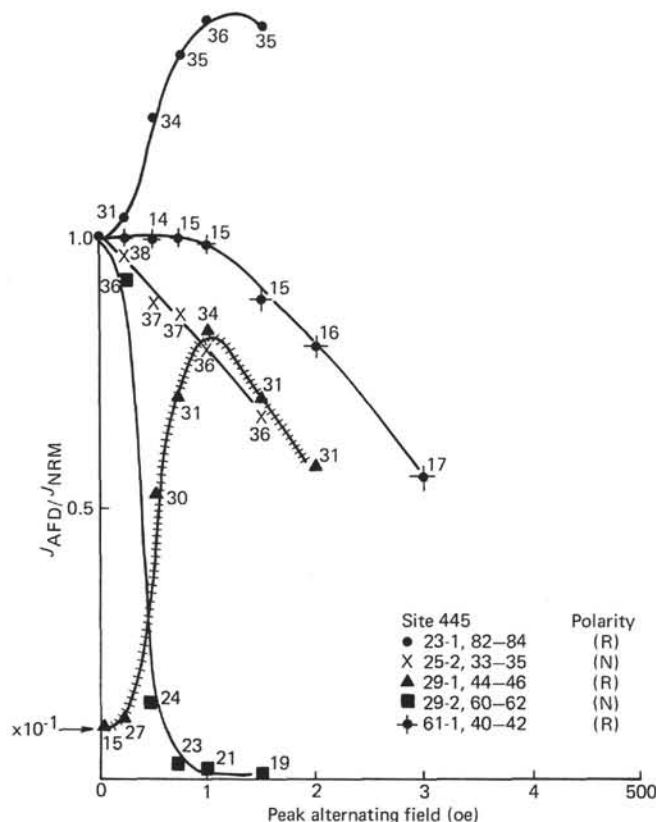


Figure 11. Change in sediment NRM intensity during AF demagnetization. Rounded numbers attached to each point represent absolute inclination values obtained after AF demagnetization.

was carried out in a 150-oe AF, decreasing to zero at a constant rate. Measuring and AF demagnetizing techniques are described in the Site 442 report. On account of a large number of samples available at this site, a rigorous examination of paleomagnetism stability was applied to distinguish paleomagnetically reliable samples. Examination was according to the following criteria: (1) intensity ratio of AF-demagnetized remanent magnetization to the initial NRM falls between 2 and 1/2; (2) angular shift of NRM during AF demagnetization is less than 5 degrees; (3) the Koenigsberger ratio is larger than 10.

Ninety-five samples of 312 were retained for further analyses. All the data on sample positions and measurements of remanent magnetization are listed in Table 5. Normal and reversed sequences of cores are diagrammatically shown in Figure 12. In the last part of Figure 12, Cores 68 through 90, the polarity column is left blank for two reasons. One, in the lowermost 200 meters of the sediment cover NRM inclination becomes smaller, and therefore the polarity change of the geomagnetic field cannot be determined through NRM inclination values alone. Two, this layer consists mostly of mudstones, and only a small portion of the recovered cores were found useful. In Figure 13A, absolute inclinations of AF-demagnetized stable NRM are plotted against sub-bottom depth. In Figure 13B, original NRM intensity values (mean for every 10 meters) are plotted against sub-bottom depth. AF-demagnetized absolute NRM inclination values are divided into four groups for 200-meter sub-bottom intervals, and statistical treatment was applied. Results are listed in Table 6. Latitude of the Daito Ridge is calculated upon the assumption that the mean virtual pole position is exactly equal to the present geographical pole position throughout the last 45 m.y.

Table 6 also includes the latitude of Site 445 in the past. The results are plotted in Figure 14. The position of the Shikoku Basin deduced from the NRM inclination data at Sites 442 through 444 is also shown in the figure for comparison. It seems likely that the scatter of data is significantly dependent on the mechanical strength of the sediment. Statistical analysis leads us to a tentative evaluation of a minimum distance of migration of the Daito Ridge during the last 45 m.y. Supposing that the ridge was situated at 5 degrees north latitude, due south of the present location, and that it continuously drifted northward to reach the present position, the distance of movement is about 2000 km (4.4 cm/yr). The other implication of Figure 13b is a high incidence of subaerial eruptions at this locality during the probable Eocene, represented by a thick mudstone layer in the bottom 200 meters of the hole. It is likely that the mudstone contains ferromagnetic minerals (probably titanomagnetite) more abundant by one order of magnitude than in the other sediments. AF demagnetization mode shows that the ferromagnetic particles are fine-grained, just as in the chilled margin of an oceanic basalt. This is possible when a tremendous amount of fine-grained materials falls to form a thick tuff layer. Actually, some of the tuffs and welded tuffs on land show a

TABLE 5  
Paleomagnetism Measurements of Sedimentary Cores from Site 445<sup>a</sup>

Sample (interval in cm)	Sub- bottom Depth (m)	$J_{\text{NRM}}$ (10 <sup>-5</sup> gauss)	Suscepti- bility (10 <sup>-5</sup> gauss/oe)	Inclination		Polar- ity
				NRM	AFD	
445-2-1, 86-88*	9.37	1.97	0.23	26.7	31.9	+
2-2, 116-118*	11.17	1.98	0.17	38.4	37.5	+
2-3, 114-116*	12.65	2.44	0.07	10.5	10.5	+
3-1, 15-17	18.16	0.27	0.11	-4.6	-1.4	-
3-2, 100-102	20.51	1.86	0.24	50.0	51.3	+
3-3, 76-78*	21.77	2.74	0.16	43.6	44.4	+
3-4, 57-59*	23.08	0.88	0.10	-31.1	-29.7	-
3-5, 21-23	24.22	0.68	0.21	-28.7	-28.5	-
3-6, 109-111*	26.60	1.14	0.17	-9.4	-14.9	-
4-1, 14-16*	27.65	2.06	0.22	22.6	25.1	+
4-4, 40-42	32.41	0.23	0.14	-45.5	-56.5	-
4-5, 52-54	34.03	0.46	0.19	-1.7	-22.2	-
4-6, 21-23	35.22	0.82	0.19	-19.0	-9.2	-
5-1, 93-95	37.94	0.94	0.22	-46.4	-45.1	-
5-2, 137-139	39.88	1.01	0.22	72.6	39.4	+
5-3, 11-13	40.12	0.99	0.25	-16.6	-23.5	-
6-1, 38-40	46.89	0.28	0.19	-20.8	-40.6	-
6-2, 137-139	49.36	0.46	0.17	74.3	30.5	+
6-3, 115-117	50.66	1.01	0.24	-1.8	-18.3	-
6-4, 86-88	51.87	0.13	0.24	-0.8	-29.2	-
6-5, 71-73	53.22	2.26	0.48	30.1	21.8	+
6-6, 79-81	54.80	1.21	0.31	16.1	4.2	+
7-6, 88-90	64.39	1.03	0.25	36.4	28.3	+
8-1, 138-140	66.89	0.57	0.26	-15.2	-15.6	-
8-2, 138-140	68.39	1.10	0.26	36.7	41.0	+
8-3, 138-140	69.89	0.05	0.25	45.3	60.4	+
8-4, 138-140	71.39	0.28	0.33	-50.1	-49.6	-
8-5, 138-140	72.89	0.51	0.29	59.0	-23.2	+/-
10-5, 144-146	91.95	1.27	0.32	55.6	51.1	+
11-2, 69-71	96.20	1.13	0.33	24.3	23.9	+
11-3, 80-82	97.81	0.11	0.33	-65.0	-80.2	-
11-4	<100	0.30	0.37	-31.1	-46.4	-
11-5	>100	0.73	0.36	-7.0	-11.6	-
12-1, 95-97	104.46	0.93	0.21	52.0	53.6	+
12-2, 95-97*	105.96	1.05	0.21	-82.6	-83.5	-
12-3, 95-97*	107.46	2.37	0.40	29.4	31.0	+
12-4, 95-97	108.96	0.76	0.26	31.7	29.7	+
13-1, 72-74	113.73	0.23	0.25	34.5	35.2	+
13-2, 72-74	115.23	0.049	0.34	37.6	-34.5	+/-
13-3, 72-74	116.73	0.51	0.23	43.2	47.5	+
13-4, 72-74	118.23	0.078	0.29	-40.0	-66.8	-
14-1, 58-60	123.09	0.010	0.16	-68.4	-50.8	-
15-1, 20-22	132.21	0.028	0.21	-17.6	-37.2	-
15-2, 20-22	133.71	1.39	0.31	-51.1	-49.8	-
16-1, 45-47	141.96	0.10	0.15	48.0	31.6	+
16-2, 45-47	143.46	0.79	0.23	22.8	13.0	+
16-3, 45-47	144.96	0.88	0.17	43.2	49.6	+
17-2, 24-26	152.75	0.40	0.16	73.6	69.7	+
18-1, 99-101	161.50	0.25	0.19	-0.0	-40.7	-
18-5, 21-23	166.72	1.31	0.17	42.2	24.0	+
19-2, 05-07	171.56	1.33	0.17	21.8	11.9	+
19-3, 05-07	173.06	0.44	0.17	21.9	-30.6	+/-
20-1, 100-102	180.49	0.35	0.18	50.8	-13.3	+/-
20-2, 100-102	181.99	0.16	0.17	30.7	-83.9	+/-
21-1, 134-136	190.35	0.19	0.19	-6.3	-38.6	-
21-2, 62-64	191.13	0.21	0.19	28.0	-73.5	+/-
22-1, 92-94	199.43	0.091	0.20	43.0	-3.5	+/-
22-2, 13-15	200.14	0.006	0.09	-9.8	-10.4	-
23-1, 82-84*	208.83	1.17	0.23	-35.1	-34.5	-
23-2, 68-70*	210.19	1.89	0.36	-52.6	-51.1	-
23-3, 103-105	212.04	0.64	0.26	44.4	41.8	+
24-1, 22-24	217.73	0.80	0.27	-40.9	-35.5	-
24-2, 16-18	219.17	1.73	0.26	51.2	43.8	+
24-3, 57-59	221.08	0.29	0.38	39.1	-20.1	+/-
24-4, 130-132	223.31	0.63	0.34	42.9	-18.2	+/-
24-5, 29-31	223.80	0.97	0.32	50.4	-2.3	+/-
25-1, 145-147	228.46	1.65	0.31	-42.0	-33.7	-
25-2, 33-35*	228.09	4.19	0.38	36.8	35.7	+
25-3, 97-99*	230.98	2.11	0.31	31.9	30.1	+
25-4, 111-113	228.87	0.14	0.42	-30.5	-40.3	-

TABLE 5 - Continued

Sample (interval in cm)	Sub- bottom Depth (m)	$J_{\text{NRM}}$ (10 <sup>-5</sup> gauss)	Suscepti- bility (10 <sup>-5</sup> gauss/oe)	Inclination		Polar- ity
				NRM	AFD	
445-25-5, 127-129	234.28	0.45	0.42	-20.6	-36.9	-
25-6, 59-61*	235.10	1.67	0.33	-23.6	-29.3	-
26-1, 139-141*	237.90	3.40	0.39	37.2	41.9	+
26-2, 140-142	239.41	1.72	0.40	33.7	32.6	+
26-3, 144-146	240.95	1.74	0.39	29.1	16.2	+
26-4, 145-147	242.46	0.47	0.27	-8.2	-33.1	-
26-5, 140-142	243.91	1.48	0.38	34.2	36.9	+
26-5, 60-62	243.11	0.46	0.31	17.1	74.7	+
27-1, 81-83*	247.57	3.96	0.50	35.0	31.2	+
27-2, 59-61*	248.10	1.71	0.31	-28.4	-29.5	-
27-3, 54-56	249.55	1.56	0.44	2.8	-17.7	+/-
27-4, 35-37	250.86	1.40	0.48	35.0	4.7	+
27-5, 50-52*	252.51	3.57	0.46	35.7	34.2	+
27-6, 41-43	253.92	0.48	0.19	10.5	-52.4	+/-
28-1, 128-130*	256.79	2.34	0.33	36.0	35.2	+
28-2, 106-108*	258.07	3.47	0.33	35.4	34.7	+
28-3, 147-149	259.98	0.03	0.27	37.6	35.0	+
28-4, 100-102*	261.01	3.19	0.50	34.2	32.1	+
28-5, 118-120	262.69	0.45	0.45	-3.0	-26.1	-
29-1, 44-46	265.45	0.11	0.40	-14.8	-31.2	-
29-2, 60-62	267.11	1.01	0.36	39.2	18.8	+
29-3, 55-57	268.56	0.98	0.40	28.4	7.9	+
29-4, 8-10	269.59	0.19	0.32	17.4	-9.3	+/-
30-1, 42-44*	274.93	2.13	0.43	38.0	31.2	+
30-4, 32-34	279.33	0.01	0.29	-57.3	-53.7	-
31-1, 22-24	284.23	1.98	0.44	37.3	31.9	+
31-4, 140-142	289.91	0.02	0.29	26.4	14.1	+
31-5, 136-138*	291.37	3.52	0.39	28.4	28.3	+
32-1, 16-18	293.67	2.07	0.34	27.7	10.2	+
33-2, 65-67	305.16	0.15	0.33	9.2	-4.5	+/-
33-3, 18-20*	306.19	1.93	0.34	-8.2	-13.0	-
33-5, 61-63*	309.62	2.20	0.27	26.9	23.7	+
34-2, 37-39	314.38	0.33	0.37	9.2	-21.6	+/-
34-4, 30-32	317.31	0.22	0.38	8.9	-39.1	+/-
34-5, 24-26	318.75	0.87	0.32	-32.4	-35.7	-
35-1, 12-14	322.13	4.86	0.44	27.0	19.2	+
35-2, 78-80	324.29	1.18	0.44	1.3	-14.3	+/-
35-3, 78-80	325.79	0.64	0.27	35.3	3.9	+
36-1, 85-87	332.36	1.41	0.47	-9.7	-30.1	-
36-2, 66-68	333.67	1.43	0.48	36.5	35.2	+
37-1, 144-146	342.45	0.75	0.36	62.2	6.7	+
37-2, 68-70	343.19	1.09	0.42	31.4	33.7	+
37-4, 23-25	345.74	0.012	0.71	38.2	21.3	+
38-1, 37-39	350.88	0.56	0.40	46.0	8.3	+
38-2, 37-39	352.38	0.26	0.40	74.2	-5.1	+/-
38-3, 39-41	353.90	0.90	0.30	39.1	50.9	+
38-4, 39-41	355.40	0.76	0.43	38.5	13.2	+
38-5, 70-72	357.21	1.03	0.32	35.7	37.3	+
39-1, 94-96	360.95	0.03	0.32	23.1	-33.3	+/-
39-3, 94-96	363.95	0.91	0.28	27.4	35.7	+
40-2, 146-148	372.47	0.001	0.32	50.8	49.8	+
40-3, 14-16	372.65	1.17	0.44	34.5	38.2	+
41-1, 60-62	379.61	0.83	0.37	25.6	33.8	+
41-2, 79-81	381.30	0.68	1.81	27.0	28.1	+
41-3, 117-119	380.93	0.00	0.31	-18.7	-27.4	-
41-4, 126-128*	384.77	1.62	0.36	34.6	28.5	+
41-5, 129-131	386.30	0.13	0.29	29.8	25.2	+
41-6, 33-35	386.84	0.28	0.24	37.3	28.5	+
42-1, 98-100	389.49	0.01	0.29	19.2	23.5	+
42-3, 44-46	391.95	0.39	0.21	-6.3	-32.0	+
42-4, 104-106	394.05	0.012	0.47	26.8	15.1	+
42-5, 75-77	395.26	1.34	0.42	55.7	66.0	+
42-6, 22-24	396.23	0.43	0.27	-10.0	-14.4	-
43-1, 23-25	398.24	0.28	0.40	39.9	35.7	+
43-2, 100-102	400.49	0.35	0.32	37.4	42.7	+
43-3, 66-68	401.67	0.74	0.34	15.5	5.0	+
44-1, 90-92	408.41	0.78	0.34	43.0	42.2	+
44-2, 63-65	408.89	1.54	0.52	35.6	21.6	+
44-3, 130-132	411.79	1.72	0.43	18.7	11.0	+
44-4, 05-07	412.06	0.76	0.39	26.9	26.9	+

TABLE 5 – *Continued*

Sample (interval in cm)	Sub- bottom Depth (m)	J <sub>NRM</sub> (10 <sup>-5</sup> gauss)	Suscepti- bility (10 <sup>-5</sup> gauss/oe)	Inclination		Polar- ity
				NRM	AFD	
445-45-1, 13-15	417.14	0.79	0.32	-4.7	-20.9	+
45-2, 73-75	419.24	0.38	0.31	-11.0	-20.8	-
45-3, 31-33	420.32	0.73	0.37	3.0	-15.2	+/-
45-4, 06-08	421.57	0.93	0.40	55.6	53.4	-
45-5, 66-68	423.67	1.31	0.45	-12.0	-19.5	-
46-1, 28-30*	426.79	1.73	0.28	34.8	35.3	+
46-2, 28-30*	428.29	2.85	0.44	24.2	23.6	+
46-3, 28-30*	429.79	3.38	0.56	23.6	21.2	+
46-4, 28-30*	431.29	6.52	0.57	30.0	29.3	+
46-5, 28-30	432.79	2.13	0.58	-80.4	-77.3	-
46-6, 28-30*	434.29	10.7	1.25	-2.9	-7.6	-
47-1, 73-75*	436.74	2.82	0.53	-19.1	-22.0	-
47-2, 73-75*	438.24	5.46	0.81	-18.2	-21.5	-
47-3, 73-75*	439.74	2.41	0.47	34.8	28.5	+
47-4, 73-75*	441.24	3.96	0.47	29.4	28.4	+
47-5, 73-75*	442.74	8.88	0.73	28.3	28.3	+
47-6, 73-75*	444.24	5.33	0.45	12.9	7.7	+
48-1, 26-28*	445.77	3.55	0.44	22.8	20.9	+
48-2, 26-28	447.27	1.93	0.40	7.4	1.5	+
48-3, 26-28*	448.77	3.52	0.52	24.2	23.2	+
48-4, 26-28	450.27	1.91	0.36	29.7	20.1	+
49-1, ?	455.75+?	1.77	0.40	35.5	30.6	+
49-2, 54-56*	457.05	5.34	0.40	27.0	26.3	+
49-3, 54-56	458.55	1.85	0.52	26.7	18.8	+
49-4, 54-56*	460.05	4.56	0.45	25.4	26.1	+
49-5, 54-56	461.55	1.68	0.44	27.9	22.9	+
50-1, 38-40	464.89	1.95	0.49	22.2	15.8	+
50-2, 38-40	466.39	1.55	0.40	-26.6	-30.9	-
51-1, 77-79	474.78	4.03	0.98	44.2	41.3	+
51-3, 108-110	478.09	0.16	0.44	11.8	-1.0	+/-
52-2, 86-88	485.87	2.33	0.42	5.2	-3.0	+/-
52-4, 64-66	488.65	1.77	0.69	20.2	12.9	+
53-1, 37-39	493.38	0.84	0.61	14.9	0.8	+
53-2, 37-39	494.89	1.12	0.46	38.0	34.6	+
53-3, 37-39	496.38	1.63	0.52	40.7	38.3	+
53-5, 23-25	499.24	3.38	1.21	38.2	39.7	+
53-5, 37-39*	499.38	2.30	0.50	-14.9	-19.6	-
54-1, 38-40*	502.89	4.22	0.69	-5.0	-10.6	-
54-2, 38-40*	504.39	2.43	0.39	32.3	30.0	+
54-3, 38-40*	505.89	3.67	0.54	4.4	-1.3	+/-
54-4, 38-40	507.39	1.91	0.55	-16.8	-29.7	-
54-5, 38-40	508.89	0.71	0.51	24.5	3.2	+
54-6, 38-40	510.39	0.16	0.38	36.2	14.7	+
55-1, 38-40	512.39	1.18	0.55	11.5	16.2	+
55-2, 38-40	513.89	1.71	0.54	-7.6	-15.5	-
55-3, 38-40	515.39	0.04	0.35	64.0	-18.2	+/-
55-4, 38-40	516.85	2.96	1.08	18.5	-0.6	+/-
55-5, 43-45*	518.44	5.60	0.54	-14.9	-18.2	-
55-6, 43-45	519.94	0.69	0.68	24.4	-7.5	+/-
56-1, 33-35*	521.84	3.31	0.57	-8.5	-13.5	-
56-2, 33-35*	523.34	8.58	1.21	-14.3	-18.7	-
56-3, 33-35*	524.84	6.49	0.88	-19.4	-22.0	-
56-4, 33-35*	526.34	5.28	1.00	-8.5	-14.4	-
56-5, 33-35	527.84	1.25	0.74	-2.1	-19.9	-
57-1, 73-75	531.74	0.64	0.31	4.1	3.7	+
57-2, 74-76*	533.27	2.10	0.48	34.0	33.7	+
57-3, 73-75*	534.74	6.48	0.64	25.1	24.2	+
57-4, 74-76	536.25	2.07	0.54	29.4	25.5	+
57-5, 73-75	537.74	1.16	0.32	23.0	17.9	+
58-1, 103-105	541.54	2.17	0.52	31.7	33.0	+
58-2, 23-25*	542.24	1.91	0.45	29.7	33.8	+
58-3, 110-112	544.61	0.92	0.35	-4.1	-10.6	-
58-4, 146-148	546.47	1.28	0.48	-7.5	-11.9	-
58-5, 93-95	547.44	1.61	0.55	-12.3	-18.2	-
59-1, 72-74	550.73	2.02	0.55	2.6	-3.2	+/-
59-2, 83-85	552.34	2.89	0.51	10.4	2.7	+
59-3, 26-28*	553.27	9.06	0.68	-25.9	-27.8	-
59-4, 145-147*	555.96	6.50	0.71	-12.1	-14.9	-
59-5, 06-08*	556.07	6.35	0.61	-4.1	-8.4	-
59-6, 28-30*	557.79	10.88	1.00	17.6	17.0	+

TABLE 5 – *Continued*

Sample (interval in cm)	Sub- bottom Depth (m)	J <sub>NRM</sub> (10 <sup>-5</sup> gauss)	Suscepti- bility (10 <sup>-5</sup> gauss/oe)	Inclination		Polar- ity
				NRM	AFD	
445-60-1, 41-43	559.92	0.05	0.78	-15.0	-18.2	-
60-2, 127-129	562.29	0.57	0.64	15.8	2.5	+
60-3, 49-51*	563.00	5.07	0.87	4.6	1.4	+
60-4, 13-15	564.14	2.22	0.81	24.5	13.9	+
61-1, 40-42*	569.41	5.70	0.95	-14.6	-15.4	-
61-2, 109-111*	571.58	8.17	0.81	16.9	16.7	+
61-3, 76-78*	572.77	10.94	0.93	20.1	20.1	+
61-4, 140-142*	574.91	6.97	0.83	19.3	19.1	+
61-5, 136-138*	576.37	5.14	0.65	20.1	17.8	+
61-6, 34-36*	576.85	7.84	0.87	22.4	22.0	+
61-7, 37-39	578.38	1.87	0.81	-2.1	-12.5	-
62-1, 143-145	579.94	1.70	0.65	9.7	-1.5	+/-
62-2, 44-46	580.45	0.88	0.56	0.4	-10.9	+/-
62-3, 143-145	582.94	2.76	0.75	29.6	26.8	+
62-4, 130-132*	584.31	7.43	1.16	29.4	29.5	+
62-5, 130-132*	585.81	5.29	0.84	22.8	20.8	+
62-6, 63-65*	586.64	9.58	0.19	31.4	31.9	+
63-1, 22-24	588.23	0.53	0.56	59.9	63.6	+
63-2, 127-129	590.78	3.29	0.74	33.7	32.0	+
63-3, 22-24	591.23	1.70	0.59	26.7	27.2	+
63-4, 116-118	593.67	1.14	0.55	22.0	17.9	+
64-1, 04-06	597.55	3.11	0.81	21.5	19.0	+
64-2, 18-20	599.19	1.68	0.52	-1.1	-3.9	-
64-3, 18-20	600.69	1.40	0.57	19.4	17.2	+
64-4, 18-20*	602.19	4.39	0.64	18.2	14.5	+
64-5, 18-20*	603.69	3.81	0.64	16.4	14.7	+
64-6, 05-07*	605.06	3.84	0.68	12.3	13.4	+
64-7, 05-07	606.56	3.16	0.78	17.2	14.9	+
65-1, 42-44	607.43	2.18	0.69	19.7	18.4	+
65-2, 42-44	608.93	0.88	0.61	34.3	-	(+)
65-3, 42-44	610.43	2.46	0.73	15.4	12.4	+
66-1, 16-18*	616.67	5.85	0.65	-13.5	-15.6	-
66-2, 16-18*	618.17	5.69	0.81	-1.0	-3.4	-
66-3, 16-18	619.67	3.16	0.88	11.0	3.6	+
66-4, 16-18*	621.17	3.95	0.55	-9.6	-11.5	-
66-5, 16-18	622.67	2.34	0.55	13.5	8.6	+
67-1, 19-21*	626.20	7.08	0.81	16.3	15.5	+
68-1, 78-80	636.29	2.39	2.05	19.5	24.0	+
68-2, 41-43	637.42	1.12	1.41	31.7	22.6	+
69-1, 24-26	645.25	1.55	0.78	2.9	2.8	+
69-2, 100-102*	647.51	18.99	2.47	-1.4	-1.7	-
69-4, 52-54	650.03	4.24	0.43	0.1	-4.2	+/-
70-2, 07-09*	656.08	27.65	1.60	7.0	6.6	+
71-1, 38-40	664.39	3.84	1.31	-18.5	-16.5	-
71-2, 06-08	665.57	14.34	2.46	6.4	7.6	+
71-3, 38-40*	667.39	15.55	2.29	6.0	4.9	+
71-4, 13-15	668.64	11.65	3.37	-36.2	-38.4	-
72-1, 43-45	673.94	14.65	3.10	2.5	2.3	-
72-2, 43-45	675.44	4.12	3.46	-33.2	-36.3	-
73-1, 84-86	683.85	11.09	6.53	-2.7	16.0	-
73-2, 84-86	685.35	11.77	5.13	16.0	18.4	-
73-3, 84-86	686.85	4.32	6.12	9.2	34.1	-
74-2, 46-48*	694.47	28.93	2.42	7.5	6.8	-
74-4, 46-48*	697.47	19.11	2.68	5.0	4.8	-
74-5, 46-48*	698.97	18.77	2.66	9.3	10.9	-
75-1, 39-41*	702.40	30.88	2.26	-10.1	-11.8	-
75-2, 39-41*	703.90	13.40	2.32	-2.8	-2.0	-
75-3, 120-122	706.21	4.73	1.71	42.6	22.3	+
75-4, 43-45	706.94	2.60	1.73	31.9	-	-
75-5, 43-45	708.44	2.66	3.20	25.0	27.8	-
76-1, 87-89*	712.38	19.52	2.49	11.6	8.1	-
76-3, 106-108*	715.57	20.42	2.78	3.9	2.7	-
76-4, 109-111*	717.10	32.21	2.20	17.9	17.7	-
76-5, 20-22	717.71	13.40	3.46	-12.6	-15.0	-
77-1, 139-141	722.40	14.05	3.13	-7.7	-3.0	-
77-2, 38-40*	722.89	24.10	4.02	5.8	4.1	-
77-3, 139-141	725.40	6.04	2.84	44.5	37.5	-
77-4, 144-146	726.95	32.84	2.77	-0.1	-2.6	-
77-5, 92-94	727.93	21.89	2.83	-24.5	-18.8	-
78-1, 84-86	731.35	14.41	2.65	2.4	-2.1	-

TABLE 5 — Continued

Sample (interval in cm)	Sub- bottom Depth (m)	$J_{\text{NRM}}$ ( $10^{-5}$ gauss)	Suscepti- bility ( $10^{-5}$ gauss/oe)	Inclination		Polar- ity
				NRM	AFD	
445-78-2, 60-62*	732.61	26.23	2.69	5.5	3.8	
78-3, 20-22*	733.71	24.49	2.99	3.5	0.9	
78-4, 25-27*	735.26	25.55	2.39	-5.3	-6.3	
78-5, 112-114	737.63	7.84	3.54	-33.8	-23.9	
79-2, 130-132*	742.81	23.43	1.81	3.2	1.4	
79-3, 52-54	743.53	19.26	3.07	-13.0	-1.6	
79-4, 54-56*	745.05	27.00	2.81	-3.4	-6.0	
79-5, 51-53*	746.52	25.49	3.39	-25.4	-27.2	
80-1, 135-137	750.86	13.14	3.70	-8.9	-13.7	
80-2, 138-140	752.39	15.38	4.45	-9.8	-13.4	
80-3, 138-140	753.89	16.10	4.19	-10.5	-10.1	
80-5, 140-142	756.89	4.23	5.76	6.3	-	
81-1, 56-58	759.57	28.79	4.12	-18.7	-20.6	
81-3, 56-58	762.57	20.12	4.29	28.8	-16.6	
81-5, 56-58	765.58	11.79	5.10	-1.9	-	
82-1, 70-72	769.19	19.59	4.97	-21.9	-	
82-3, 70-72	772.21	7.48	6.23	-26.6	-	
83-1, 76-78	778.77	23.01	4.94	-2.9	-4.4	
83-3, 76-78	781.77	10.87	6.52	48.4	-	
83-5, 76-78	784.77	9.93	4.67	-24.6	-	
84-1, 100-102	788.51	17.09	4.37	9.8	-	
84-3, 100-102	791.51	18.38	4.59	3.3	-	
85-3, 97-99	800.98	7.44	4.41	-31.1	-	
85-4, 94-96*	802.45	32.57	5.52	10.3	11.9	
86-2, 56-58	808.57	7.99	2.61	11.1	-	
86-4, 56-58	811.57	14.17	5.97	13.8	-	
87-4, 40-42	820.91	7.03	4.86	19.9	-	
87-6, 40-42	823.91	17.48	3.68	59.1	-	
88-3, 58-60	829.09	13.37	3.46	18.5	-	
88-4, 58-60	830.59	14.83	3.67	8.6	-	
89-4, 89-91	840.40	9.89	4.60	23.6	-	
89-6, 64-66*	843.15	43.88	2.59	14.1	12.7	

<sup>a</sup> $J_{\text{NRM}}$  is the intensity of NRM given in  $10^{-5}$  gauss/cm<sup>3</sup>; AFD is obtained by peak alternating demagnetizing field of 150 oe, decreasing to zero at a constant rate of 20 milligauss/cycle; polarity shows whether the inclination of NRM is positive (+) or negative (-); an asterisk at the sample number denotes good reliability for the sample.

highly stable and strong NRM similar to this mudstone (for example, lapilli tuff at Oshima, Japan, and welded tuff of Aso volcano, Kyushu, Japan).

In addition a fine, continuous record of polarity reversal was found. Between 142 and 551 meters, the sedimentary cores consist mainly of milky-colored nannofossil chalk. This part of the sediment was consolidated and was easily drilled, with a lesser number of fractures within each 9-meter core. Routine measurement in minicores every 1.5 meters showed that between Cores 45 and 47 there are more than two reversals. It is interesting to note that these cores cover only about 1 m.y. (24-25 Ma).

After rechecking the magnetic stability of NRM, it appears that the reversal would not have taken place due to a secondary effect, but must have recorded a real reversal of the geomagnetic field (at least the direction of a local magnetic field).

### PHYSICAL PROPERTIES

Site 445 sediments and sedimentary rocks vary considerably in composition, texture, porosity, and water content. This variability is directly reflected in physical properties (Table 7).

Figure 15 shows the variation of sonic velocity and wet-bulk density with depth for Hole 445. Velocities range from about 1.5 to 1.6 km/s in sub-units Ia, Ib, and IIa, and systematically increase from 1.7 to 2.6 km/s within sub-units IIb and IIc. Velocities decrease near the base of sub-unit IIc, and generally continue to decrease through units IIId and III. This decrease may be related to the decrease of calcium carbonate relative to silica in unit III. Unit IV is characterized by high and relatively variable sonic velocities. This is caused by the interlayering of various lithologies, including chert, radiolarite, mudstone, and conglomerate. The mudstones of unit V show monotonous velocities somewhat above 2 km/s and lower than those of the overlying unit. Velocities again increase in the conglomerates and sandstones of unit VI. Wet-bulk density follows the same pattern with respect to depth as sonic velocity (Figure 15).

There are several features of note in the velocity-depth profile presented in Figure 15. First, a clear velocity gradient distinguishes the upper 450 meters of the velocity column. Velocity gradients such as this will not necessarily be evident in marine reflection and refraction data and, if present, will tend to cause overestimation of layer thickness. Second, there are two distinctive low-velocity layers within the sediment column at Site 445. One occurs between the base of unit IIc and the top of unit IV. Unit V comprises the second velocity inversion. The presence of low-velocity layers will also lead to miscalculation of layer thicknesses using marine seismic data.

The variation of thermal conductivity with depth (Figure 16) is the same as the variations of sonic velocity and wet-bulk density (Figure 15). Conductivity increases with depth in units I, IIa, IIb, and IIc. This trend is similar to that observed by Hyndman et al. (1974) for Leg 26 sediments. The increase in conductivity with depth is directly related to the decreasing water content of the nannofossil oozes and chalks. This inverse relationship is shown in Figure 17 for sediments in Cores 1 through 58. The decrease of conductivity with increasing water content conforms to the general relationship established by Ratcliffe (1960). Site 445 data for carbonates, however, plot consistently above Ratcliffe's empirically determined curve (Figure 17).

Conductivity values decrease from a maximum in sub-unit IIc to a minimum in unit III. The decrease reflects the down-hole decrease in the amount of calcite, a mineral with a high thermal conductivity. Conductivities increase to an average of 3.2 mcal/cm<sup>2</sup>°C-s below unit III, and remain nearly constant for the mudstones, sandstones, and conglomerates of the lower portion of the recovered sedimentary section.

In general, conductivities for Site 445 sediments are higher than those reported for the clay-dominated lithologies of Sites 442, 443, and 444. This difference is caused by the generally higher thermal conductivity of the calcite-bearing lithologies of Site 445 and the generally lower water content of the carbonates. Bimodal distributions of thermal conductivities for oceanic sedi-



■ Normal    □ Reversed    X Not Sampled or No Recovery

Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity	Core	NRM	AFD	Polarity				
1			X	13			X	25			X	37			X	49			X	61			X	73			X	85			X
2	++	++		14			X	26	++	++		38	++	++		50			X	62	++	++		74	++	++		86			X
3	++	++		15			X	27	++	++		39	++	++		51	++	++		63	++	++		75	++	++		87	+		X
4	++	++		16	++	++		28	++	++		40	++	++		52	++	++		64	++	++		76	++	++					X
5	++	++		17	++	++		29	++	++		41	++	++		53	++	++		65	++	++		77	++	++					X
6	++	++		18	++	++		30	++	++		42	++	++		54	++	++		66	++	++		78	++	++					X
7	++	++		19	++	++		31	++	++		43	++	++		55	++	++		67	++	++		79	++	++					X
8	++	++		20	++	++		32	++	++		44	++	++		56	++	++		68	++	++		80	++	++					X
9	++	++		21	++	++		33	++	++		45	++	++		57	++	++		69	++	++		81	++	++					X
10	++	++		22	++	++		34	++	++		46	++	++		58	++	++		70	++	++		82	++	++					X
11	++	++		23	++	++		35	++	++		47	++	++		59	++	++		71	++	++		83	++	++					X
12	++	++		24	++	++		36	++	++		48	++	++		60	++	++		72	++	++		84	++	++					X

Figure 12. Results listed in Table 5 illustrated in descending order of cores and sections. Lowermost part of the polarity column is left blank because of poor resolution of NRM inclinations (see text).

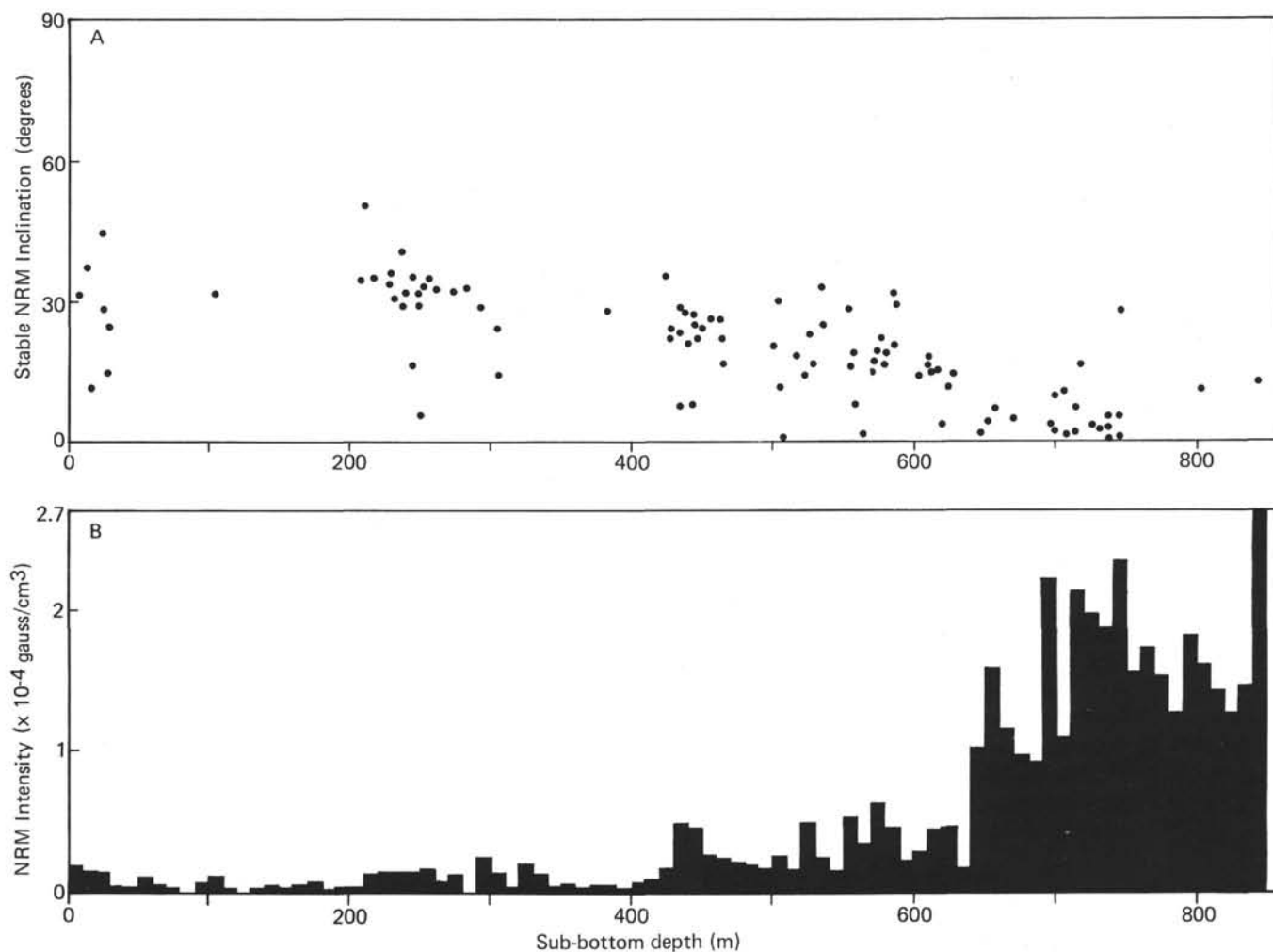


Figure 13. Paleomagnetism parameters plotted against sub-bottom depth. A. Absolute value of inclination of stable NRM of sediments. B. NRM intensity values averaged for every 10 meters.

TABLE 6  
Statistical Treatment of Absolute Values of NRM Inclination<sup>a</sup>

Sub-bottom Depth (m)	0-200	200-400	400-600	600
Number of Reliable Samples	9	18	41	27
Mean Inclination (degrees)	34.3	32.1	20.8	8.9(?) <sup>b</sup>
Standard Deviation	21.2	7.6	8.3	6.3
Expected Latitude (degrees)	18.8 <sup>+</sup> 17.2 <sup>-</sup> - 7.0	17.4 <sup>+</sup> 5.3 <sup>-</sup> - 5.0	10.8 <sup>+</sup> 4.3 <sup>-</sup> - 4.5	4.5 <sup>+</sup> 2.6 <sup>-</sup> - 3.3

<sup>a</sup>Expected latitude values are deduced from the mean of NRM inclinations.

<sup>b</sup>Poor resolution power of inclination at lower latitude.

ments (e.g., Erickson, 1973; Marshall and Erickson, 1974; Hyndman et al., 1974) apparently are created by these fundamental differences between lithologies dominated by clay minerals and carbonate minerals.

Shear strength ranges from 0.0 to 3.0 dynes/cm<sup>2</sup>; two measurements exceed this range. The nearly constant shear strength below 40 meters (Figure 18) reflects the constant lithology of the nannofossil ooze. Because of

the lithologic differences between Site 445 and Sites 442, 443, and 444, the previously reported variations of shear strength with depth are not in evidence at Site 445.

#### CORRELATION OF GEOPHYSICAL DATA WITH DRILLING RESULTS

Site 445 is 1 nautical mile northwest of shot point 3960 of line 3-2, multichannel seismic reflection profile of S/V *Kaiyo-Maru* (Figure 19). The uppermost layer, with a two-way normal time thickness of 0.4 seconds, is semi-transparent, and the underlying 0.3-second section is moderately stratified. Below this, two discrete, sharp reflectors can be observed at 0.68 seconds and 0.72 seconds, respectively.

Lithology and sonic velocity of the recovered cores can be correlated with these acoustic reflectors as follows:

1. The uppermost, semi-transparent layer corresponds to nannofossil ooze and the upper portion of nannofossil chalk sections lying above a sub-bottom depth of 360 meters, at which sonic velocity increases from 1.6 to 1.8 km/s to 2.0 km/s.

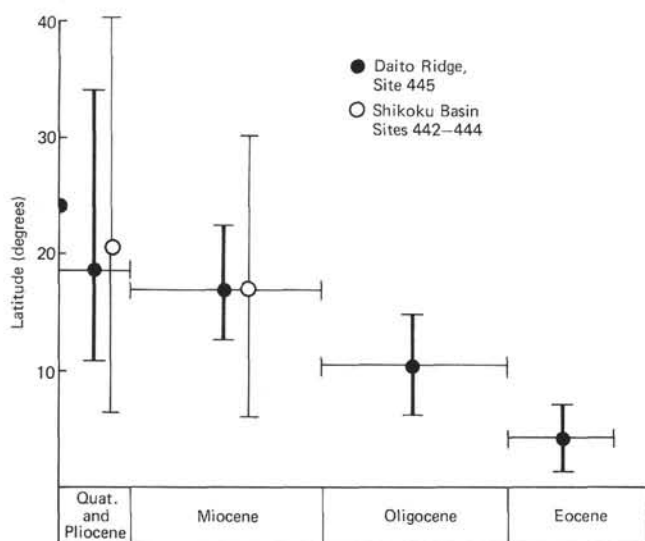


Figure 14. Latitude of Site 445 in the geological past. Solid circles are calculated and plotted from mean values of the absolute NRM inclination, as listed in Table 6. Horizontal bars show time spans for which inclination values are averaged. Vertical bars show probable errors corresponding to the standard deviations of NRM inclination in Table 6, asymmetrical around the expected geographic latitude. Open circles are plots for Sites 442 through 444 (mixed), for comparison. A solid semicircle on the left of the figure indicates the present latitude.

2. There are a few levels at which the sonic velocity sharply increases and then gradually decreases downhole; this is possibly associated with increase in clay content in chalk. The stratified structure of the lower sediment sections seen in the acoustic record is due to such a repetition of composition in chalk downhole.

3. Occurrence of chert and radiolarite, which show high sonic velocity of 3.0 km/s, is indicated by the sharp reflector observed at 0.68 seconds.

4. Conglomerate occurring at the hole bottom shows sonic velocity higher than 3.0 km/s. The boundary between the conglomerate and the overlying mudstone ( $V_p = 2.1$ ) probably marks the discrete reflector at 0.72 seconds.

5. Lithified basement is not clearly recognized in the seismic profile. An indistinct reflector at 0.91 seconds is likely the basement. If the average velocity of the overlying layers is assumed to be 2.1 km/s, the sub-bottom depth of basement is 956 meters below the mudline.

Sonobuoy site-survey measurements show that a 310-meter-thick layer ( $V_p = 1.85$ ) overlies a 240-meter-thick layer ( $V_p = 2.11$ ) and a 890-meter-thick layer ( $V_p = 3.9$ ), underlain by the basement with  $V_p = 5.7$ . Thickness and sonic velocity of the upper two layers are roughly consistent with those of the recovered cores. Sonic velocity of the underlying layer, in contrast, seems to be overestimated; thus thickness of the layer is likewise overestimated. If the interval velocity is assumed to be 2.5 km/s, the layer apparently becomes as thin as 580

meters, consistent with the depth of basement estimated by other methods. Sonobuoy data may be less precise here because of rough basement topography.

## SUMMARY AND CONCLUSIONS

### Summary

The stratigraphic section at Site 445 consists of six sedimentary units. Some of these units have been subdivided further on the basis of biogenic content, color, and minor compositional changes.

The total penetration at Site 445 was 892.0 meters, and the entire section was sedimentary. Acoustic basement was not reached, but may not be very far below the depth of maximum penetration.

The preservation of foraminifers and nannofossils was excellent in samples recovered from Pleistocene, Pliocene, and upper-Miocene rocks. This fauna is tropical. However, in the early Miocene and late Oligocene, the foraminifer and nannofossil species represent colder water. The Eocene fauna is tropical. The depth of deposition was well above the CCD from the late Eocene onward, and probably above the lysocline from latest Eocene to the present (Figure 20). The depth of deposition during the middle Eocene was most likely at or slightly below the CCD, apparently following the general Pacific CCD curve of van Andel et al. (1975).

Discontinuities in floral associations in the middle to early Oligocene chalks of unit IVd suggest that a hiatus may occur between Cores 57-5 and 59-4, encompassing a major part of the early and middle Oligocene. The discontinuity consists of an abrupt occurrence of well-developed late-Eocene assemblages below Core 59-5, whereas no Eocene assemblage exists in the overlying sediment. A similar discrepancy was also observed in cores between 57-4 and 57-5, where an early-Oligocene assemblage directly underlies an assemblage of the late middle Oligocene. Because this sequence is heavily reworked, the paleontological evidence indicates a hiatus lasting at least 8 m.y.

Sedimentation at Site 445 was mainly resedimentation of biogenic pelagic oozes by turbidity currents and slumping, and of terrigenous and volcanoclastic sands and conglomerates by turbidity currents, fluidized-sediment flow, debris flow, and slumping. Evidence for resedimentation by turbidity currents includes graded terrigenous and volcanoclastic foraminifer sands; graded sequences, T(a,b,e); and sharp scours. Most of the graded beds contain fossils reworked from shallower water depths. The interbedded pelagic sediments contain deeper-water fossils. Evidence for fluidized-sediment flow is found in conglomerate deposits organized into a framework conglomerate with a fabric. Many of these conglomerates contain specimens of the large shallow-water foraminifer *Nummulites*; these platy tests and other platy clasts define the fabric of such conglomerates. Evidence for deposition by debris flow consists again is found in conglomerates, but such conglomerates are supported by a mudstone or muddy-sandstone matrix, and the gravel-sized clasts display a dispersed fabric. Much of the sequence, particularly the sandstones, conglomerates, and mudstones, and chalks and

TABLE 7  
Summary of Physical Properties of Sediments, Site 445

Sample (interval in cm)	Lithology	Sonic Velocity (km/s)	Thermal Conductivity (mcal/cm-s-°C)	Shear Strength ( $\times 10^{-5}$ dynes/cm <sup>2</sup> )	Wet-Bulk Density (g/cm <sup>3</sup> )	Porosity (%)	Water Content (%)
445-1-2, 125-128	nannofossil ooze	—	—	0.05	—	—	—
1-3, 96-99	"	—	—	0.00	—	—	—
1-4, 83-86	"	—	—	0.38	—	—	—
1-4, 144-150	"	—	—	—	1.54	69.43	46.09
1-5, 71-73	"	—	—	0.10	1.58	70.62	45.78
1-6, 77-90	"	—	2.542	0.10	—	—	—
2-1, 90-92	"	—	—	0.02	1.52	72.85	49.08
2-2, 110-112	"	—	—	1.34	—	—	—
2-3, 136-138	"	—	—	2.11	—	—	—
2-4, 110-112	"	—	—	0.67	—	—	—
3-1, 120-122	"	—	—	0.96	—	—	—
3-2, 120-122	"	—	—	1.24	—	—	—
3-3, 69-71	"	—	—	0.62	—	—	—
3-4, 69-71	"	—	—	1.24	—	—	—
3-5, 30-32	"	—	—	1.72	1.60	67.16	42.92
3-6, 120-122	"	—	—	0.86	—	—	—
3-7, 20-22	"	—	—	2.49	—	—	—
4-1, 20-22	"	—	—	1.15	—	—	—
4-2, 110-112	"	—	—	0.10	—	—	—
4-4, 44-46	"	—	—	0.86	—	—	—
4-5, 42-52	"	1.568	2.811	0.77	1.54	59.71	39.67
4-6, 44-46	"	—	—	0.48	—	—	—
6-5, 144-150	"	—	—	—	1.63	63.41	39.75
6-6, 64-74	"	1.527	2.878	2.39	1.68	65.51	39.83
7-5, 90-100	"	1.509	2.867	2.39	1.63	65.32	41.03
8-3, 79-81	"	—	—	1.39	—	—	—
8-5, 110-120	"	1.517	2.875	2.49	1.58	68.41	44.39
8-5, 110-120	"	1.508*	—	—	—	—	—
10-5, 130-140	"	1.592	2.917	4.98	1.66	62.22	38.40
11-4, 87-89	"	—	—	1.53	—	—	—
11-4, 90-100	"	1.553	3.222	6.70	1.74	67.28	39.70
11-4, 144-150	"	—	—	—	1.70	62.79	37.93
12-2, 40-50	"	1.618	2.875	—	1.76	79.98	46.45
13-3, 62-72	"	1.571	2.742	—	1.60	68.14	43.73
14-1, 122-132	"	1.490	2.878	3.26	1.55	68.68	45.39
15-1, 126-136	"	1.535	2.900	—	1.72	64.26	38.21
16-2, 84-94	"	1.552	3.306	—	1.60	57.38	36.76
17-1, 40-50	"	1.577	3.411	—	1.72	61.20	36.56
17-1, 140-150	"	—	—	—	1.76	57.48	33.40
18-3, 69-79	"	1.579	—	—	1.70	53.84	32.44
18-4, 136-146	nannofossil chalk	1.552	—	—	1.70	61.34	36.95
20-2, 87-97	"	1.533	3.167	2.68	1.75	62.99	36.91
21-1, 62-72	"	1.550	3.389	—	—	—	—
22-1, 144-150	"	—	—	—	1.66	52.70	32.51
22-1, 144-150	"	—	—	—	1.78	54.32	31.30
22-2, 62-75	"	1.554	3.294	—	—	—	—
23-1, 70-71	"	—	—	—	1.62	61.49	38.98
23-2, 35-45	"	1.543	3.028	—	—	—	—
24-2, 78-79	"	—	—	—	1.79	55.07	31.49
24-3, 124-134	"	1.659	3.422	—	—	—	—
25-1, 24-25	"	—	—	—	1.77	53.20	30.86
25-2, 47-57	"	1.604	2.933	—	—	—	—
26-3, 102-103	"	—	—	—	1.74	56.32	33.24
27-4, 140-150	"	—	—	—	1.68	66.20	40.37
28-2, 5-15	"	1.600	2.781	—	1.68	64.87	39.66
29-4, 94-104	"	1.712	3.067	—	1.84	55.05	30.72
30-3, 14-29	"	1.657	3.236	—	1.72	57.38	34.25
31-1, 85-95	"	—	3.097	—	—	—	—
31-4, 110-121	"	1.663	3.394	—	1.75	53.46	31.36
32-1, 48-58	"	1.807	2.892	—	1.50	69.50	47.56
32-1, 72-82	"	1.736	3.983	—	1.67	59.28	36.36
32-2, 140-150	"	—	—	—	1.65	62.66	38.93



TABLE 7 – Continued

Sample (interval in cm)	Lithology	Sonic Velocity (km/s)	Thermal Conductivity (mcal/cm-s-°C)	Shear Strength ( $\times 10^{-5}$ dynes/cm <sup>2</sup> )	Wet-Bulk Density (g/cm <sup>3</sup> )	Porosity (%)	Water Content (%)
445-33-2, 50-60	"	1.733	3.467	—	—	—	—
34-2, 37-40	"	1.799	—	—	—	—	—
34-2, 46-56	"	—	—	—	1.79	54.32	31.13
34-2, 58-59	nannofossil chalk	—	—	—	1.76	55.49	32.36
35-2, 78-81	"	1.716	—	—	—	—	—
35-2, 83-93	"	—	3.511	—	1.79	57.14	32.62
36-2, 66-76	"	1.705	3.364	—	1.73	63.09	37.32
36-2, 114-124	"	—	2.903	—	—	—	—
37-2, 68-78	"	1.730	3.700	—	1.91	50.11	26.95
37-4, 23-33	"	1.865	3.389	—	1.95	53.13	27.92
37-4, 140-150	mudstone	—	—	—	1.73	60.90	36.04
38-3, 39-49	nannofossil chalk	1.802	4.111	—	1.96	51.32	26.80
39-1, 9-19	"	—	3.969	—	—	—	—
39-1, 94-97	"	1.892	—	—	1.94	58.69	30.94
40-2, 146-150	"	1.929	4.414	—	2.06	40.91	20.32
41-6, 33-43	"	1.999	4.081	—	2.11	41.57	20.18
42-4, 104-114	"	1.780	3.236	—	1.91	54.25	29.11
42-5, 90-100	"	—	—	—	1.79	55.33	31.68
43-2, 100-103	"	2.018	—	—	2.21	43.26	20.09
44-3, 130-133	"	1.958	—	—	2.23	46.06	21.18
45-3, 31-34	"	2.143	—	—	2.33	47.78	21.00
46-3, 28-30	"	—	—	—	1.91	44.09	23.67
46-4, 28-31	"	1.996	—	—	2.05	49.71	24.86
47-3, 73-76	"	1.937	—	—	2.13	47.76	22.92
47-4, 90-100	"	—	—	—	2.01	44.75	22.75
48-2, 26-36	"	2.524	5.144	—	2.14	41.56	19.89
49-3, 54-64	"	2.252	4.511	—	2.31	48.61	21.56
50-1, 38-48	"	2.358	4.492	—	2.26	47.08	21.36
51-3, 108-118	"	2.086	3.728	—	1.86	45.30	24.97
52-2, 86-96	"	2.275	4.997	—	1.99	36.65	18.84
52-2, 90-100	"	—	—	—	1.90	48.42	26.17
53-2, 34-44	"	2.111	4.094	—	1.92	47.10	25.12
54-2, 38-48	"	2.110	4.178	—	1.93	42.36	22.47
55-3, 38-48	"	2.065	4.228	—	1.95	46.89	24.59
56-3, 33-43	"	2.081	3.928	—	1.92	48.98	26.12
57-3, 73-83	"	2.124	3.381	—	1.76	53.45	31.04
58-1, 103-106	"	2.376	—	—	—	—	—
58-2, 23-33	"	2.407	5.036	—	—	—	—
59-4, 139-149	"	1.988	3.386	—	1.85	47.45	26.25
60-1, 38-48	"	2.106	2.878	—	1.92	47.80	25.45
61-3, 67-69	"	1.848	2.569	—	1.59	59.31	38.30
62-2, 44-47	"	1.948	—	—	1.77	55.33	32.02
62-2, 143-150	"	—	2.422	—	—	—	—
62-2, 140-150	"	—	—	—	1.75	56.95	33.38
63-2, 127-140	"	1.884	2.486	—	1.64	59.79	37.35
64-3, 18-28	"	1.823	2.594	—	1.61	59.06	37.51
65-1, 42-52	"	1.773	2.464	—	1.55	66.27	43.83
65-2, 90-100	radiolarite	—	—	—	1.46	66.24	46.60
65-3, 42-52	radiolarian ooze	1.736	2.575	—	1.46	61.04	42.92
65-3, 92-95	chert	2.160	—	—	—	—	—
66-2, 16-26	radiolarite	1.839	2.906	—	1.68	63.00	38.44
67-1, 19-22	mudstone	3.321	—	—	2.33	24.73	10.89
68-1, 78-88	siltstone	2.329	3.450	—	2.08	43.11	21.26
68-2, 41-51	"	2.538	2.989	—	2.10	37.52	18.26
69-1, 24-26	mudstone	—	—	—	2.19	32.86	15.37
69-2, 100-103	"	2.088	—	—	2.02	44.86	22.75
69-3, 28-38	sandstone	3.030	3.142	—	2.22	36.53	16.84
70-2, 7-17	claystone	2.065	3.275	—	2.78	47.18	24.78
71-1, 38-45	mudstone	2.537	3.375	—	1.93	41.21	21.92
72-1, 43-46	"	2.265	—	—	2.08	39.66	19.54
73-1, 80-83	"	2.078	—	—	2.01	42.52	21.68
73-3, 84-94	sandstone	2.061	6.531	—	1.96	42.87	22.39
74-2, 46-49	mudstone	2.244	—	—	2.01	40.16	20.42

TABLE 7 - Continued

Sample (interval in cm)	Lithology	Sonic Velocity (km/s)	Thermal Conductivity (mcal/cm-s-°C)	Shear Strength ( $\times 10^{-5}$ dynes/cm <sup>2</sup> )	Wet-Bulk Density (g/cm <sup>3</sup> )	Porosity (%)	Water Content (%)
445-75-2, 39-42	mudstone	2.152	-	-	2.03	38.72	19.59
75-5, 43-53	"	2.182	3.156	-	1.93	42.72	22.66
76-3, 106-116	"	2.282	3.211	-	2.11	36.75	17.82
77-1, 137-147	"	2.257	3.436	-	2.09	36.64	17.99
78-4, 25-35	"	2.067	3.194	-	2.06	48.48	24.14
79-2, 125-135	"	2.049	3.236	-	2.01	47.47	24.25
80-1, 135-145	"	2.037	3.231	-	1.96	47.25	24.68
80-4, 90-100	"	-	-	-	1.95	47.57	24.95
81-3, 56-66	"	1.994	3.033	-	2.03	56.91	29.69
82-1, 70-80	"	2.045	3.142	-	1.97	46.40	24.10
83-3, 76-79	"	2.062	-	-	1.97	45.99	23.91
84-1, 100-110	"	2.058	3.250	-	1.97	47.71	24.80
85-3, 97-107	"	2.110	3.239	-	2.03	42.35	21.38
86-2, 56-59	"	2.076	-	-	1.97	45.76	23.75
87-4, 40-50	siltstone	2.162	3.328	-	2.04	40.66	20.43
88-3, 58-68	mudstone	2.305	2.719	-	2.04	43.94	22.10
89-3, 75-85	"	2.092	3.158	-	2.02	44.46	22.59
90-1, 25-35	sandstone	2.949	4.428	-	2.18	32.06	15.05
91-2, 50-66	"	2.561	3.089	-	2.07	33.41	16.51
92-3, 95-105	"	3.192	2.783	-	2.32	28.44	12.55
93-3, 90-100	conglomerate	3.246	3.558	-	2.30	26.64	11.85
94-2, 96-106	"	3.059	3.597	-	2.31	26.30	11.68

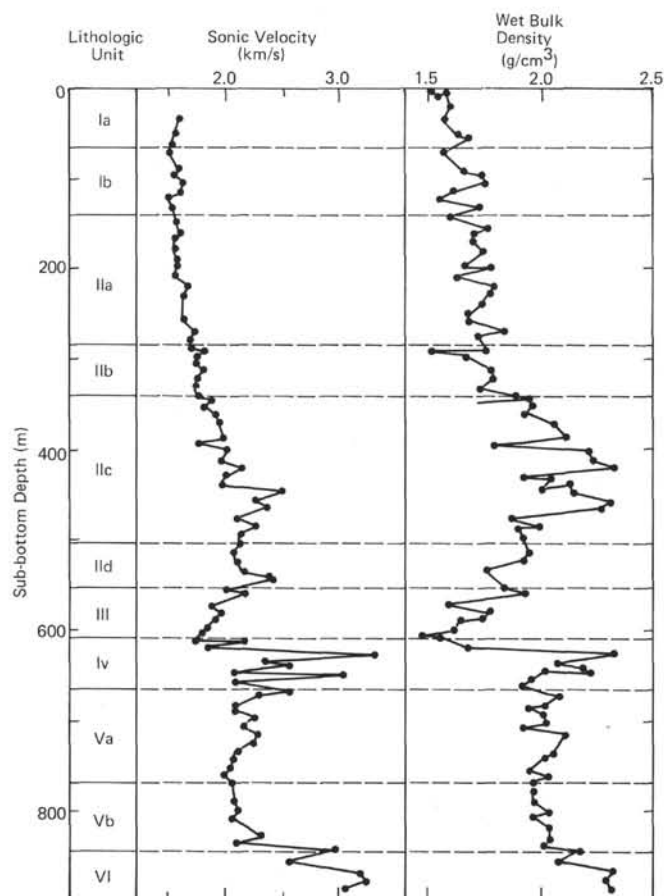


Figure 15. Sonic velocity and wet-bulk density versus depth for Site 445 sediments.

limestones of units II, V, and VI, show features clearly indicative of slumping, including rotation of large blocks (some as thick as 5 m), microfaults, pull-apart fractures, and well-developed slump folds. Slumping of sediments involved both the chalks and the volcanoclastic sandstones. This style of sedimentation has been observed before both in the South Fiji and New Hebrides marginal basins (Klein, 1975a,b), and off the Line Islands in the central Pacific (Cook et al., 1976).

The rates of sediment accumulation reflect this style of sedimentation. Rates are very low for the purely pelagic intervals, and moderate for the resedimented biogenic pelagic carbonates. Rates of sediment accumulation are very high for the sandstones and conglomerates of unit V. This pattern of sedimentation rates (an early history of rapid sedimentation, and then a decrease with time) is identical to a similar change in sediment accumulation rates reported by Andrews, Packham, et al. (1975) from the New Hebrides Basin (Site 286), and from the South Fiji Basin (Site 285).

The volume of sediment deposited by resedimentation processes is high, perhaps as much as 70 per cent for unit V, and perhaps as much as 80 per cent for the upper carbonate units. (See White et al., this volume.)

Organic-carbon content decreases exponentially to a depth of about 100 meters, below which it remains constant throughout the pelagic sequence. In the deepest part of the section, organic-carbon content increases again, probably reflecting the contribution of resedimented terrigenous organic material carried in the debris flows.

The observed decrease in organic carbon with sediment age and depth in these pelagic sediments is analogous to the trends observed in the hemipelagic sediments

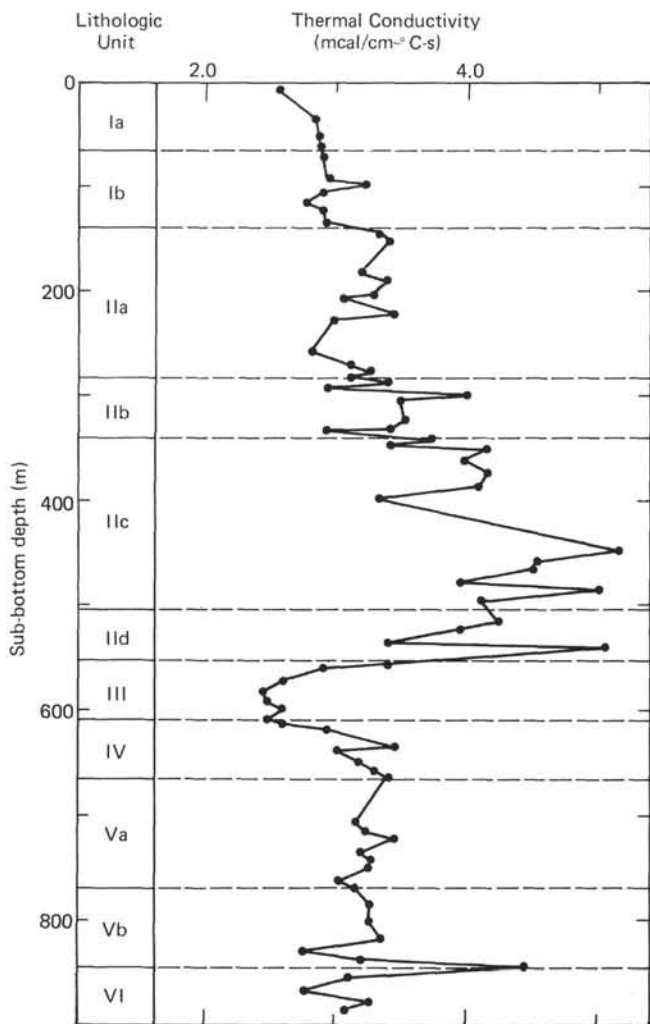


Figure 16. Thermal conductivity versus depth for Site 445 sediments.

at Sites 442, 443, and 444. Approximately equal lengths of time (about 4 m.y) are required for the degradation process to be completed in the pelagic and hemipelagic environments. Both the initial and final organic-carbon contents are only about half of those in the relatively pure hemipelagic clays at Sites 442 and 443; the explanation for this probably lies in the absence of organic matter adsorbed on clay surfaces at Site 445.

The pH of the sediments averages 7.46, alkalinity averages 1.6 meq/kg, salinity averages 35.7 per mill, and chlorinity averages 19.96 per mill.

Table 8 summarizes average physical properties of the sediments recovered at Site 445. Sonic velocity increases in the older units with depth, and from this a minimum elevation of the sediment/basement boundary was estimated at approximately 950 meters sub-bottom.

Paleomagnetism measurements of sediment samples recovered at Site 445 show that there is a systematic change in magnetic inclination with depth. The change indicates that Site 445 has drifted in a net northerly

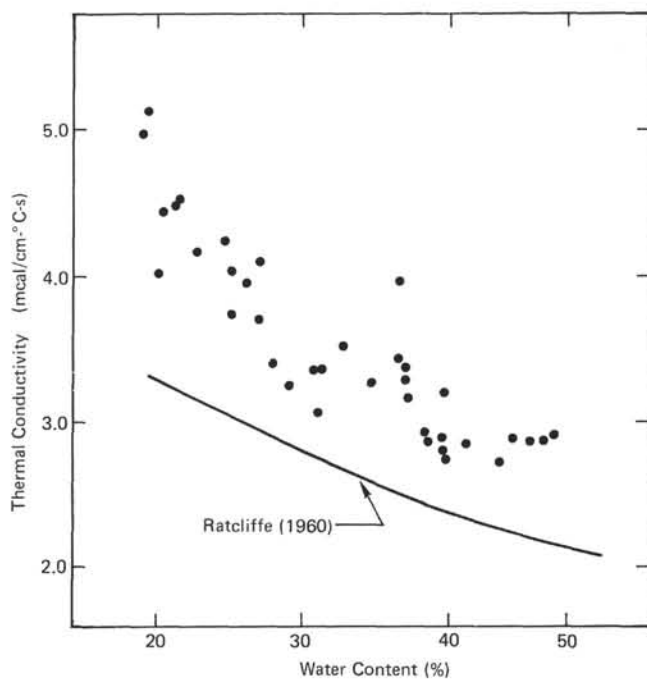


Figure 17. Relationship between thermal conductivity and water content for nannofossil oozes and chalks in Cores 1 through 58, Site 445. Also shown is the relationship for deep-ocean sediments established by Ratcliffe (1960).

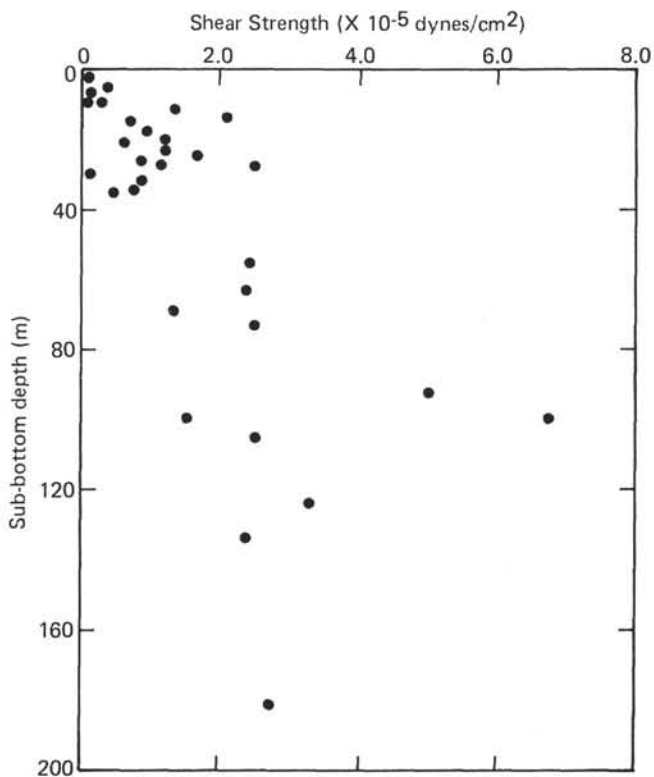


Figure 18. Shear strength versus depth for Site 445 sediments.

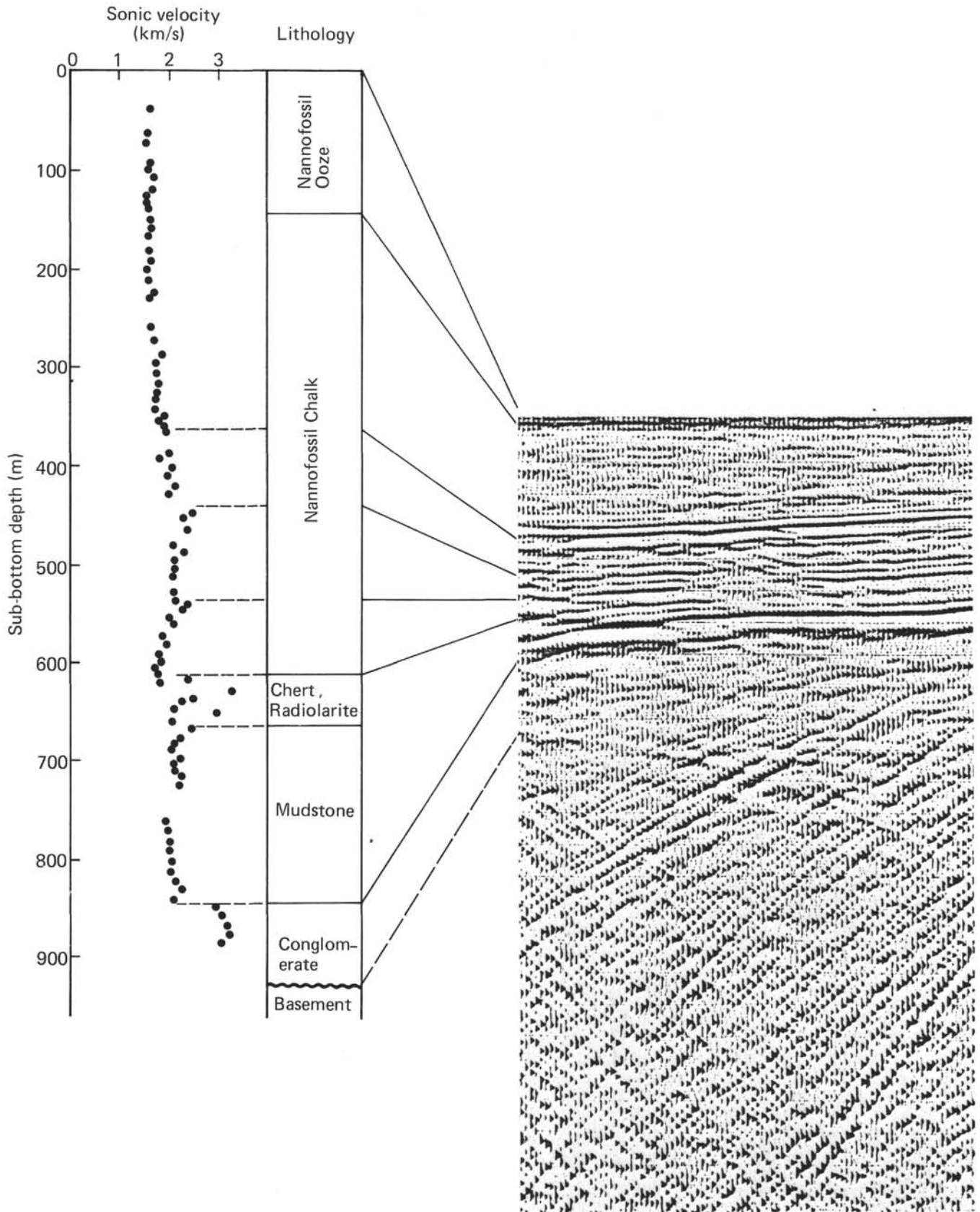


Figure 19. Site 445 stratigraphic section and corresponding seismic profile of R/V Kaiyo-Maru.



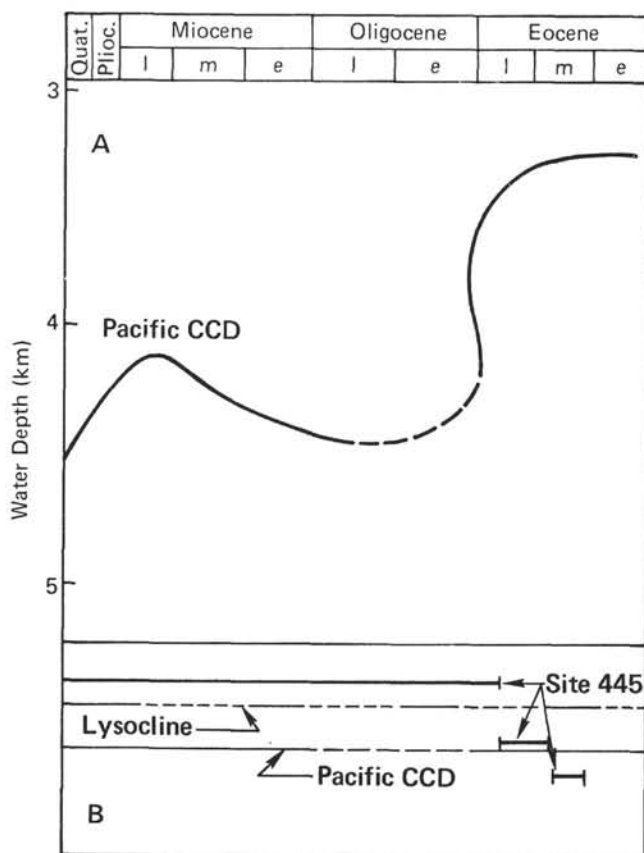


Figure 20. A. General curve showing estimated water depth of CCD in Pacific Ocean (after van Andel et al., 1975, p. 47, fig. 29). B. Relative depth of deposition at Site 445 compared to CCD curve for Pacific Ocean.

TABLE 8  
Average Values and Ranges of Sonic Velocity, Shear Strength, and Thermal Conductivity for Sediments, Site 445

	Sonic Velocity (km/s)	Shear Strength ( $\times 10^5$ dynes/cm <sup>2</sup> )	Thermal Conductivity (mcal/cm-s-°C)
Average	2.007	1.60	3.421
Range	1.490-3.321	0.0-6.70	2.422-6.531

direction over the last 47 m.y. At 47 Ma, Site 445 was close to the equator; it has migrated nearly 2000 km to its present position at an average rate of 4.4 cm/yr. These data are in agreement with results obtained by Loudon (1976, 1977) in the West Philippine Basin. Intensity of natural remanent magnetization is greater in deeper parts, which appear to contain higher amounts of volcanogenic magnetite.

### Conclusions

Our findings permit us to draw the following conclusions about the geological history of Site 445:

1. Although we did not reach acoustic basement, the systematic increase in sonic velocity with depth suggests that basement may occur not far below our maximum depth of penetration.

2. The depositional surface at Site 445 was well above the CCD during most of its depositional history.

3. The dominant motif of sedimentation at Site 445 was one of resedimentation of pelagic biogenic sediment, terrigenous sediment, and volcanoclastic sediment by turbidity currents, slump processes, fluidized-sediment flows, and debris flows.

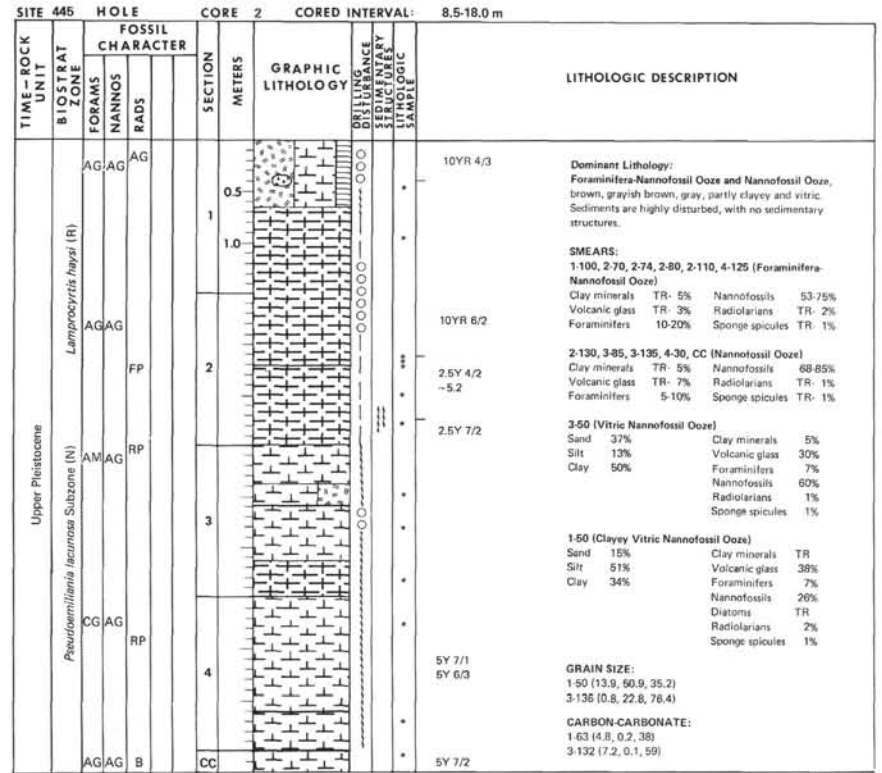
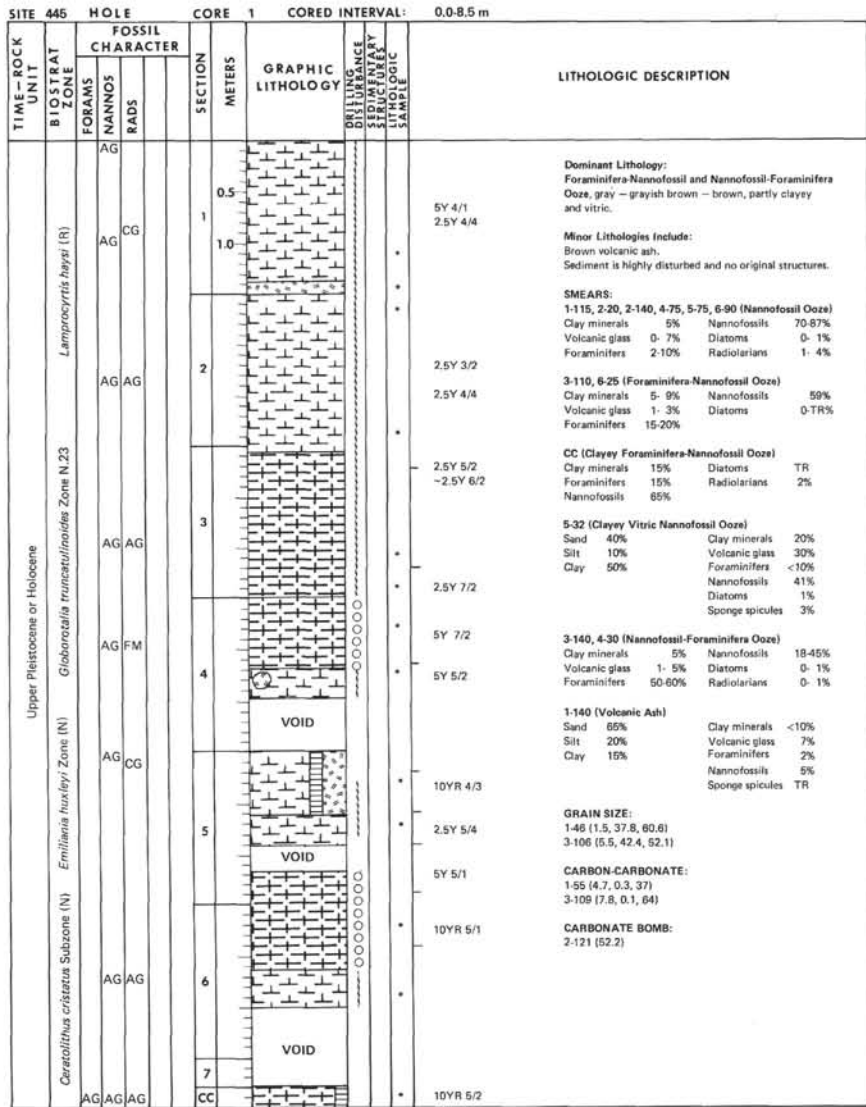
4. The occurrence of *Nummulites* in debris flow and fluidized-sediment-flow conglomerates indicates clearly that it and associated shallow-water bivalves, bryozoans, and gastropods were emplaced by resedimentation processes from shallower water. This origin for *Nummulites* is significant to consideration of dredge recovery of *Nummulites*-bearing rocks from the Amami Plateau and the Daito Ridge (Mizuno et al., 1975). The combined evidence from dredge sampling and our drilling indicates that the Daito Ridge, at present nearly 1300 meters below sea level, probably subsided no more than 1200 meters between Eocene time and the present.

5. Paleomagnetism data indicate that Site 445 and the surrounding Daito Ridge were around the equator at 47 Ma. Since then, northward drift averaging 4.4 cm/yr has occurred.

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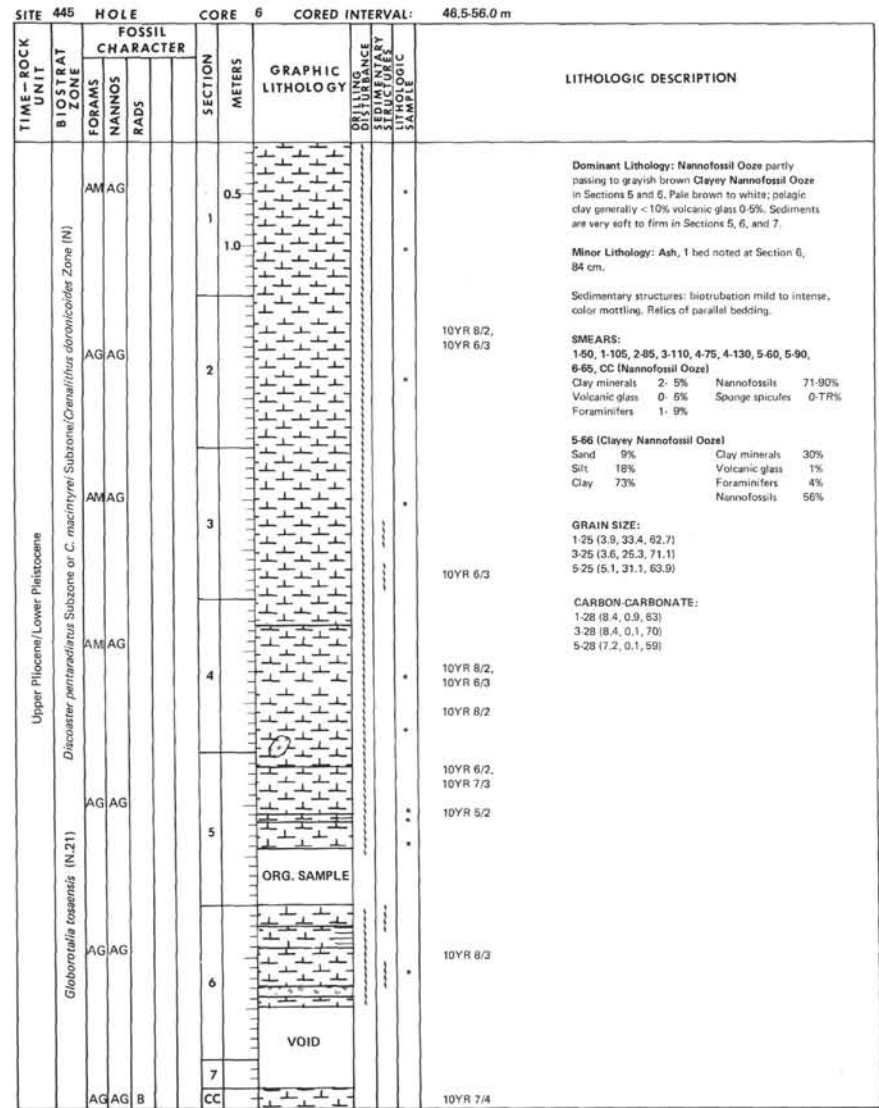
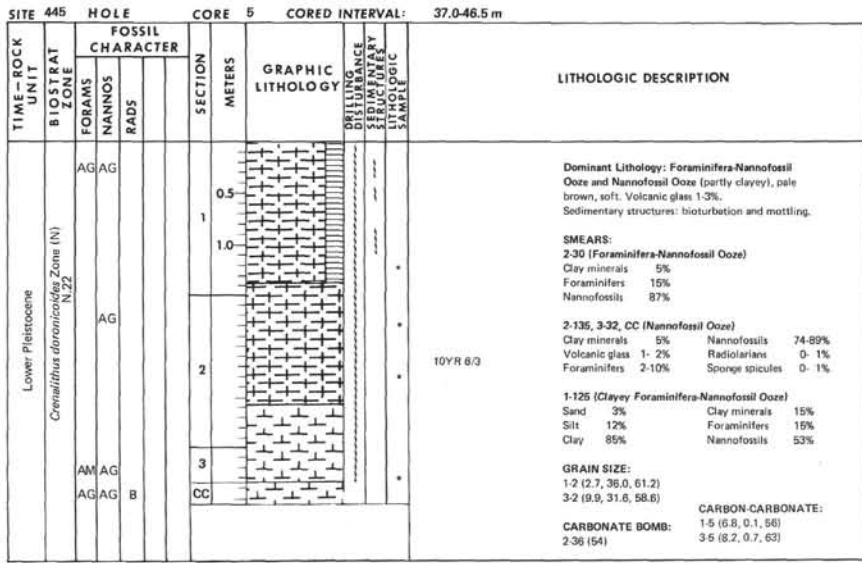
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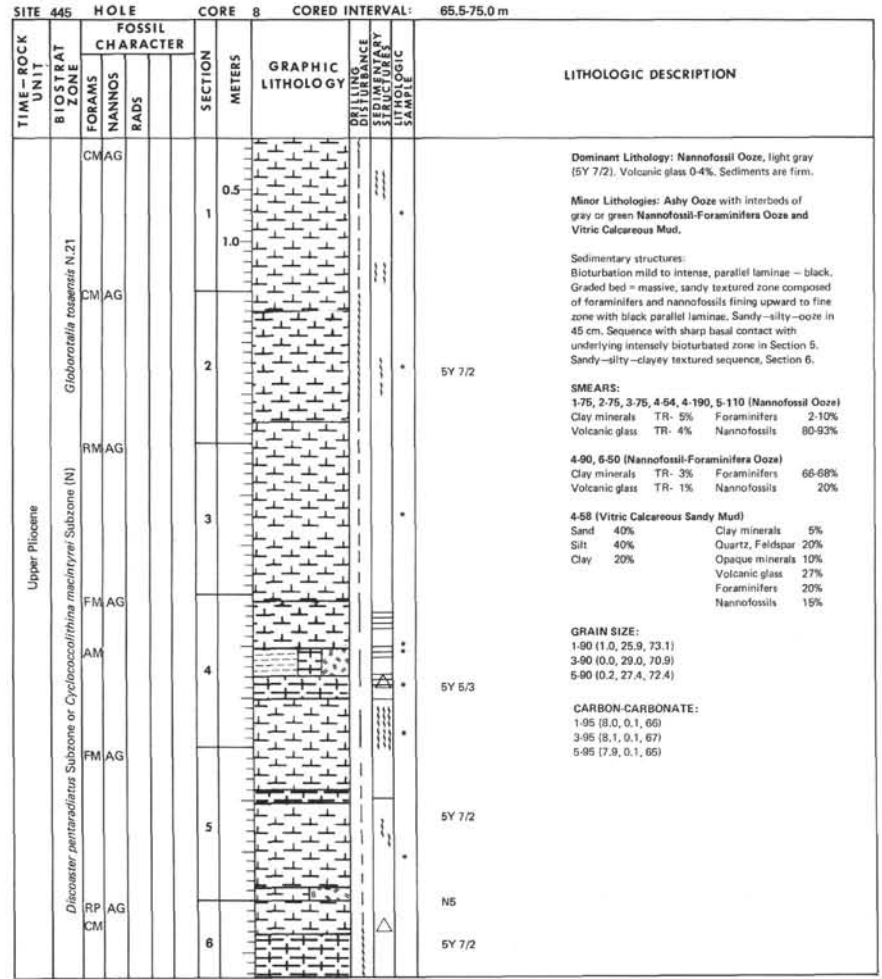
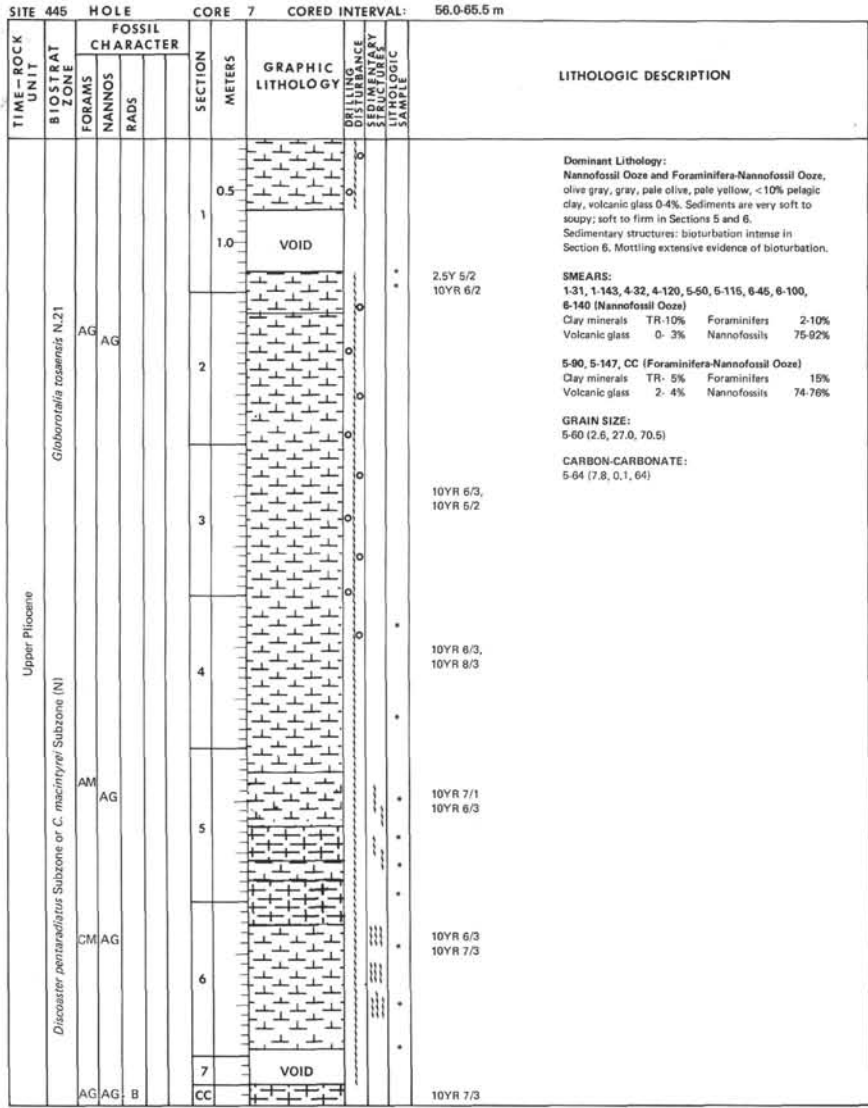


SITE 445		HOLE		CORE 3		CORED INTERVAL: 18.0-27.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Pleistocene	N23	AG	AG	1	0.5		5Y 6/1 5Y 7/1
		AG	AG	2	1.0		10YR 8/2
		AG	AG	3			
	N22	AG	AG	4			
		CG	AG	5			
	Pseudonellina's Incurvosa Subzone (N)	AG	AG	6			
		AM	AG	7			
	AG	AG	CC				
<p><b>Dominant Lithology:</b> Foraminifera-Nannofossil Ooze, Nannofossil Ooze, with Nannofossil-Foraminifera Ooze, partly vitric. Lower part of the core is clayey (up to 15%). Colors are light gray, light brownish gray, light olive gray, light brown. The sediments are firm to stiff. Sedimentary structures: Parallel bedding obvious even in intensely disturbed sediments. Intense to moderate bioturbation in Sections 3 and 4. Graded bed fining upward in Section 3, 102-117 cm: sand-sized at base grading upward into silt then white clay.</p> <p><b>SMEARS:</b> 1-75, 2-110, 3-75, 3-130, 4-75, 4-130, 6-120 (Nannofossil Ooze) Clay minerals TR: 5% Nannofossils 69-70% Volcanic glass TR: 5% Radiolarians TR: 1% Foraminifers 5-10% Sponge spicules 0-1%</p> <p>1-130, 2-39, 3-95, 3-115 (Foraminifera-Nannofossil Ooze) Clay minerals 5% Nannofossils 55-71% Volcanic glass TR: 2% Radiolarians 0-TR% Foraminifers 15% Sponge spicules TR</p> <p>5-105, 5-135, 7-10, CC (Clayey Nannofossil Ooze) Sand 2-5% Clay minerals 10-15% Silt 8-15% Volcanic glass TR: 2% Clay 80-90% Foraminifers 3-7% Nannofossils 65-88% Radiolarians 0-TR% Sponge spicules 0-1%</p> <p>3-115 (Nannofossil-Foraminifera Ooze) Clay minerals TR Carbonate unspecified 19% Foraminifers 50% Nannofossils 30%</p> <p>1-20 (Vitric Nannofossil Ooze) Sand 40% Clay minerals 5% Silt 20% Volcanic glass 30% Clay 40% Foraminifers 5% Nannofossils 60% Radiolarians TR Sponge spicules 1%</p> <p>5-30 (Vitric Foraminifera-Nannofossil Ooze) Sand &lt;10% Clay minerals &lt;10% Silt &gt;45% Volcanic glass 15% Clay &gt;40% Foraminifers 15% Nannofossils 50% Sponge spicules 1%</p> <p><b>GRAIN SIZE:</b> 1-75 (1.7, 35.1, 63.2) 3-81 (1.8, 31.2, 67.3) 5-46 (6.4, 49.9, 43.7)</p> <p><b>CARBON-CARBONATE:</b> 1-78 (5.1, 0.1, 42) 3-84 (6.4, 0.1, 52) 5-49 (3.2, 0.1, 26)</p>							

SITE 445		HOLE		CORE 4		CORED INTERVAL: 27.5-37.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Lower Pleistocene	N22	CM	AG	1	0.5		10YR 5/3
		CM	AG	2	1.0		10YR 8/3
		CM	AG	3			10YR 5/3, 10YR 8/3
	Crenalithus doronicoides Zone (N)	AG	AG	4			10YR 8/3
		AM	AG	5			10YR 7/3 10YR 7/3, 10YR 6/2 10YR 8/4
	Pseudonellina's Incurvosa Subzone (N)	CM	AG	6			10YR 8/1 -8/4
		AG	AG	7			10YR 6/2
	AG	AG	CC				
<p><b>Dominant Lithology:</b> Nannofossil-Foraminifera Ooze (upper sections) and Nannofossil Ooze (lower sections). Colors are light brown, brownish gray, brown. Volcanic glass rare. Sediments are soft to firm. Color mottling. Sedimentary structures not present except for bioturbation.</p> <p><b>SMEARS:</b> 1-15, 1-125, 4-65, 4-130, 5-17, 5-130, 6-75, 7-25, CC (Nannofossil Ooze) Clay minerals 0-8% Nannofossils 71-90% Volcanic glass 0-3% Sponge spicules 0-1% Foraminifers 1-10%</p> <p>2-38, 2-45, 2-115, 3-135, 4-20 (Foraminifera-Nannofossil Ooze) Clay minerals TR: 5% Nannofossils 60-83% Volcanic glass TR: 3% Sponge spicules 0-TR% Foraminifers 10-25%</p> <p>5-10 (Nannofossil-Foraminifera Ooze) Clay minerals 0% Nannofossils 45% Volcanic glass 5% Radiolarians TR Foraminifers 50%</p> <p><b>GRAIN SIZE:</b> 1-20 (1.2, 36.2, 62.6) 3-67 (2.7, 36.1, 61.2) 5-66 (0.0, 34.0, 65.0)</p> <p><b>CARBONATE BOMB:</b> 4-59 (58.5)</p> <p><b>CARBON-CARBONATE:</b> 1-23 (4.1, 0.1, 33) 3-71 (6.4, 0.1, 53) 5-59 (8.2, 1.0, 60)</p>							







SITE 445		HOLE		CORE 9		CORED INTERVAL: 75.0-84.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Pliocene	<i>Globobulimina esmerina</i> N.21 <i>Discoaster pentaradiatus</i> Subzone or <i>Cyclocoelithina machilynii</i> Subzone (N)	CMAG	B	1	0.5		5Y 7/2  Dominant Lithology: Nannofossil Ooze, light gray with intense bioturbation.  SMEARS: CC (Nannofossil Ooze) Clay minerals 5% Foraminifers 7% Volcanic glass 2% Nannofossils 80%

SITE 445		HOLE		CORE 10		CORED INTERVAL: 84.5-94.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Pliocene	<i>Discoaster tamalis</i> Subzone or <i>D. surculus</i> Subzone (N)	AG	AG	1	0.5		5Y 5/3  Dominant Lithology: Foraminifera-Nannofossil Ooze and Nannofossil Ooze, light gray to white. Volcanic glass 1-2%. Sediments are soft to firm.  Minor Lithology: Clayey Nannofossil Ooze (Section 1) olive color with 20% clay.  Sedimentary structures: mottling.
		AM	AG	2	1.0		SMEARS: 1-54, 1-75, 2-75, 3-75, 4-75 (Foraminifera-Nannofossil Ooze) Clay minerals 5-8% Foraminifers 10-18% Volcanic glass TR: 2% Nannofossils 71-87%
		AM	AG	3	2.0		5Y 7/2, 5Y 8/2  4-147, 5-75, CC (Nannofossil Ooze) Clay minerals TR: 9% Nannofossils 64-83% Volcanic glass 1-2% Quartz, Feldspar (4-147) 9% Foraminifers 5-10% Opaque minerals (4-147) 10%
		AM	AG	4	3.0		1-20 (Clayey Nannofossil Ooze) Sand 2% Clay minerals 20% Silt 10% Volcanic glass 1% Clay 88% Foraminifers 2% Nannofossils 21%
		AM	AG	5	4.0		5Y 7/2, 5Y 8/1-2
		AM	AG	B	6	5.0	

SITE 445		HOLE		CORE 11		CORED INTERVAL: 94.0-103.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Pliocene	<i>Globobulimina tosaensis</i> N.21	AG	AG	1	VOID	5Y 7/2, 5Y 5/2	Dominant Lithology: Nannofossil Ooze and minor amounts of Foraminifera-Nannofossil Ooze, light gray to white with olive gray. Minor amounts of clay and volcanic glass. Sediments are soupy to firm.
		AG	AG	2	VOID	5Y 7/2	Minor Lithologies: Vitric Nannofossil Ooze and Clayey Nannofossil Ooze. Sedimentary structures: bioturbation (mild to intense), mottling with olive gray, parallel laminae white and light gray.
		CM	AG	3	VOID	5Y 7/2	<b>SMEARS:</b> 1-116, 2-75, 3-75, 4-100, 5-41, CC (Nannofossil Ooze) Clay minerals 5-10% Foraminifers 1- 8% Volcanic glass TR- 5% Nannofossils 62-85%  4-75 (Foraminifera-Nannofossil Ooze) Clay minerals 9% Foraminifers 15% Volcanic glass 2% Nannofossils 69%  1-96, 3-10 (Vitric Nannofossil Ooze) Sand 5-25% Clay minerals 5- 9% Silt 15-20% Volcanic glass 13-18% Clay 55-80% Foraminifers 3- 4% Nannofossils 62-74%  1-137 (Clayey Nannofossil Ooze) Sand 1% Clay minerals 25% Silt 10% Volcanic glass 1% Clay 89% Foraminifers 1% Nannofossils 82%
		FM	AG	4		5Y 7/2, 5Y 8/1	<b>GRAIN SIZE:</b> 2-62 (3.8, 28.7, 67.7) 4-62 (4.8, 31.9, 63.3)  <b>CARBON-CARBONATE:</b> 2-66 (8.6, 0.1, 79) 4-66 (8.1, 0.1, 67)
		AM	AG B	5		5Y 7/2	
				CC			

SITE 445		HOLE		CORE 12		CORED INTERVAL: 103.5-113.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Pliocene	<i>Globobulimina tosaensis</i> N.21	FM	AG	1		5Y 5/2	Dominant Lithology: Nannofossil Ooze and lesser amounts of Clayey Nannofossil Ooze, associated with Foraminifera-Nannofossil Ooze. Alternating olive gray (clayey nannofossil ooze) and light gray (nannofossil ooze and foraminifera-nannofossil ooze); gray at base, clay <20% and glass 1-5%. Sediments are firm to very firm.
		AG	AG	2		5Y 7/1	Minor Lithologies: Nannofossil-Foraminifera Ooze, Volcanic Ash in Section 3. 2 and 3.  Sedimentary structures: bioturbation: mild to intense graded bed; scoured basal content; fining upward. NOTE: beds thinner, less massive than in higher cores.
		CM	AG	3		5Y 5/2	<b>SMEARS:</b> 1-80, 2-60, 3-100, 4-80, CC (Nannofossil Ooze) Clay minerals 7- 9% Foraminifers TR-10% Volcanic glass TR- 3% Nannofossils 81-92%  1-21, 4-70 (Clayey Nannofossil Ooze) Sand 1% Clay minerals 20% Silt 9% Volcanic glass 0-TR% Clay 90% Foraminifers 1- 2% Nannofossil 69-76%
		CM	AG	4		5Y 7/1	2-140 (Foraminifera-Nannofossil Ooze) Clay minerals 7% Foraminifers 12% Volcanic glass 4% Nannofossils 73%
		CM	AG	5		5Y 7/1	2-74, 3-16 (Nannofossil-Foraminifera Ooze) Clay minerals 4- 5% Nannofossils 15-20% Volcanic glass 2- 3% Carbonate Foraminifers 30-62% unspecified 15-43%
		CM	AG B	6		5Y 7/2	3-33 (Volcanic Ash) Sand 15% Volcanic glass 67% Silt 70% Foraminifers 1% Clay 15% Nannofossils 6% Quartz, Feldspar 20%
		CC		7		5Y 7/1	<b>GRAIN SIZE:</b> 1-54 (0.3, 24.5, 75.2) 3-54 (0.8, 29.2, 70.2)  <b>CARBON-CARBONATE:</b> 1-58 (8.4, 0.1, 69) 3-58 (7.9, 0.1, 65)



SITE 445		HOLE		CORE 13		CORED INTERVAL: 113.0-122.5 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS RADS						
Lower Pliocene	N.21	Gibborotalia marginata Zone N.19	AG	1				5Y 5/2	Dominant Lithology: Nannofossil Ooze and interbedded Clayey Nannofossil Ooze partly Vitric Nannofossil Ooze and olive gray Clayey Nannofossil Ooze, gray - light gray, rare volcanic glass. Sediments are firm to very firm.
								5Y 7/1	
								5Y 5/2	Minor Lithology: Foraminifera Ooze, Nannofossil-Foraminifera Ooze and Calcareous Ooze at a base of a graded bed. Nannofossil Ash in Section 4, dark olive gray.
								5Y 7/1	Sedimentary structures: bioturbation, parallel laminae with graded beds with coarse sand-sized foraminifers grading up to silt size to calcareous ooze. Color change light to dark upward.
								5Y 5/2	
								5Y 7/1	
								5Y 7/1	SMEARS: 1-74, 1-139, 2-69, 3-30, 3-140 (Nannofossil Ooze) Clay minerals 0-10% Foraminifers TR- 3% Volcanic glass 0- 4% Nannofossils 74-88%
								5Y 7/1	2-78, 3-100, CC (Clayey Nannofossil Ooze) Sand TR- 4% Clay minerals 15% Silt 12-20% Volcanic glass 1- 5% Clay 80-87% Nannofossils 71-79%
								5Y 7/1	4-140 (Clayey Vitric Nannofossil Ooze) Sand 10% Clay minerals 15% Silt 15% Volcanic glass 15% Clay 75% Foraminifers 2% Nannofossils 57%
								5Y 7/1	4-130 (Nannofossil Ash) Sand 70% Clay minerals 5% Silt 20% Volcanic glass 72% Clay 10% Foraminifers 3% Nannofossils 10%
								5Y 7/2	1-100 (Nannofossil-Foraminifera Ooze) Foraminifers 78% Nannofossils 15%
								5Y 7/2	3-95 (Foraminifera Ooze) Foraminifers 80% Nannofossils <10%
								5Y 7/2	4-107 (Calcareous Ooze) Clay minerals 2% Carbonate unspecified 36% Foraminifers 15% Rock fragments 30% Nannofossils 10%
								5Y 7/2	GRAIN SIZE: CARBON-CARBONATE: 1-48 (0.0, 27.2, 72.8) 1-52 (9.2, 0.1, 76) 3-47 (0.0, 31.9, 68.1) 3-52 (9.0, 0.1, 74)
				3					
				4					
				5					
				CC					

SITE 445		HOLE		CORE 14		CORED INTERVAL: 122.5-132.0 m										
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION								
		FORAMS	NANNOS RADS													
Lower Pliocene	Sphaerofhrus meekales Subzone (N) Forams N.19	FM	AG	1				5Y 7/1	Dominant Lithology: Nannofossil Ooze, light gray with olive gray mottling. Sediments are very firm.							
								7	Minor Lithology: Vitric Nannofossil Ooze at base of graded bed (Section 1).							
								5Y 7/1	Sedimentary structures: bioturbation, mottling load casts at base of graded(?) ashy nannofossil ooze.							
								5Y 7/1	SMEARS: 1-75, CC (Nannofossil Ooze) Clay minerals 5- 8% Nannofossils 87-91% Foraminifers 1- 2%							
								5Y 7/1	1-139 (Vitric Nannofossil Ooze) Sand 15% Clay minerals 7% Silt 20% Volcanic glass 30% Clay 65% Foraminifers 2% Nannofossils 56%							
								5Y 7/1	GRAIN SIZE: CARBON-CARBONATE: 1-78 (0.1, 27.8, 72.2) 1-82 (9.7, 0.1, 81)							
												2				
												CC				

SITE 445		HOLE		CORE 15		CORED INTERVAL: 132.0-141.5 m										
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION								
		FORAMS	NANNOS RADS													
Lower Pliocene	Sphaerofhrus meekales Subzone (N) N.19 Gibborotalia marginata	FM	AG	1				5Y 7/1	Dominant Lithology: Nannofossil Ooze and interbedded Clayey Nannofossil Ooze, light gray. Alternating beds of light gray (5Y 7/1) Nannofossil Ooze and light gray (5Y 7/2) Clayey Nannofossil Ooze, have ~10% pelagic clay. Glass very rare. Sediments are very firm.							
								5Y 7/2	Minor Lithology: Clayey Nannofossil Ooze, Foraminifera Ooze in parallel laminae at base of bed (Section 1).							
								5Y 7/1	Sedimentary structures: moderately intense bioturbation. Parallel lamination of foraminifera ooze above sharp contact (Section 1, 104 cm). Chondrites.							
								5Y 7/1	SMEARS: 1-120, 2-35, 2-70, CC (Nannofossil Ooze) Clay minerals 7-10% Foraminifers TR- 1% Volcanic glass 0- 1% Nannofossils 87-90%							
								5Y 7/1	1-50 (Clayey Nannofossil Ooze) Sand 1% Clay minerals 12% Silt 8% Foraminifers 1% Clay 93% Nannofossils 85%							
								5Y 7/1	1-104 (Nannofossil-Foraminifera Ooze) Foraminifers 70% Carbonate unspecified 10% Nannofossils 15%							
								5Y 7/1	GRAIN SIZE: CARBON-CARBONATE: 1-63 (0.6, 30.0, 69.4) 1-67 (8.1, 0.1, 67)							
												2				
												CG				
												AG				
												B				
												CC				

SITE 445		HOLE		CORE 16		CORED INTERVAL: 141.5-151.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Miocene - Lower Pliocene	Ceratolithus acutus Subzone (N1) / Ceratolithus rugosus Subzone N18(?)	B	AG	1	0.5		<p>Dominant Lithology: Nannofossil Chalk, light gray with minor amounts of clay. 0-33 cm - firm nannofossil ooze, light gray associated with nannofossil-foraminifera ooze and calcareous sand at base of graded bed (30 and 32 cm). Very firm (change from ooze to chalk at 33 cm).</p> <p>Minor Lithology: Nannofossil-Foraminifera Chalk and Sandy Foraminifera-Nannofossil Chalk at lower part of graded bed (Section 3).</p> <p>Sedimentary structures: parallel laminae-grayish green. Bioturbation. Slump structures and inclined bedding Sections 3 and 4. Graded bed with sharp scoured basal contact: dark gray; abundant sand-sized grains; foraminifera, sediment clasts - dark gray, black, green. Contorted bedding in slump.</p> <p>SMEARS:</p> <p>1-10 (Nannofossil Ooze) and 1-80, 1-88, 1-127, 2-75, 3-70, 3-129, 3-140, CC (Nannofossil Chalk)            Clay minerals 0-8% Foraminifera 1-5%            Volcanic glass 0-1% Nannofossils 85-91%</p> <p>1-30 (Nannofossil-Foraminifera Ooze) and 3-29 (Nannofossil-Foraminifera Chalk)            Clay minerals 0% Foraminifera 60-80%            Volcanic glass TR-5% Nannofossils 12-23%</p> <p>3-86 (Sandy Foraminifera-Nannofossil Chalk)            Sand 50% Clay minerals 2%            Silt 25% Foraminifera 10%            Clay 25% Nannofossils 67%            Volcanic glass 4%            Opaque minerals 5%            Rock fragments 10%</p> <p>1-32 (Calcareous Sand)            Sand 85% Carbonate unspecified 10%            Silt 10% Foraminifera 15%            Clay 5% Nannofossils 8%            Rock fragments 60%</p> <p>GRAIN SIZE: 1-76 (0.0, 28.1, 71.9) CARBON-CARBONATE: 1-80 (10.1, 0.0, 84)            3-76 (35.8, 38.3, 25.9) 3-80 (11.0, 0.1, 91)</p>
		B	AG	2	1.0		
		AG	AG	3	1.0		
		AG	AG	4	1.0		
		CM	AG	CC			

SITE 445		HOLE		CORE 17		CORED INTERVAL: 151.0-160.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Miocene	Ceratolithus acutus Subzone (N1) / Ceratolithus rugosus Subzone N18(?)	CG	AG	1	0.5		<p>Dominant Lithology: Nannofossil Chalk, light gray with white. No glass, but some lapilli (Section 2). Little clay. Sediments are very, very firm - hard.</p> <p>Sedimentary structures: bioturbation, slump structures, inclined bedding.</p> <p>SMEARS:</p> <p>1-58 (light), 1-58 (dark), CC (Nannofossil Chalk)            Clay minerals TR-5% Nannofossils 90-93%            Foraminifera 4-8%</p> <p>GRAIN SIZE: 1-10 (0.8, 27.5, 71.7)</p> <p>CARBON-CARBONATE: 1-74 (10.1, 0.1, 83)</p>
		CM	AG	2	1.0		
				CC			

SITE 445		HOLE		CORE 18		CORED INTERVAL: 160.5-170.0 m		
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS RADS					
Upper Miocene	N18	AG	AG	1	0.5		6Y 7.5/1	Dominant Lithology: Nannofossil Chalk and interbedded olive gray Clayey Nannofossil Chalk. Light gray to white. Very firm - hard sediments.
		CG	AG		1.0		6Y 7.5/1, 7/1	Minor Lithology: Nannofossil Foraminifera Chalk at base of bed (Section 2).  Sedimentary structures: slump structures with contorted and inclined bedding in Section 1. Mild to intense bioturbation. Chondrites and <i>Zoophycus</i> . Fining upward beds. Color change to very pale brown in Section 4.
		CG	AG	2			6Y 5/2	SMEARS: 1-75, 2-30, 2-100, 3-75, 3-100, 4-130, CC (Nannofossil Chalk) Clay minerals TR-10% Nannofossils 85-96% Foraminifers TR: 4%
		CG	AG		6Y 7/1			
		CG	AG	3			6Y 7/1, 7.5/1	4-36 (Clayey Nannofossil Chalk) Sand 0% Clay minerals 15% Silt 10% Nannofossils 80% Clay 90%
CG	AG	6Y 5/2						
FG	AG	4			6Y 7.5/1	2-56 (Nannofossil-Foraminifera Chalk) Foraminifers 59% Carbonate unspecified 20% Nannofossil 20%		
RP	AG		6Y 7/2		GRAIN SIZE: 1-43 (1.9, 30.3, 67.9) 3-43 (0.1, 31.5, 68.4)			
FM	AG	5			5Y 5/3	CARBONATE BOMB: 3-77 (88)		
FM	AG		10YR 7/3		CARBON-CARBONATE: 1-40 (8.7, 0.1, 80) 3-40 (7.3, 0.1, 61) 5-40 (9.3, 0.0, 77)			
		CC					10YR 7/3	

SITE 445		HOLE		CORE 19		CORED INTERVAL: 170.0-179.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Miocene	N17(?)	CG	AG	1	0.5		10YR 6/2, 10YR 7/3
		CG	AG		1.0		Sedimentary structures: bioturbation - mild. Long massive section in Section 2 with no sedimentary structures but faint mottling to intense mottling with abundant deformation and contorted bedding Section 3. Parallel laminae contorted. Slump structures in Section 4.
		CM	AG	2			10YR 7/3, 7/2
		FM	AG		10YR 7/3, -8/3, -8/2		
		CM	AG	3			10YR 7/3, -8/3, -8/2
FM	AG	10YR 7/3, -8/3, -8/2					
CM	AG	FP	CC				10YR 7/3, -8/3, -8/2

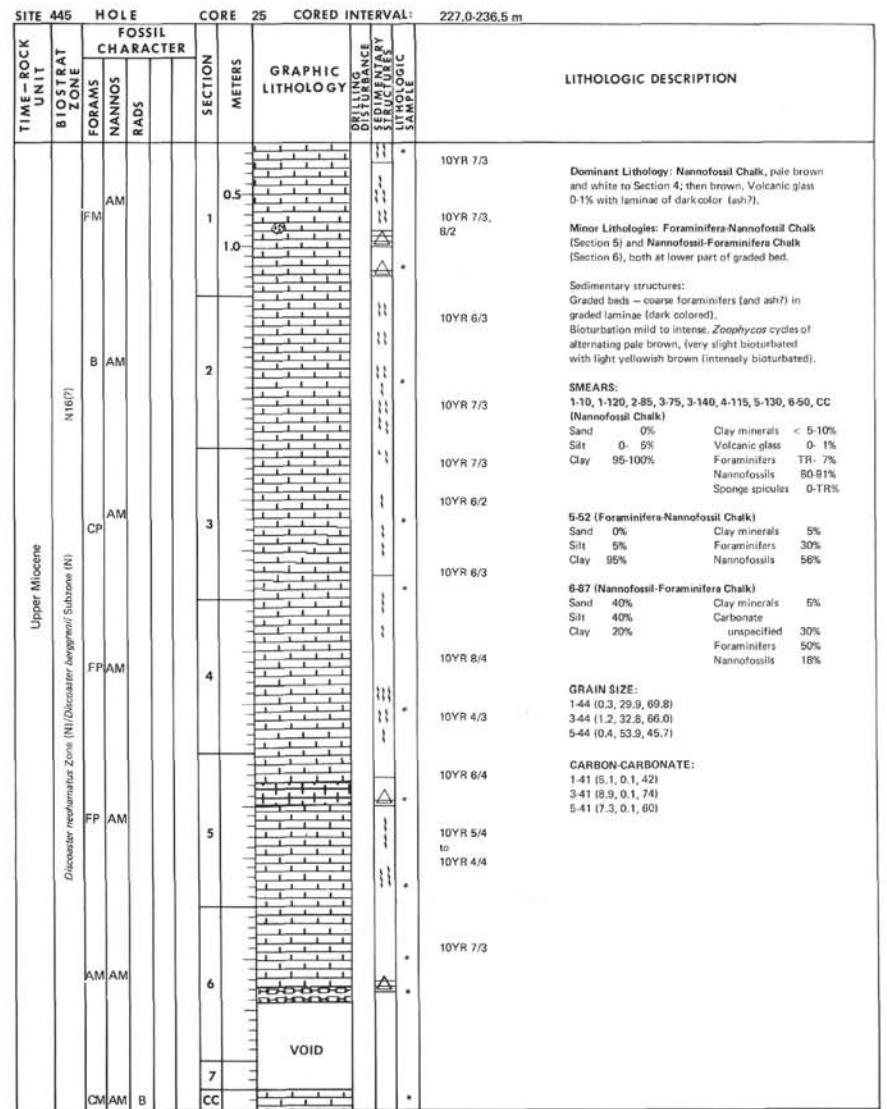
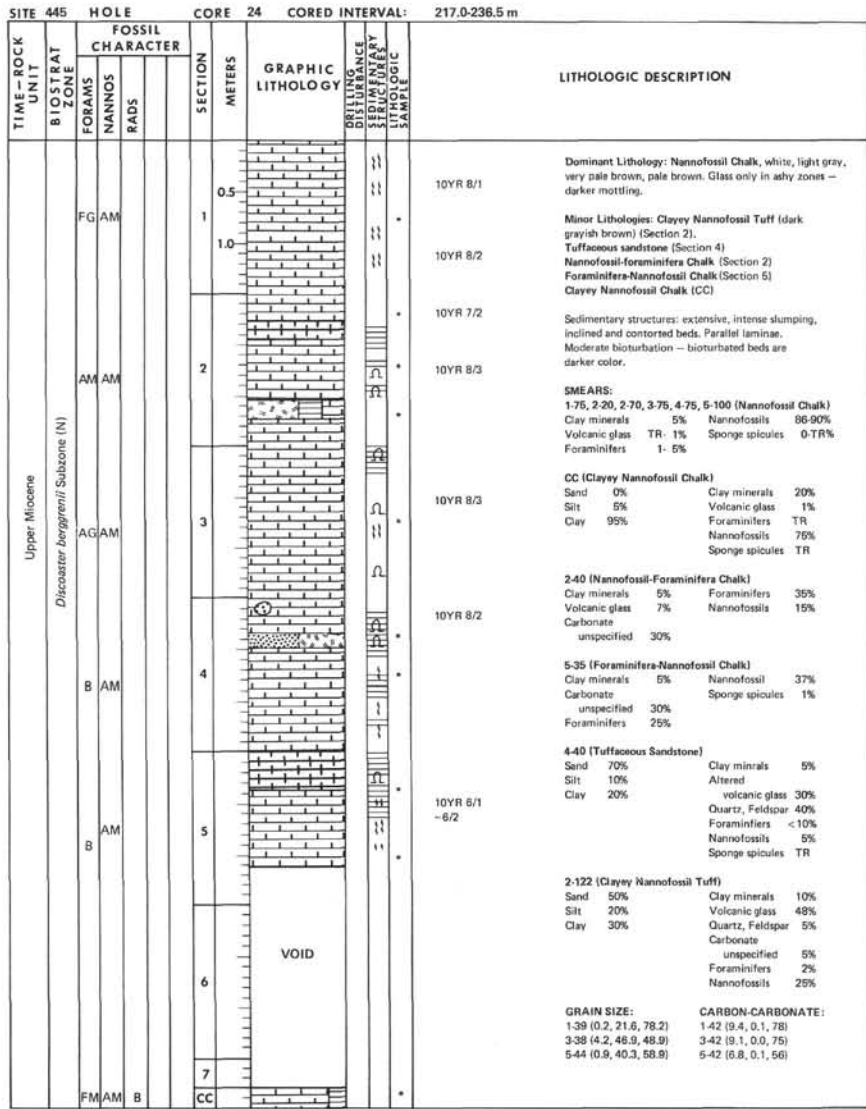
SITE 445		HOLE		CORE 20		CORED INTERVAL: 179.5-189.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Miocene	Amaurolithus primus Subzone (N)	CG	AG	1		10YR 7/3 Dominant Lithology: Nannofossil Chalk, very pale brown with minor amounts of pelagic clay. Volcanic glass 0-2% with an increase in ash zones.  Minor Lithology: Vitric Nannofossil Chalk (Section 1) pale brown.  Sedimentary structures: slump structures - contorted and inclined bedding. Bioturbation.  <b>SMEARS:</b> 1-90, 2-90, 3-15, CC (Nannofossil Chalk) Clay minerals TR- 5% Nannofossils 81-91% Volcanic glass 0- 2% Sponge spicules 0- 1% Foraminifers 1- 5%  1-56 (Vitric Nannofossil Chalk) Sand 50% Clay minerals < 5% Silt 20% Quartz, Feldspar 7% Clay 30% Volcanic glass 32% Foraminifers 5% Nannofossils 55% Sponge spicules TR  <b>GRAIN SIZE:</b> CARBON-CARBONATE: 1-24 (1.5, 33.5, 65.0) 1-29 (9.5, 0.0, 79) 3-24 (3.9, 39.4, 56.7) 3-29 (9.5, 0.0, 78)	
		AG	AG	2			
		CM	AG RP	CC			

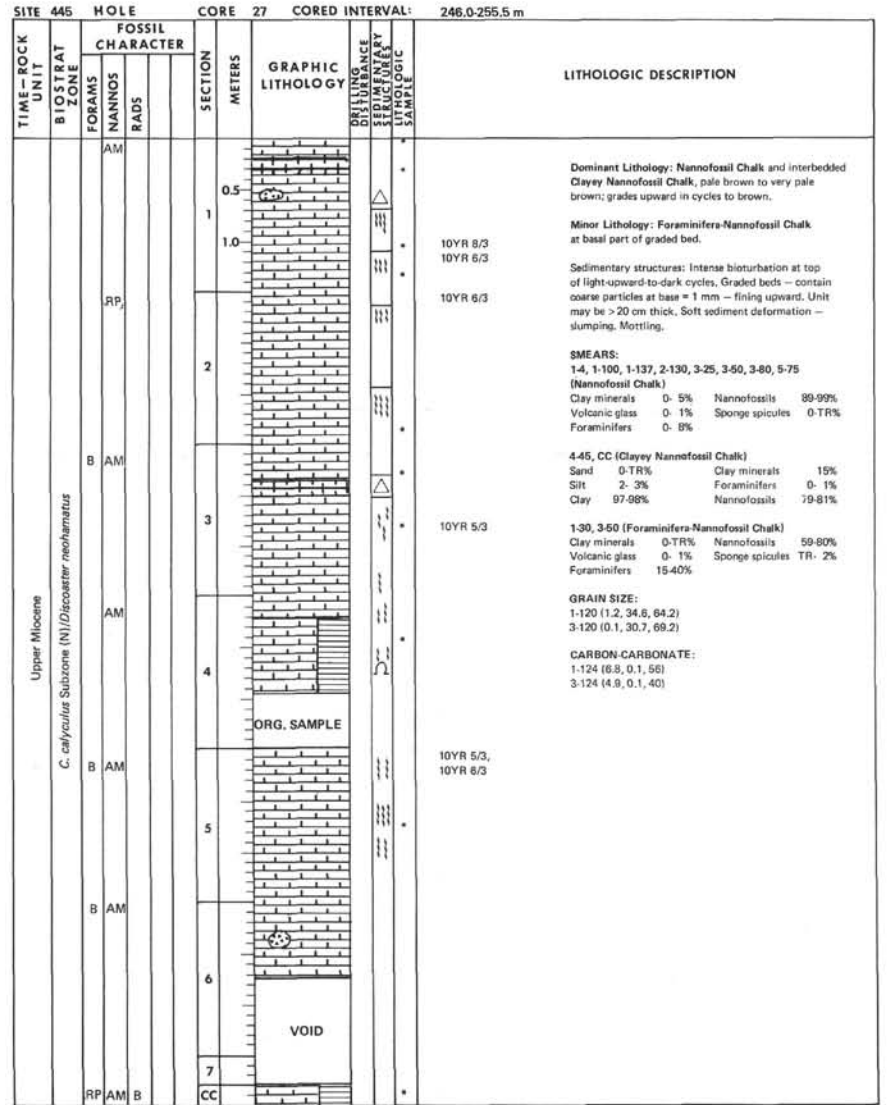
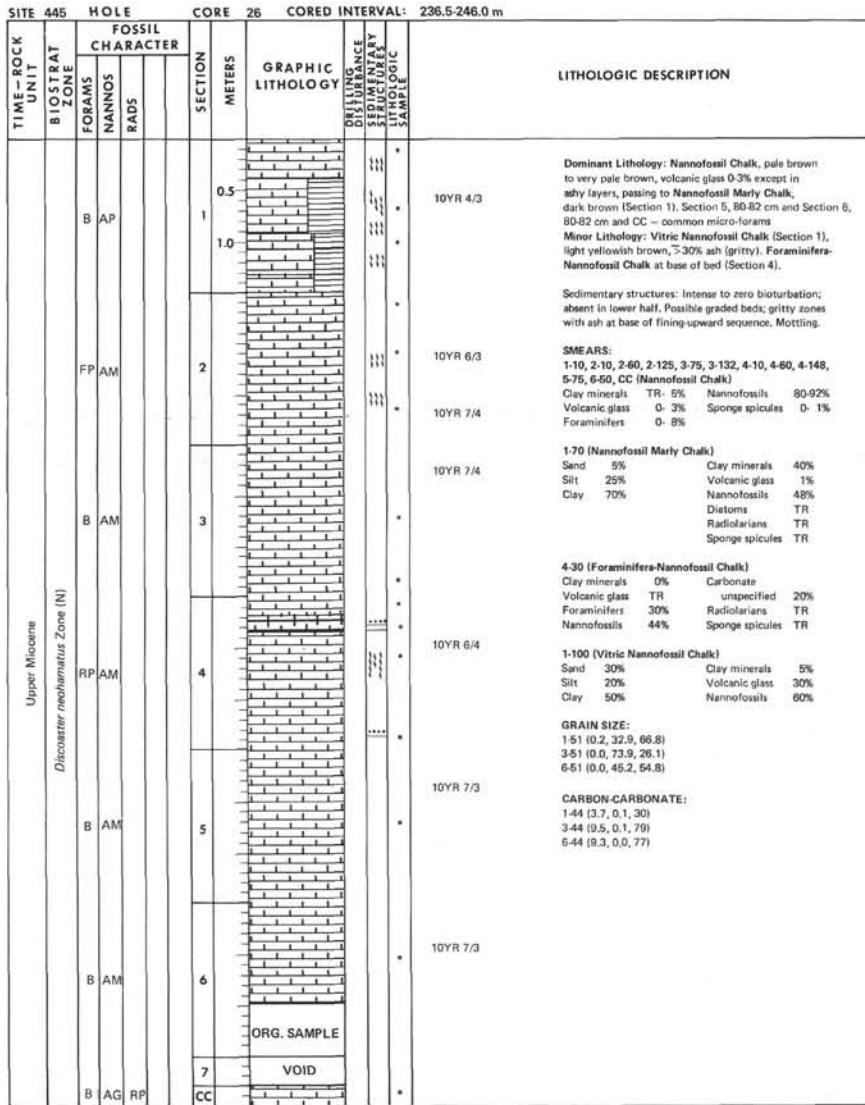
SITE 445		HOLE		CORE 21		CORED INTERVAL: 189.0-198.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Miocene	Amaurolithus primus Subzone (N)	AG		1		10YR 7/3, 10YR 6/3 Dominant Lithology: Nannofossil Chalk, pale brown to very pale brown (alternation).  Sedimentary structures: parallel laminae. Intensive slumping, contorted bedding.  <b>SMEARS:</b> 1-48, 1-75, 1-100, 2-65, 2-100, CC (Nannofossil Chalk) Clay minerals TR-10% Nannofossils 80-85% Volcanic glass TR- 5% Sponge spicules TR- 1% Foraminifers 2- 5%  <b>GRAIN SIZE:</b> 1-60 (1.6, 33.1, 65.3)  <b>CARBON-CARBONATE:</b> 1-64 (9.1, 0.0, 75)	
		CM	AG	2			
		FM	AG RP	CC			

SITE 445		HOLE		CORE 22		CORED INTERVAL: 198.5-208.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Miocene	Discoaster berggrenii Subzone (N) Amaurolithus primus Subzone (N)	CM	AG	1		10YR 8/1 10YR 8/2 10YR 7/2 10YR 8/1 Dominant Lithology: Nannofossil Chalk, white to light gray (alternating), 0-1% volcanic glass.  Sedimentary structures: parallel laminae, slump structures.  <b>SMEARS:</b> 1-75, 1-95, 2-10, 2-29, CC (Nannofossil Chalk) Clay minerals TR-10% Nannofossils 81-93% Volcanic glass 0- 1% Sponge spicules 0-TR% Foraminifers 0- 3%  <b>GRAIN SIZE:</b> 1-39 (1.0, 27.8, 71.3)	
		CG	AG	2			
		RM	AG B	CC			

SITE 445		HOLE		CORE 23		CORED INTERVAL: 208.0-217.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS				
Upper Miocene	Discoaster berggrenii Subzone		AM	1		10YR 6/3 10YR 8/3 10YR 8/2 10YR 8/1 10YR 8/2 10YR 8/2, -7/2 Dominant Lithology: Nannofossil Chalk associated with Nannofossil Marly Chalk at top of core. Pale brown, very pale brown, and white, light gray lower.  Minor Lithology: Nannofossil Tuff (Sections 2 and 3), brown; dark lamination. Nannofossil-Foraminifera Chalk (Section 3).  Sedimentary structures: parallel laminations in beds with sharp basal contacts. Ash beds laminated. Graded beds; laminae. Bioturbation.  <b>SMEARS:</b> 1-75, 2-120, 3-10, 3-75, CC (Nannofossil Chalk) Clay minerals TR- 5% Foraminifers 0- 3% Volcanic glass 0- 2% Nannofossil 71-90%  1-2 (Nannofossil Marly Chalk) Sand 10% Clay minerals 30% Silt 20% Volcanic glass <10% Clay 20-30% Nannofossils 48% Sponge spicules 1%  2-56, 3-15 (Nannofossil Tuff) Sand 50-60% Clay minerals 5-10% Silt 20% Volcanic glass 52-55% Clay 20-30% Foraminifers 0- 1% Nannofossils 29-30%  3-120 (Nannofossil-Foraminifera Chalk) Clay minerals 5% Nannofossils 34% Foraminifers 50%  <b>GRAIN SIZE:</b> CARBON-CARBONATE: 1-29 (0.9, 40.0, 59.2) 1-33 (4.2, 0.1, 34) 3-29 (0.6, 28.1, 71.3) 3-33 (8.9, 0.1, 74)	
		FP	AG	2			
		AM	AG	3			

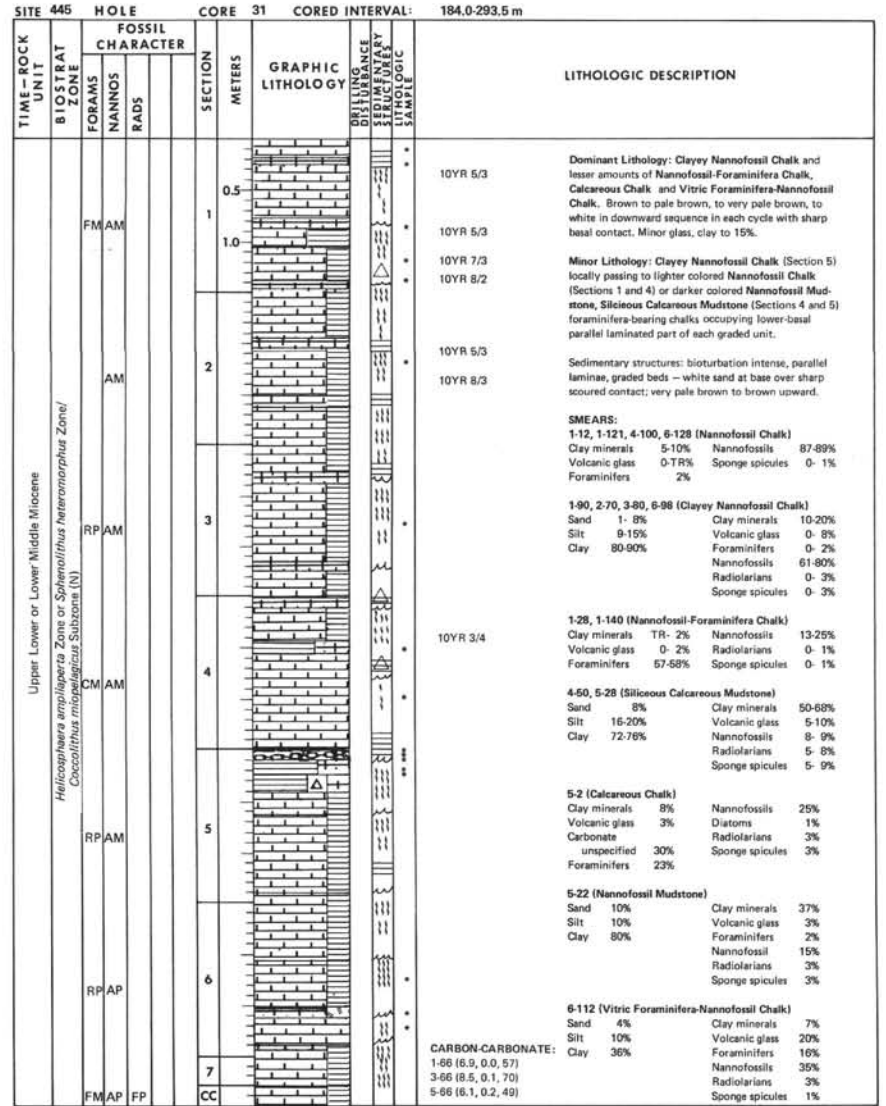
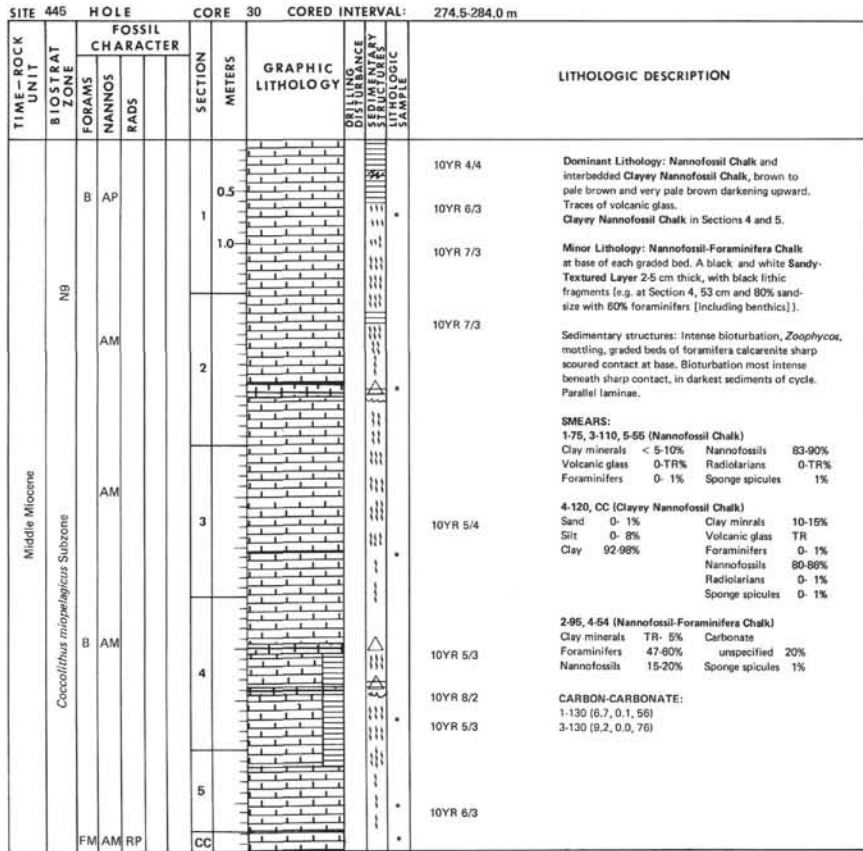




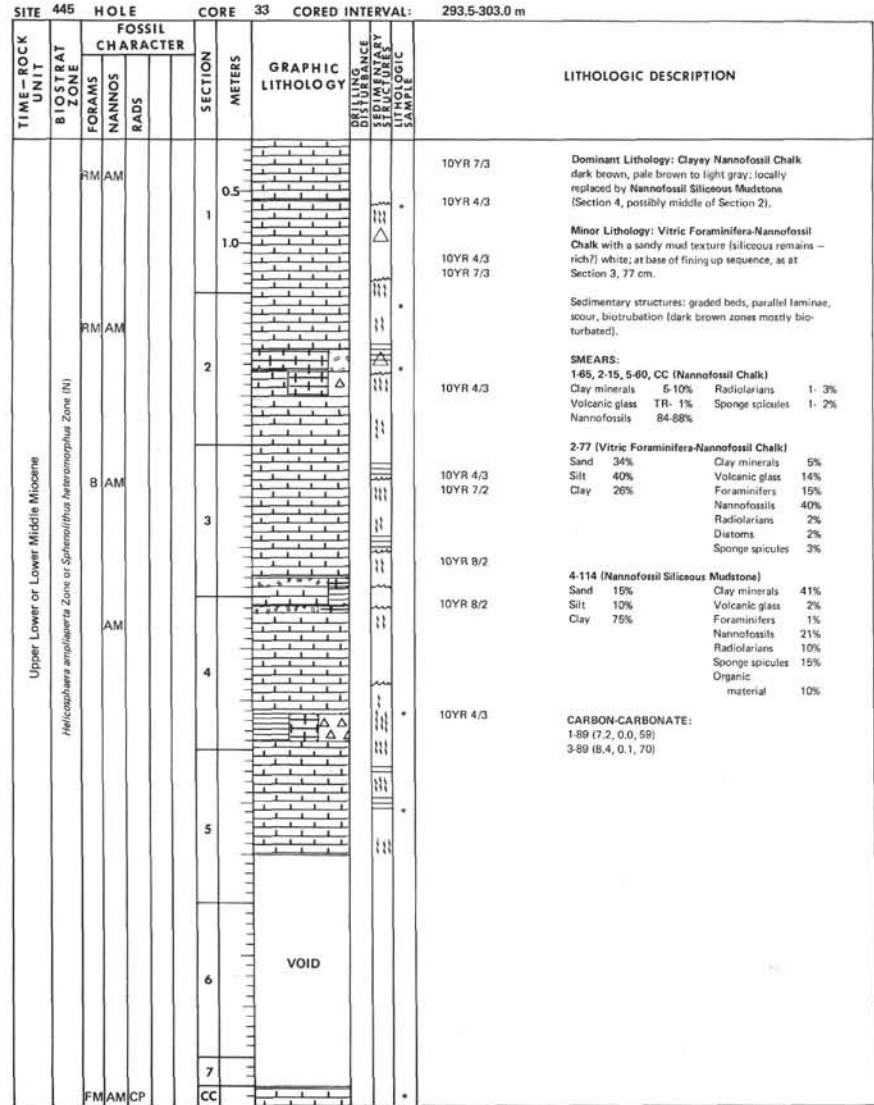
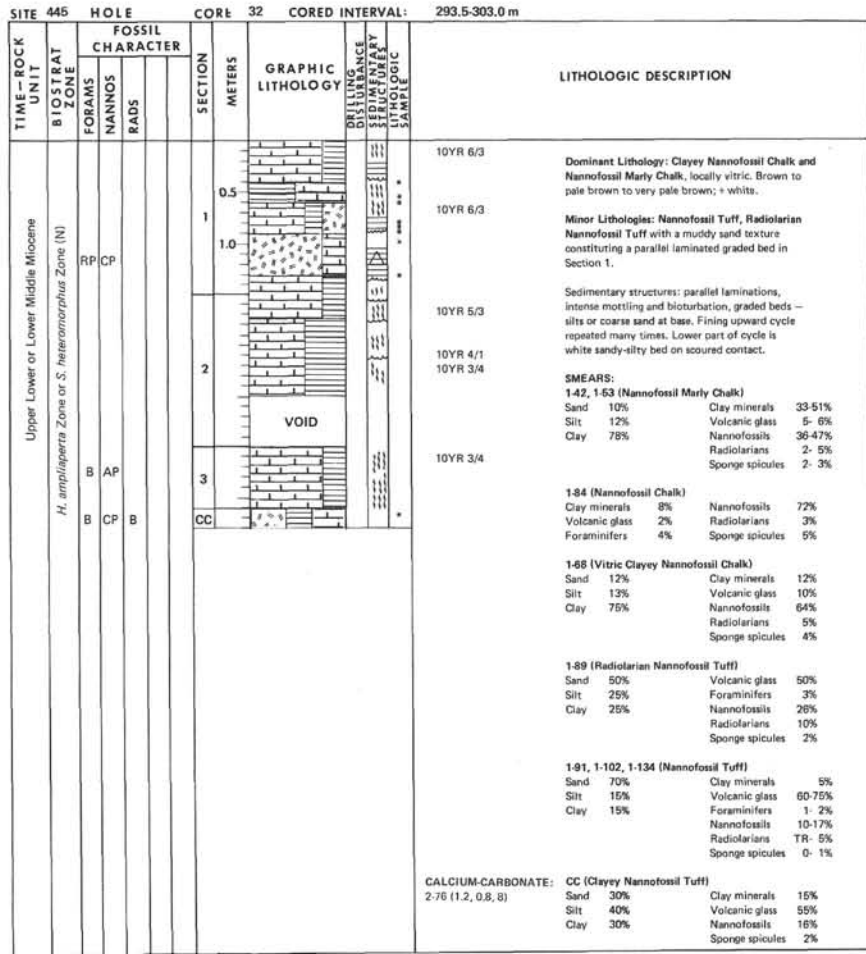


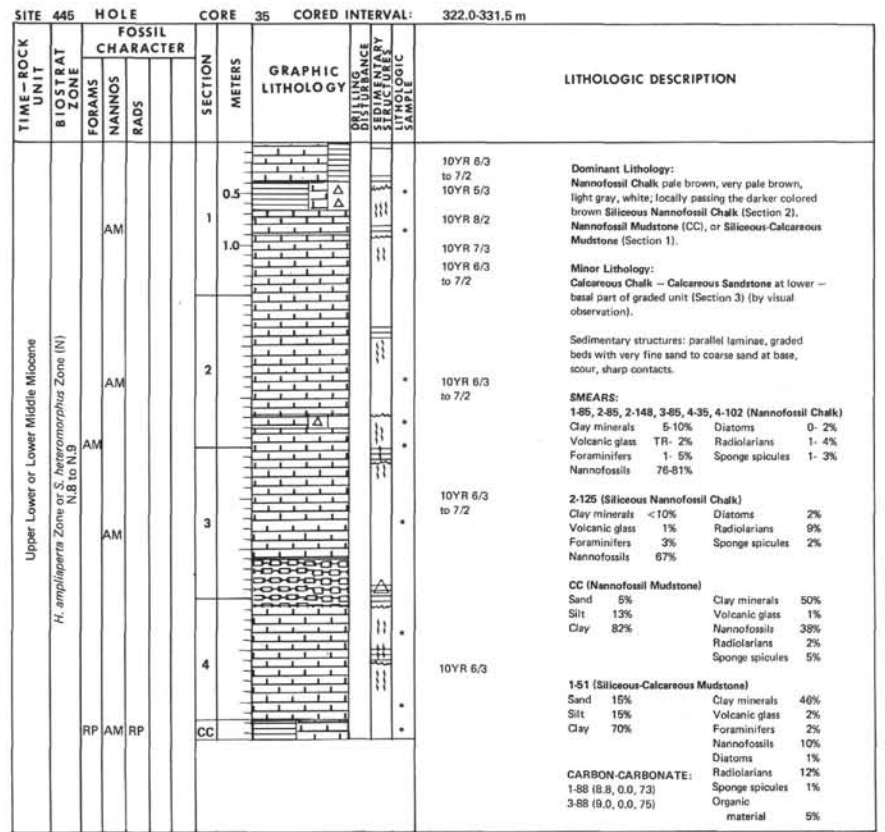
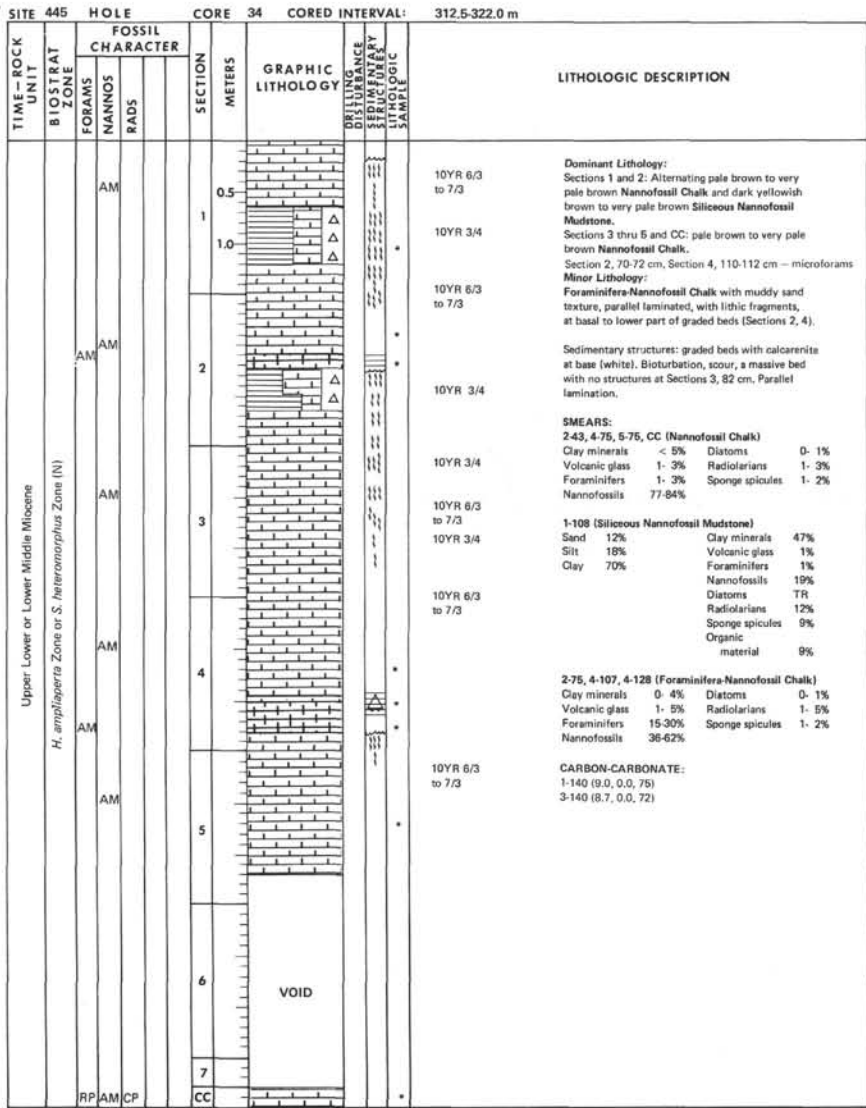
SITE 445		HOLE		CORE 28		CORED INTERVAL: 255.5-265.0 m						
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	ORILLING DISTURBANCE	SEDIMENTARY STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION		
		FORAMS	NANNOS	RADS								
Upper Miocene	Helicospira carteri Subzone (N)/Caryoceras calyculus Subzone (N)	B	AM		1					10YR 5/3 10YR 8/3	<p><b>Dominant Lithology:</b> (1) <b>Nannofossil Chalk</b>, colors in cycles, grading upward from pale brown to very pale brown. Cycles darker in lower part, from dark brown grading up to brown to pale brown; locally passing to <b>Clayey Nannofossil Chalk</b> (CC).</p> <p>(2) <b>Nannofossil Mudstone</b> (Sections 4 and 5), dark brown to pale brown, darkens toward base, clay up to 70%.</p> <p><b>Minor Lithology:</b> Foraminifera-Nannofossil Chalk above base of bed, Section 2.</p> <p><b>Sedimentary structures:</b> Intense mottling, graded beds in cycles and subcycles (e.g. clay-ooze to ooze in large cycle; smaller ones within). Sharp base. Intense bioturbation. Parallel laminae (faint).</p> <p><b>SMEARS:</b> 1-21, 2-75, 3-75 (Nannofossil Chalk) Clay minerals 5-10% Nannofossils 82-88% Volcanic glass TR- 1% Sponge spicules TR- 2% Foraminifers TR- 5%</p> <p><b>CC (Clayey Nannofossil Chalk)</b> Sand 0% Clay minerals 20% Silt 5% Volcanic glass 1% Clay 95% Foraminifers 2% Nannofossils 71%</p> <p><b>4-95, 5-75 (Nannofossil Mudstone)</b> Sand 0- 2% Clay minerals 52-70% Silt 5- 9% Volcanic glass 1% Clay 91-93% Nannofossils 25-40% Sponge spicules 1%</p> <p><b>2-93 (Foraminifera-Nannofossil Chalk)</b> Clay minerals &lt;10% Nannofossils 41% Volcanic glass 2% Carbonate Foraminifers 20% unspecified 26%</p> <p><b>GRAIN SIZE:</b> 1-114 (0.1, 28.0, 71.8)</p> <p><b>CARBONATE BOMB:</b> 1-40 (69) 5-12 (11.5)</p> <p><b>CARBON-CARBONATE:</b> 1-120 (5.5, 0.1, 45)</p>	
		B	AG		2					10YR 6/4 10YR 6/3		
		RP	AM		3						10YR 5/3 10YR 6/3 10YR 8/3	
		FP									10YR 5/3 10YR 6/3	
		AM			4						10YR 6/3 10YR 5/3 10YR 4/3	
Middle Miocene		B	AM		5					10YR 5/3 10Y 4/3 3/3		
					CC							

SITE 445		HOLE		CORE 29		CORED INTERVAL: 265.0-274.5 m						
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	ORILLING DISTURBANCE	SEDIMENTARY STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION		
		FORAMS	NANNOS	RADS								
Middle Miocene	Disconer rugifer Subzone (N)/Caryoceras calyculus Zone or Helicospira carteri Subzone (N)	AM			1					10YR 5/3 ~7/3	<p><b>Dominant Lithology:</b> Nannofossil Chalk, brown to pale brown to very pale brown. Volcanic glass - trace; locally passing to <b>Clayey Nannofossil Chalk</b> (Section 1).</p> <p><b>Minor Lithology:</b> Foraminifera Chalk.</p> <p><b>Sedimentary structures:</b> bioturbation mild to intense. Possible graded beds - coarse bed of foraminifera-chalk at base of Section 1, 25-43 cm. Mottled. Parallel laminae (faint).</p> <p><b>SMEARS:</b> 2-75, 3-75, 4-75, CC (Nannofossil Chalk) Clay minerals &lt; 10% Nannofossils 84-89% Volcanic glass 0-TR% Sponge spicules TR Foraminifers 0- 1%</p> <p><b>1-75 (Clayey Nannofossil Chalk)</b> Sand TR Clay minerals 20% Silt 9% Nannofossils 75% Clay 91% Radiolarians TR Sponge spicules TR</p> <p><b>GRAIN SIZE:</b> 3-52 (0.5, 33.3, 66.2)</p> <p><b>CARBON-CARBONATE:</b> 2-47 (7.0, 0.1, 57)</p>	
		RP	AM		2						10YR 6/3	
		AM			3						10YR 5/3	
		FP	AM		4						10YR 5/3	
		RM	AP	B	CC							



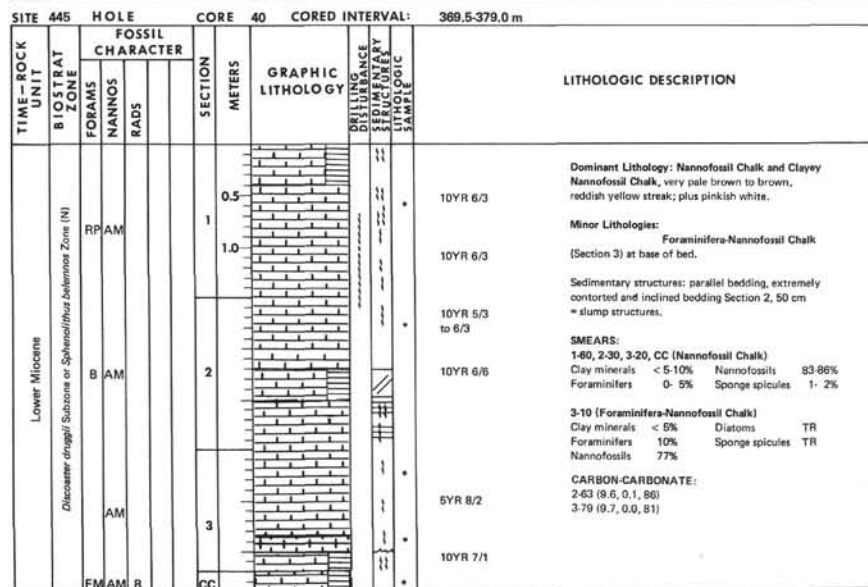
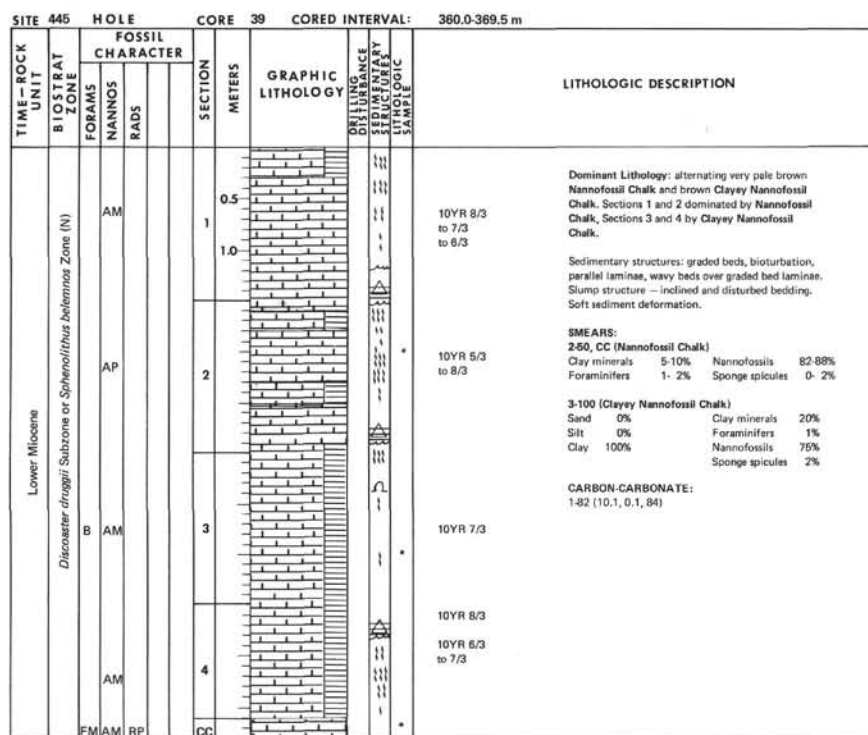
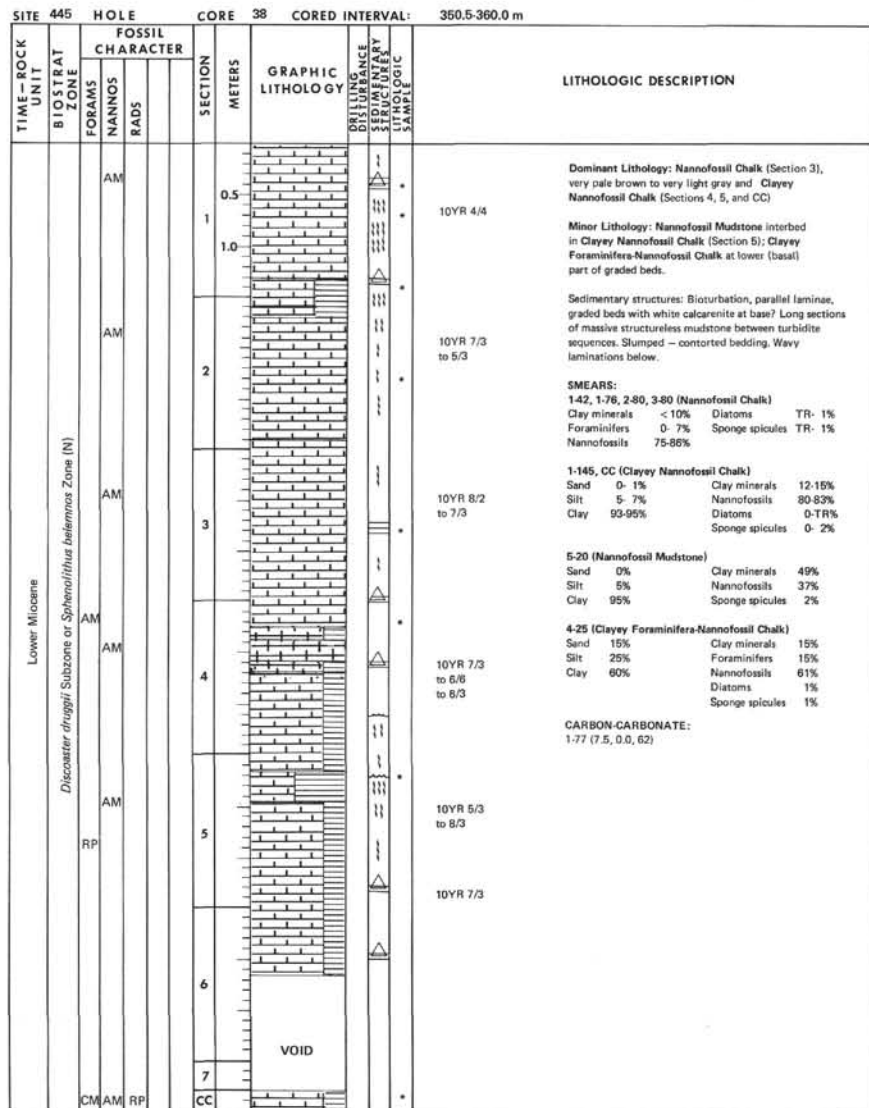


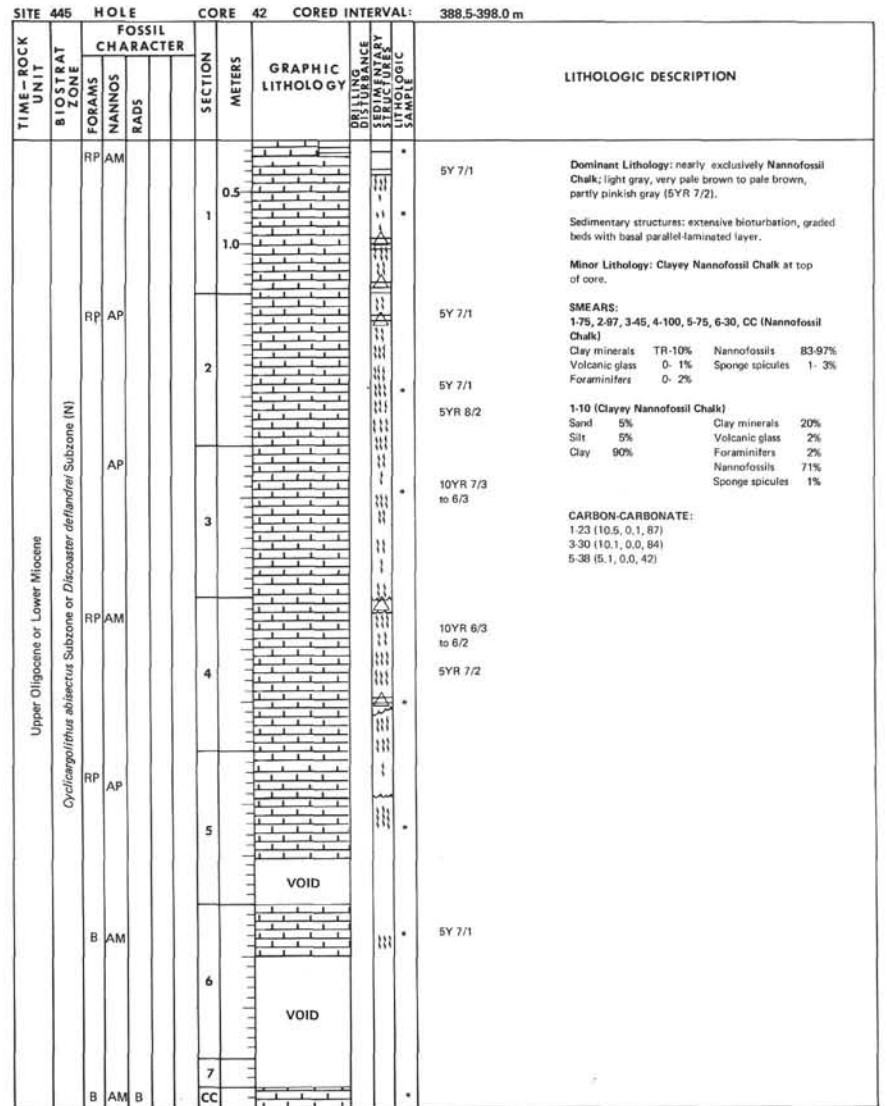
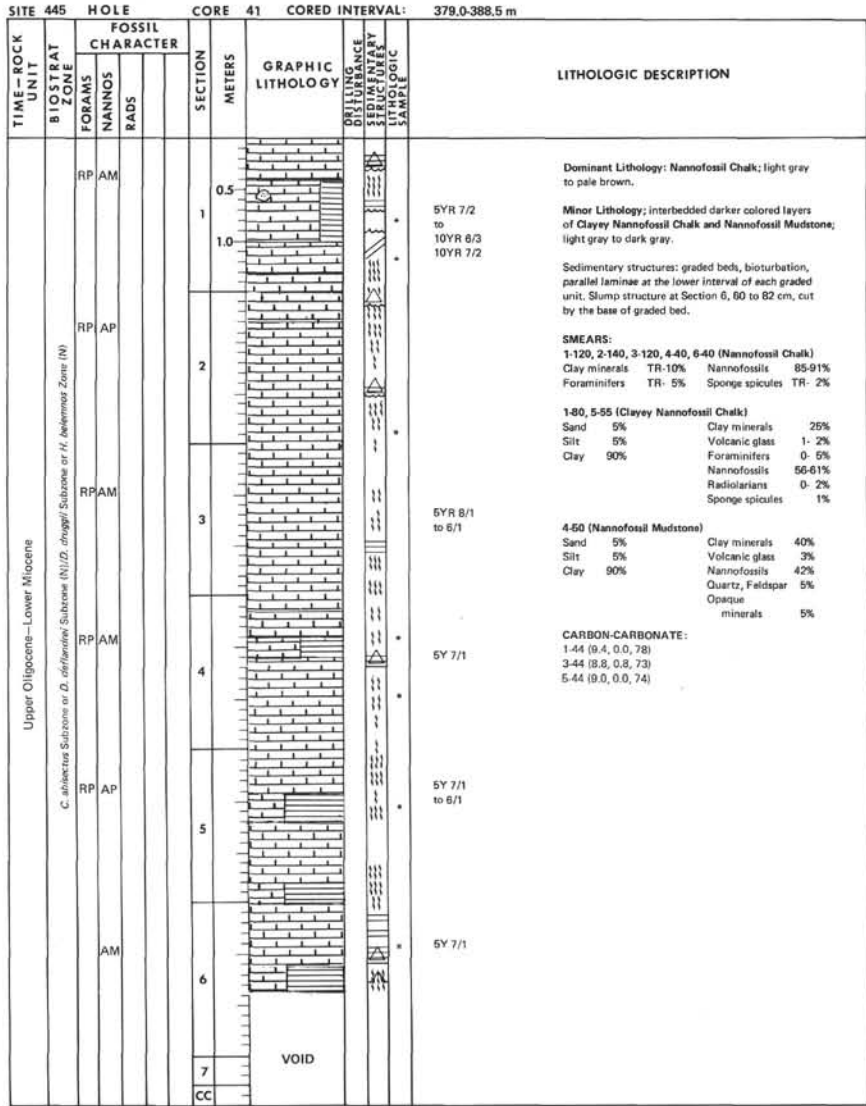




TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS				
Lower Miocene	<i>Discoaster druggii</i> Subzone or <i>Sphenolithus belemnos</i> Zone (N)	AM			0.5		10YR 6/3	Dominant Lithology: Alternating Nannofossil Chalk and Clayey Nannofossil Chalk, pale brown to very pale brown to light gray. Clay 4-22%; volcanic glass 0-5%.
		CM			1.0		10YR 6/3 to 7/2	Minor Lithology: (1) Calcareous Mudstone (Section 5, 52-150 cm), pale brown, 72% clay. (2) Siliceous Nannofossil Mudstone (Section 1), reddish yellow, dark brown. (3) Calcareous Chalk with muddy sand to sandy mud texture at base of sequences.
		AM			2		10YR 6/3 to 7/2	Sedimentary structures: bioturbation, parallel laminae, graded beds, soft sedimentary deformation = microfault, weak stratification in the more massive layers in mid-cycle.
		AM			3		10YR 4/3 to 6/3	SMEARS: 1-75, 2-75, 4-75 (Nannofossil Chalk) Clay minerals <10% Diatoms TR- 2% Foraminifers 1- 3% Radiolarians 0-13% Nannofossils 77-83% Sponge spicules 1- 2% 3-75, CC (Clayey Nannofossil Chalk) Sand 2- 4% Clay minerals 16-22% Silt 8-17% Foraminifers 1% Clay 79-90% Nannofossils 64-77% Diatoms 0- 1% Radiolarians 0- 3% Sponge spicules 1- 2%
		AP			4		10YR 6/3 to 4/3	1-57, 5-75 (Calcareous Mudstone) Sand 3% Clay minerals 72-77% Silt 9% Foraminifers 0- 1% Clay 90% Nannofossils 5- 8% Carbonate unspecified 5- 7% Diatoms 1% Radiolarians 3- 4% Sponge spicules 2- 4%
		AP			5		10YR 6/3	1-54 (Siliceous Nannofossil Mudstone) Sand 4% Clay minerals 65% Silt 15% Nannofossils 10% Clay 81% Diatoms 2% Radiolarians 4% Sponge spicules 4%
		AP			6		10YR 4/3	1-91 (Calcareous Chalk) Clay minerals 4% Foraminifers 10% Volcanic glass 5% Nannofossils 19% Carbonate Radiolarians 9%
		RM	AP	RP	CC		CARBON-CARBONATE: 1-39 (4.4, 0.1, 36) unspecified 37%	

TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS				
Lower Miocene	<i>Discoaster druggii</i> Subzone or <i>Sphenolithus belemnos</i> Zone (N)	CP			0.5		10YR 4/3	Dominant Lithology: (very hard) Nannofossil Chalk and less amounts of Clayey Nannofossil Chalk, Brown, dark brown - pale to very pale brown. Clay materials more dominant in Sections 4 and 5. Silty claystone interbed, pink (Section 4).
		AP			1.0		10YR 6/3	Sedimentary structures: parallel laminae, closely spaced. Graded beds. Horizontal layering of burrows. Bioturbation.
		AP			2		10YR 6/3 to 7/2	Minor Lithology: Foraminifera-Nannofossil Chalk, parallel laminae, with sandy-mud texture at lower part of graded bed (Section 4).
		AM			3		10YR 6/3 to 7/2	SMEARS: 1-75, 2-75, 3-71, 3-133, CC (Nannofossil Chalk) Clay minerals 5- 9% Radiolarians 0- 1% Foraminifers 0- 3% Diatoms 0-TR% Nannofossil 81-87% Sponge spicules 0- 2% 1-99, 4-61, 4-73, 5-75 (Clayey Nannofossil Chalk) Sand 2% Clay minerals 15-25% Silt 7-10% Foraminifers 0- 2% Clay 88-90% Nannofossils 65-78% Diatoms 0-TR% Radiolarians 0- 1% Sponge spicules 0- 2%
		AM			4		7.5YR 7/4	Color change
		CP			5		7.5YR 7/4	4-10, 4-36 (Foraminifera-Nannofossil Chalk) Clay minerals 6-10% Foraminifers 10% Carbonate Nannofossils 61-68% unspecified 10-16% Diatoms 0-TR%
		AP			6		VOID	4-85 (Silty Claystone) Sand 2% Clay minerals 78% Silt 16% Nannofossils 7% Clay 82%
		AP			7		VOID	CARBON-CARBONATE: 1-90 (7.3, 0.0, 60)
		FM	AP	RP	CC		7.5YR 6/3 to 7/2	

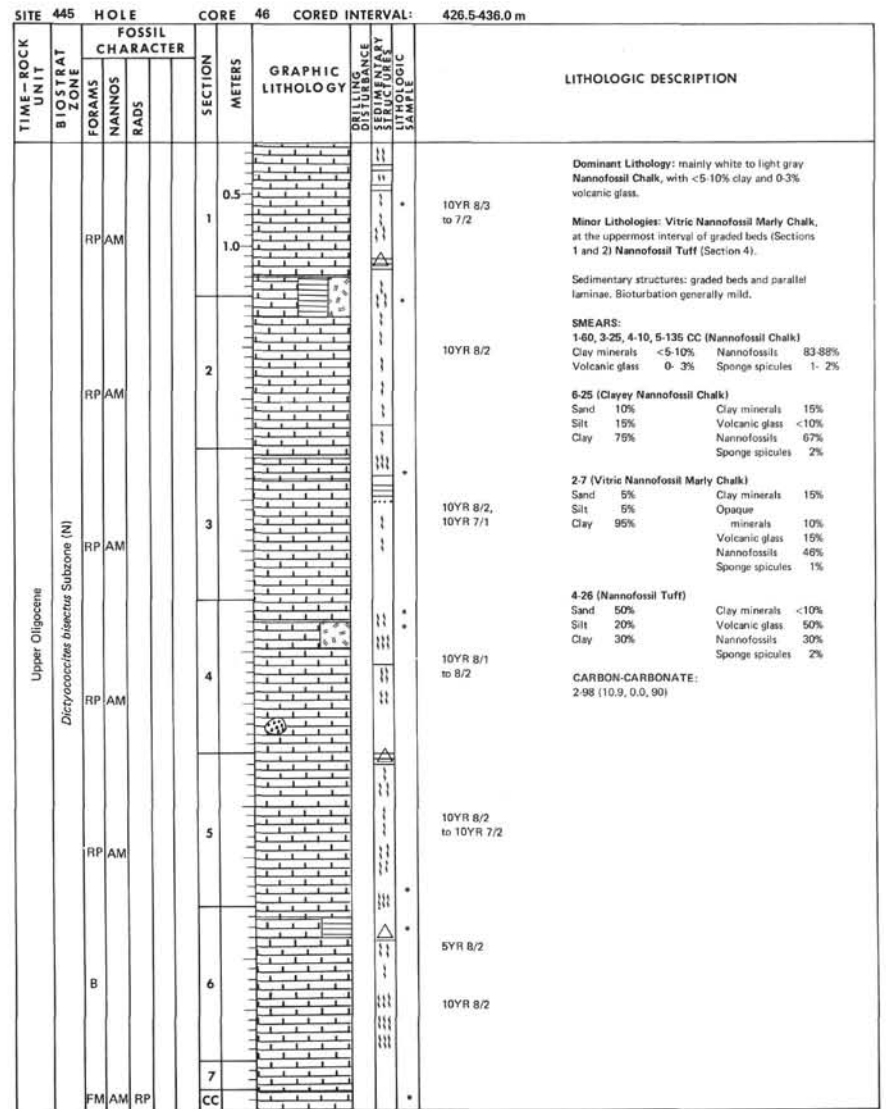
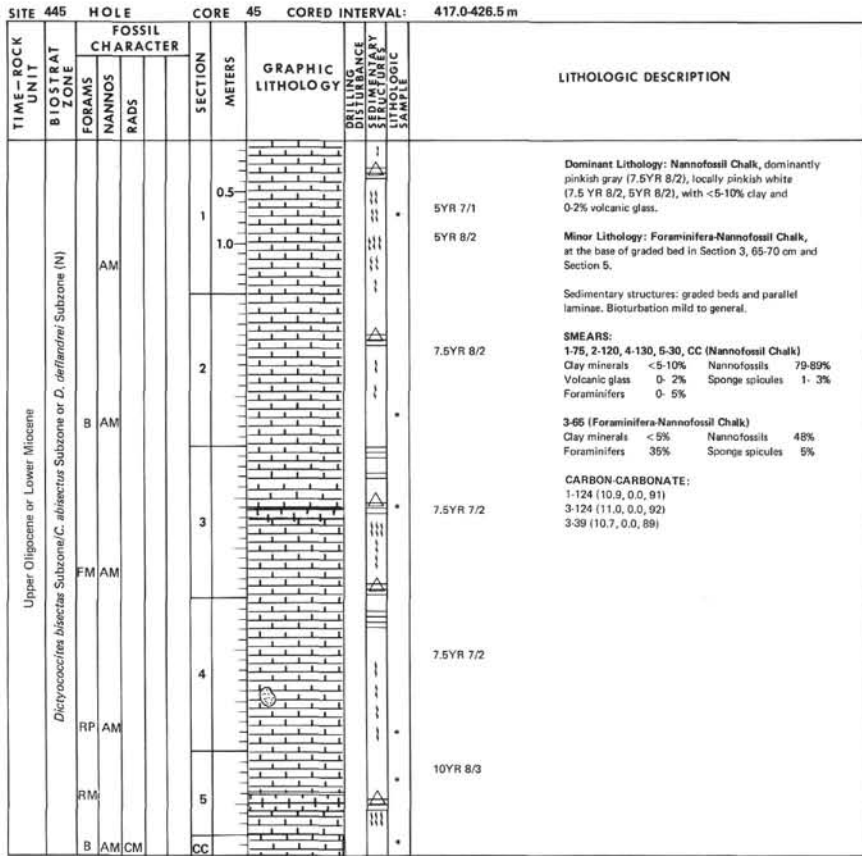


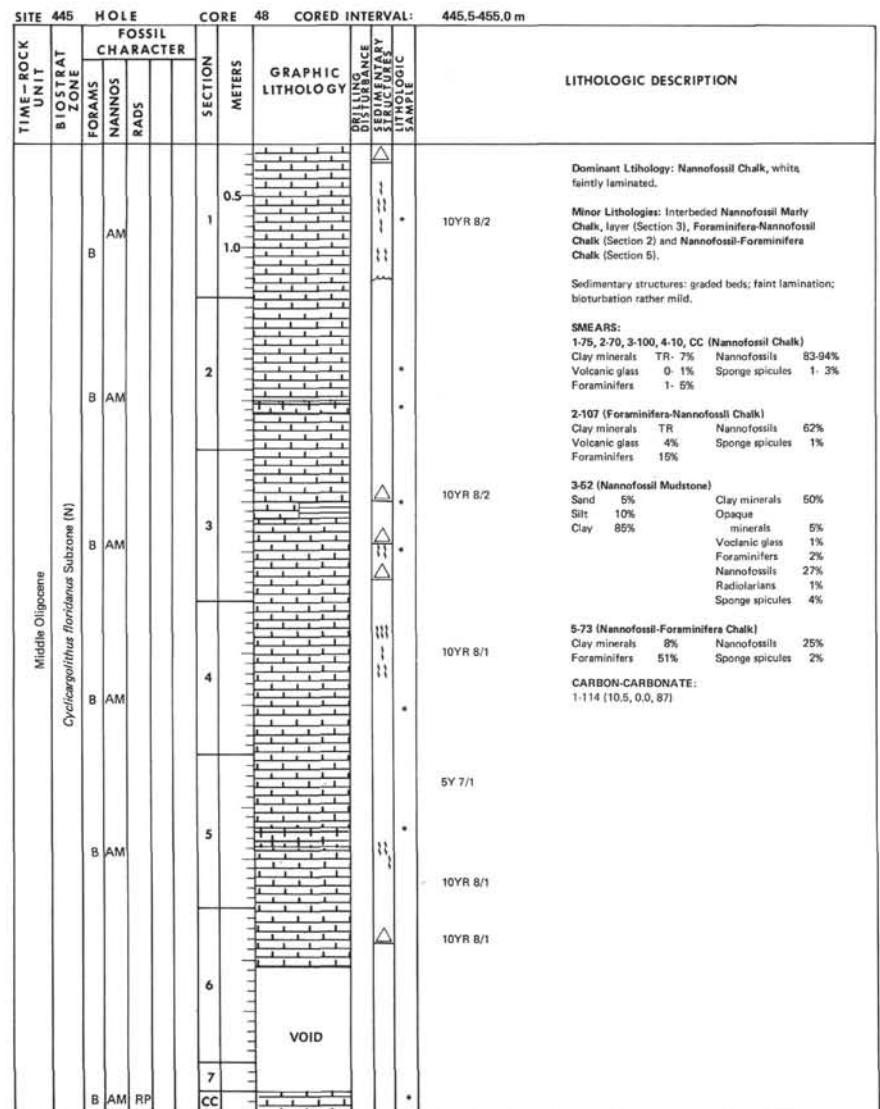
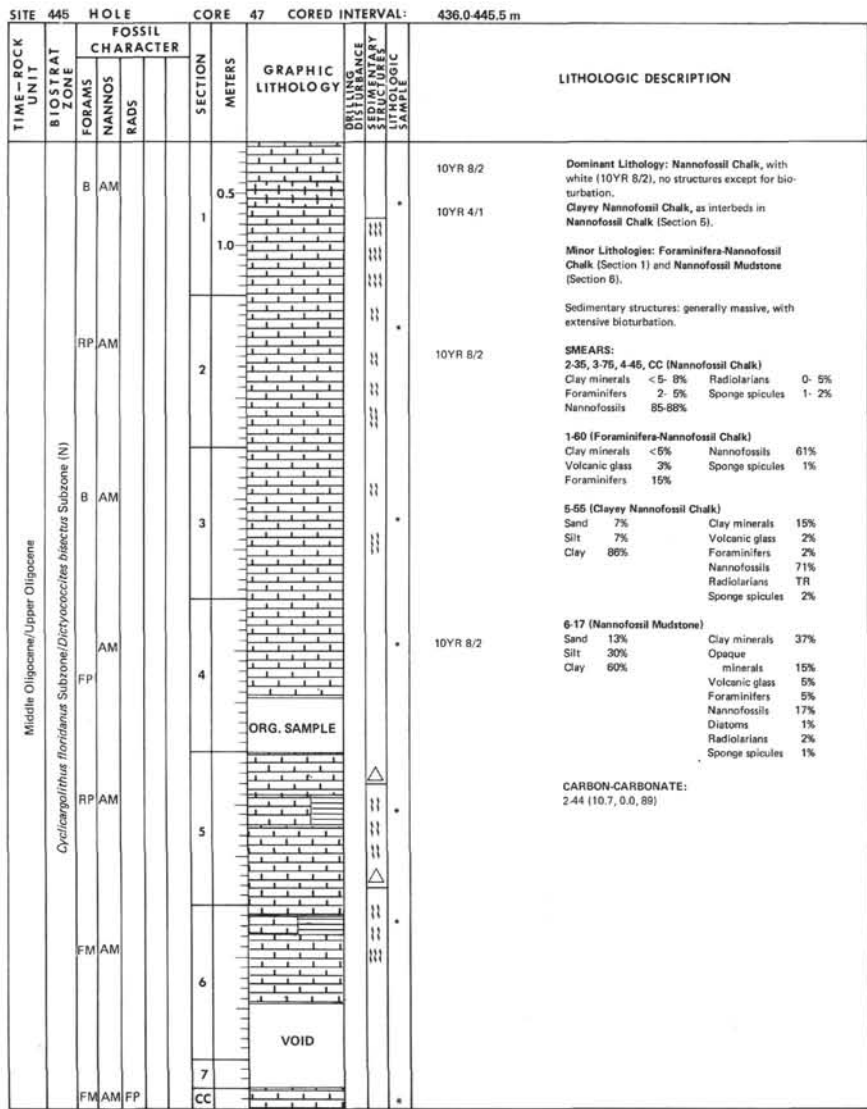




SITE 445		HOLE		CORE 43		CORED INTERVAL: 398.0-407.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Oligocene or Lower Miocene <i>Cyclargophilus abisectus</i> Subzone or <i>Discoaster delalandei</i> Subzone (N)	RP	AM		1	0.5		5GY 7/1
				1.0			
	B	AP		2			5GY 7/1
B	AM		3			5Y 7/1	
B	AP		4			10YR 4/3	
FM	AM	B	CC				

SITE 445		HOLE		CORE 44		CORED INTERVAL: 407.5-417.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Upper Oligocene or Lower Miocene <i>Cyclargophilus abisectus</i> Subzone (N) or <i>D. delalandei</i> Subzone (N)	RP	AM		1	0.5		10YR 7/3 to 6/2
				1.0			
	B	AM		2			10YR 6/2
B	AM		3			10YR 6/2	
B	AM		4			10YR 8/2	
CM	AM	B	CC				





SITE 445		HOLE		CORE 49		CORED INTERVAL: 455.0-464.5 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SERVICED LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS	RADS					
Middle Oligocene <i>Cyclacargolithus floridanus</i> Subzone (N)	N3/P22(7)	AM			0.5			10YR 8/1 to 7/1	
					1.0			10YR 7/1	
		CM	AM			2			10YR 8/1
		CM	AM			3			10YR 8/1
		AM				4			10YR 7/1
		AM				5			10YR 8/1
		B7	AM	RP		CC	6	VOID	

SITE 445		HOLE		CORE 50		CORED INTERVAL: 464.5-474.0 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SERVICED LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS	RADS					
Middle Oligocene <i>Cyclacargolithus floridanus</i> Subzone (N)		RM	AM		0.5			10YR 8/1	
					1.0			N4	
		AM				2			10YR 8/1
		RM	AM	RP		CC			10YR 7/1

SITE 445		HOLE		CORE 51		CORED INTERVAL: 474.0-483.5 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SERVICED LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS	RADS					
Middle Oligocene <i>Cyclacargolithus floridanus</i> Subzone (N)		B	AM		0.5			10YR 7/1	
					1.0			10YR 8/1	
		RP	AM			2			10YR 8/1
		FM	AP	CP		3			10YR 8/1
					4			10YR 8/1	

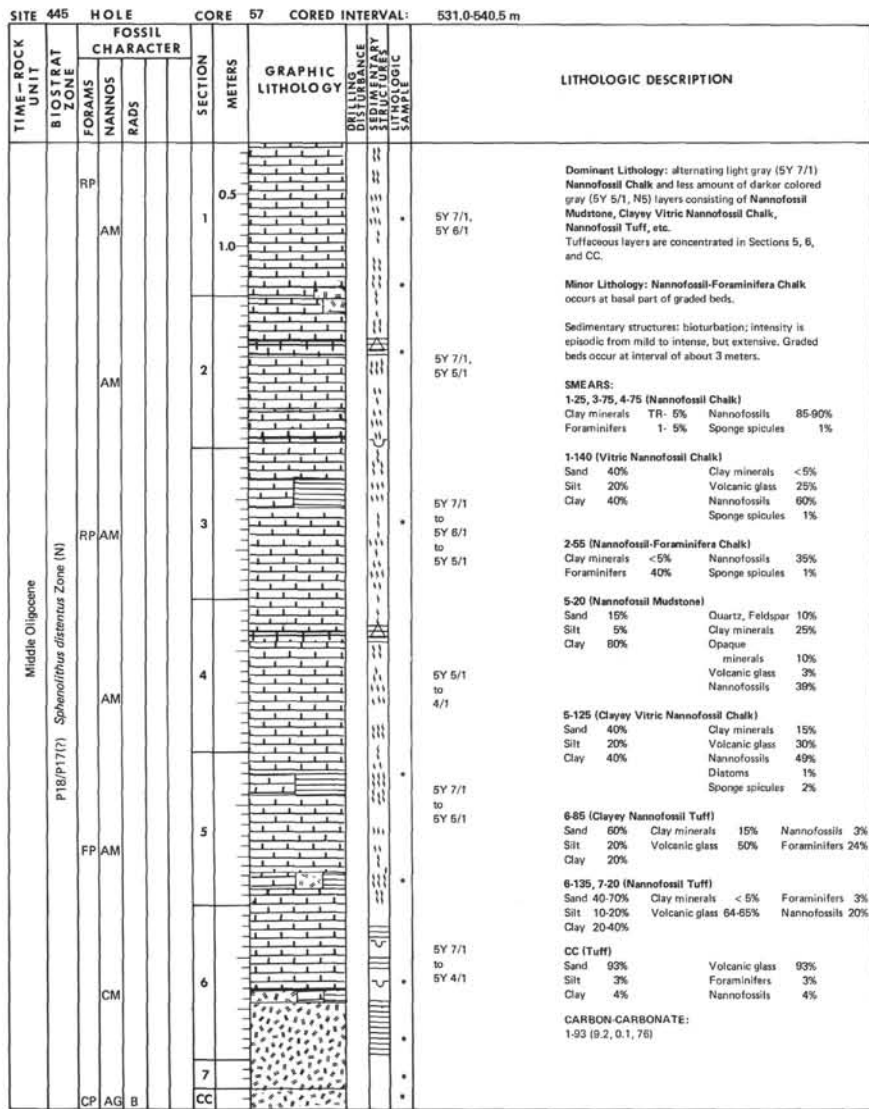
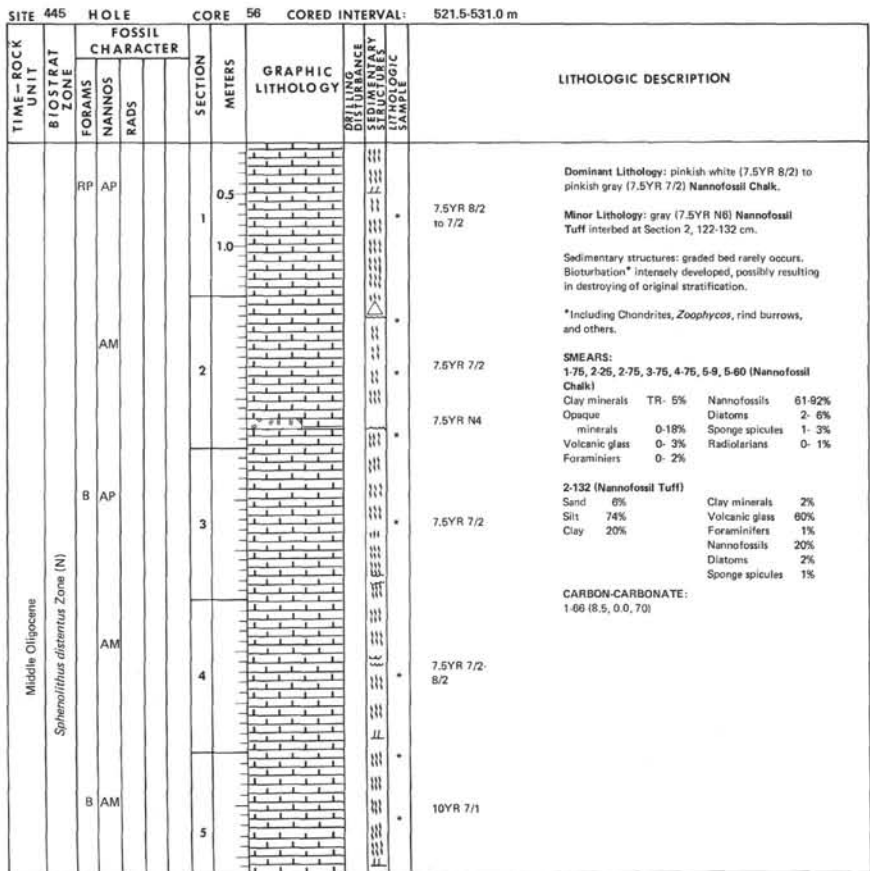
SITE 445		HOLE		CORE 52		CORED INTERVAL: 483.5-493.0 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Middle Oligocene <i>Sphenolithus distentus</i> Zone (N)C. <i>Iloridanus</i> Subzone (N)	RP	AM			0.5		7.5YR 8/2 to 7/2  Dominant Lithology: pinkish white (7.5YR 8/2) to pinkish gray (7.5YR 7/2) Nannofossil Chalk.  Minor Lithology: Pinkish white Siliceous Nannofossil Chalk (coarse grained, with black lithic fragments) layer at lower part of a graded bed (Section 2, 40-80 cm), associated with basal Calcareous Chalk (visual observation).
					1.0		
	AM				2		7.5YR 8/2  Sedimentary structures: graded beds with parallel lamination and bioturbation. Bioturbation intense and extensive through entire core. Slumping structure with slump ball and contorted bedding, at around Section 2, 70 cm and Section 3, 70-80 cm.  SMEARS: 1-100, 2-100, 3-81, 4-50 (Nannofossil Chalk) Clay minerals TR 5% Radiolarians TR Foraminifers 0-2% Sponge spicules 0-2% Nannofossils 85-94%
					3		
	B AM				3		7.5YR 7/2 to 8/2  2-40 (Siliceous Nannofossil Chalk) Quartz, Feldspar 9% Nannofossils 57% Clay minerals <10% Diatoms 4% Volcanic glass 2% Radiolarians 3% Foraminifers 4% Sponge spicules 5%
					4		
	AM				4		10YR 6/3  CC (Calcareous Chalk) Quartz, Feldspar 15% Foraminifers 20% Clay <5% Nannofossils 10% Volcanic glass 8% Diatoms 1% Carbonate unspecified 30% Radiolarians 6%
5							
FP	AM	AP			5		7.5YR 8/2 to 7/2  VOID
					6		
FP	AM	AP			7		7.5YR 7/2  VOID
					CC		

SITE 445		HOLE		CORE 53		CORED INTERVAL: 493.0-502.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Middle Oligocene <i>Sphenolithus distentus</i> Zone (N)	N2/P21(?)	RP	AP		0.5		10YR 8/2  Dominant Lithology: Nannofossil Chalk, pinkish gray (10YR 8/2) to light gray (5Y 7/2).  Minor Lithologies: Calcareous Chalk at Section 4 gradually passing upwards to Nannofossil Chalk; including many lithic angular pebbles and granules. Diatomaceous Nannofossil Chalk (Section 5), parallel laminated, showing interbedded light gray (dominant) white (minor) and grayish green (minor) color attenuation (thin).
					1.0		
	AP				2		5G 7/1.5  SMEARS: 1-75, 2-75, 3-75, 4-113, 5-95 (Nannofossil Chalk) Clay minerals <5% Nannofossils 85-93% Volcanic glass 0-2% Radiolarians 0-6% Foraminifers 0-5% Sponge spicules 0-2%
					3		
	RP	AP			3		4-35 (Calcareous Chalk) Clay minerals <5% Foraminifers 15% Volcanic glass 2% Nannofossils 8% Carbonate unspecified 55%
					4		
	CP	AP			4		5-95 (Diatomaceous Nannofossil Chalk) Clay minerals <5% Diatoms 20% Volcanic glass 4% Radiolarians 3% Nannofossils 67% Sponge spicules 5%
5							
B				5		CARBON-CARBONATE: 1-70 (9.7, 0.0, 81)	
				N7			



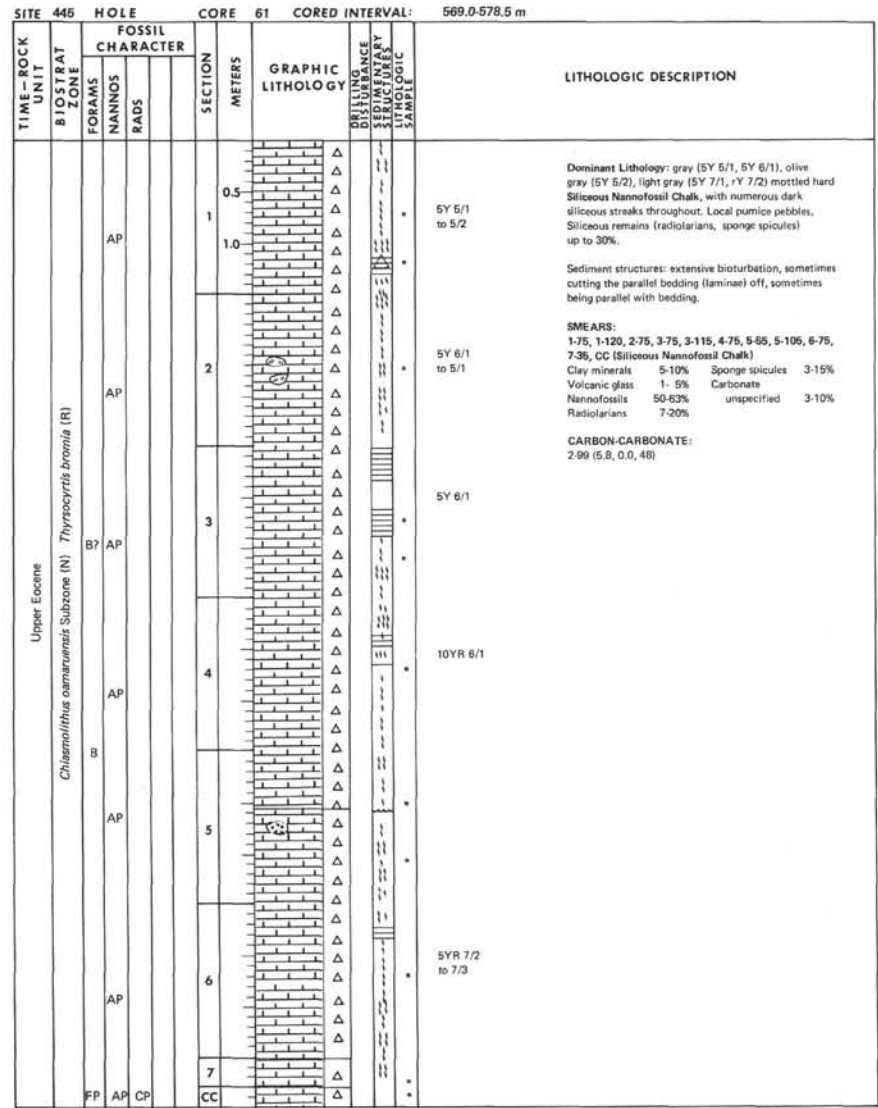
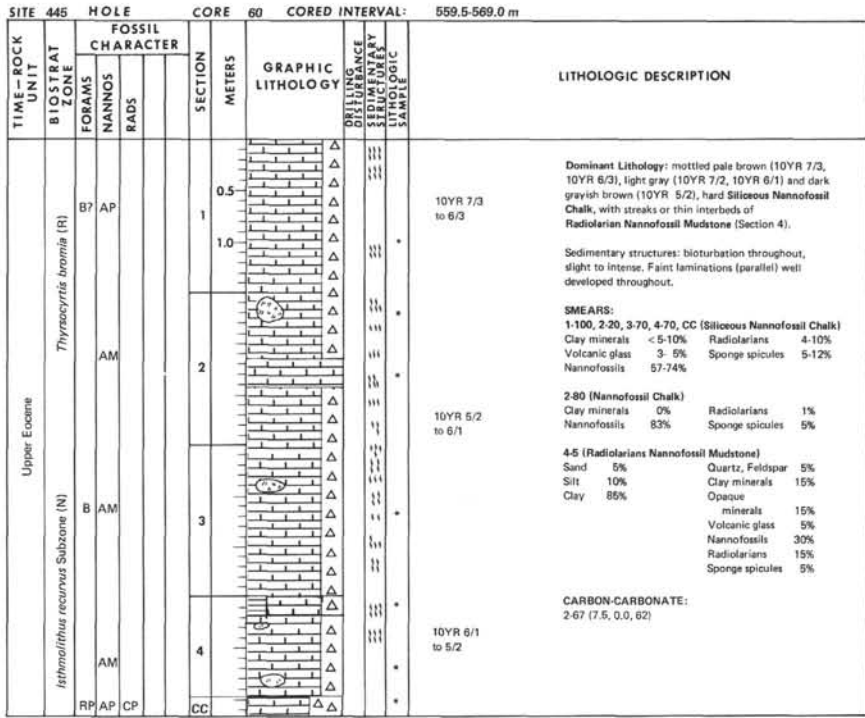
SITE 445		HOLE		CORE 54		CORED INTERVAL: 502.5-512.0 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS RADS						
Middle Oligocene <i>Sphaerolithus distentus</i> Zone (N)	B	AP		0.5				5Y 7/1	Dominant Lithology: <b>Nannofossil Chalk</b> , with dominant color of pinkish white (7.5YR 8/2) to pinkish gray (7.5YR 7/2) and white (5Y 8/2) to light gray (5Y 7/2).
				1				10R 8/1	
	AM			1.0				7.5YR 8/3	Minor Lithologies: <b>Calcareous Silty Sandstone</b> , at base of parallel-laminated and micro-cross-laminated <b>Nannofossil Chalk</b> sequence (Section 3, 103-118 cm).
				2					Sedimentary structures: No definite graded beds; parallel laminations, wavy laminations, and micro-cross-laminations; bioturbation locally intense.
	B	AM						7.5YR 7/2	SMEARS: 1-70, 2-75, 3-75, 4-75, 5-75, 6-75, CC (Nannofossil Chalk) Clay minerals TR- 5% Diatoms 0- 2% Foraminifers 0- 5% Sponge spicules 1- 5% Nannofossils 78-96%
									3
	AP							7.5YR 8/2	CARBON-CARBONATE: 1-61 (7.8, 0.0, 65)
4									5Y 8/2 to 7/2
B	AM						5Y 7/2 to 8/2	5Y 7/2 to 8/2	
									5
AM							7.5Y 8/2, 7/2	7.5Y 8/2, 7/2	
									6
FP	AM	CP		7			7.5Y 8/2	7.5Y 8/2	
				CC					

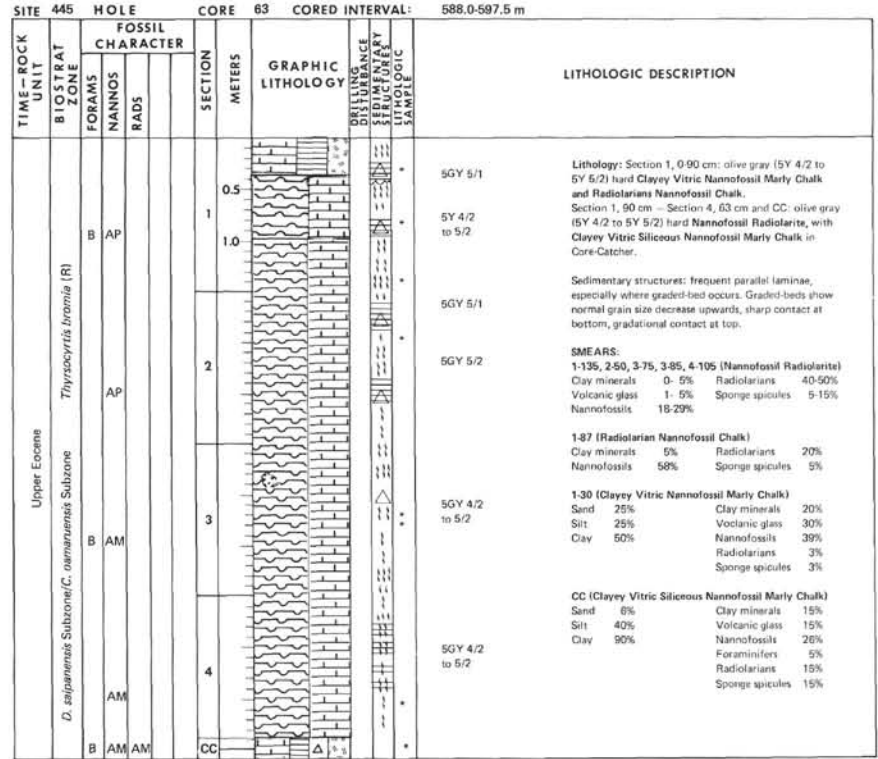
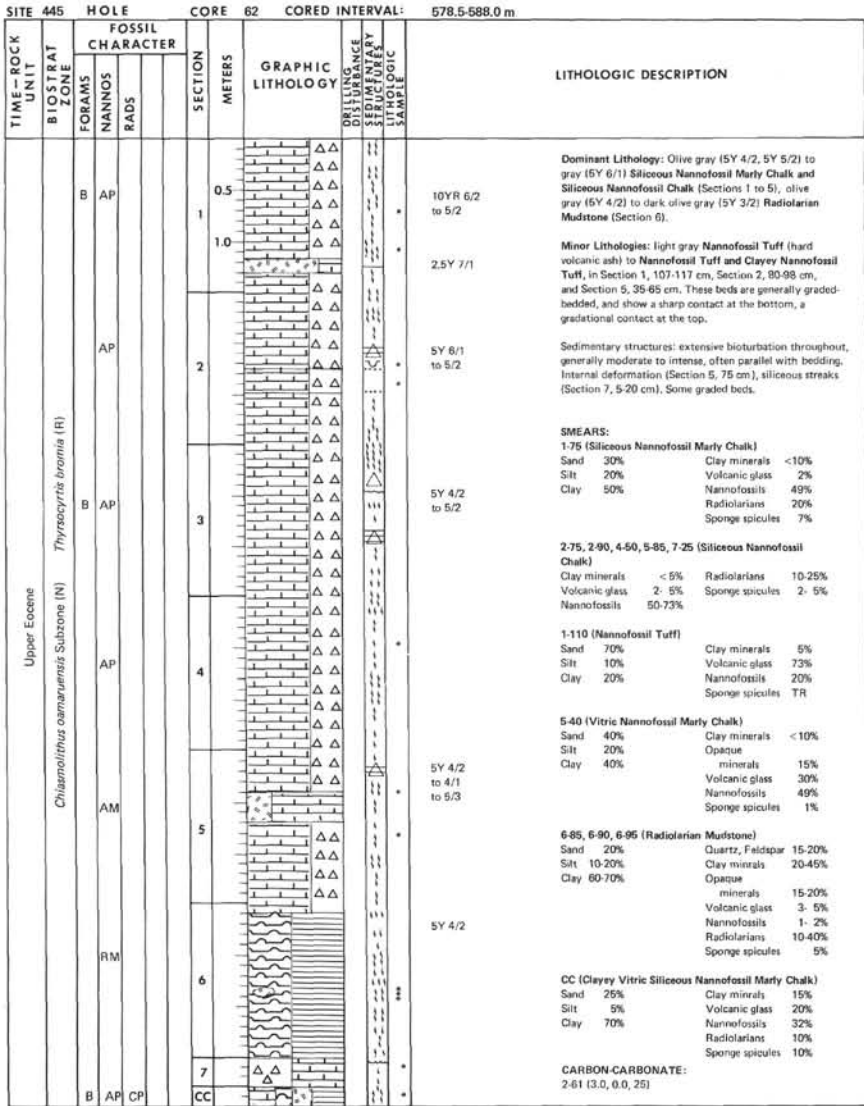
SITE 445		HOLE		CORE 55		CORED INTERVAL: 512.0-521.5 m			
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
		FORAMS	NANNOS RADS						
Middle Oligocene <i>Sphaerolithus distentus</i> Zone (N)	AM	RP		0.5				7.5YR 8/2	
				1				7.5YR 7/2	
	AM							7.5YR 7/2	
									2
	FP	AM						7.5Y 7/2	6-134 (Vitric Nannofossil Chalk) Sand 5% Volcanic glass <10% Silt 60% Nannofossils 60% Clay 35% Diatoms 2% Opaque minerals 8%
									3
	AM							7.5YR 7/2	6-63 (Nannofossil Tuff) Sand 5% Clay minerals 3% Silt 65% Opaque minerals 5% Clay 30% Volcanic glass 54% Nannofossils 25% Diatoms 2%
4									7.5YR 8/2
B	AM						7.5Y 7/2	CARBON-CARBONATE: 1-74 (9.6, 0.0, 80)	
								5	7.5YR 8/2
AM							7.5Y 7/2	7.5YR 8/2	
									6
FM	AM	CP		7			7.5Y 7/2	7.5YR 8/2	
				CC					



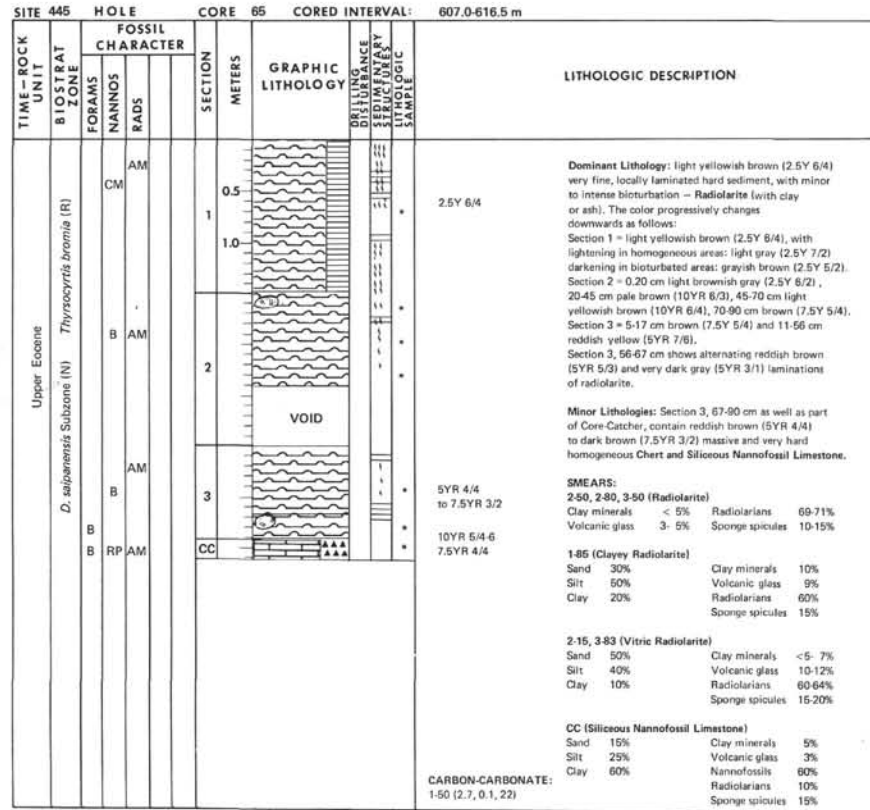
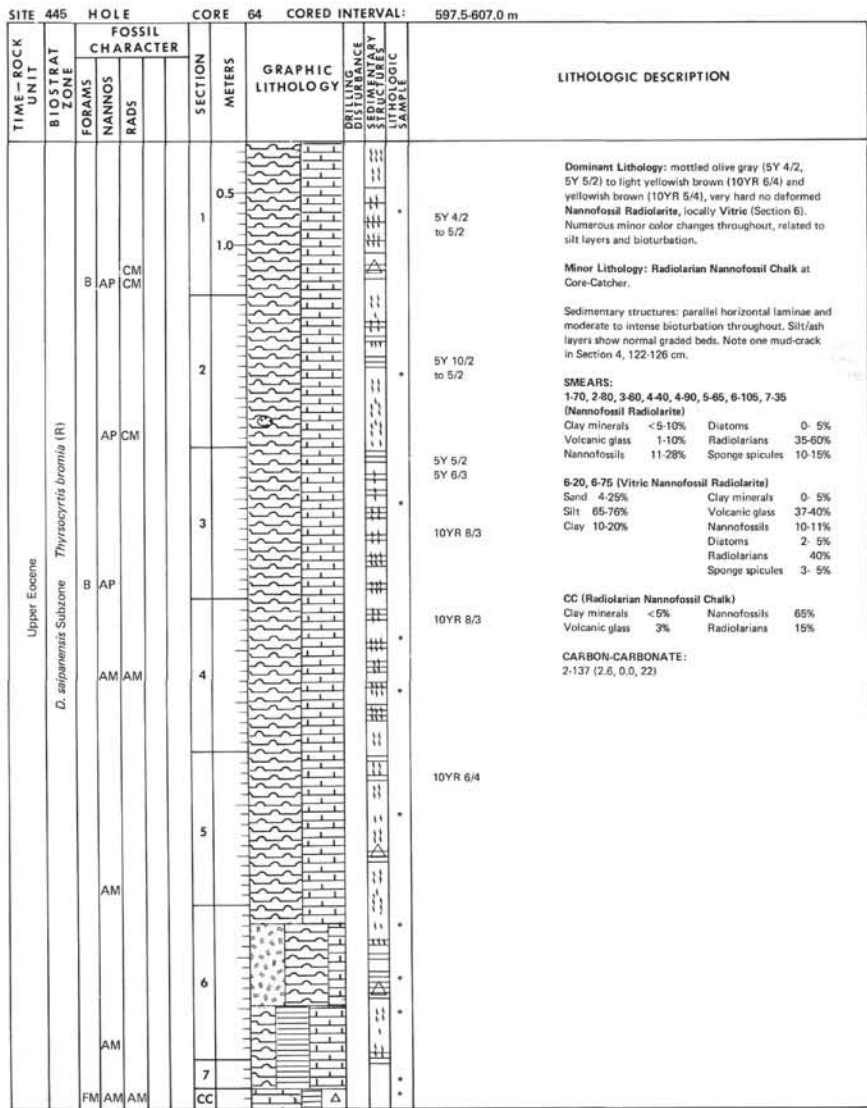
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG BIOTURBATION SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE	CORED INTERVAL: 540.5-550.0 m	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS					
		SECTION	METERS	GRAPHIC LITHOLOGY					
Lower or Lower Middle Oligocene <i>Helicosphaera reticulata</i> Zone or <i>Sphenolithus praedistantis</i> Zone (N)	RP	AM		1	0.5		5Y 5/1 to 4/1	<p>Dominant Lithology: Nannofossil Chalk with light gray (5Y 7/1) to gray (5Y 5/1) color, locally passing to Foraminifera-Nannofossil Chalk (Section 1); generally rich in volcanic ash grains throughout the core.</p> <p>Minor Lithologies: Vitric Nannofossil Chalk interbeds and Foraminifera-Nannofossil Marly Chalk. The latter occurs at the basal part of graded bed (Section 2, 90-100 cm).</p>	
					1.0				
	AM		2	2	5Y 5/1	<p>Sedimentary structures: graded bed is seen in Section 2. Most part of Clayey Nannofossil Chalk is intensely to mildly bioturbated and without original stratification.</p> <p>Other structures: fault at Section 3, 75-80 cm, dipping about 45°</p>			
				3	5Y 7/1 to 5Y 5/1	<p><b>SMEARS:</b> 2-50, 3-85, 4-75, 5-75, 6-15, CC (Nannofossil Chalk) Clay minerals 0-5% Nannofossils 85-95% Volcanic glass 0-2% Diatoms 0-5% Foraminifera 1-5% Sponge spicules 0-4%</p> <p><b>1-42, 2-80 (Foraminifera-Nannofossil Chalk)</b> Clay minerals 0-TR% Foraminifera 15-25% Volcanic glass 1-2% Nannofossils 60-80%</p> <p><b>1-125 (Vitric Nannofossil Marly Chalk)</b> Sand 30% Clay minerals 20% Silt 20% Volcanic glass 30% Clay 5% Foraminifera 2% Nannofossils 31%</p> <p><b>CARBON-CARBONATE:</b> 2-55 (10.5, 0.0, 87)</p>			
	AM		4	4	5Y 6/1 to 5/1		5Y 6/1 to 5/1		
				5	5Y 7/1				
	FP	AM	FP	CC	7	VOID			

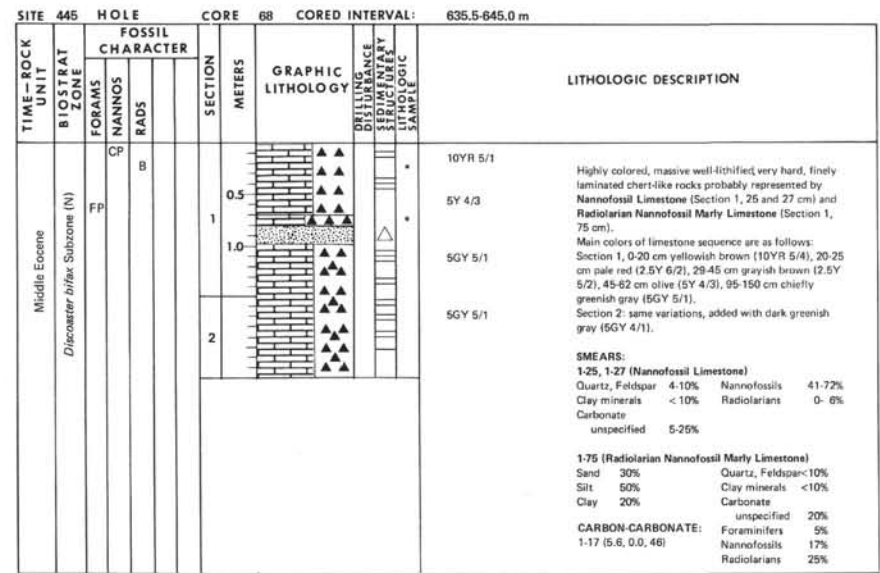
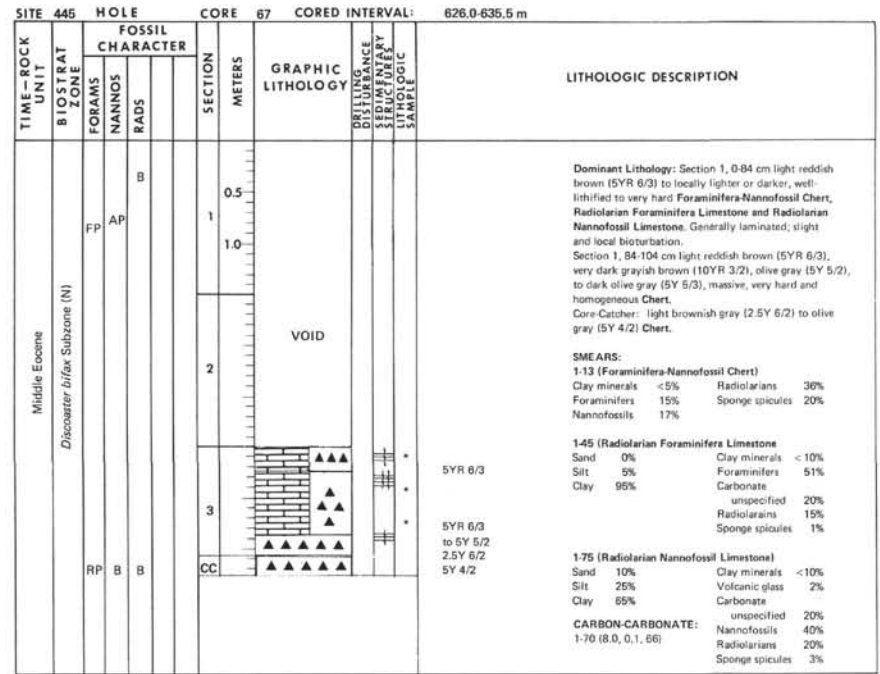
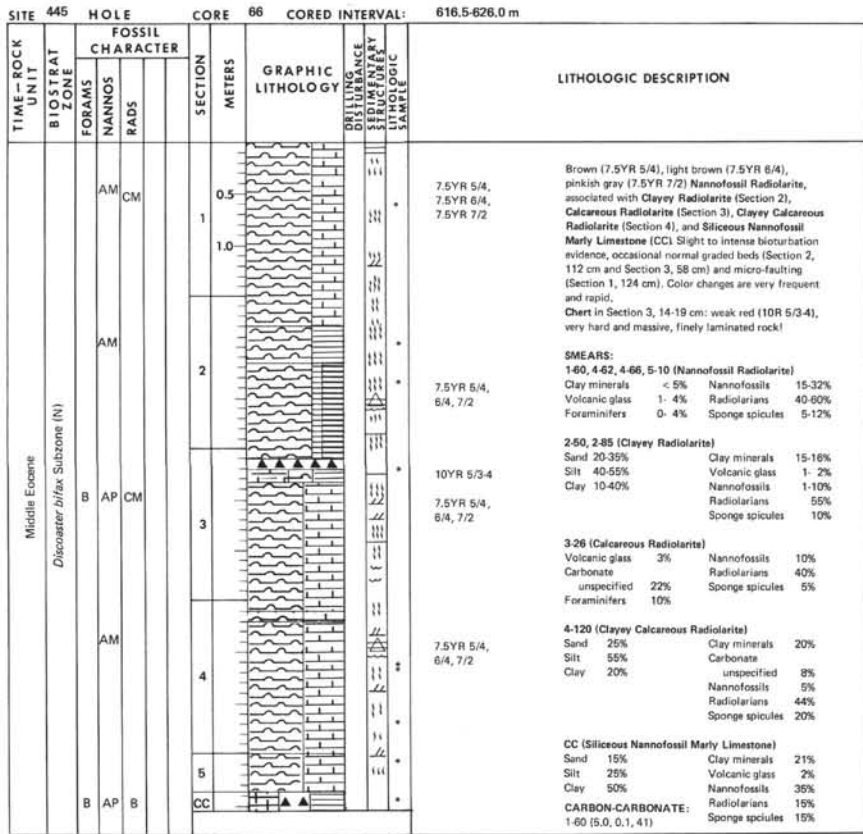
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING LOG BIOTURBATION SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE	CORED INTERVAL: 550.0-559.5 m	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS					
		SECTION	METERS	GRAPHIC LITHOLOGY					
Upper Eocene <i>Idmonolithus recurvus</i> Subzone (N) <i>Thyracystis bromia</i> (R) P.17 Foraminifera Zone or <i>S. praedistantis</i> Zone (N) P.18 Foraminifera Zone	RP	AP		1	0.5		5Y 7/1 to 5/1	<p>Dominant Lithology: Section 1, 0 cm to Section 2, 20 cm Nannofossil Chalk (light gray - 5Y 7/1 to gray - 5Y 5/1). Section 2, 20 cm to CC Siliceous Nannofossil Chalk (pinkish gray - 7.5YR 7/2 to 6/2). Color change from 5Y series to 7.5YR series is gradual.</p> <p>Minor Lithology: Clayey Nannofossil Chalk at Section 4, 95 cm.</p>	
					1.0				
	AM		2	2	7.5YR 7/2 to 6/2	<p>Sedimentary structures: extensive to intense to mild bioturbation. Faint parallel laminations throughout the entire core.</p> <p><b>SMEARS:</b> 2-15, 2-85, 3-60, 4-140, 5-75, 6-110, 7-20, CC (Siliceous Nannofossil Chalk) Sand 3-15% Clay minerals &lt;5-10% Silt 7-30% Volcanic glass 1-9% Clay 60-86% Nannofossils 64-75% Radiolarians 5-8% Sponge spicules 3-16%</p>			
				3	7.5YR 7/2 to 4/2	<p><b>1-75, 4-95 (Nannofossil Chalk)</b> Sand 3-4% Clay minerals &lt;5% Silt 12-16% Volcanic glass 1-3% Clay 80-85% Nannofossils 84-88% Radiolarians TR Sponge spicules 2%</p> <p><b>CARBON-CARBONATE:</b> 2-17 (7.4, 0.0, 61)</p>			
	AP		4	4	7.5YR 6/1 to 7/1		7.5YR 6/1 to 7/1		
				5	7.5YR 7/2 to 4/2				
	AP		6	6	10YR 6/2 to 7/2		10YR 6/2 to 7/2		
7									
B	AP	CP	CC	7					









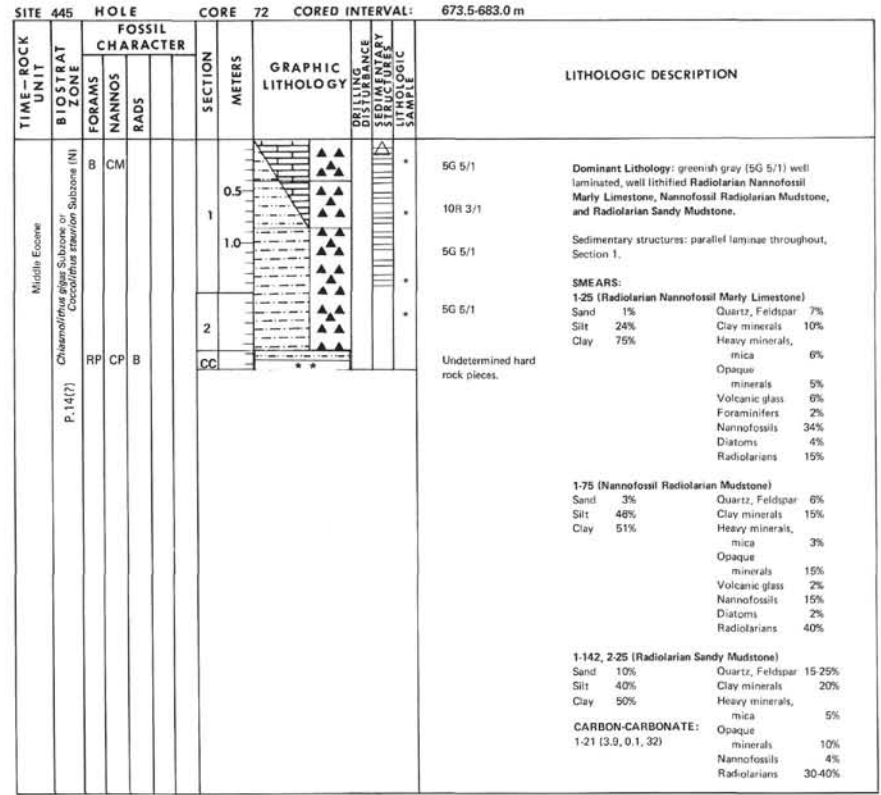
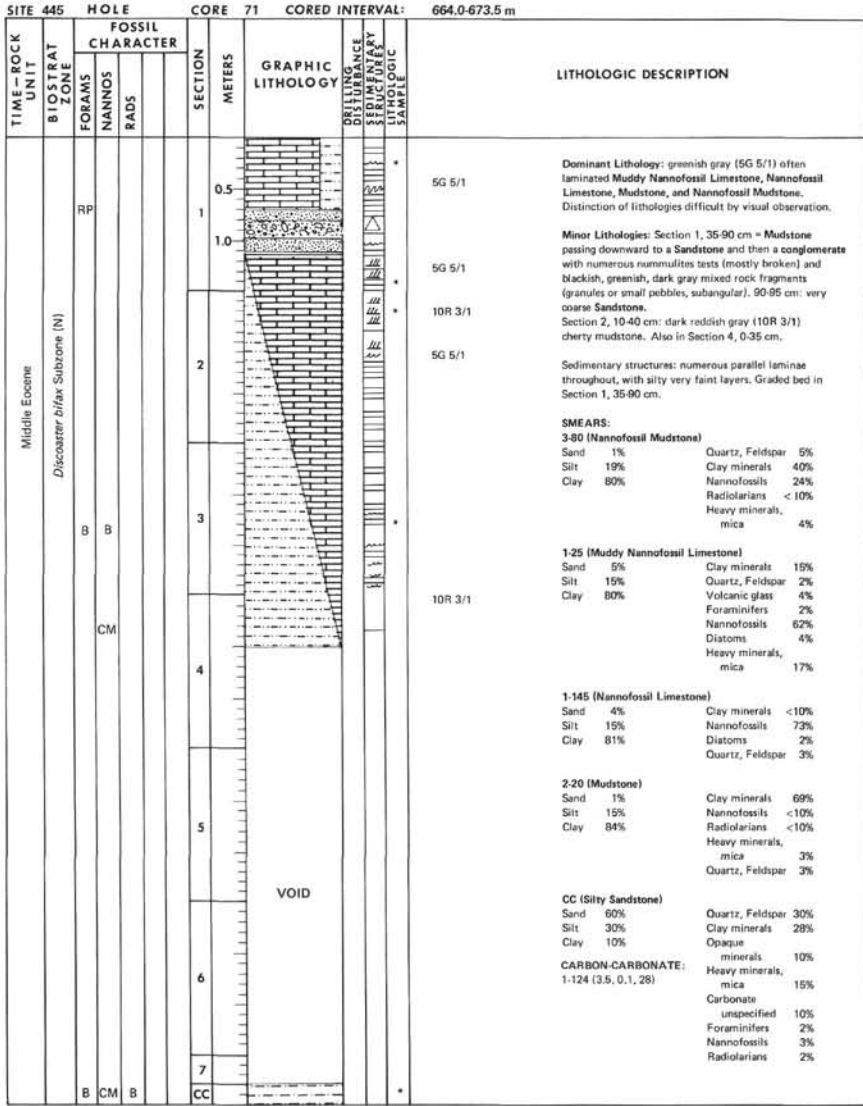


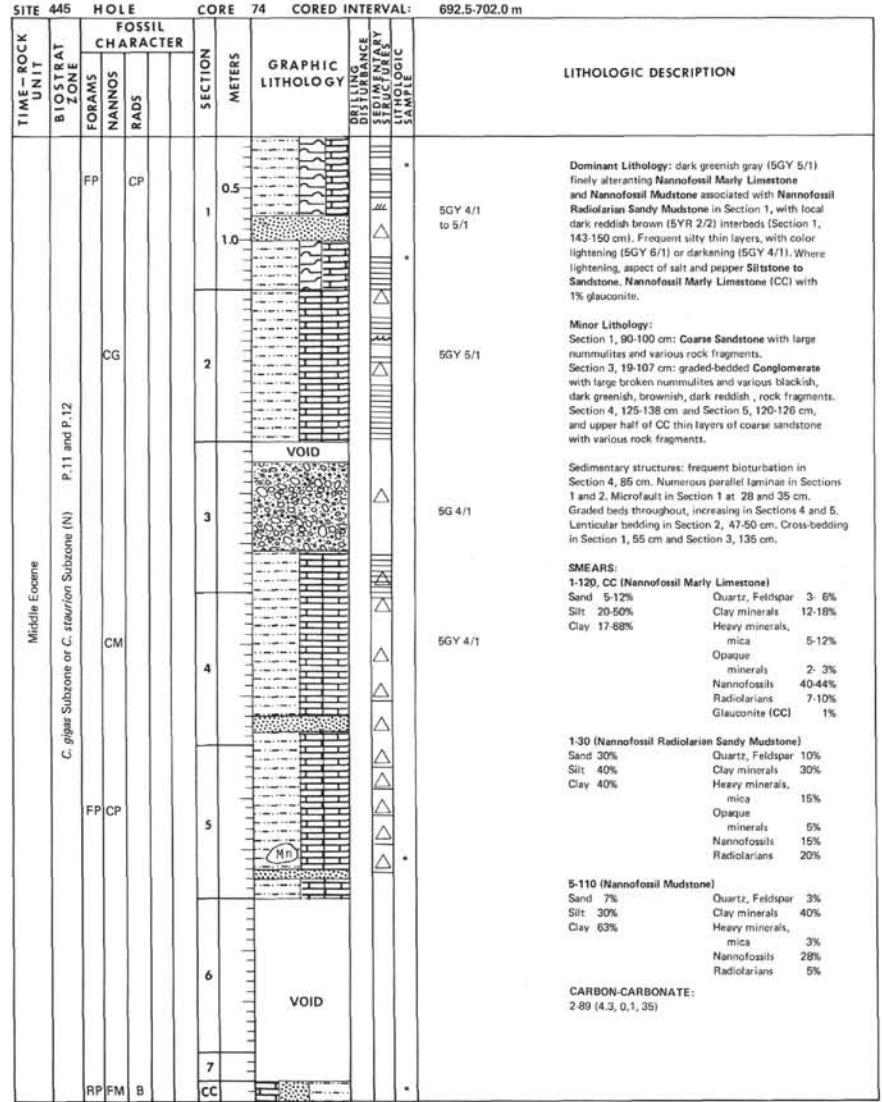
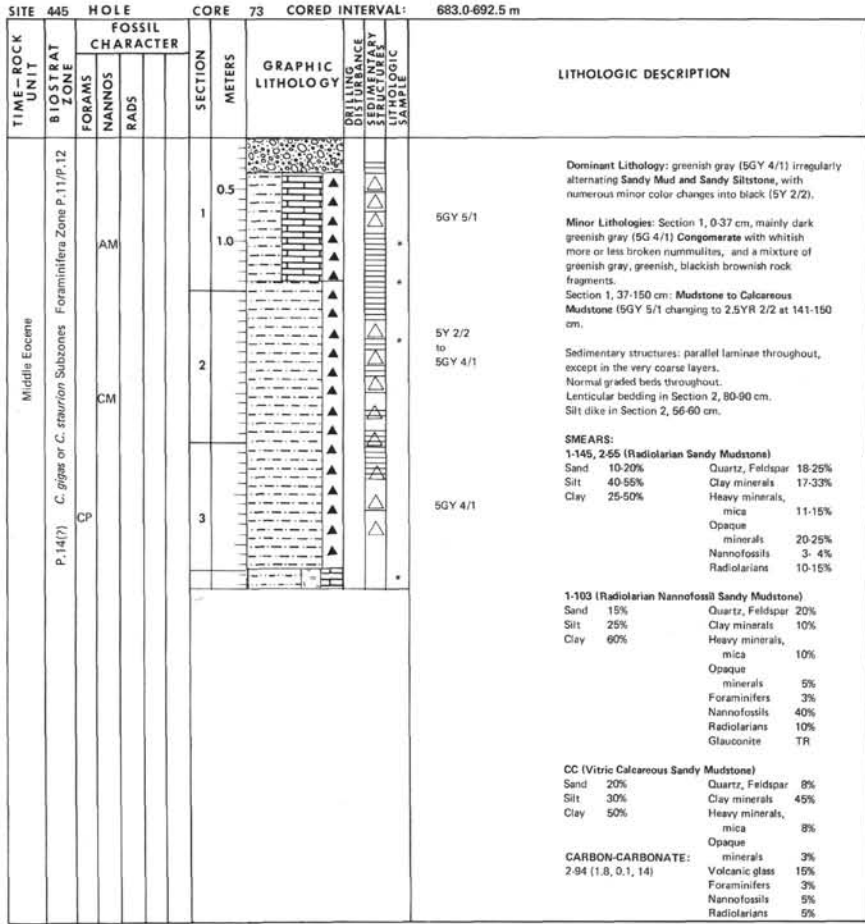
SITE 445 HOLE CORE 69 CORED INTERVAL: 645.0-654.0 m

TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING PERFORMANCE LOG	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS					
Middle Eocene					0.5 1 1.0			#	5GY 5/1 Varying Lithology: Radiolarian Calcareous Mudstone and Chert (Section 1, 0-106 cm) greenish gray (5GY 5/1) to grayish green (5G 5/2) and dark greenish gray (5GY 4/1) alternating massive hard laminated chert and normal-graded silty Mudstone. Mudstone is laminated and presents local micro-fault (12). Litic Conglomerate (Section 1, 106-150 cm, and Section 2, 0-92 cm) breccias subangular mixture of large nummulites (to 3 cm) and greenish, brownish, reddish and blackish rock fragments. White to dark greenish gray (5GY 4/1), changing. Thin section: foraminifers, basalt, quartz, mica, mudstone, sandstone - matrix of nannofossils, foraminifers, clay, glauconite extensive, small cherty areas. Also in Section 3, 0-5 cm. Radiolarian Nannofossil Marly Limestone to Calcareous Sandy Mudstone to Sandstone (Section 2, 92 cm to Section 3, 56 cm), dark greenish gray (5GY 4/1) fine to coarse, graded-bedded, irregularly laminated and interbedded sediment. Mudstone clasts and grain size increase toward bottom. Radiolarian Nannofossil-Marly Limestone (Section 3, 56 cm to Section 4, 92 cm), greenish gray mudstone with laminae, conglomerate subangular pebbles, hard mudstone.
					2			*	5GY 4/1 Radiolarian Nannofossil Marly Limestone to Calcareous Sandy Mudstone to Sandstone (Section 2, 92 cm to Section 3, 56 cm), dark greenish gray (5GY 4/1) fine to coarse, graded-bedded, irregularly laminated and interbedded sediment. Mudstone clasts and grain size increase toward bottom. Radiolarian Nannofossil-Marly Limestone (Section 3, 56 cm to Section 4, 92 cm), greenish gray mudstone with laminae, conglomerate subangular pebbles, hard mudstone.
					3			*	5GY 4/1 Radiolarian Nannofossil-Marly Limestone (Section 3, 56 cm to Section 4, 92 cm), greenish gray mudstone with laminae, conglomerate subangular pebbles, hard mudstone.
					4			*	5GY 4/1 SMEARS: 1-11 (Radiolarian Calcareous Mudstone) Sand 10% Clay minerals 30% Silt 20% Volcanic glass 2% Clay 70% Carbonate unspecified 15% Foraminifers 2% Nannofossils 10% Radiolarians 25% Sponge spicules 2%  1-12 (Radiolarian Mudstone) Sand 7% Clay minerals 65% Silt 15% Volcanic glass 1% Clay 78% Radiolarians 25%  2-136, 4-35 (Radiolarian Nannofossil Marly Limestone) Sand 5-8% Clay minerals 15-16% Silt 15-28% Carbonate Clay 70-80% unspecified 20-28% Foraminifers 3% CARBON-CARBONATE: 1-77 (2.2, 0.0, 18) Nannofossils 10-20% Radiolarians 30%

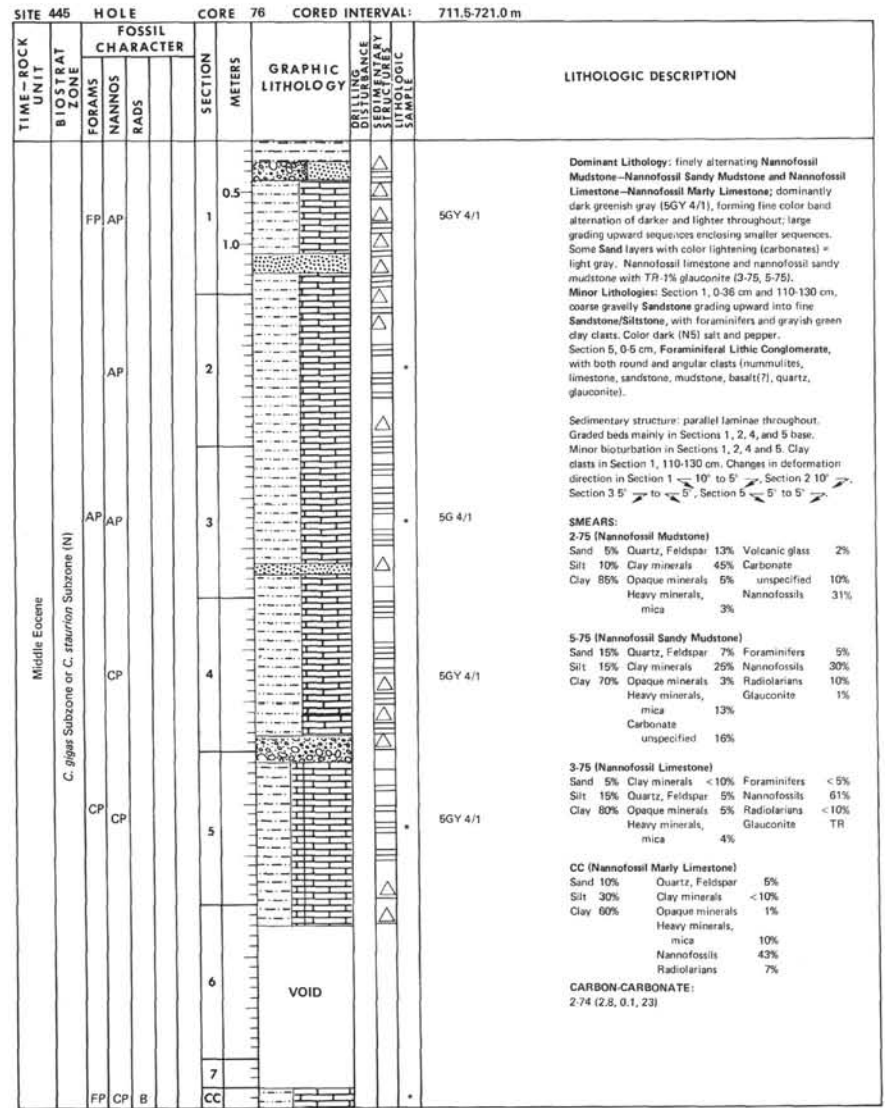
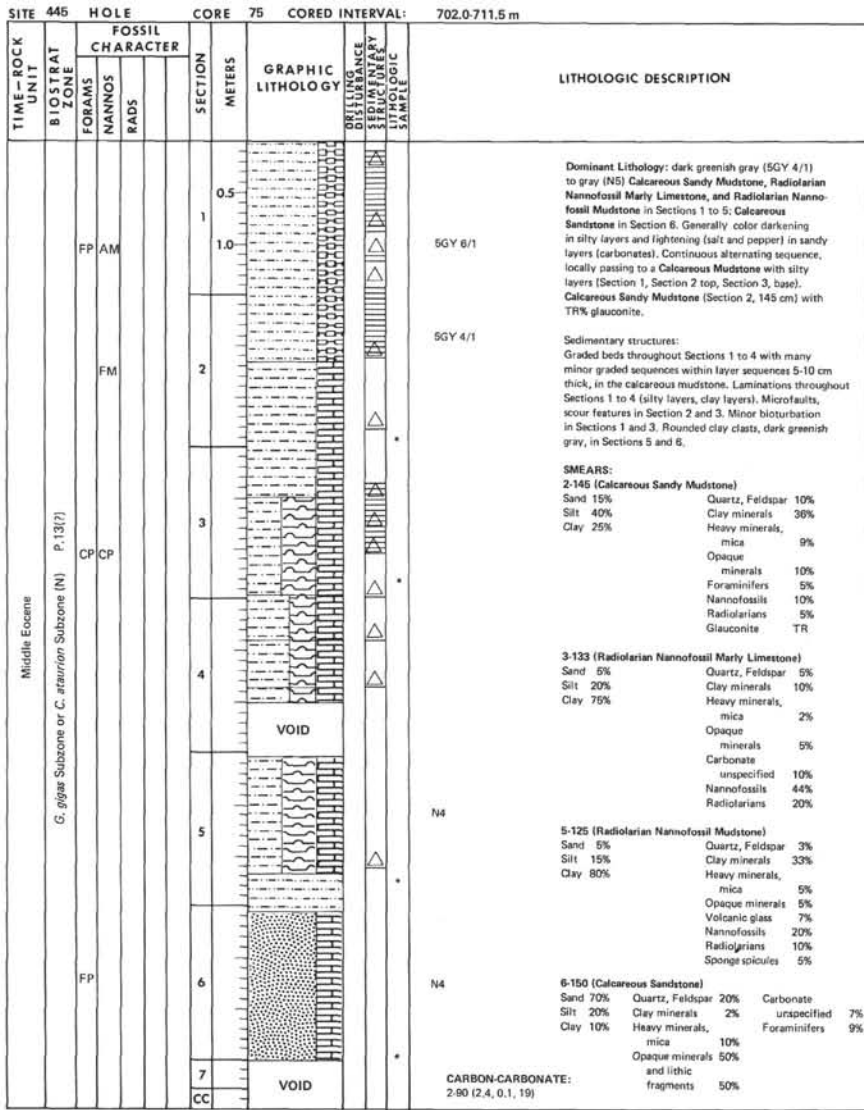
SITE 445 HOLE CORE 70 CORED INTERVAL: 654.5-664.0 m

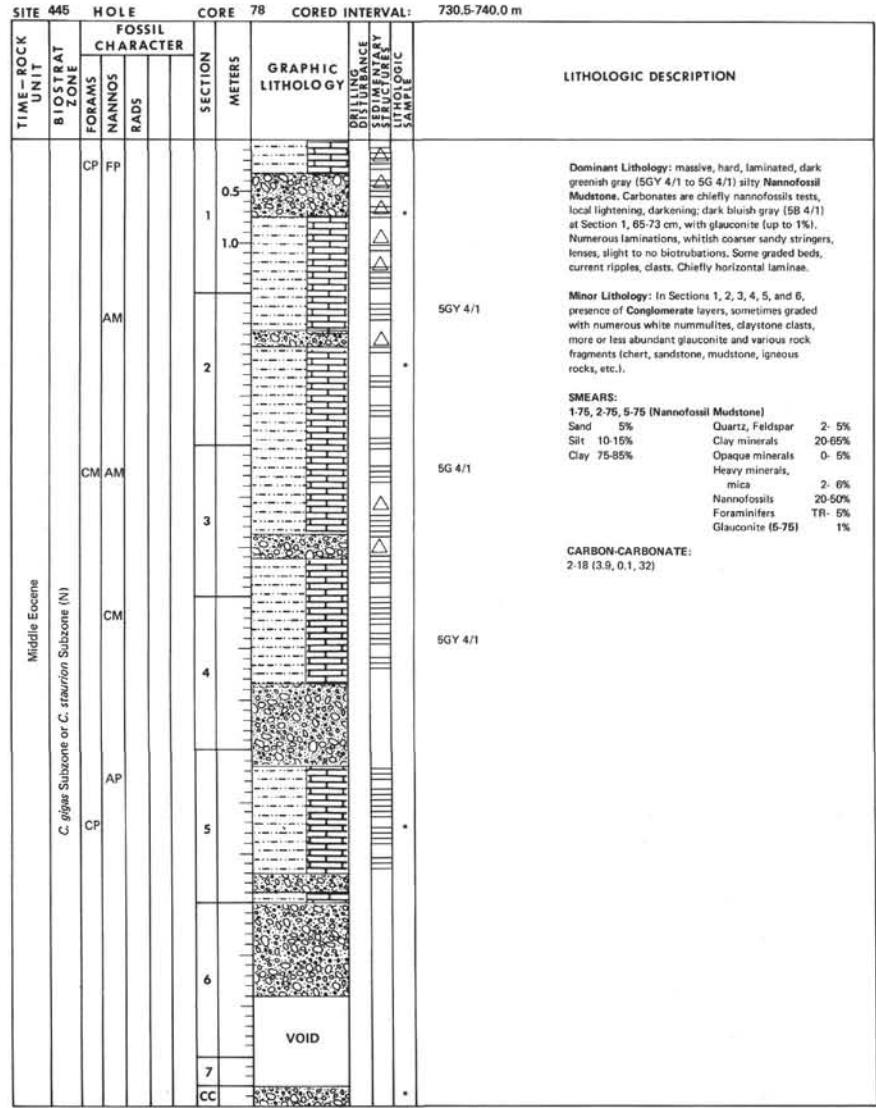
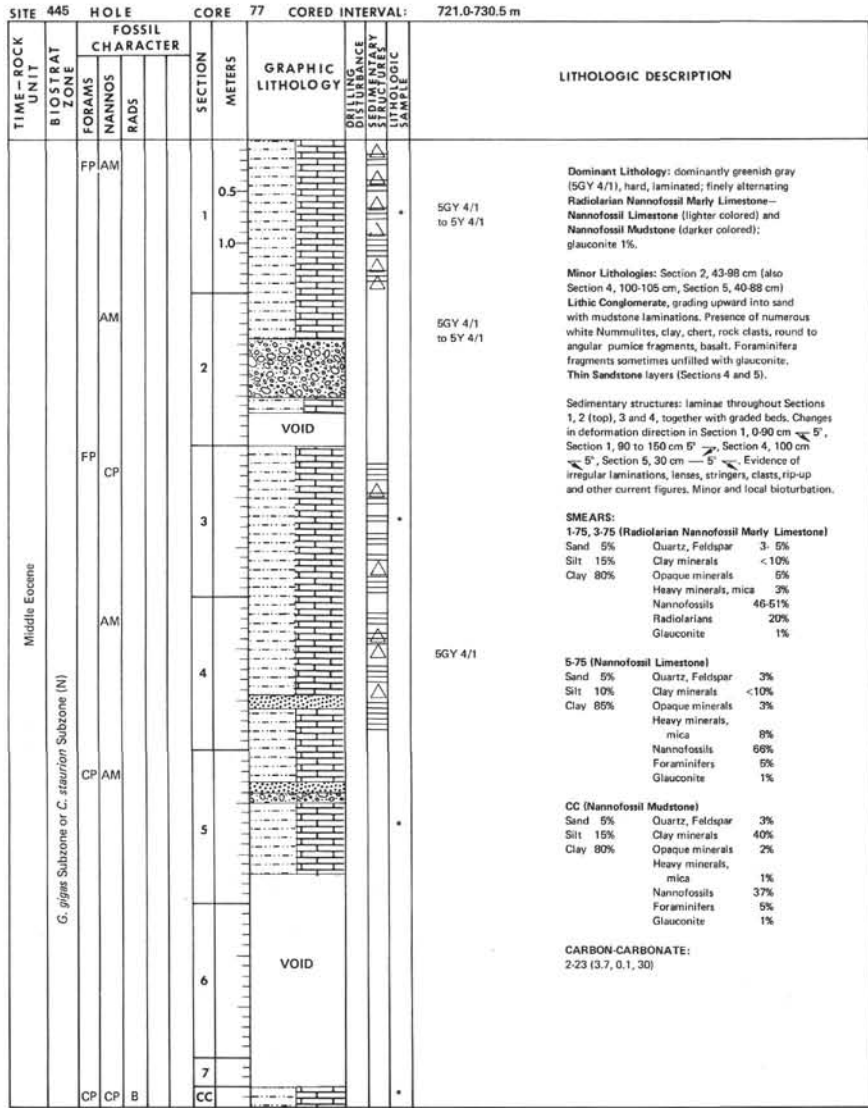
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION METERS	GRAPHIC LITHOLOGY	DRILLING PERFORMANCE LOG	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS					
Middle Eocene					0.5 1 1.0			*	5GY 4/1 Dominant Lithology: Muddy Chert (Dominant) to Cherty Mudstone alternating with a Calcareous Mudstone to Radiolarian Mudstone and Vitric Nannofossil Clayey Radiolarite. Dominant color is dark greenish gray (5GY 4/1) for both chert and mudstone, but chert often shows color changes: dark reddish brown (5YR 3/2) in Section 1, 66-71 cm and 120-125 cm, dark gray (10YR 4/1) in Section 1, 125-144 cm, greenish gray (5GY 6/1) in Section 1, 144-150 cm. Lamination throughout, more abundant in mudstone.  SMEARS: 1-137, 2-10 (Calcareous Mudstone) Sand 2-15% Quartz, Feldspar 5-10% Silt 9-25% Clay minerals 31-52% Clay 62-89% Volcanic glass 5-10% Carbonate unspecified 0-12% Nannofossils 5-40% Radiolarians 3%  1-137 = sandy mudstone  1-15 (Radiolarian Calcareous Mudstone) Sand 2% Clay minerals 36% Silt 13% Carbonate Clay 85% unspecified 20% Nannofossils 15% Diatoms 3% Radiolarians 20% Sponge spicules 1%  1-77 (Vitric Nannofossil Clayey Radiolarite) Sand 8% Clay minerals 16% Silt 22% Volcanic glass 13% Clay 70% Nannofossils 15% Radiolarians 40%  CARBON-CARBONATE: 1-62 (2.8, 0.1, 23)
					2			*	5GY 4/1 Dominant Lithology: Muddy Chert (Dominant) to Cherty Mudstone alternating with a Calcareous Mudstone to Radiolarian Mudstone and Vitric Nannofossil Clayey Radiolarite. Dominant color is dark greenish gray (5GY 4/1) for both chert and mudstone, but chert often shows color changes: dark reddish brown (5YR 3/2) in Section 1, 66-71 cm and 120-125 cm, dark gray (10YR 4/1) in Section 1, 125-144 cm, greenish gray (5GY 6/1) in Section 1, 144-150 cm. Lamination throughout, more abundant in mudstone.  SMEARS: 1-137, 2-10 (Calcareous Mudstone) Sand 2-15% Quartz, Feldspar 5-10% Silt 9-25% Clay minerals 31-52% Clay 62-89% Volcanic glass 5-10% Carbonate unspecified 0-12% Nannofossils 5-40% Radiolarians 3%  1-137 = sandy mudstone  1-15 (Radiolarian Calcareous Mudstone) Sand 2% Clay minerals 36% Silt 13% Carbonate Clay 85% unspecified 20% Nannofossils 15% Diatoms 3% Radiolarians 20% Sponge spicules 1%  1-77 (Vitric Nannofossil Clayey Radiolarite) Sand 8% Clay minerals 16% Silt 22% Volcanic glass 13% Clay 70% Nannofossils 15% Radiolarians 40%  CARBON-CARBONATE: 1-62 (2.8, 0.1, 23)





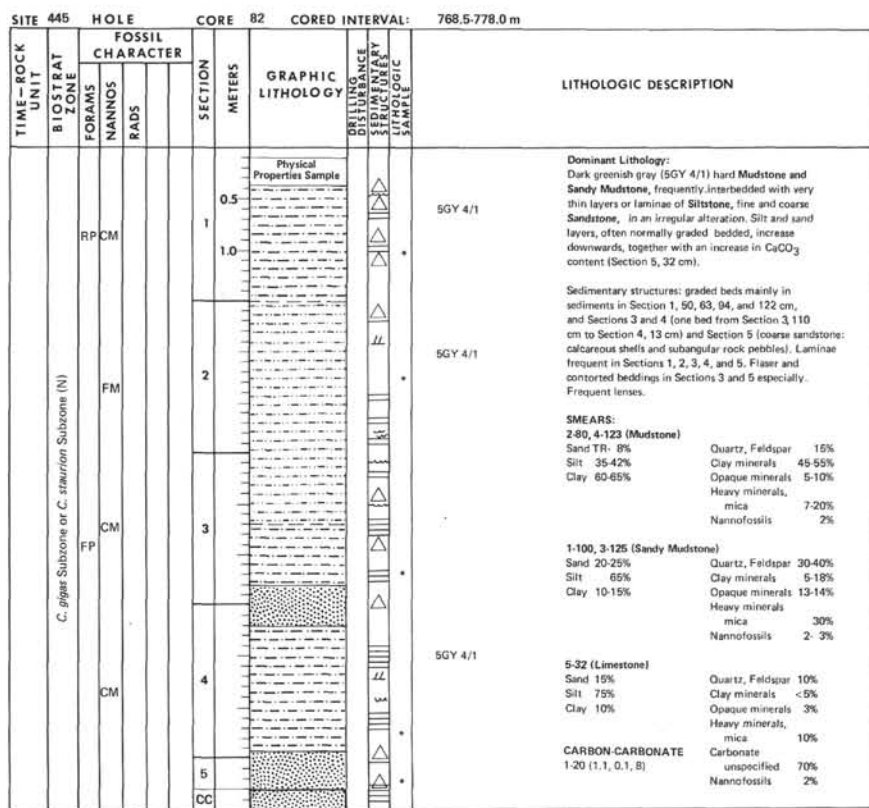
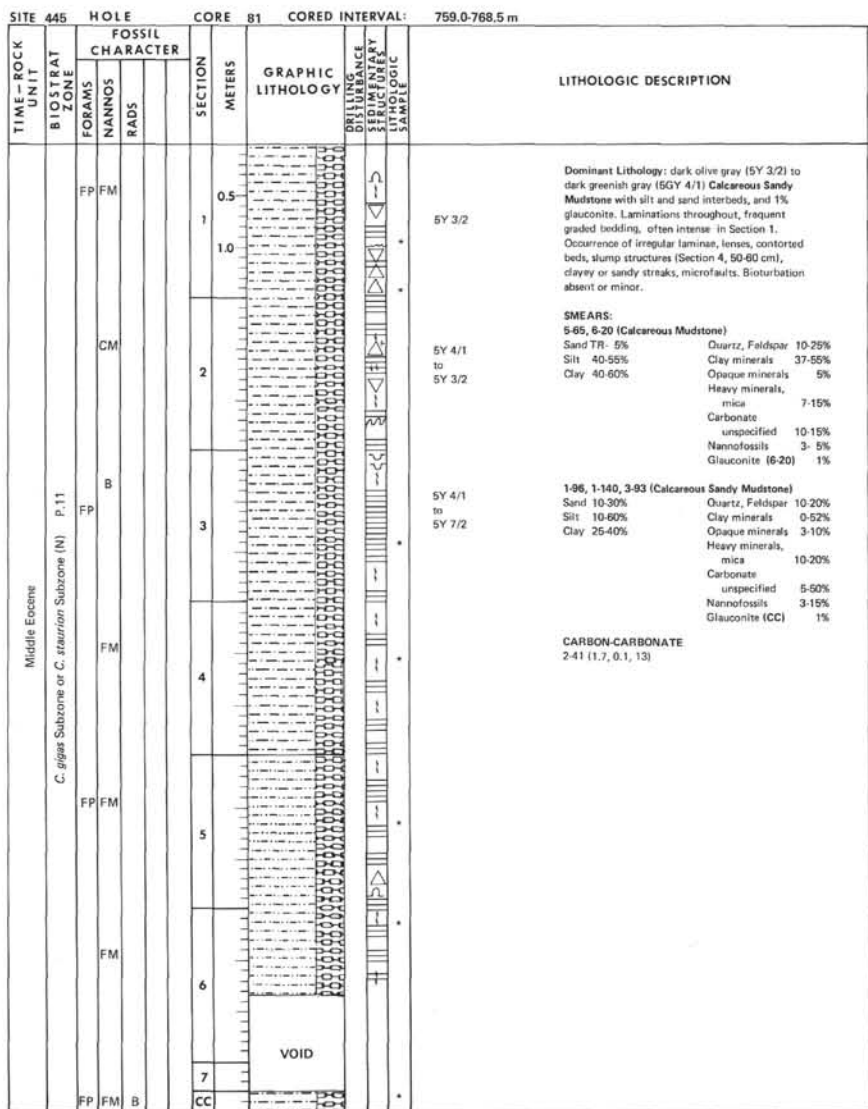


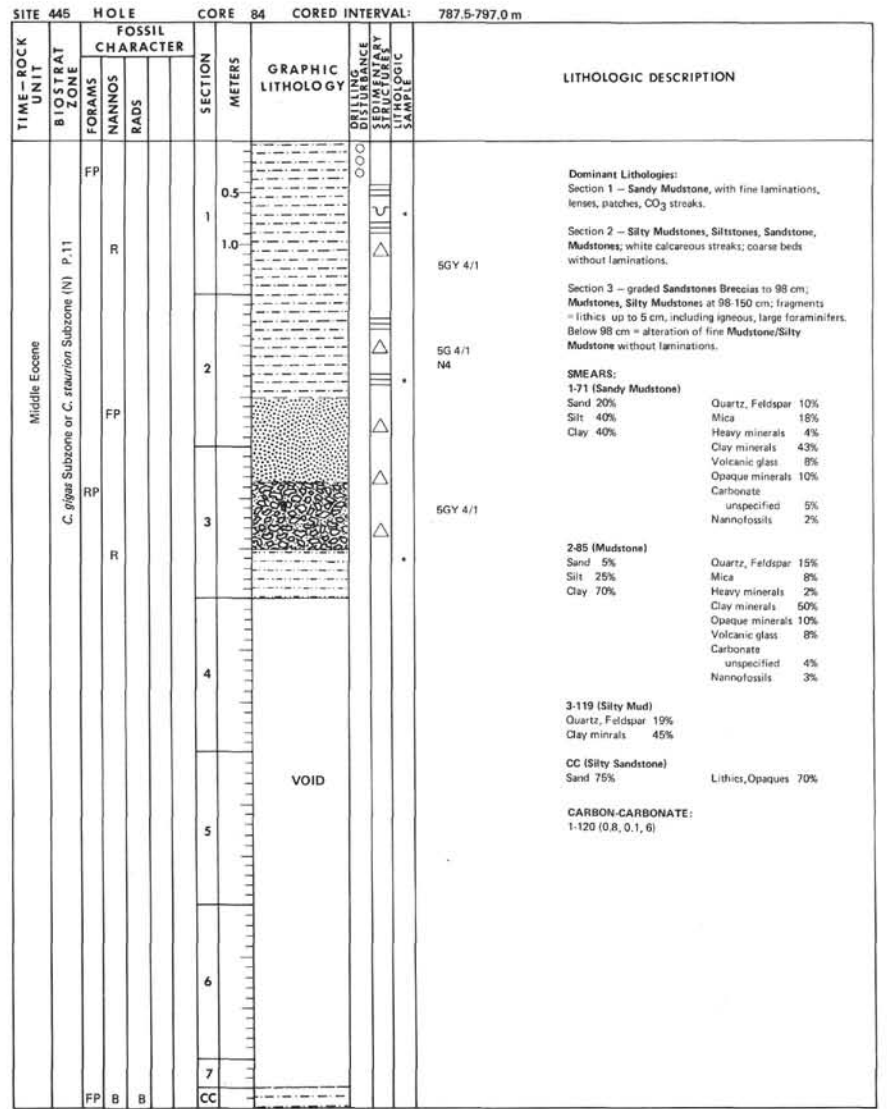
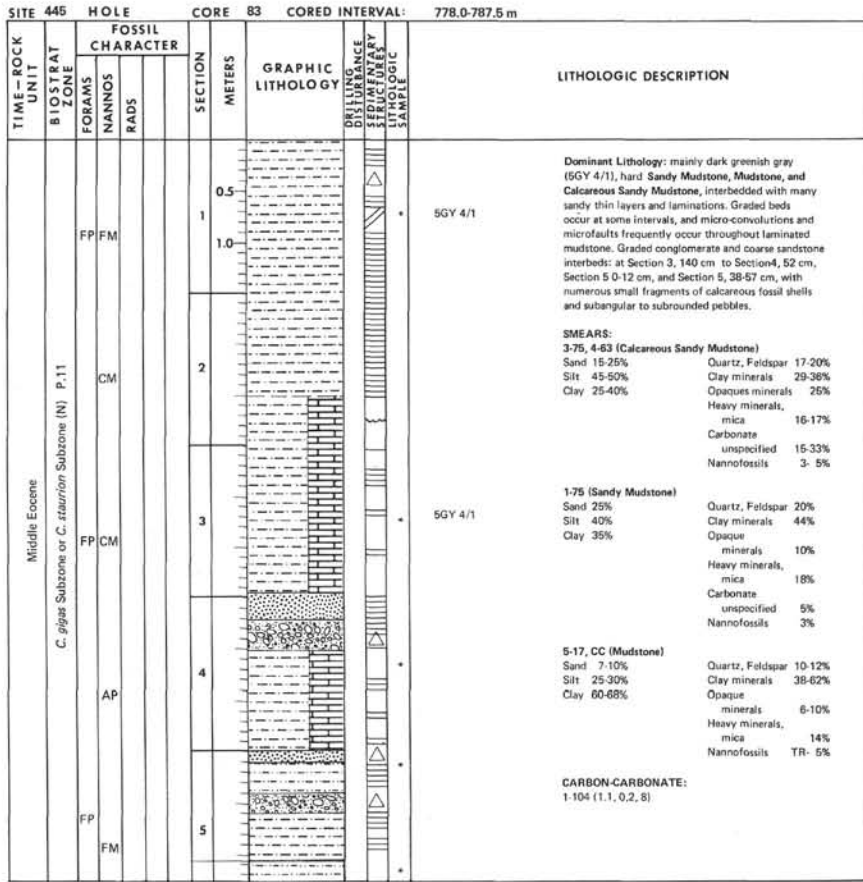


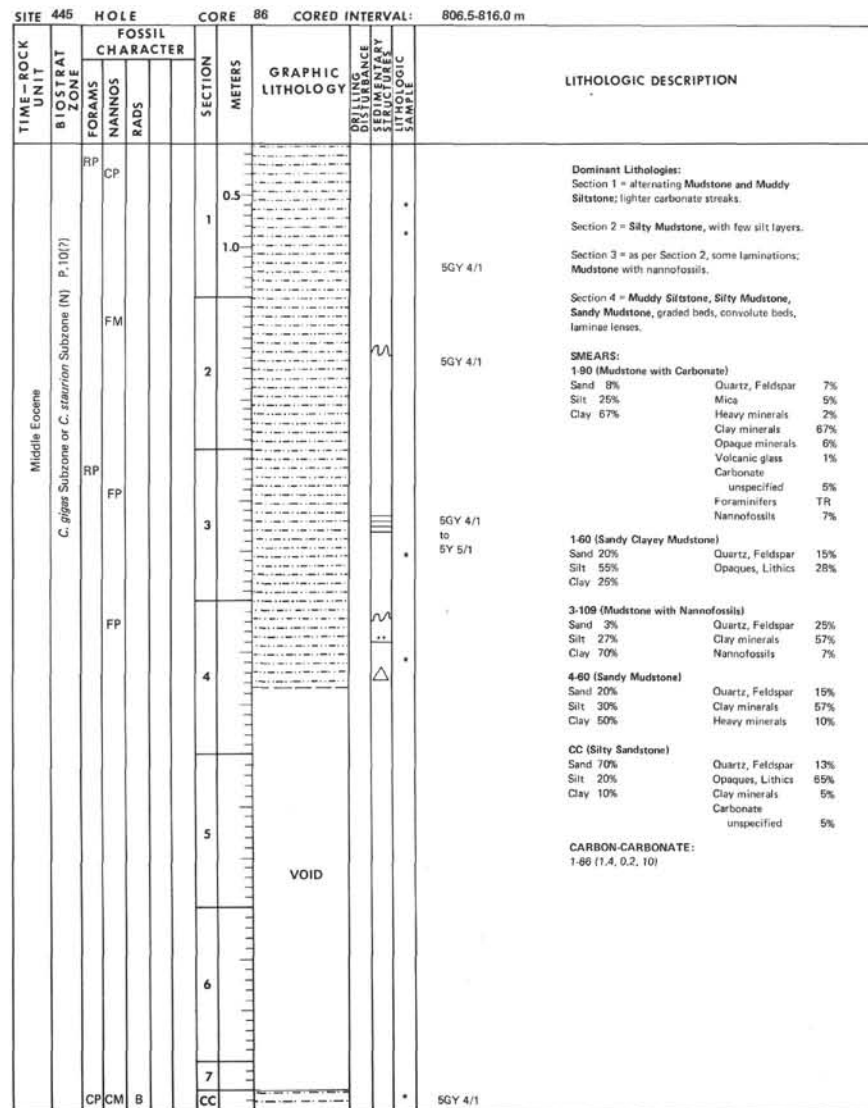
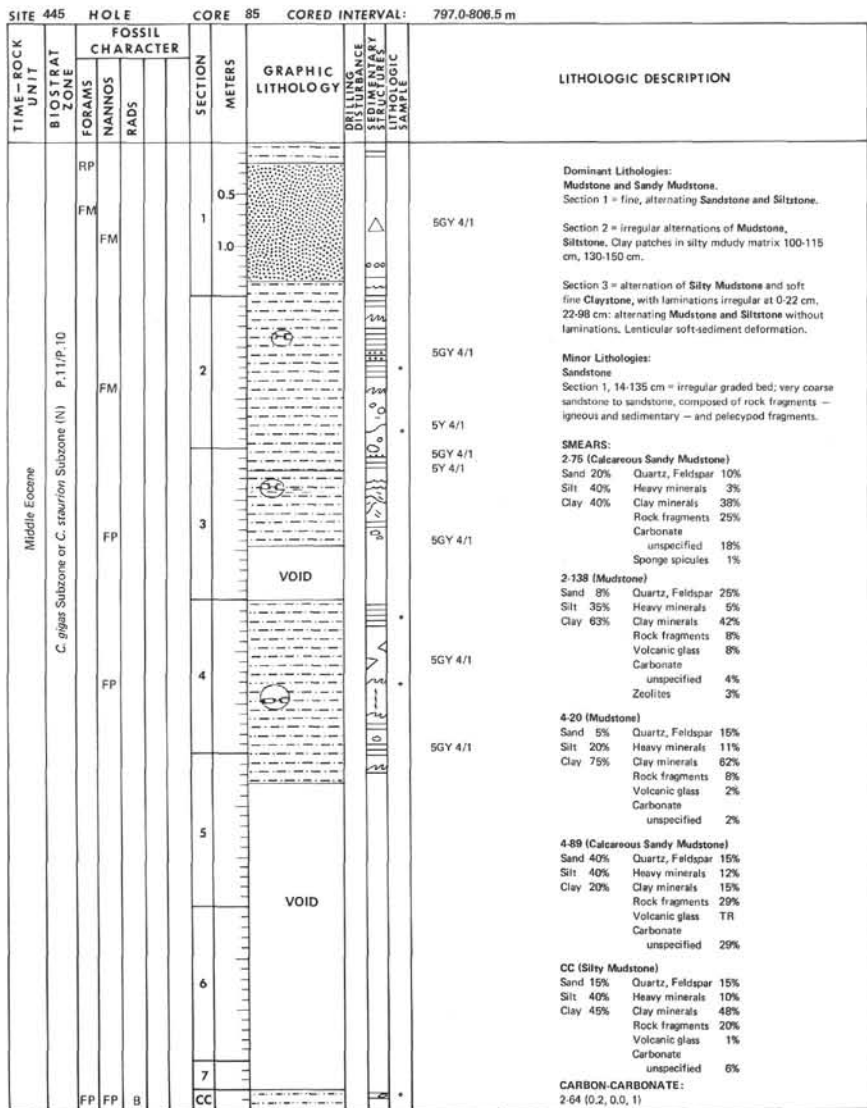


SITE 445		HOLE		CORE 79		CORED INTERVAL: 740.0-749.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Middle Eocene <i>C. gigas</i> Subzone or <i>C. staurion</i> Subzone (N) P.12/P.13(?)	CP	AP	FP	CP	1	VOID	5GY 4/1 Dominant Lithology: dominantly dark greenish gray (5GY 4/1); finely alternating, laminated <b>Nannofossil Limestone</b> (lighter) and <b>Nannofossil Mudstone</b> (darker), with coarser grained white <b>Limestone</b> (foraminifera-rich) (Sections 1 to 5); rather massive darker colored <b>Nannofossil Mudstone</b> (Sections 6 and 7, CC); all the sediments with 1% glauconite, sharp lower boundaries often gradational upwards. Evidence of soft sediment deformation and moderate/minor bioturbation. Current scours at base of coarse units, sometimes rip-up clasts, clay clasts, laminations throughout.
						2	5G 4/1 Minor Lithology: Section 2, 145-150 cm = <b>Conglomerate</b> grading upward to sand. Red and black mudstone chunks and several other kinds of lithic fragments. <b>SMEARS:</b> 2-75, 3-75 ( <b>Nannofossil Limestone</b> ) Quartz, Feldspar 2-3% Nannofossils 60-70% Clay minerals 10-17% Glauconite 1% Opaque minerals 2-5% Heavy minerals, mica 5%
						3	5GY 4/1 2-95 ( <b>Limestone</b> ) Quartz, Feldspar 1% Carbonate unspecified 67% Clay minerals 10% Foraminifers 20% Opaque minerals 1% Glauconite 1% Heavy minerals, mica TR
						4	4-75, 6-120, CC ( <b>Nannofossil Mudstone</b> ) Sand 3-10% Quartz, Feldspar 2-15% Silt 7-15% Clay minerals 10-40% Clay 75-90% Opaque minerals 5-10% Heavy minerals, mica 1-5% Foraminifers 3-5% Nannofossils 25-40% Glauconite (6-120, CC) 1%
						5	5G 4/4 <b>CARBON-CARBONATE:</b> 3-46 (1.7, 0.1, 13)
						6	7.5YR N/3
						7	
CC	CM	B		CC			

SITE 445		HOLE		CORE 80		CORED INTERVAL: 749.0-749.5 m	
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER		SECTION	METERS	GRAPHIC LITHOLOGY	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS RADS				
Middle Eocene <i>C. gigas</i> Subzone or <i>C. staurion</i> Subzone (N) P.11	CP	FM	CP	FP	1	5GY 5/1	Dominant Lithology: dark greenish gray (5G 4/1 to 5G 4/1) nannofossil rich <b>Mudstone</b> . <b>Sandy Mudstone</b> with numerous silt layers and laminations throughout, glauconite streaks.  Silt layers generally show a sharp contact at base, and a gradational one upwards. Sequence and structure identical to Core 79. Some lenses, clay streaks, carbonate-rich sediment streaks; generally normal graded beds, but sometimes intense (Section 2, 145 cm, Section 6, 25 and 53 cm). Bioturbation minor to moderate.
						2	10R 2/2 <b>SMEARS:</b> 1-75, 2-95, 2-105 ( <b>Nannofossil Mudstone</b> ) Sand 7-20% Quartz, Feldspar 3-7% Silt 13-20% Clay minerals 30-48% Clay 20-60% Opaque minerals 0-20% Heavy minerals, mica 2-10% Foraminifers 3-5% Nannofossils 10-40% Glauconite 1-2%
						3	<b>CARBON-CARBONATE</b> 2-60 (2.5, 0.1, 20)
						4	5GY 4/1, 5G 4/1, 5B 4/1
						5	VOID
						6	VOID
						7	VOID
CC	CM			RM			



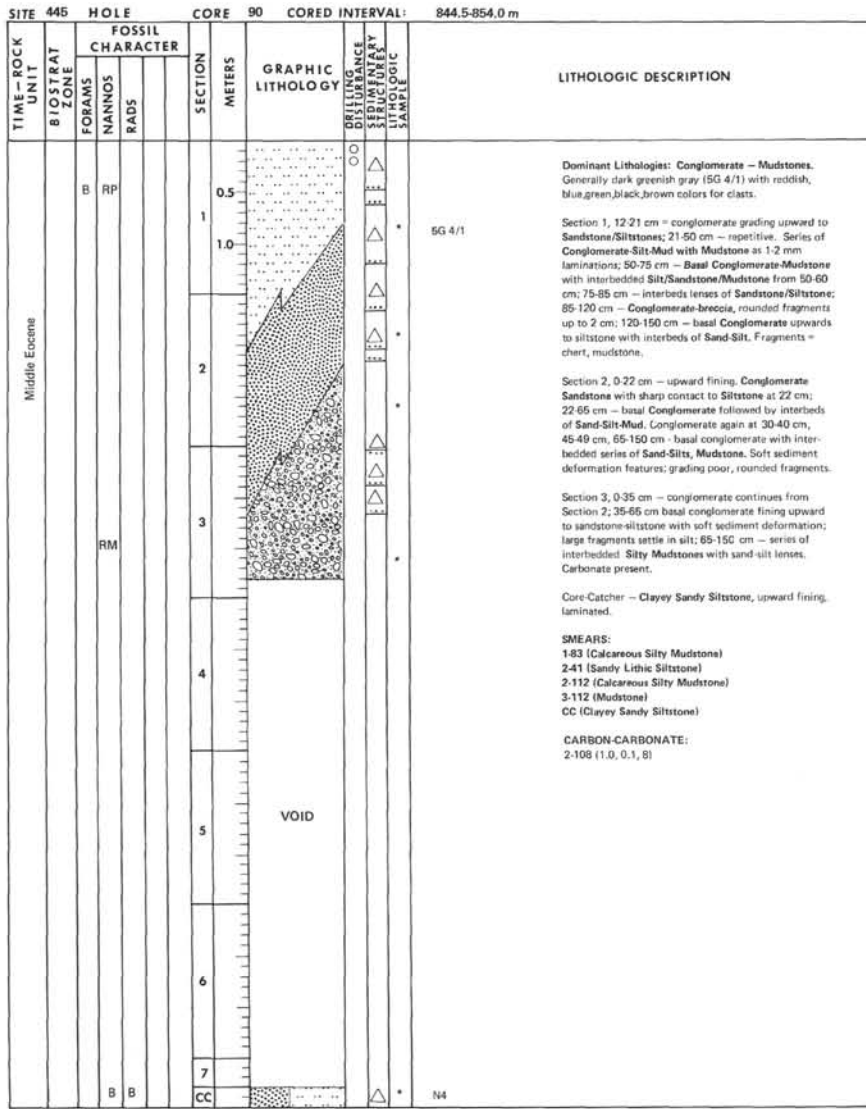
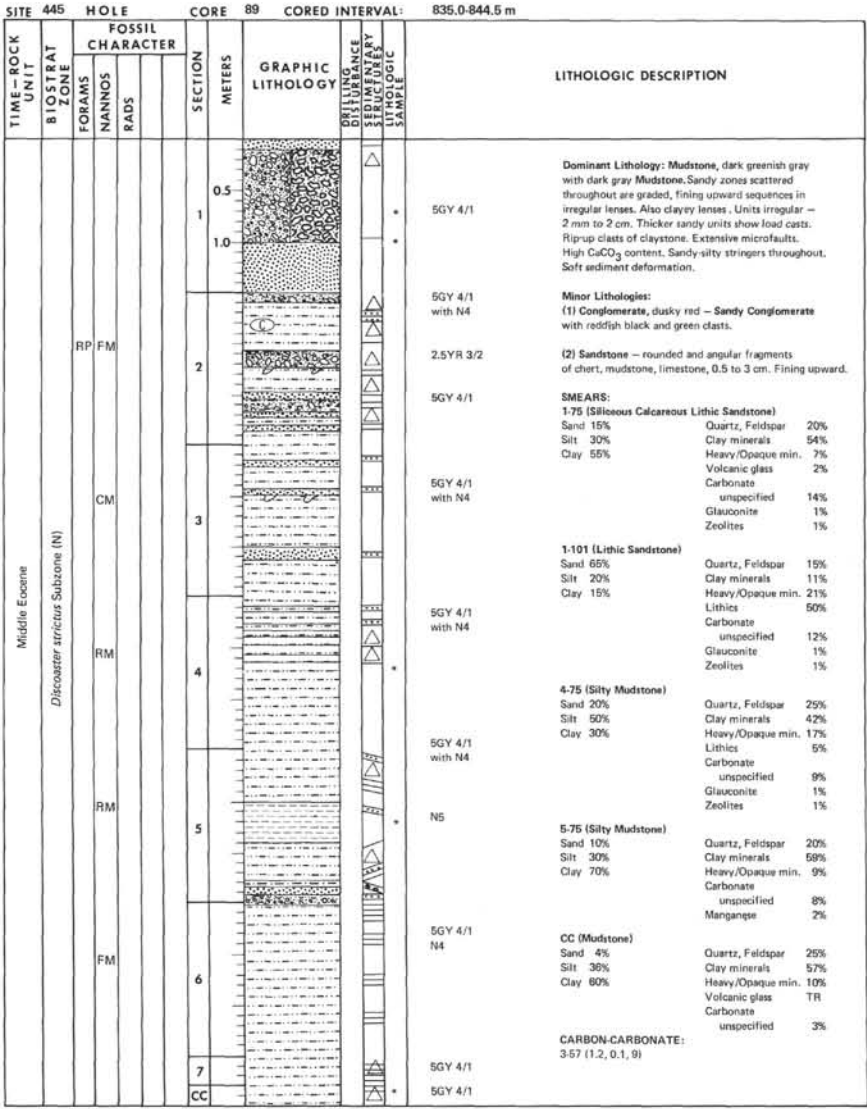


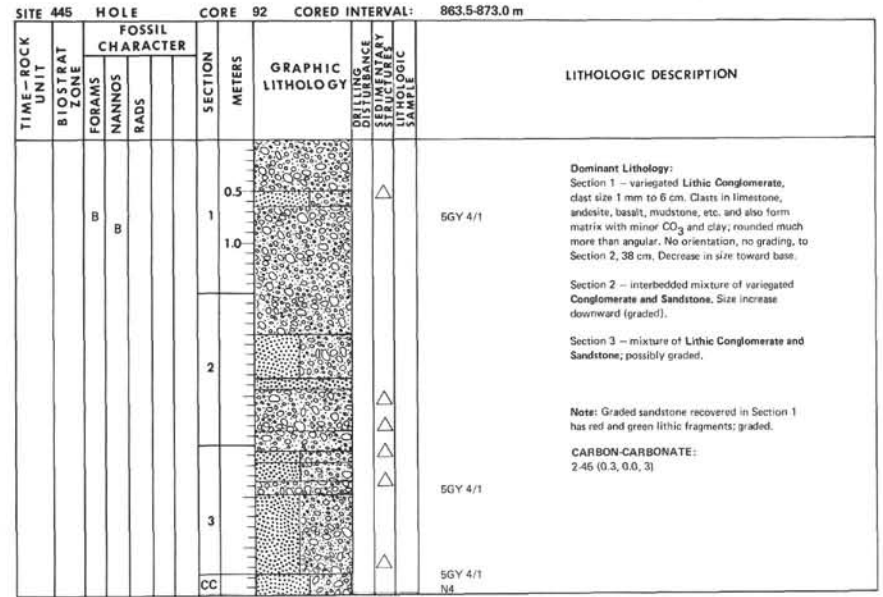
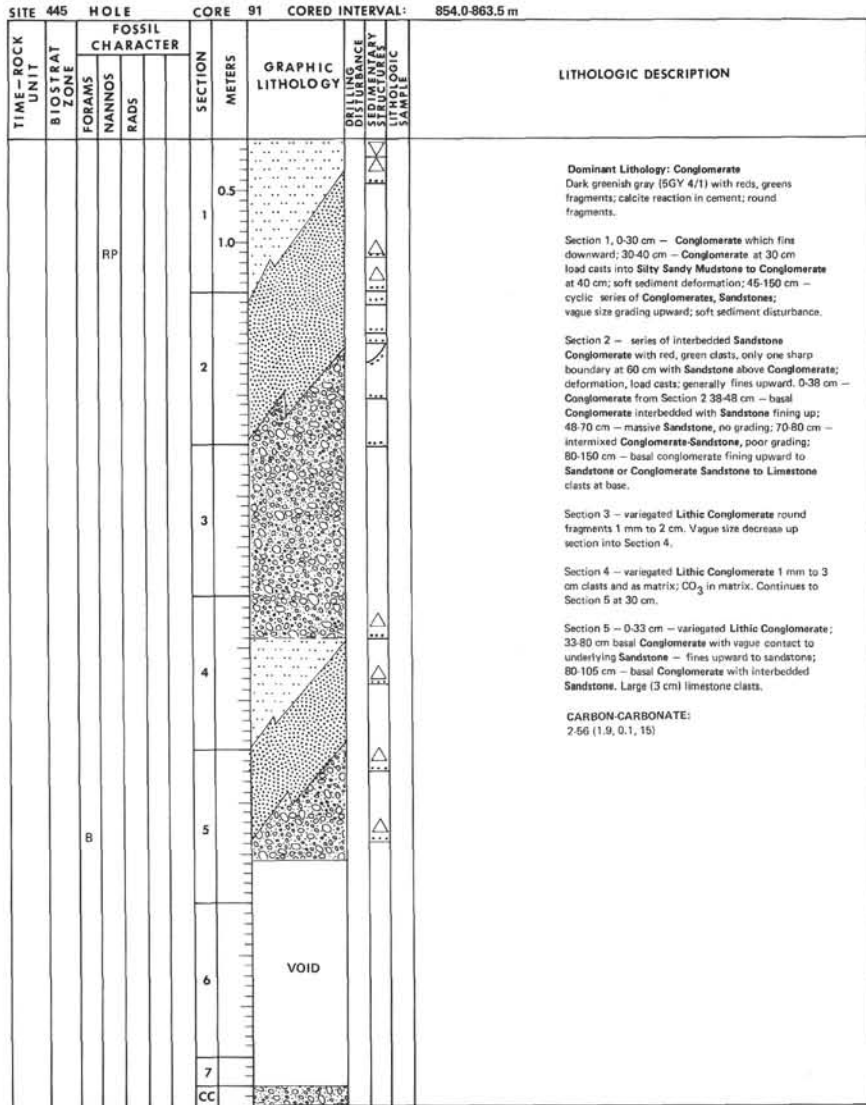




SITE 445		HOLE			CORE 87		CORED INTERVAL: 816.0-825.5 m				
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS							
Middle Eocene	C. gigas Subzone or C. flauvifolius Subzone (N) P.10(?)	B			1	0.5 1.0		OO	△	●	<p>Dominant Lithologies:</p> <p>Section 1 = alternating Silty Mudstone and Muddy Siltstone, grading.</p> <p>Section 2 = alternating Silty Mudstone, Muddy Siltstone, Sandstone. Numerous sediment structures.</p> <p>Section 3 = Calcareous Mudstone with sediment structures.</p> <p>Section 4 = Clayey Siltstone, with grading, slump structures.</p> <p>Section 5 = Silty Mudstone, white CO<sub>2</sub> laminations, local color darkening; microfaults.</p> <p>Section 6 = Silty Sandstone</p> <p><b>SMEARS:</b>                      1-64 (Mudstone)                      Sand 2% Quartz, Feldspar 16%                      Silt 20% Clay minerals 76%                      Clay 78%</p> <p>1-75 (Mudstone)                      Sand 25% Quartz, Feldspar 15%                      Silt 40% Heavy minerals 5%                      Clay 35% Opaques, Lithics 15%                      Carbonate unspecified 10%</p> <p>3-79 (Calcareous Mudstone)                      Sand 5% Quartz, Feldspar 12%                      Silt 30% Clay minerals 53%                      Clay 65% Carbonate unspecified 20%</p> <p>4-75 (Clayey Siltstone)                      Sand 4% Quartz, Feldspar 50%                      Silt 60% Clay minerals 38%                      Clay 36% Volcanic glass 5%</p> <p>5-75 (Silty Mudstone)                      Sand 5% Quartz, Feldspar 18%                      Silt 45% Clay minerals 52%                      Clay 50% Carbonate unspecified 5%                      Foraminifers 2%                      Nannofossils 3%</p> <p>6-80 (Silty Sandstone)                      Sand 55% Quartz, Feldspar 15%                      Silt 30% Lithics 60%                      Clay 15% Carbonate unspecified 5%</p> <p><b>CARBON-CARBONATE:</b>                      1-70 (1.3, 0.1, 10)</p>
					2	2					
		3	3		OO	△	●				
		4	4						OO	△	●
		5	5		OO	△	●				
		6	6						OO	△	●

SITE 445		HOLE			CORE 88		CORED INTERVAL: 825.5-835.0 m				
TIME-ROCK UNIT	BIOSTRAT ZONE	FOSSIL CHARACTER			SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY STRUCTURE	LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION
		FORAMS	NANNOS	RADS							
Middle Eocene	Discoaster stricus Subzone (N) P.10(?)	RP			1	0.5 1.0		OO	△	●	<p>Dominant Lithologies:</p> <p>Section 1 = alternating Silty Mudstone/Muddy Siltstones. Convolute beds, graded, slump structures.</p> <p>Section 2 = alternation of Silty Mudstone, Muddy Siltstones with convolute microfaults, Nummulites fragments.</p> <p>Section 3 = Silty Mudstones/Muddy Siltstones; with laminae. Convolute, graded beds, microfaults - local darkening.</p> <p>Section 4 = Muddy Siltstone, Silty Mudstone; coarser lenses, stringers, laminations, microfaults rip-up clasts.</p> <p>Sections 5 and 6 = dips. Mudstone with coarser laminations, stringers, lenses. White CO<sub>2</sub> laminations = calcarenites; microfaults, current scour.</p> <p><b>SMEARS:</b>                      1-100 (Sandy Silty Mudstone)                      Sand 30% Quartz, Feldspar 20%                      Silt 40% Clay minerals 53%                      Clay 30% Nannofossils 7%</p> <p>1-90 (Silty Mudstone)                      Sand 20% Quartz, Feldspar 25%                      Silt 40% Clay minerals 42%                      Clay 40% Lithics 5%                      Nannofossils 10%                      Carbonate unspecified 7%</p> <p>2-70 (Silty Mudstone)                      Sand 15% Quartz, Feldspar 15%                      Silt 45% Clay minerals 54%                      Clay 40%</p> <p>6-115 (Silty Sandy Mudstone)                      Sand 50% Quartz, Feldspar 40%                      Silt 30% Clay minerals 51%                      Clay 20%</p> <p><b>CARBON-CARBONATE:</b>                      1-94 (0.4, 0.1, 3)</p>
					2	2					
		3	3		OO	△	●				
		4	4						OO	△	●
		5	5		OO	△	●				

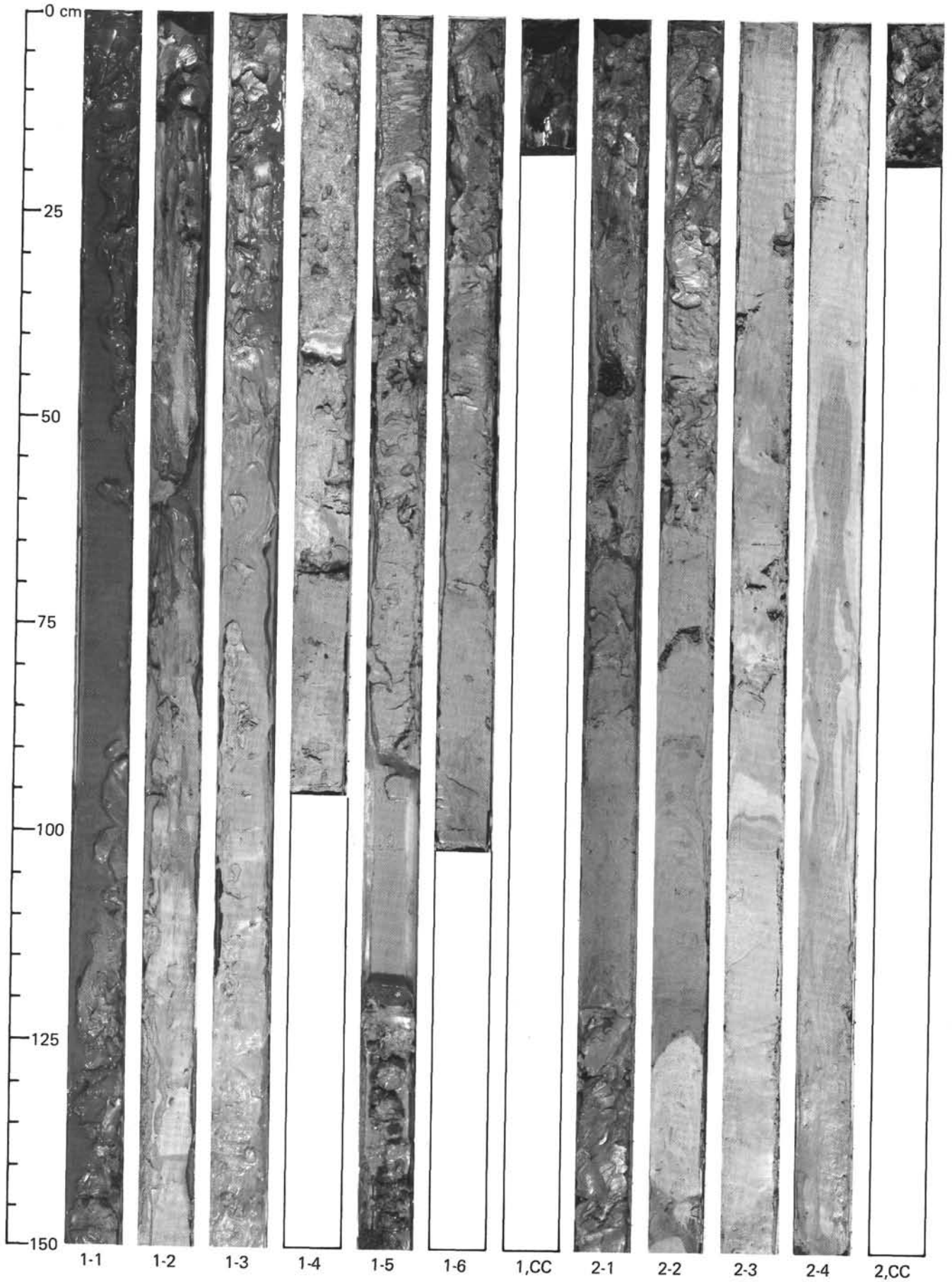


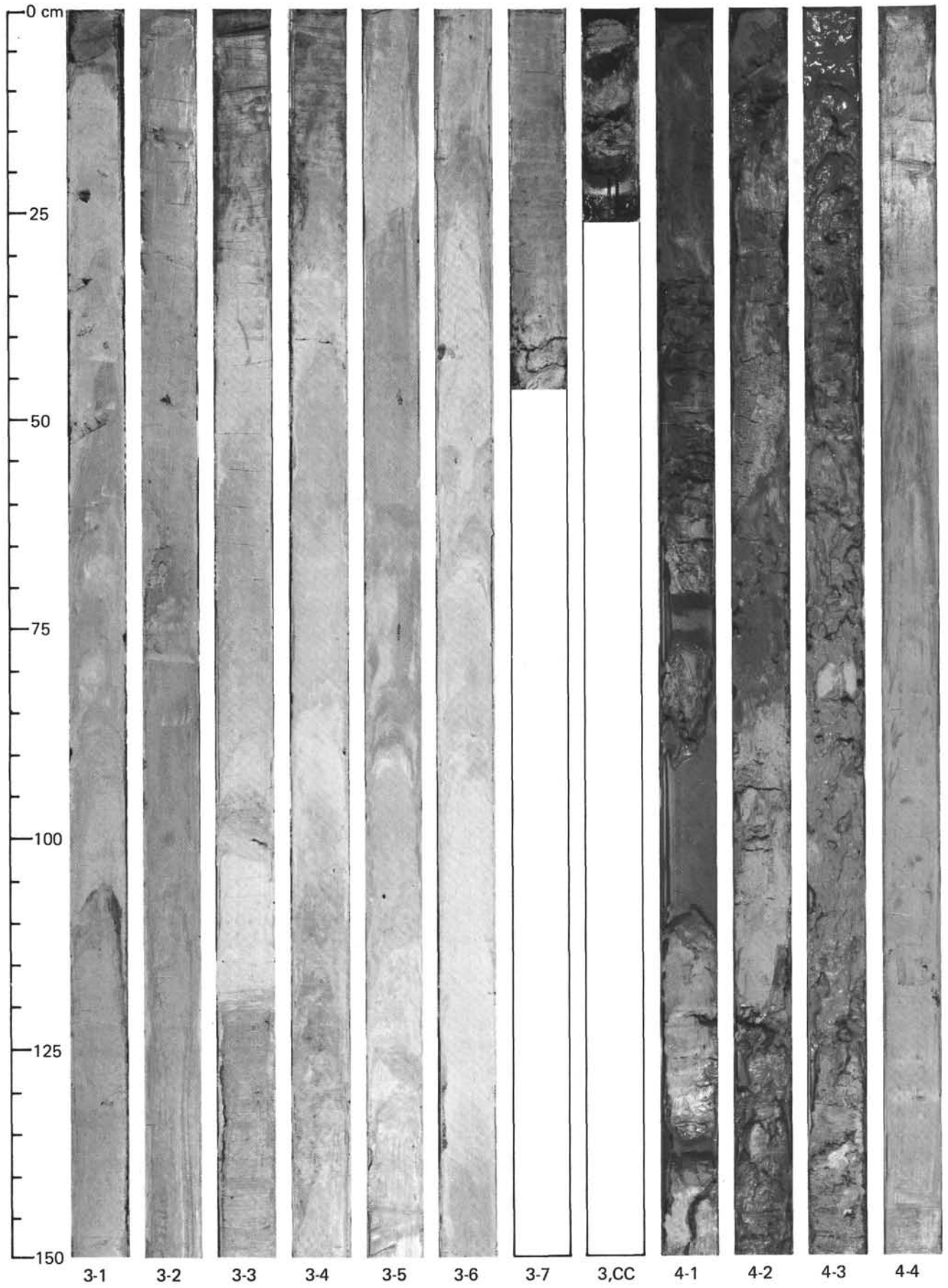


SITE 445 HOLE		CORE 93 CORED INTERVAL: 873.0-882.5 m		FOSSIL CHARACTER	SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
TIME-ROCK UNIT	BIOSTRAT ZONE	FORAMS	NANNOS							RADS
		B			0.5 1				SGY 4/1	
					1.0 2				Section 3 - beginning at 140 cm in Section 2, Conglomerate becomes unimodal (no large clasts) (matrix only). Dark greenish gray with red, green and white clasts. CO <sub>2</sub> cement approximately 1-2 mm. No grading. Size increase in Section 4.	
		B			3				Section 4 - Bimodal Conglomerate up to 35-55 mm clasts to approximately 3 cm. Angular fragments increase grading-fining up from base (coarse clasts at base). Core-Catcher - variegated pebble cycles.	
					4				CARBON-CARBONATE: 2-140 (0.4, 0.0, 3)	
					5 CC					

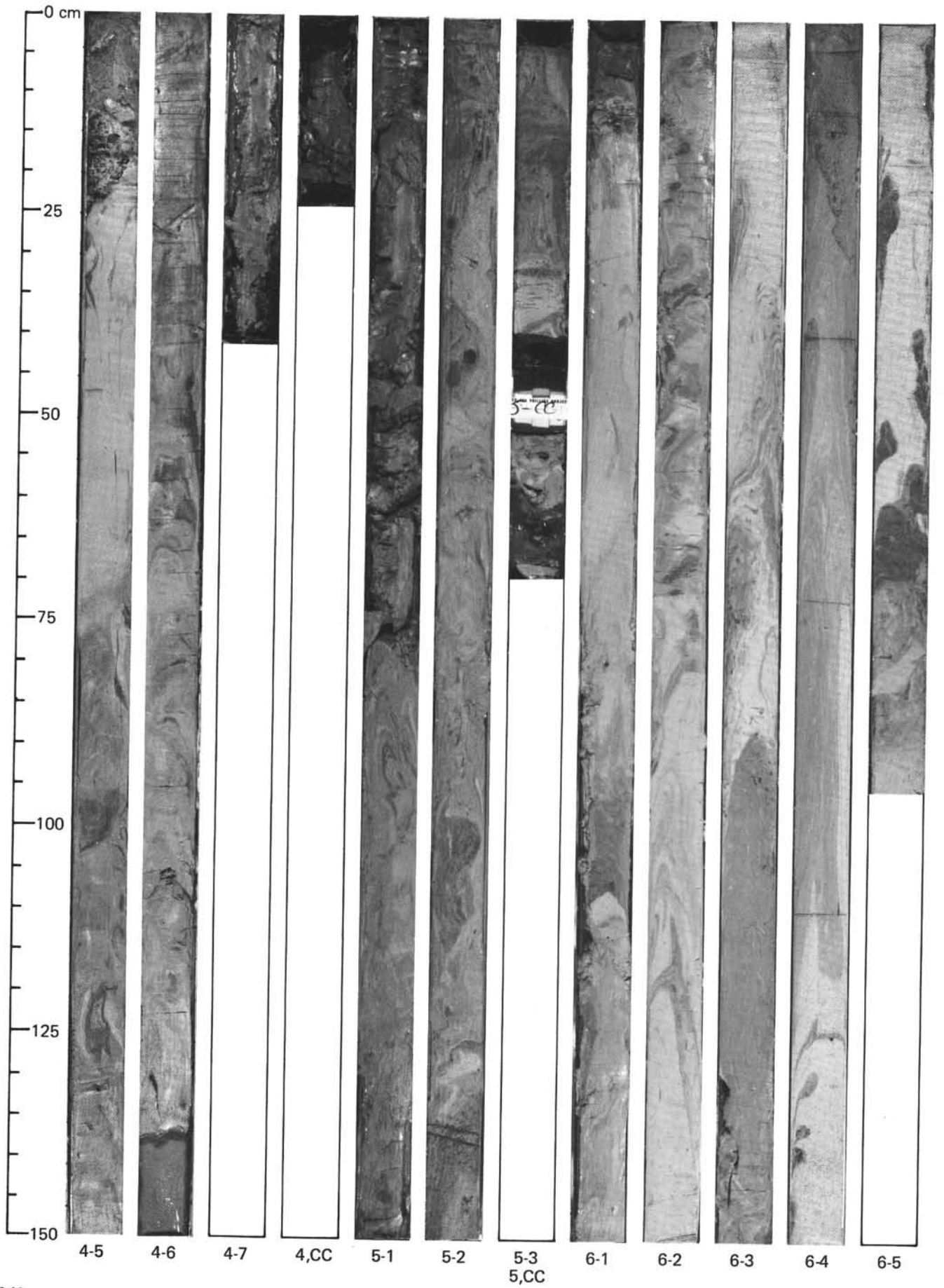
SITE 445 HOLE		CORE 94 CORED INTERVAL: 882.5-892.0 m		FOSSIL CHARACTER	SECTION METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE	SEDIMENTARY LITHOLOGIC SAMPLE	LITHOLOGIC DESCRIPTION	
TIME-ROCK UNIT	BIOSTRAT ZONE	FORAMS	NANNOS							RADS
		B			0.5 1				Dominant Lithology: Conglomerate. Basic dark greenish gray (SG 4/1) with red and green clasts.	
					1.0 2				Section 1, 0-55 cm - basal Conglomerate with load cast boundary to Sandstone at 55 cm. Decrease size upward to Sandstone at 30 cm followed by intermixed zone to 10 cm, high and bedding dip; 55-150 cm - basal medium bimodal Conglomerate, fining upward into Conglomerate Sandstone at 110 cm (55 to 100) with intermixing of Conglomerate Sandstone to 55 cm and appearance of grading with intermixing rounded fragments. Dip below 70 cm is with soft sediment microfaults.	
		B			3				Section 2, 0-15 cm - continuation of Conglomerate at Section 1. Clast size up to 0.5 mm. 15 to 70 or 80 = intermixed series of Sandstone, Conglomerate Sandstone merging into Conglomerate at 57 cm; bimodal. Conglomerate clast size to 2 cm, soft sediment load casts. 60-70 cm = bimodal Conglomerate. 70, 80 to 150 cm = basal bimodal Conglomerate fins upward to Conglomerate at 123 cm with horizontal bedding and to 70 or 80 cm. Basal Conglomerate to 35 cm of Section 3.	
					4				Section 3, 0-40 cm - bimodal variegated Conglomerate fins down; vague boundary at 40 cm. 40-67 cm - dominant Sandstone down boundary gradational to conglomerate. 67-140 cm - bimodal Conglomerate (3 cm clasts) fins down and upward. 140-150 cm - Sandstone-Conglomerate.	
					5 CC				Section 4, 0-30, 40 cm - intermixed; 30, 40-75 cm = intermixed to Conglomerate at 75 cm; pebble conglomerate starts at 150 cm into Section 5. CARBON-CARBONATE: 2-102 (0.4, 0.0, 3)	

Site 445

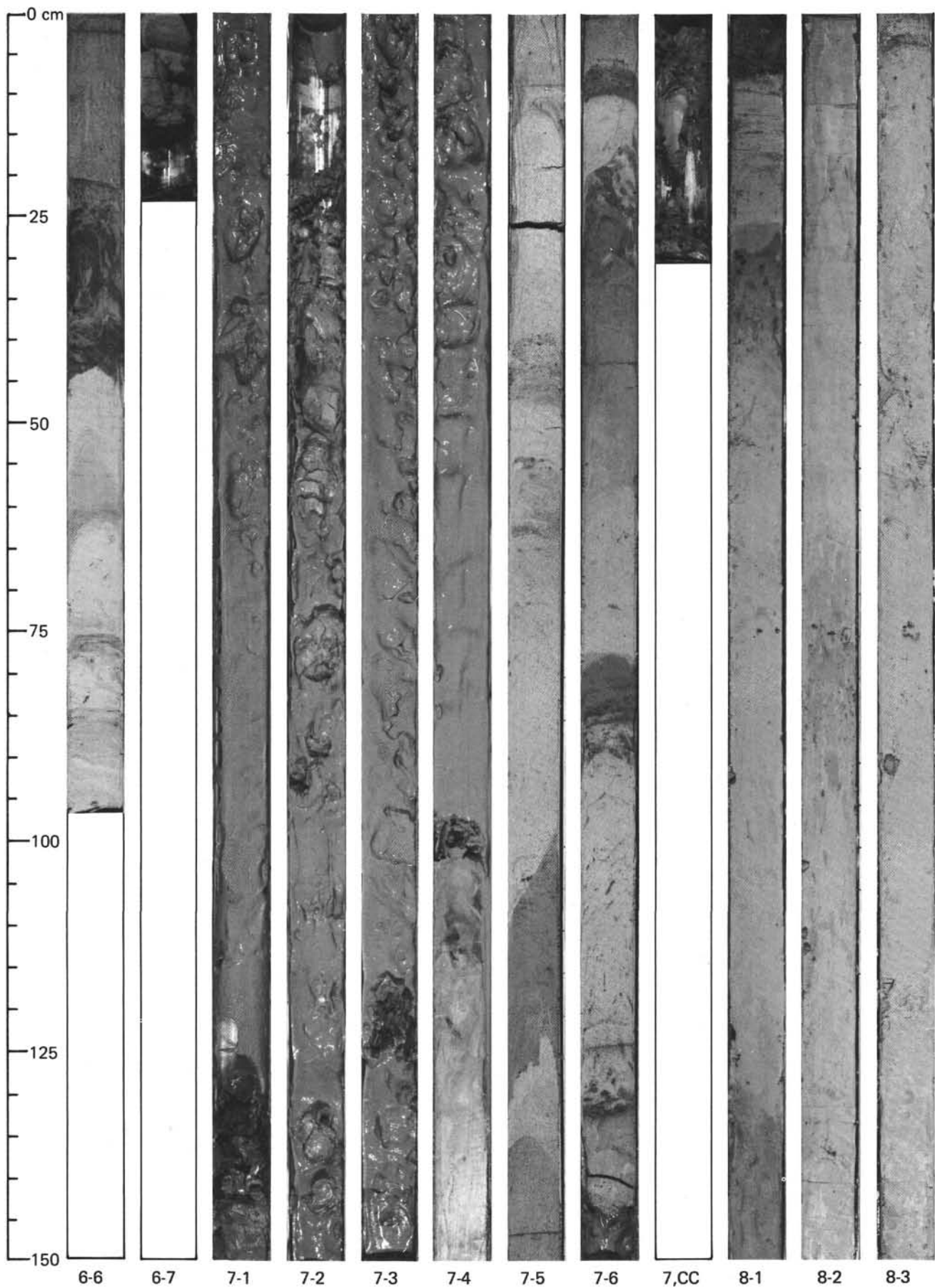




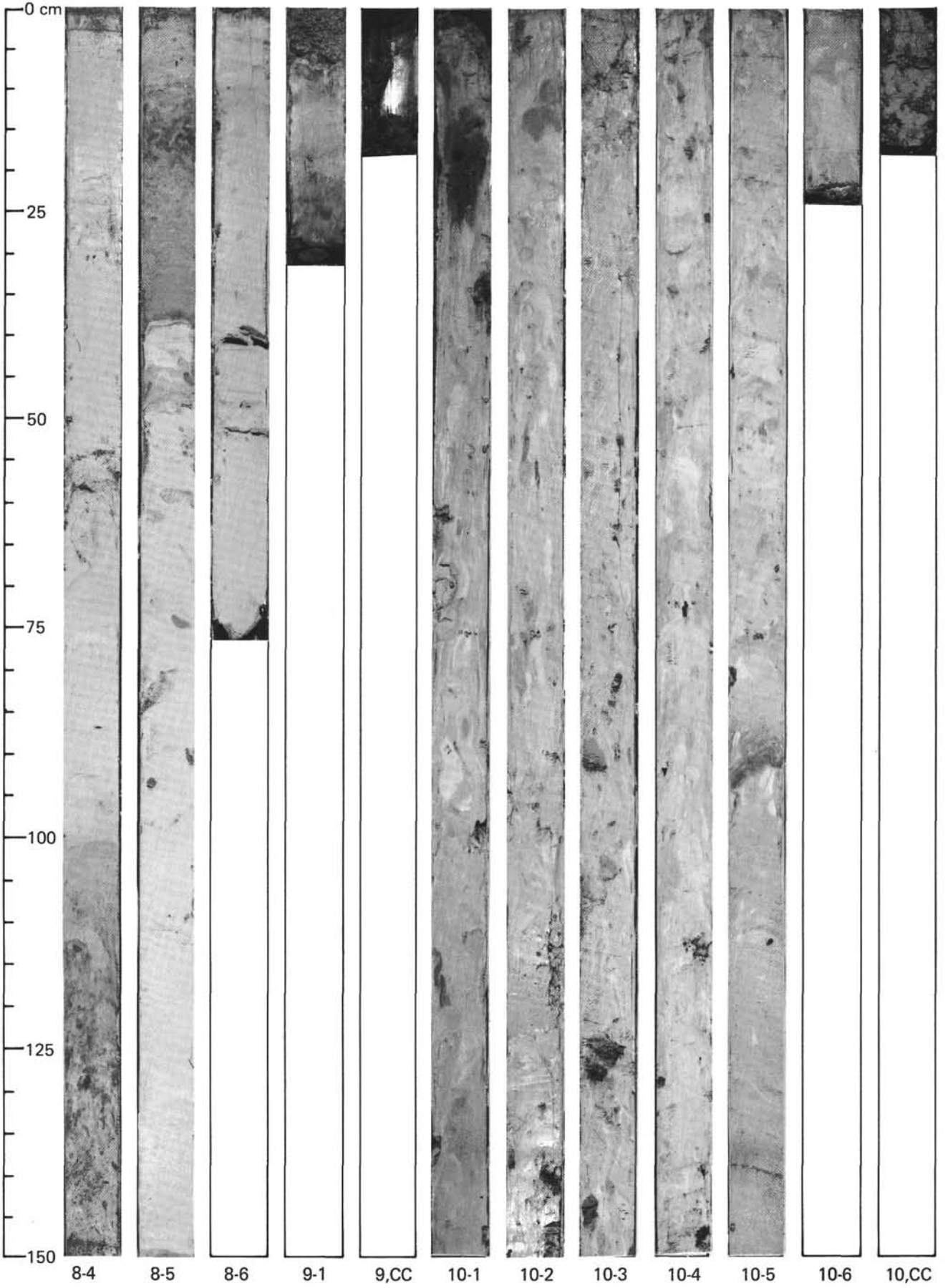




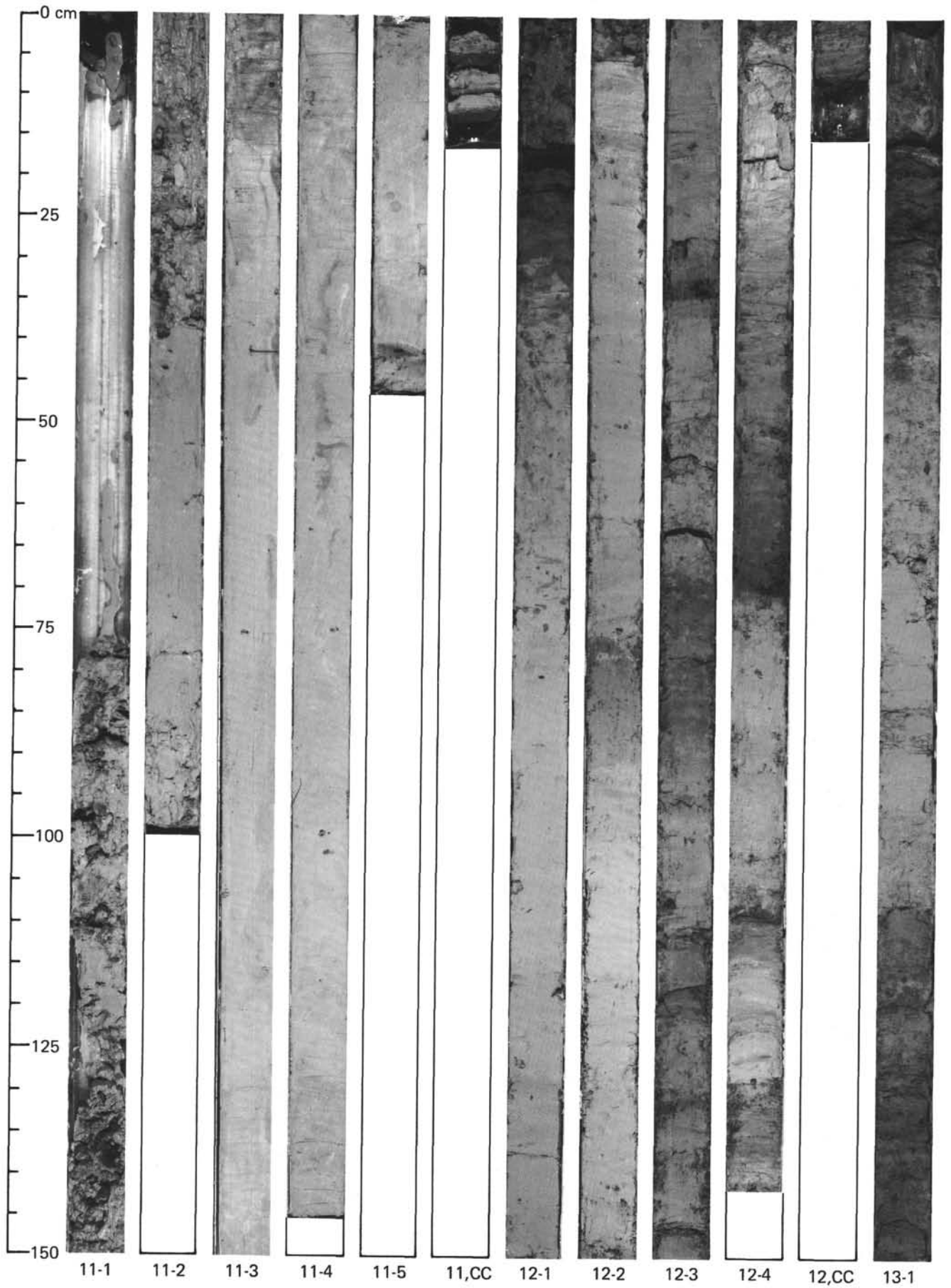
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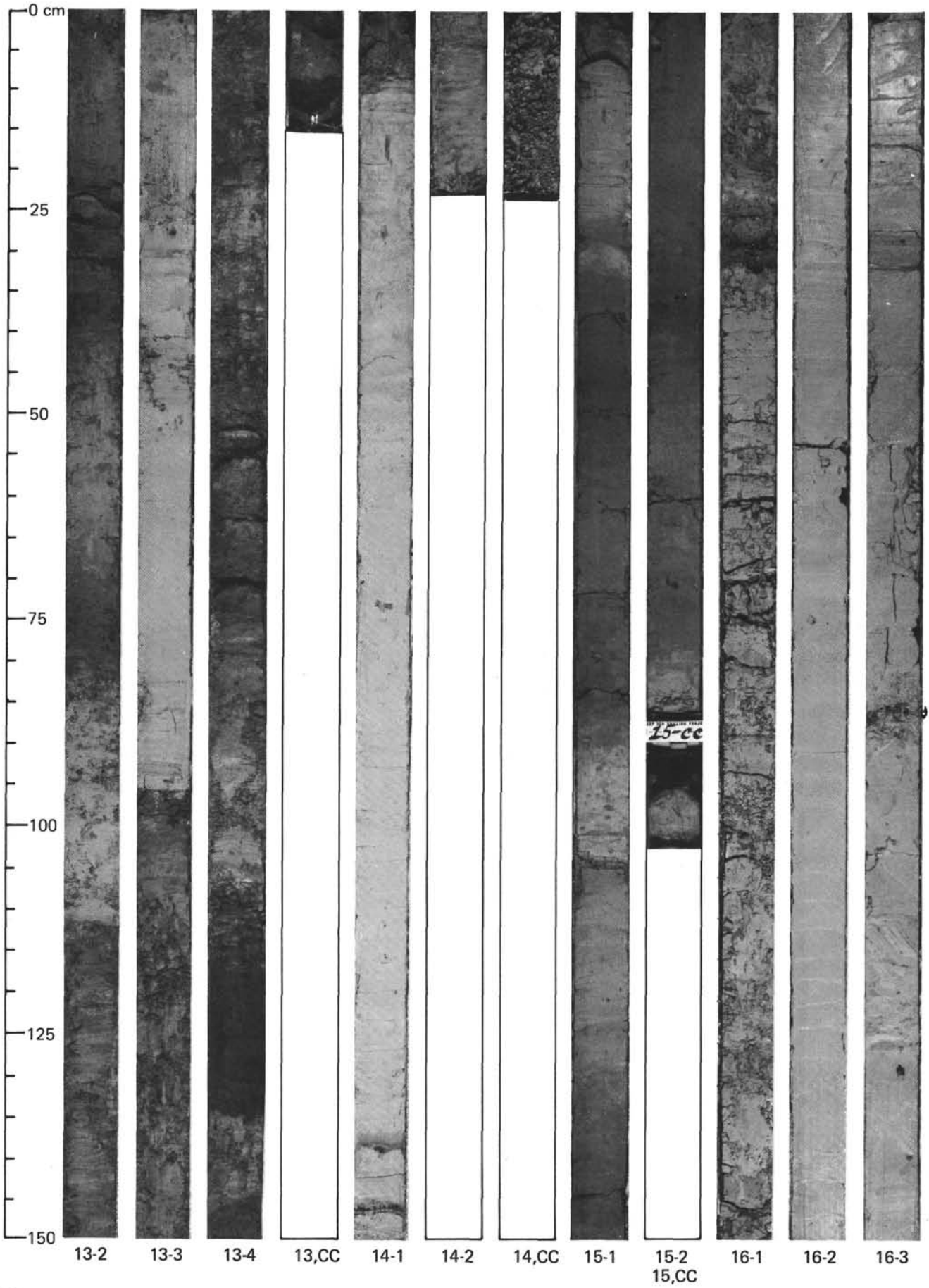
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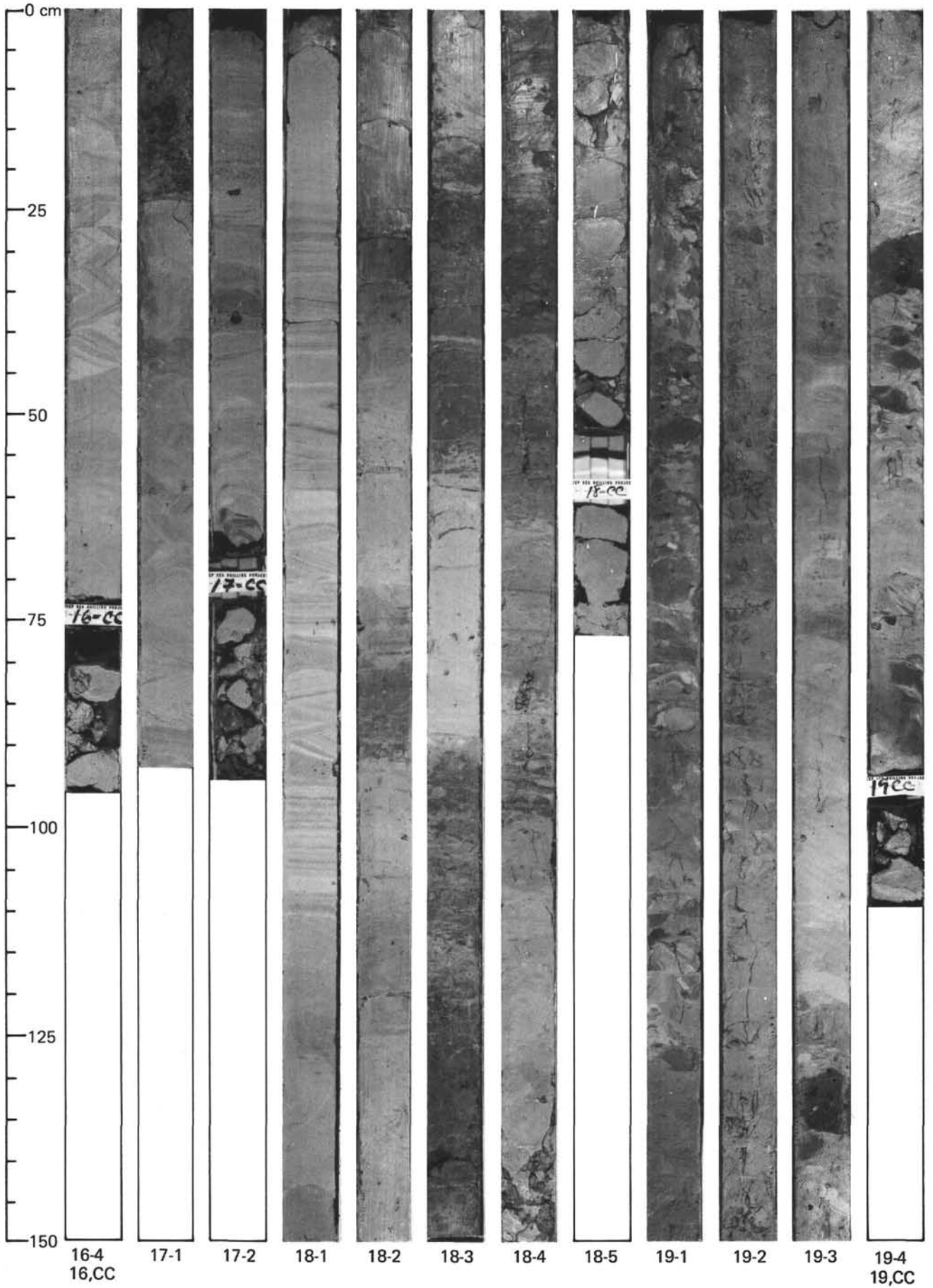


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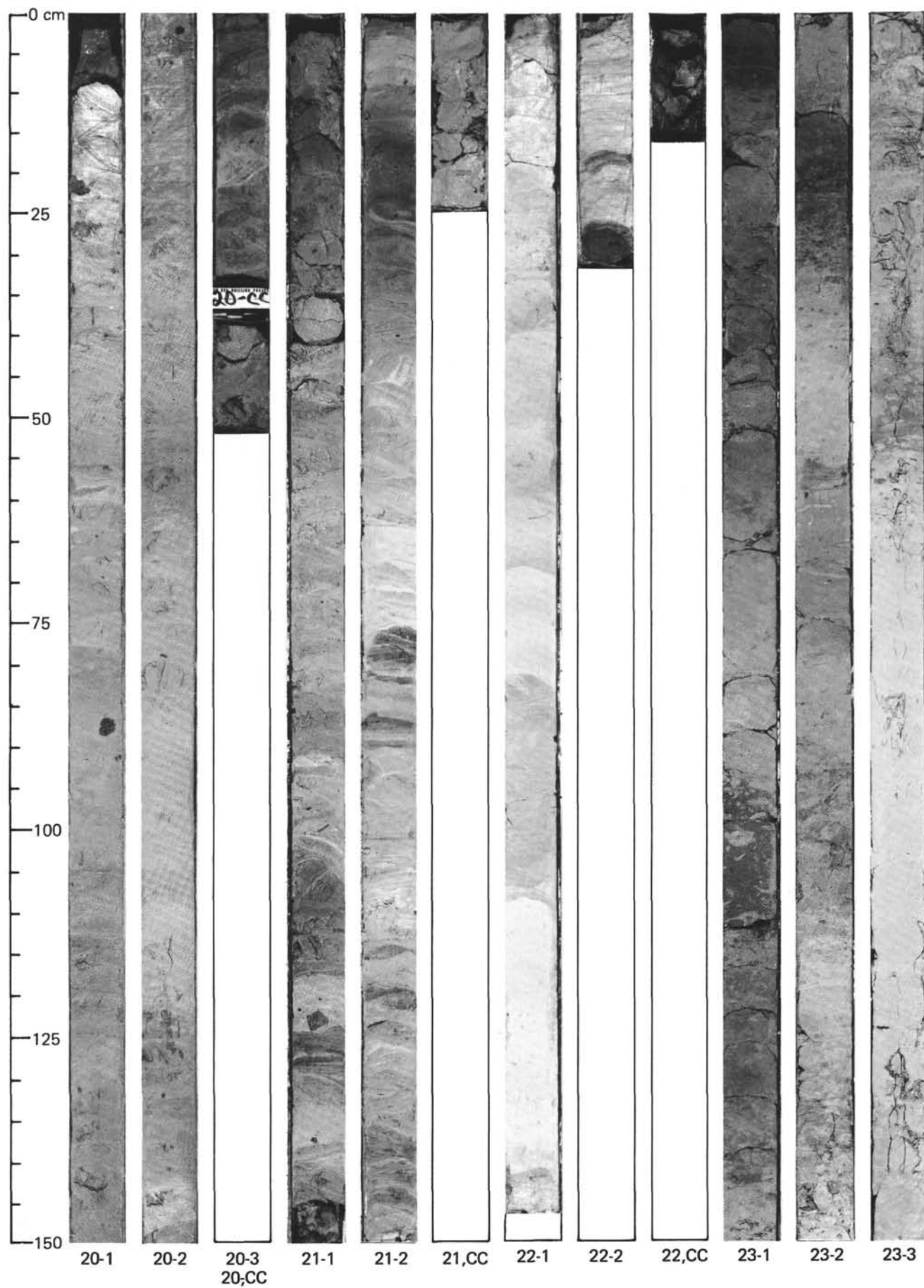


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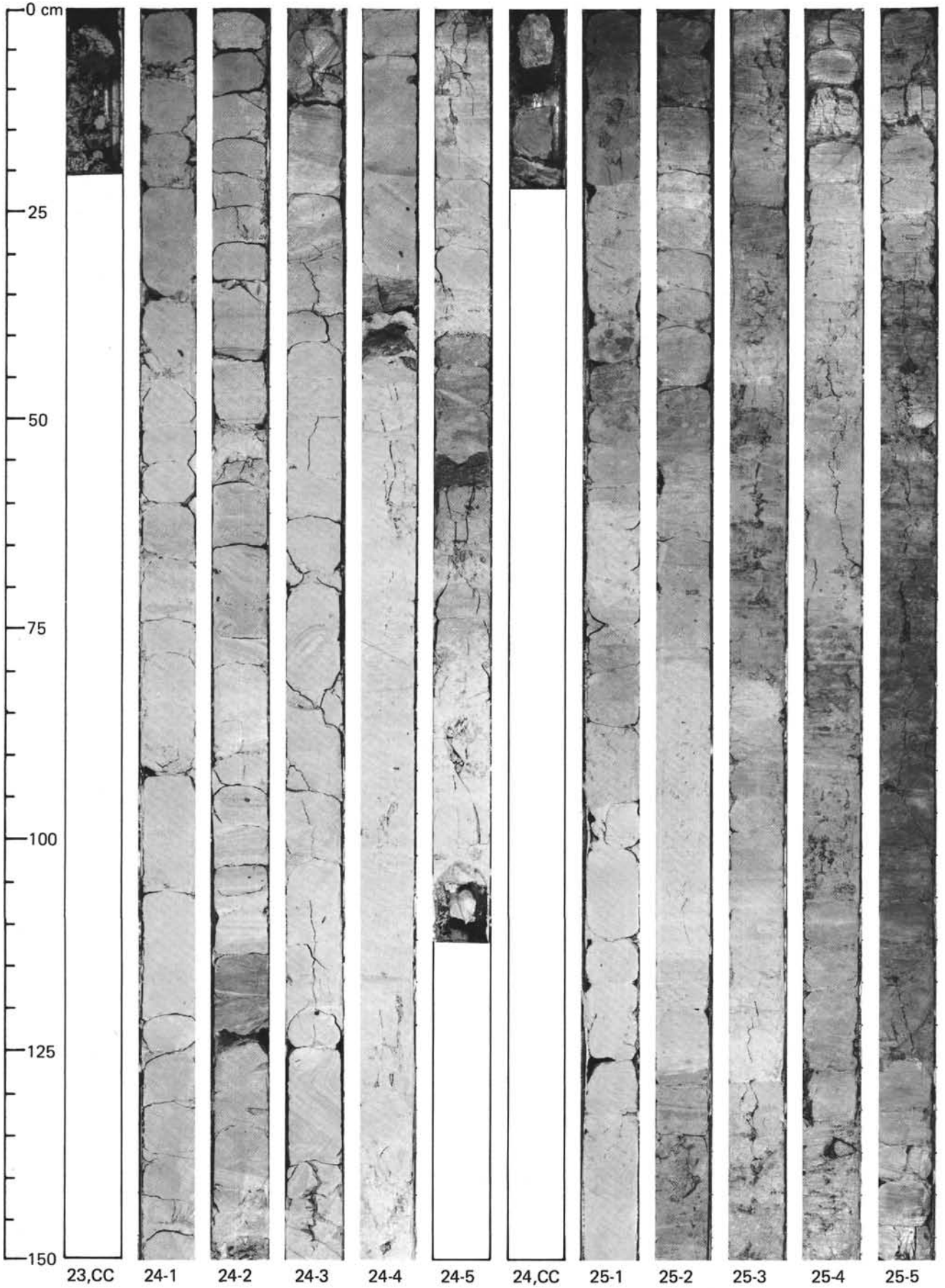




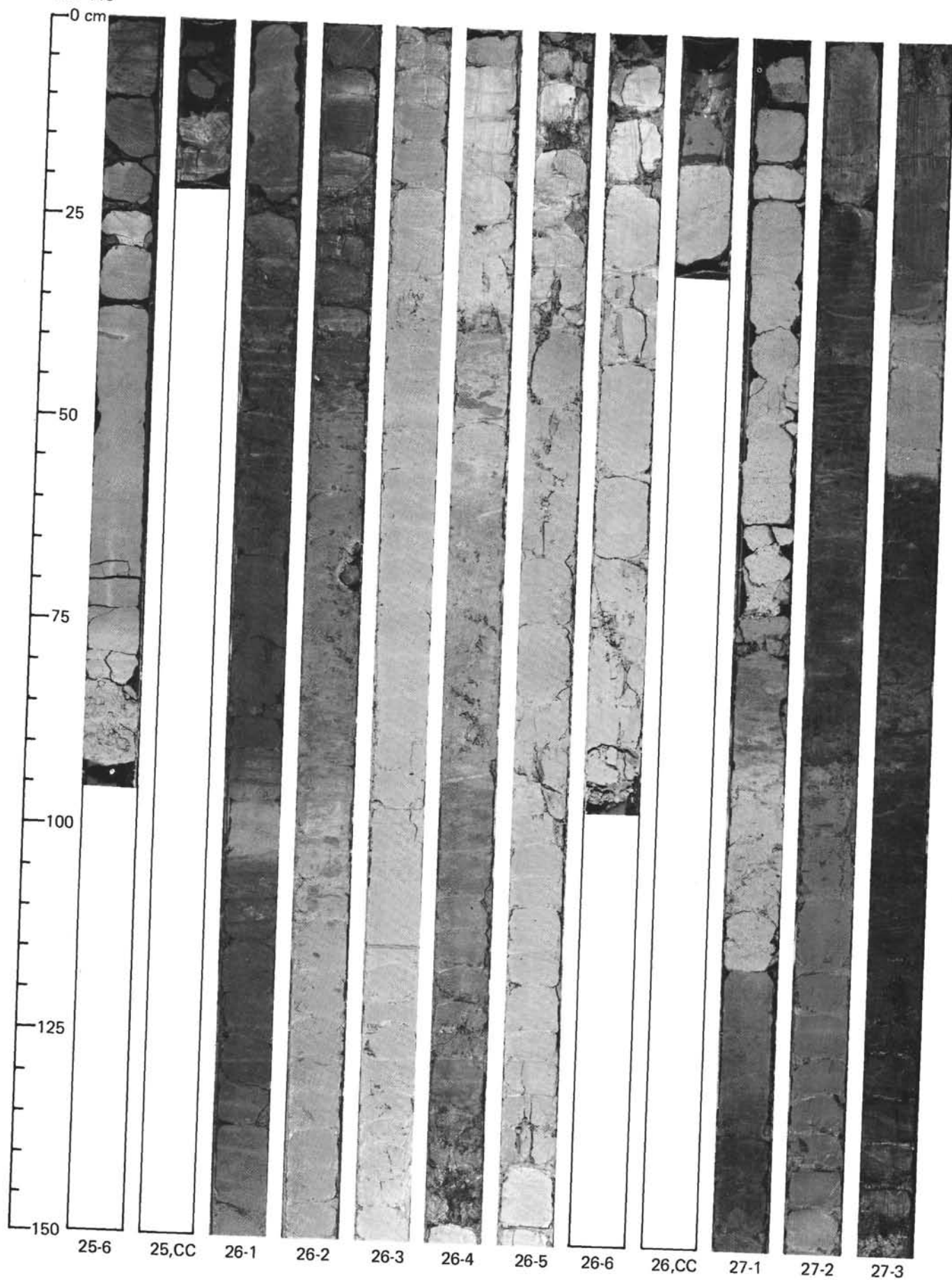
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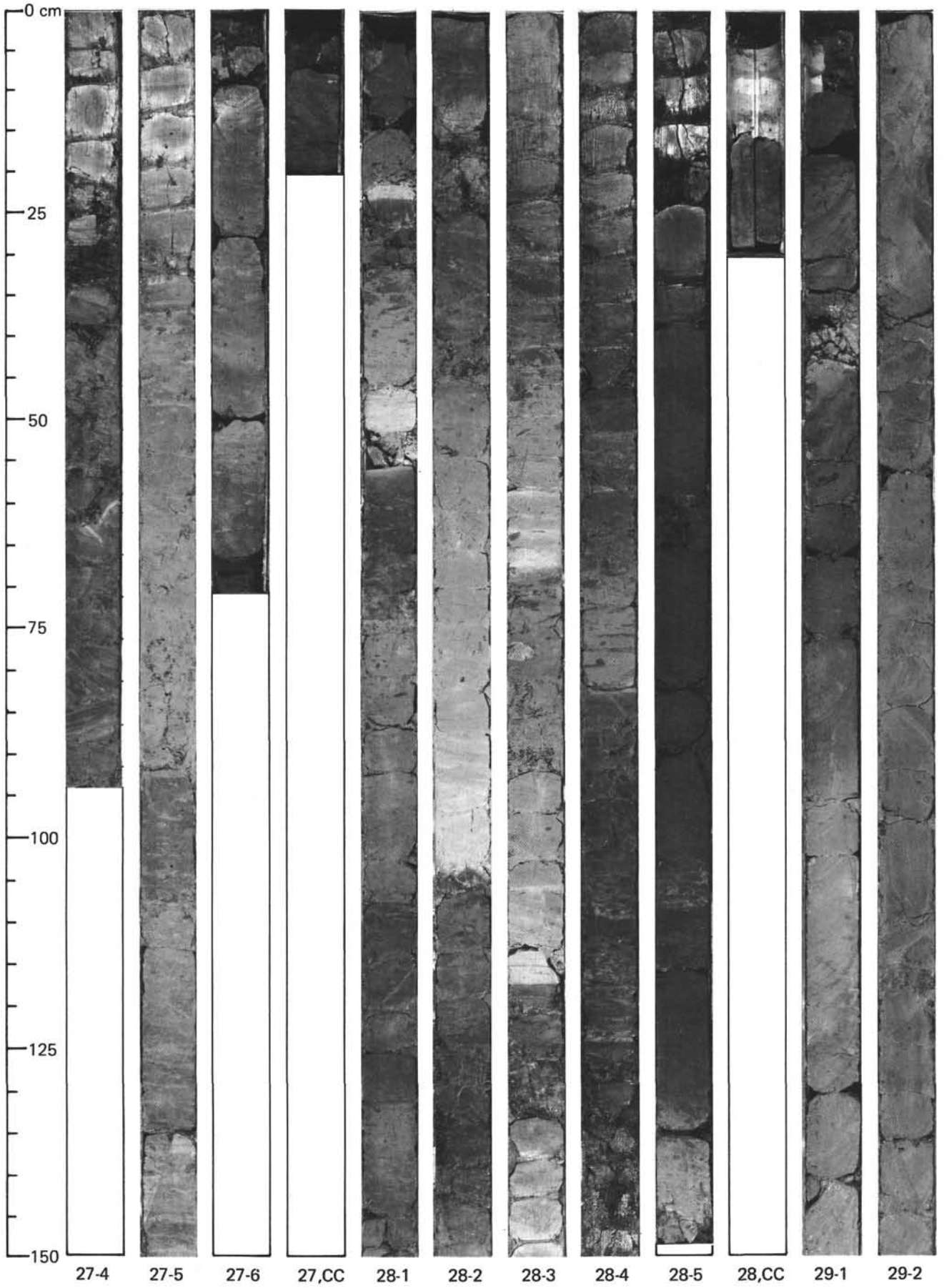
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Site 445



Site 445





Site 445

0 cm

25

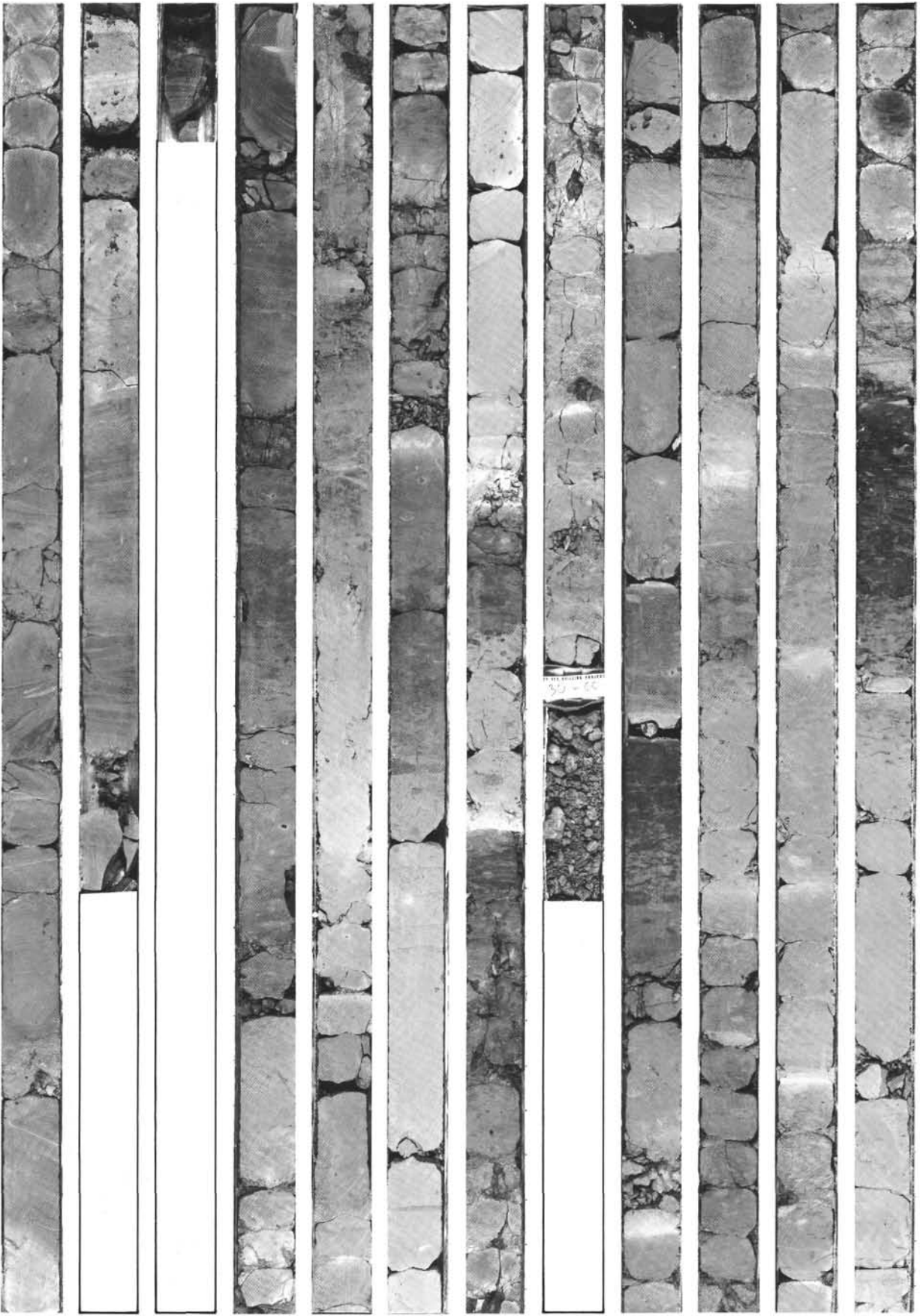
50

75

100

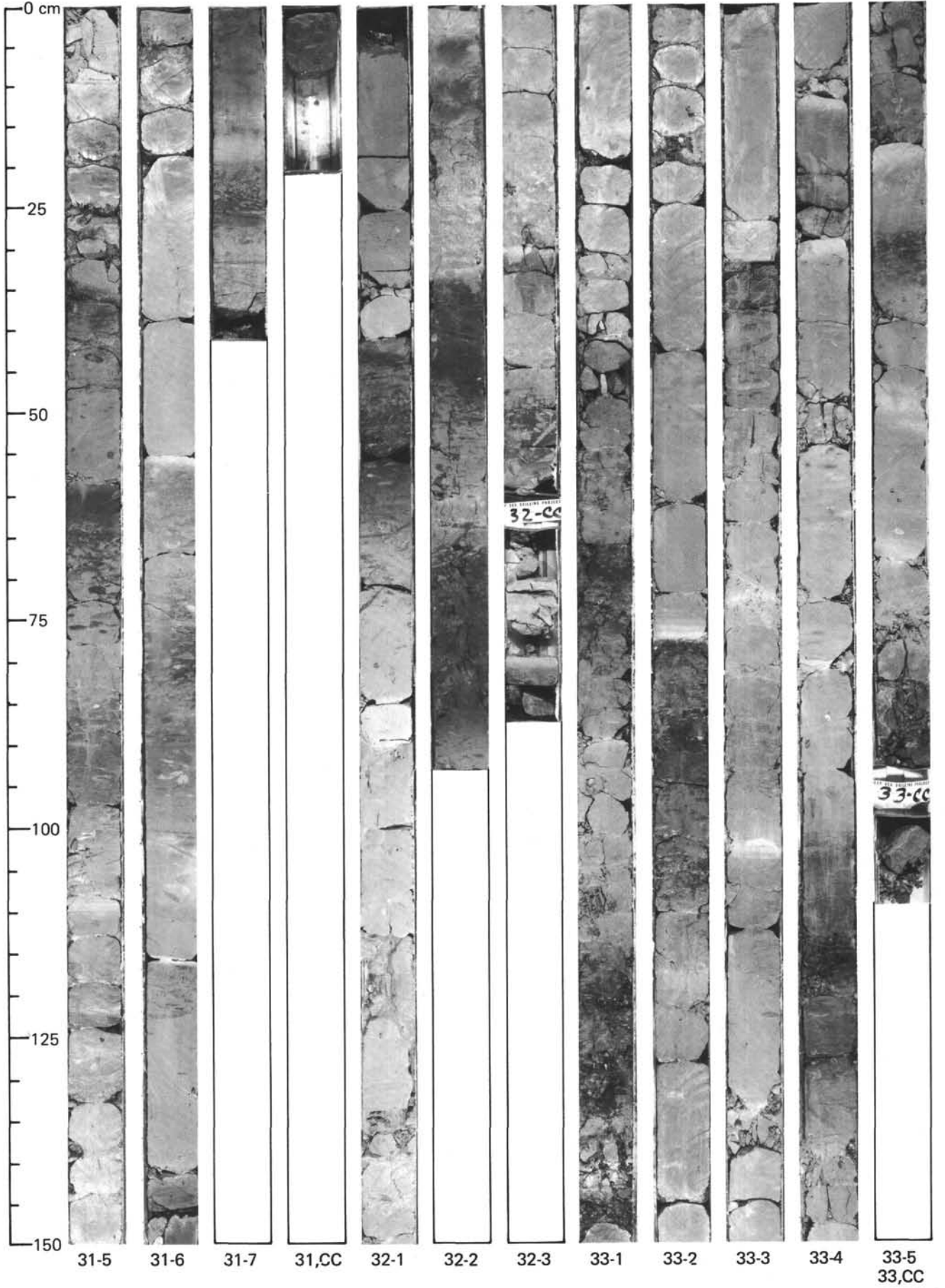
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150



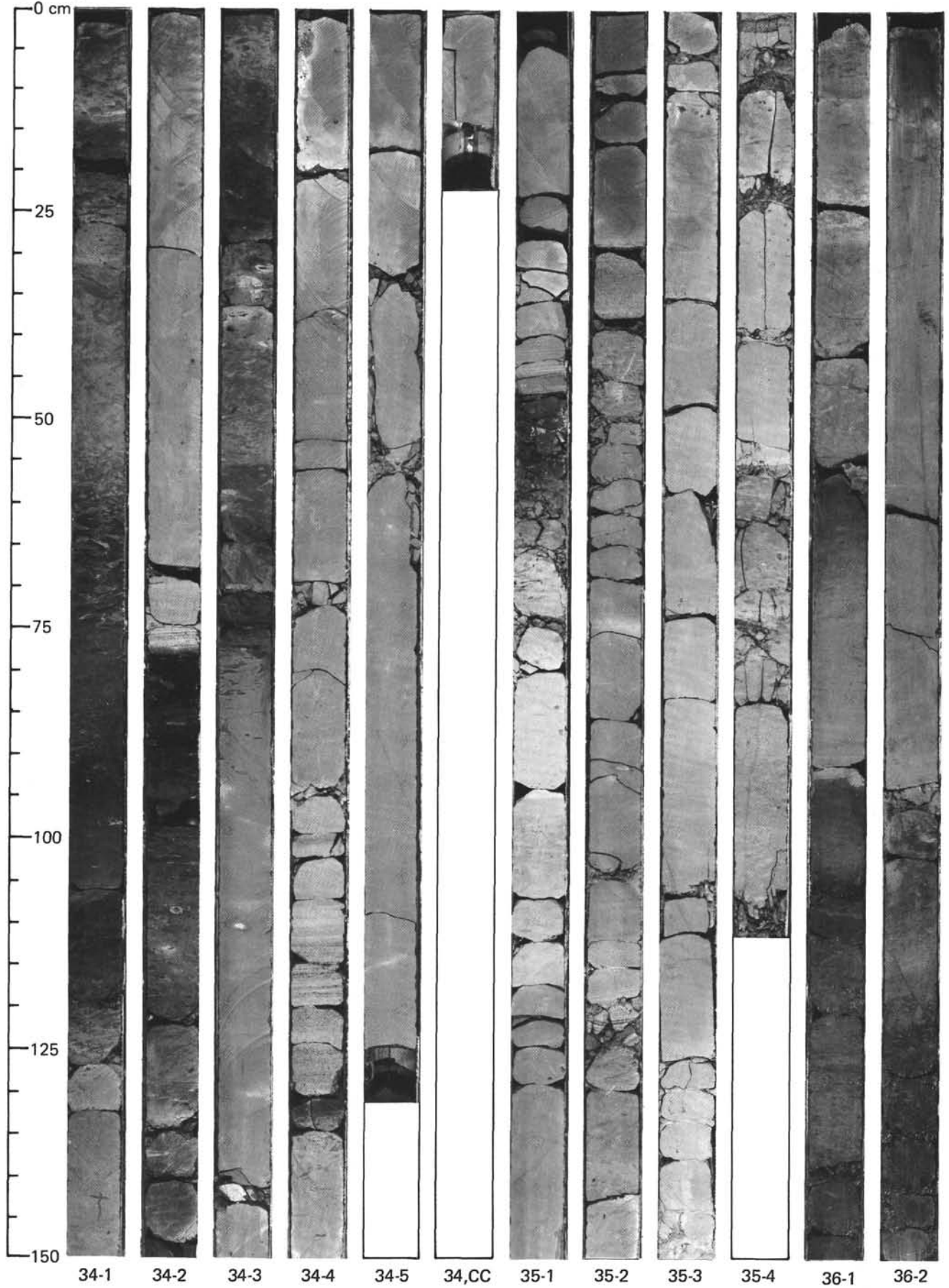
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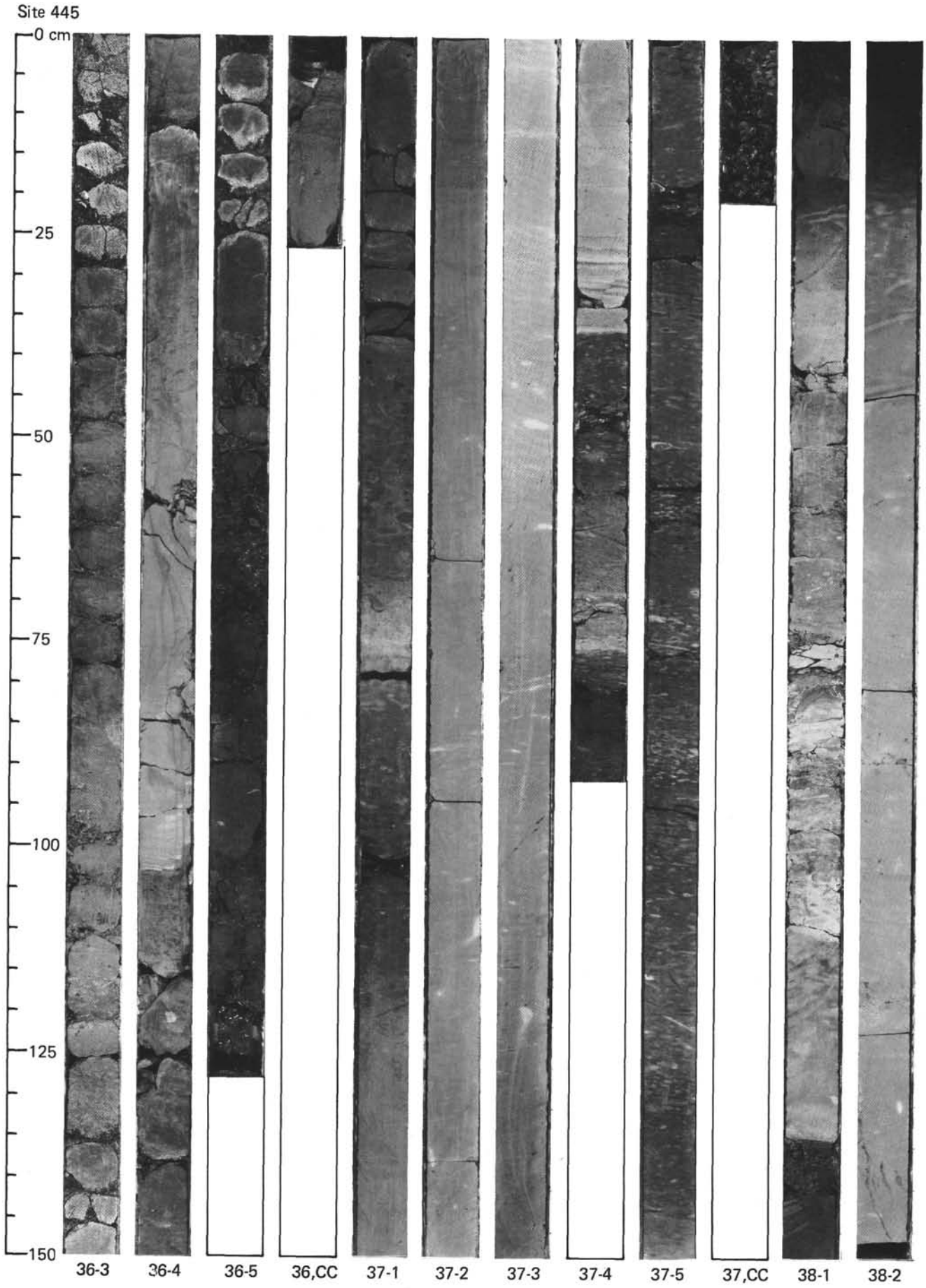
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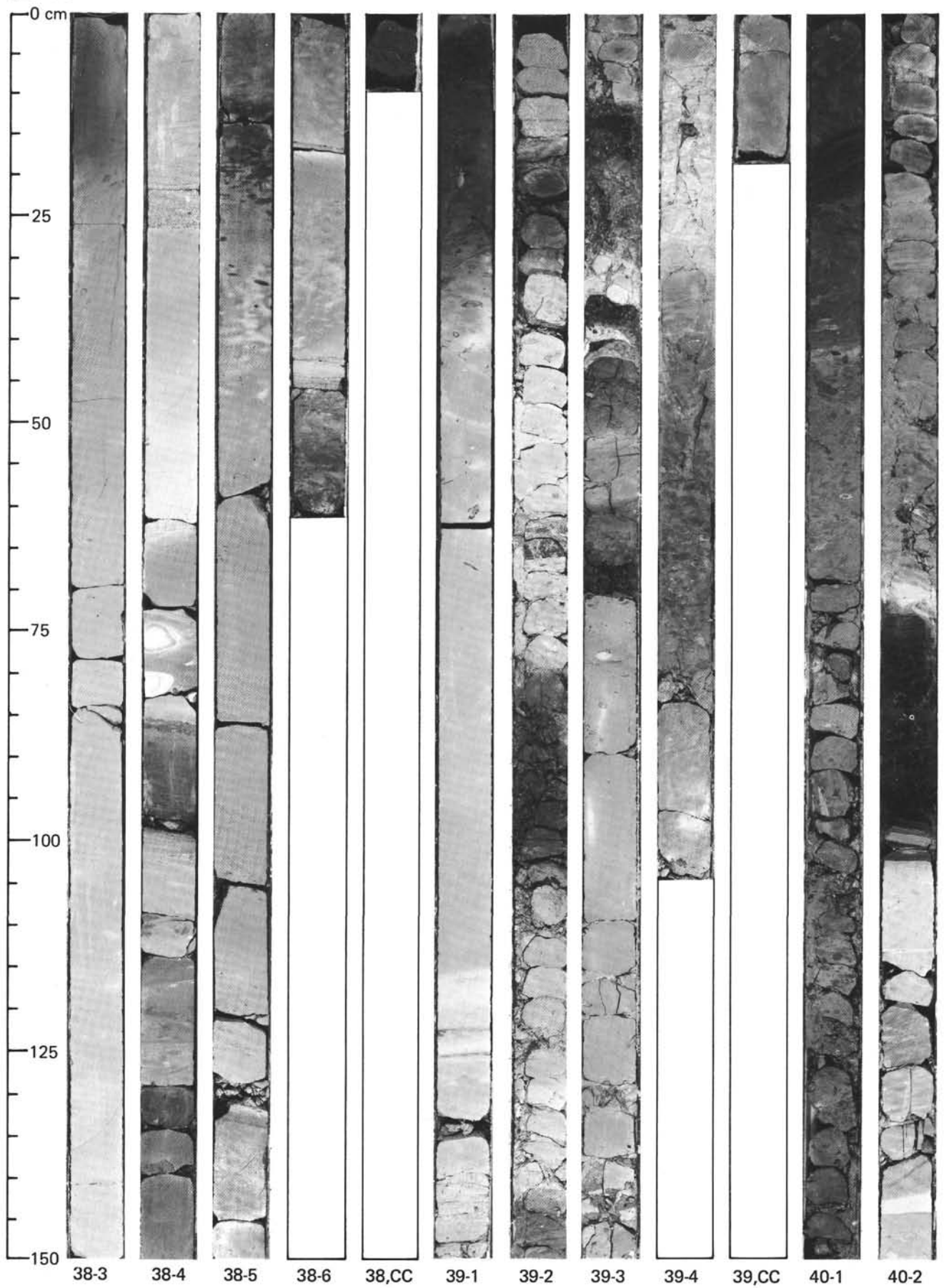


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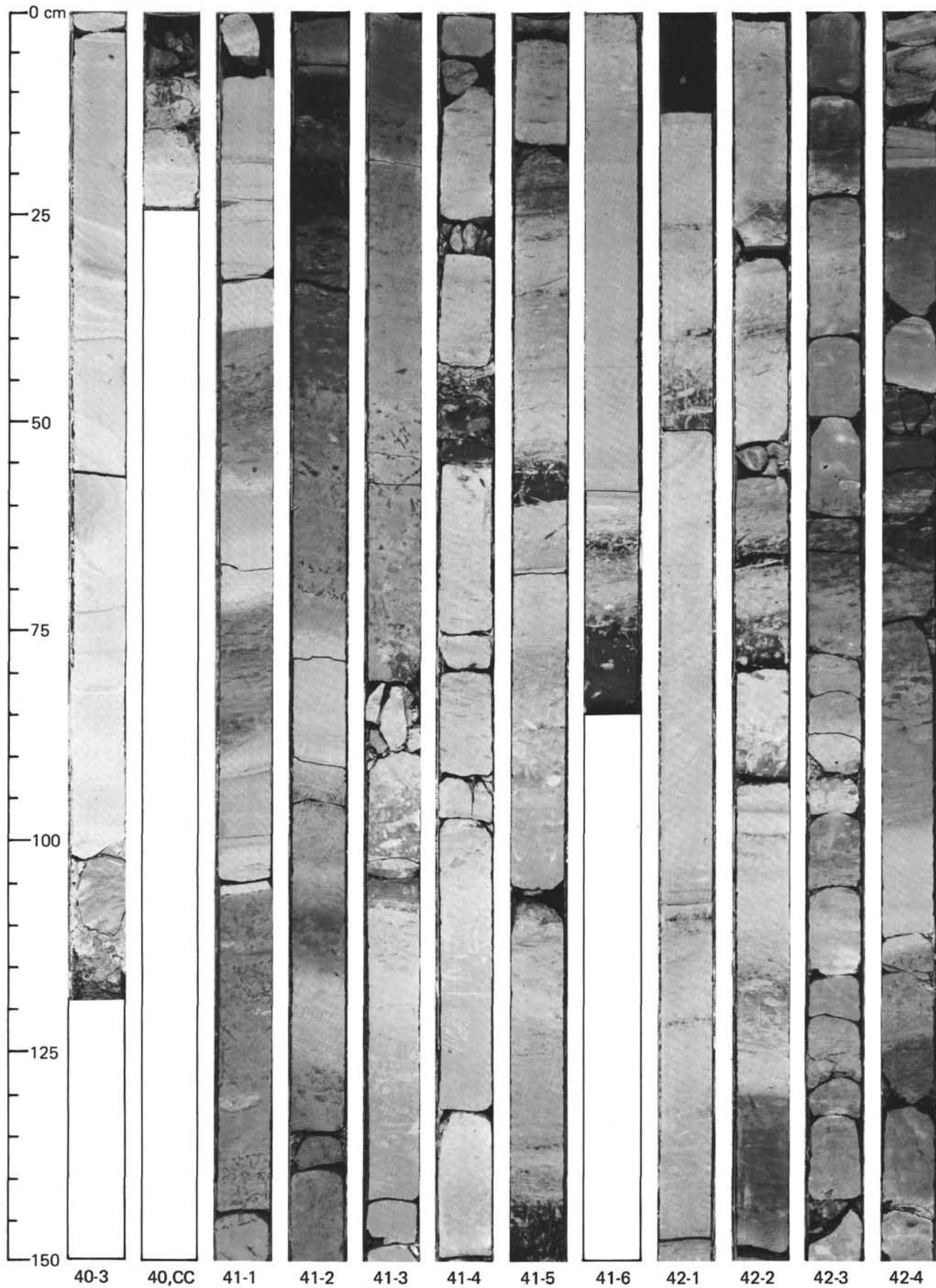




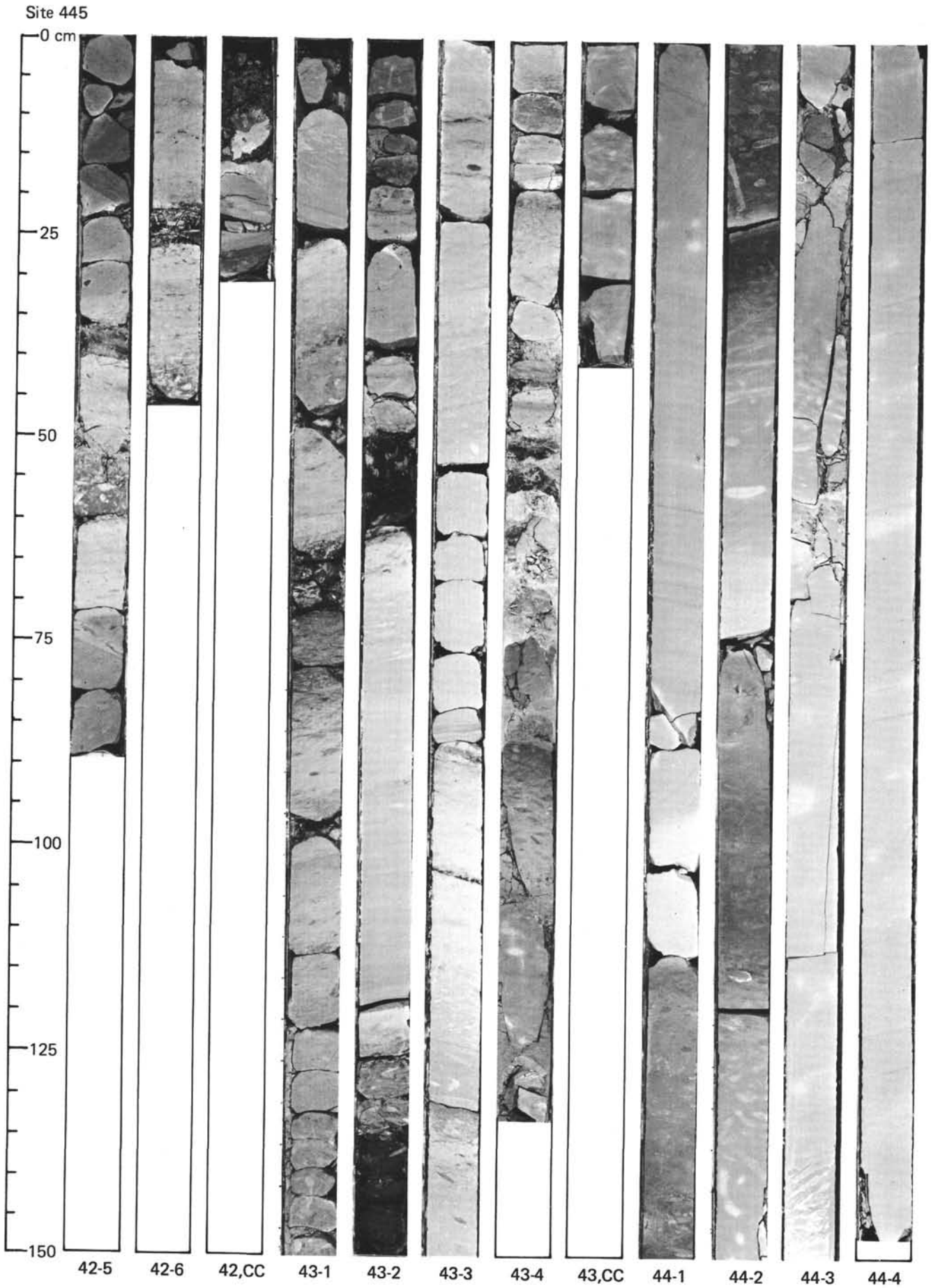
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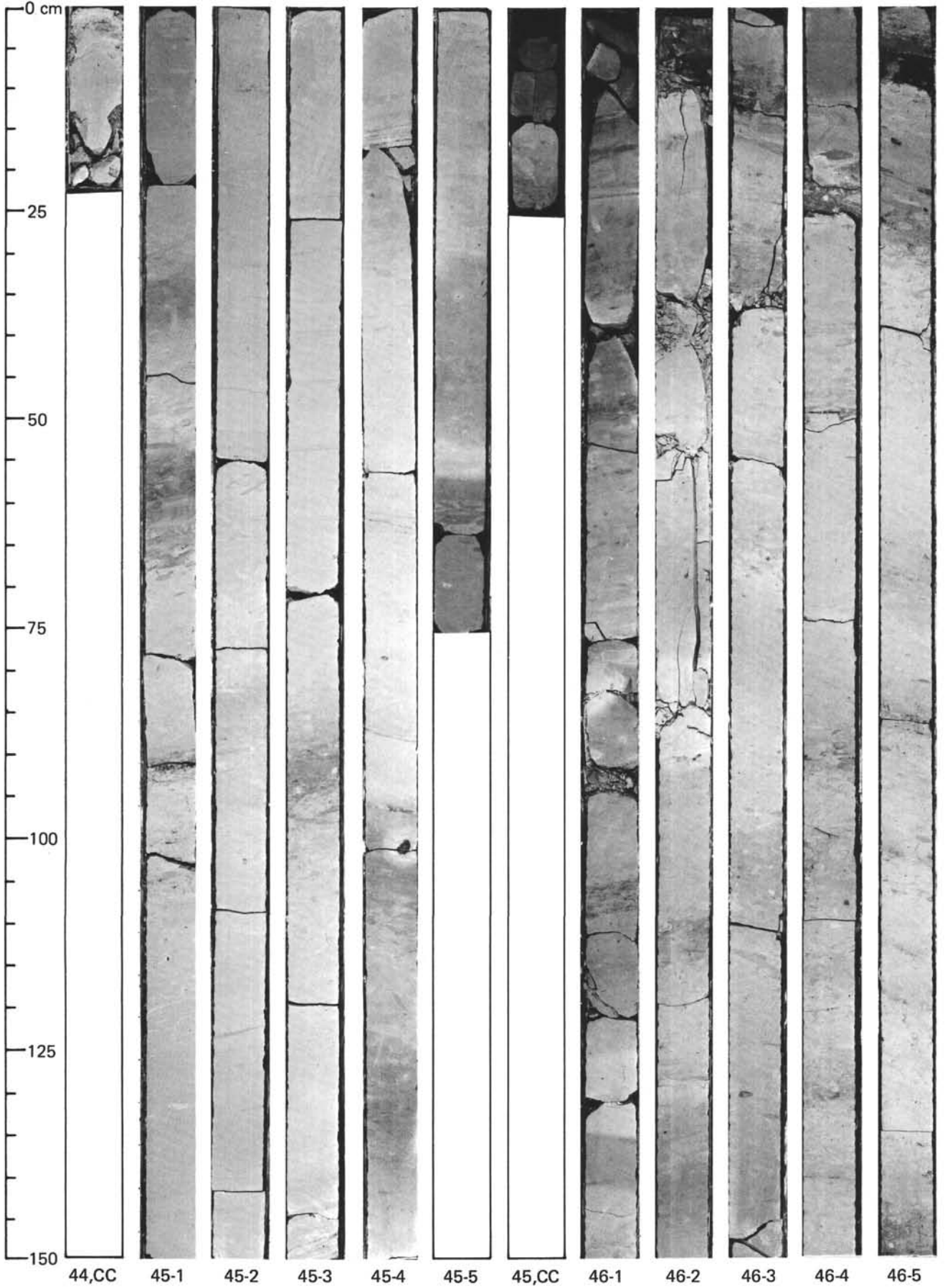
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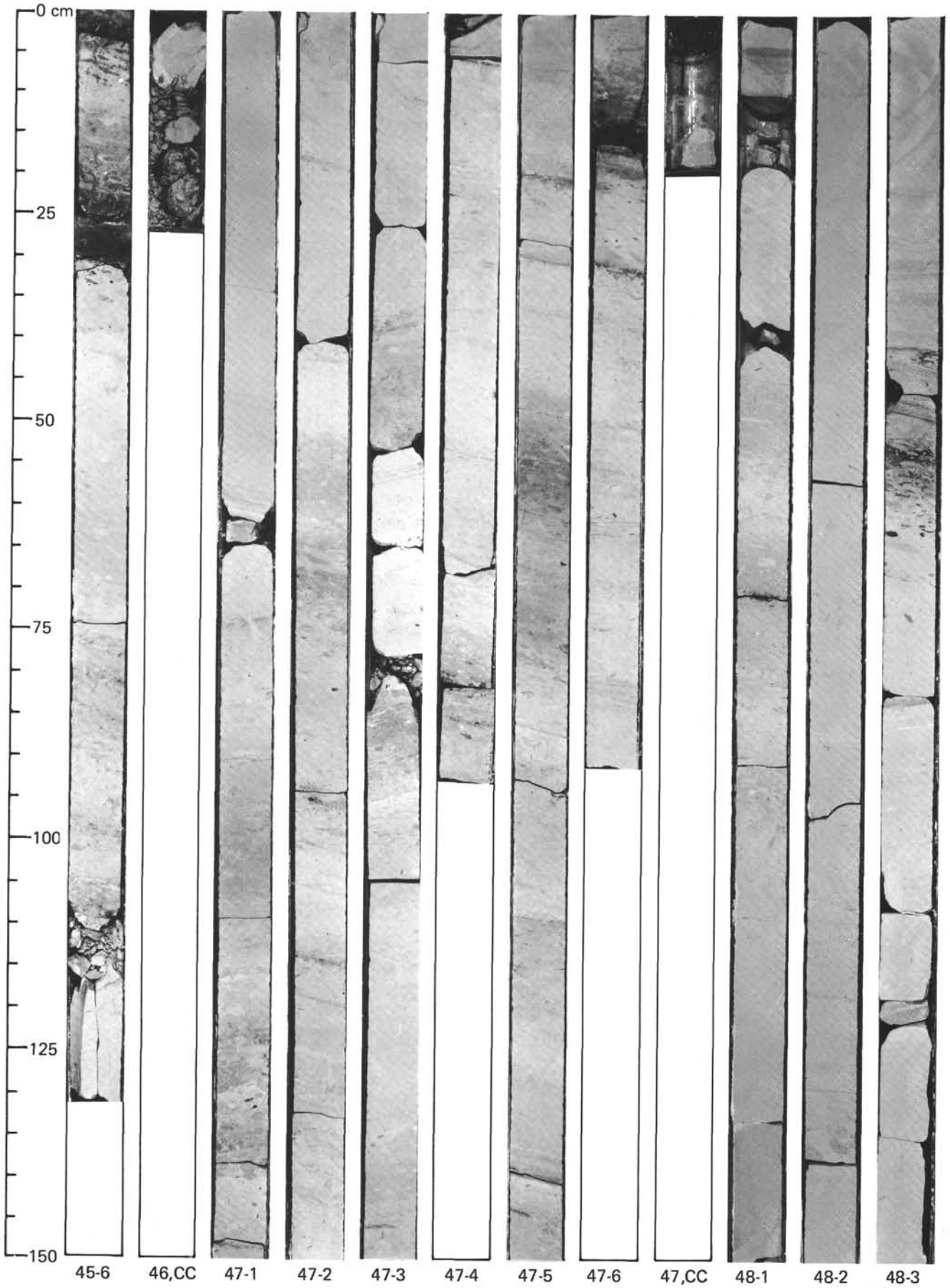


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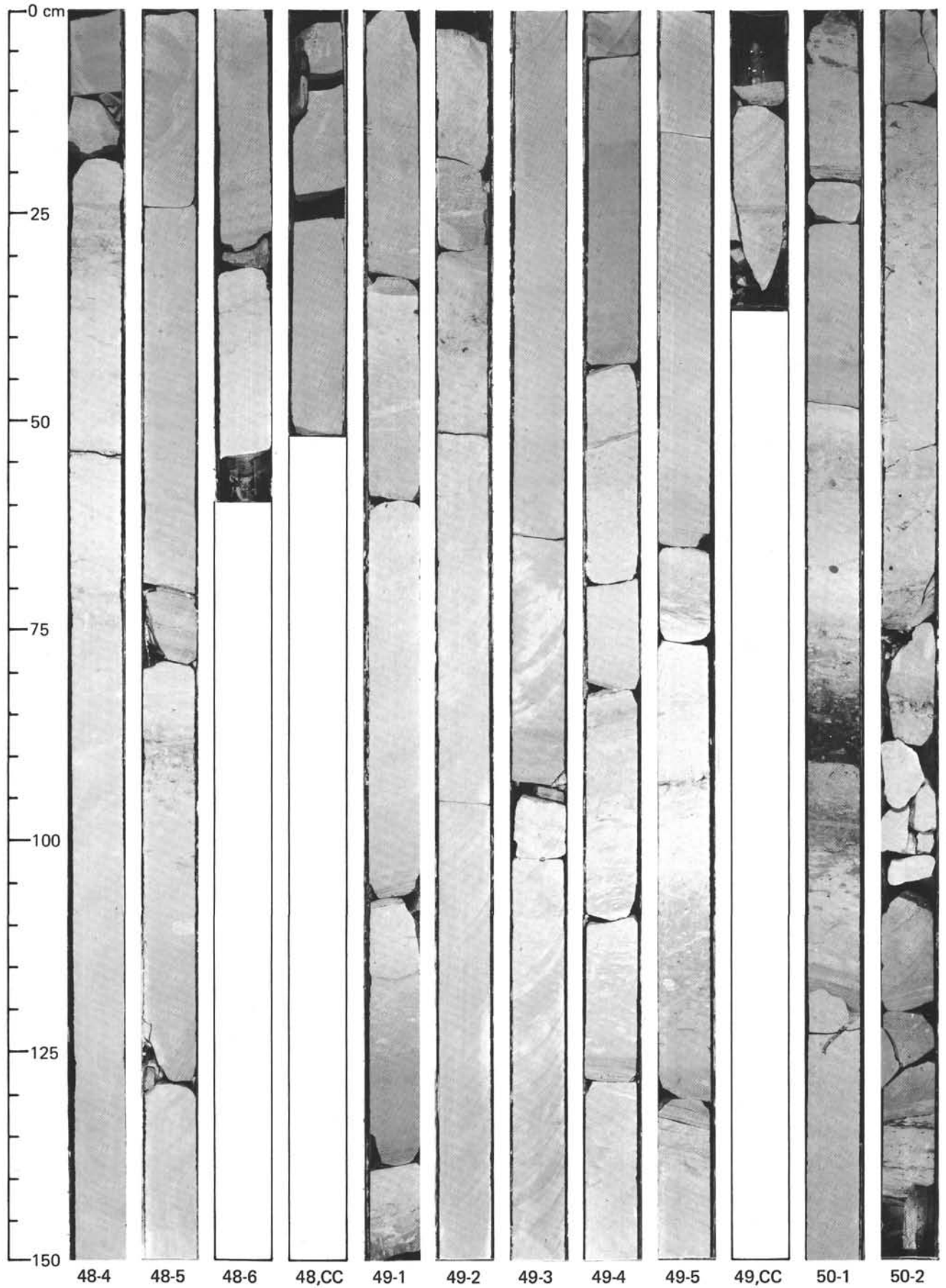




Site 445



Site 445



Site 445

0 cm

25

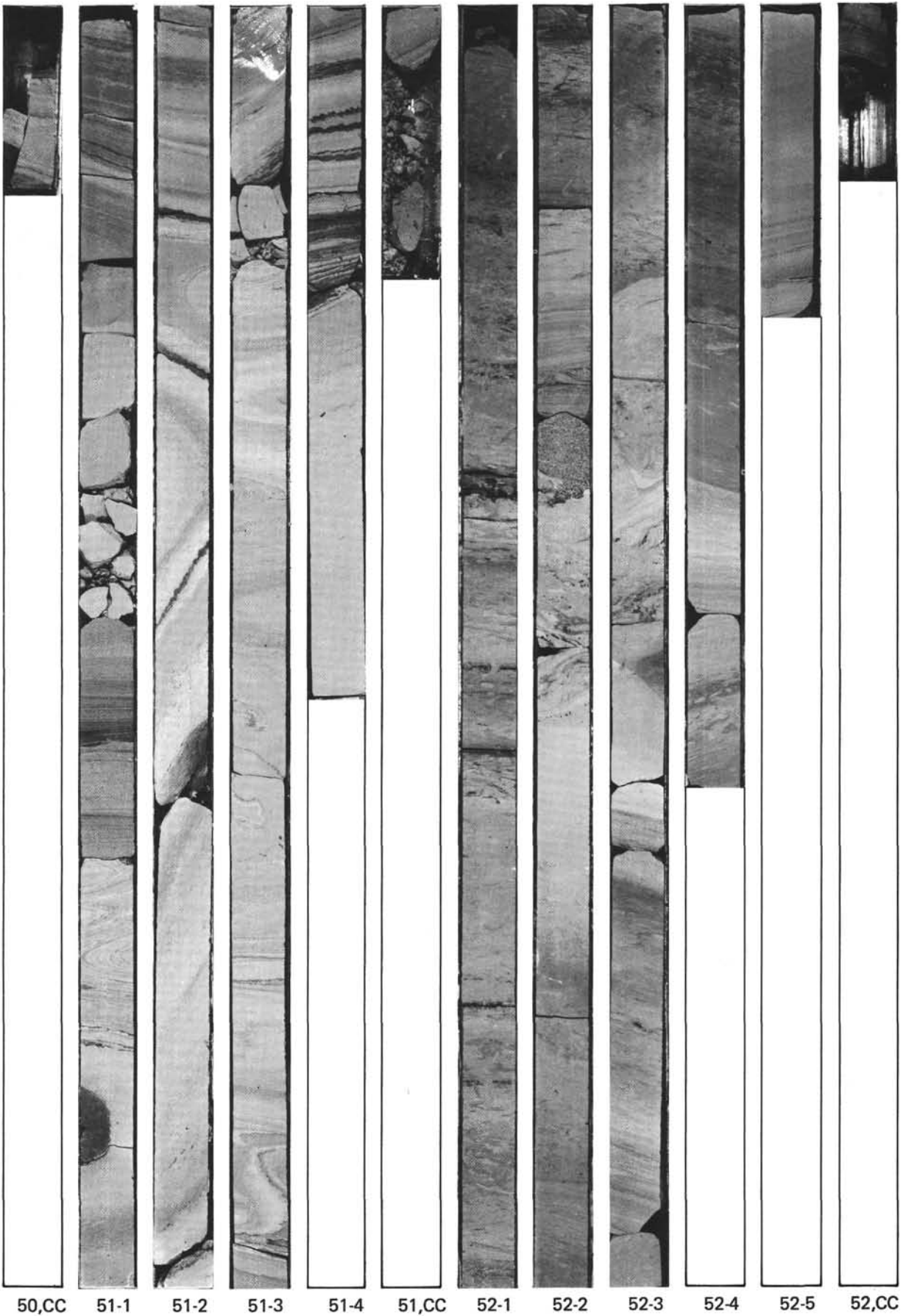
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75

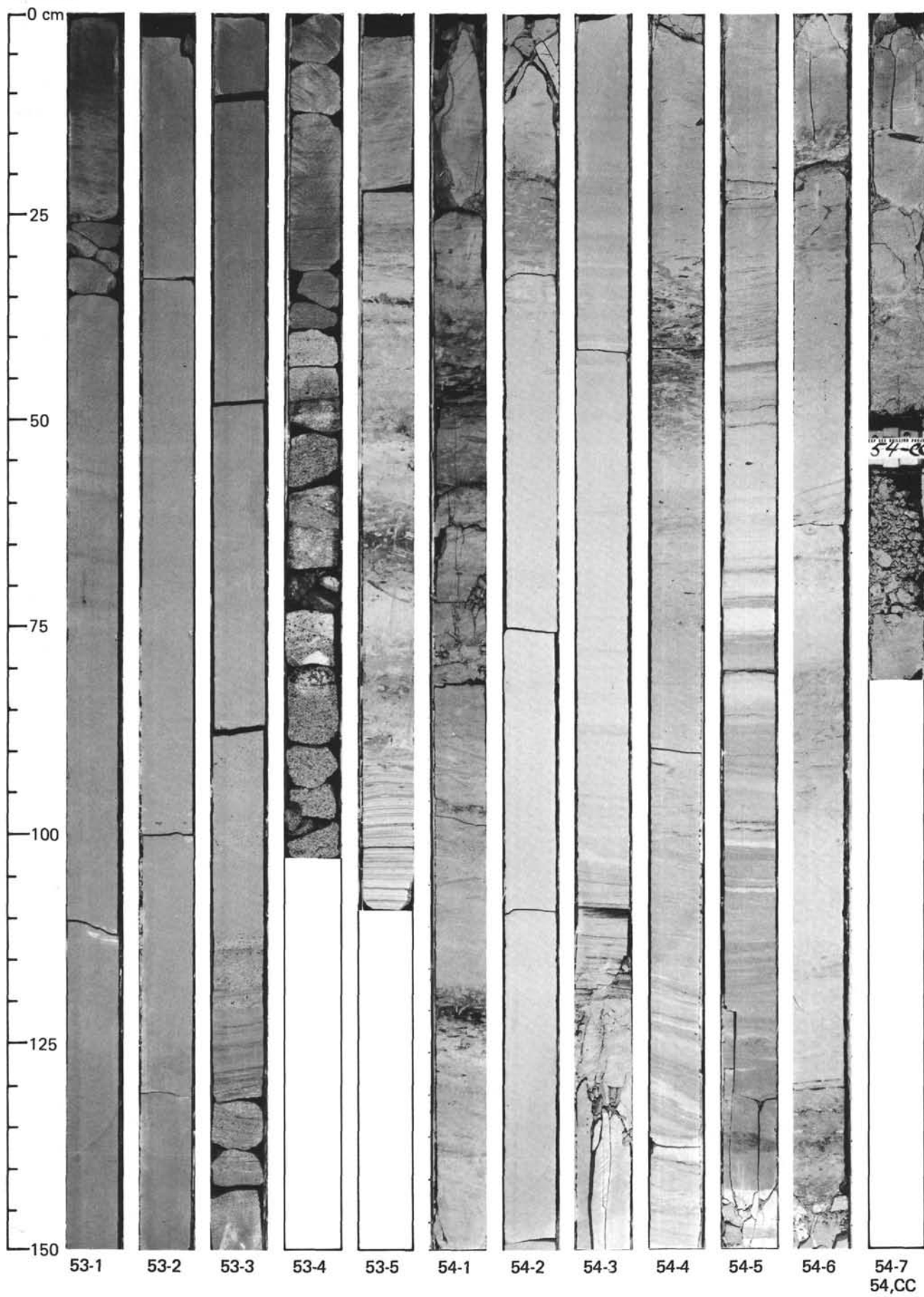
100

125

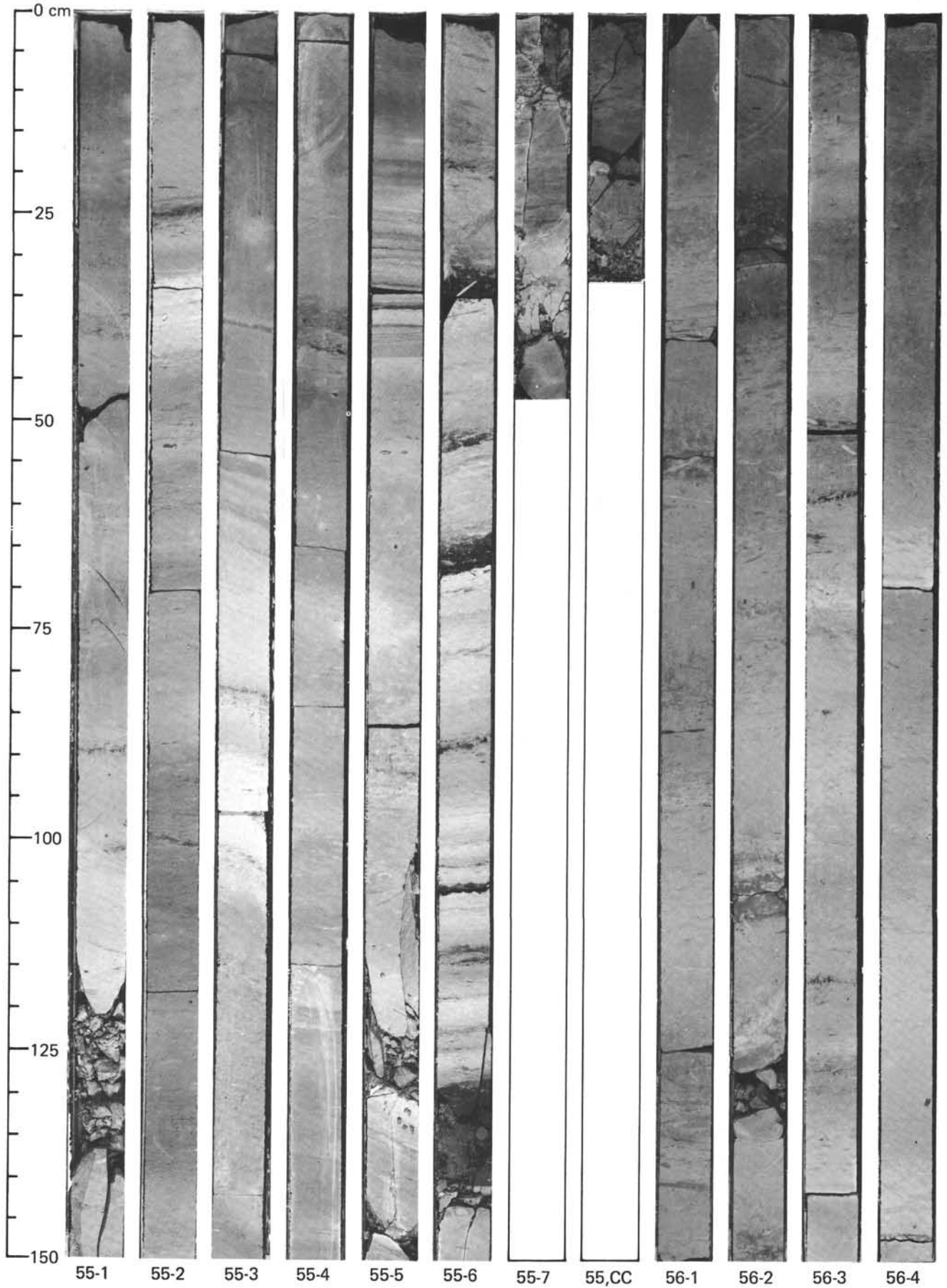
150



Site 445



Site 445





Site 445

0 cm

25

50

75

100

125

150



56-5



57-1



57-2



57-3



57-4



57-5



57-6

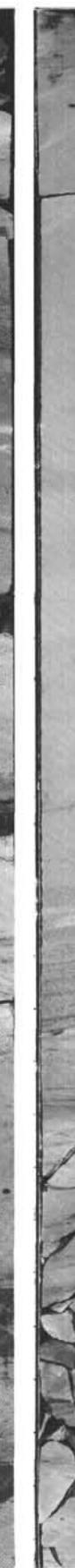


57-7

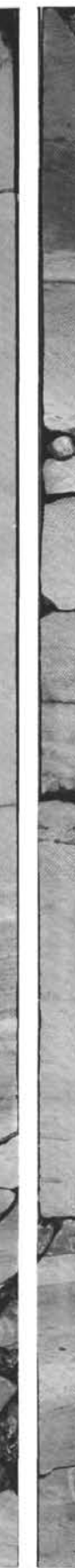
57,CC



58-1



58-2



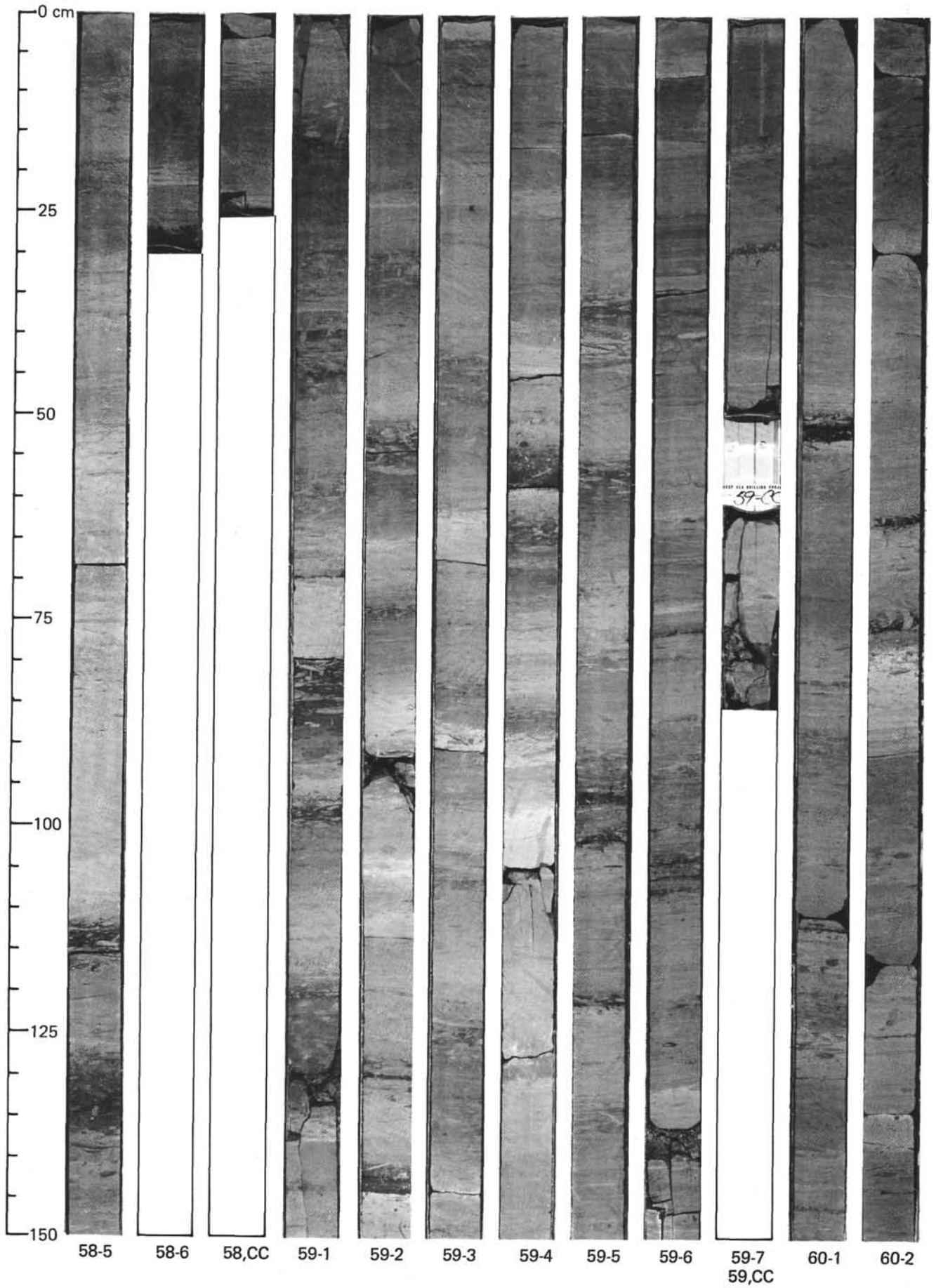
58-3



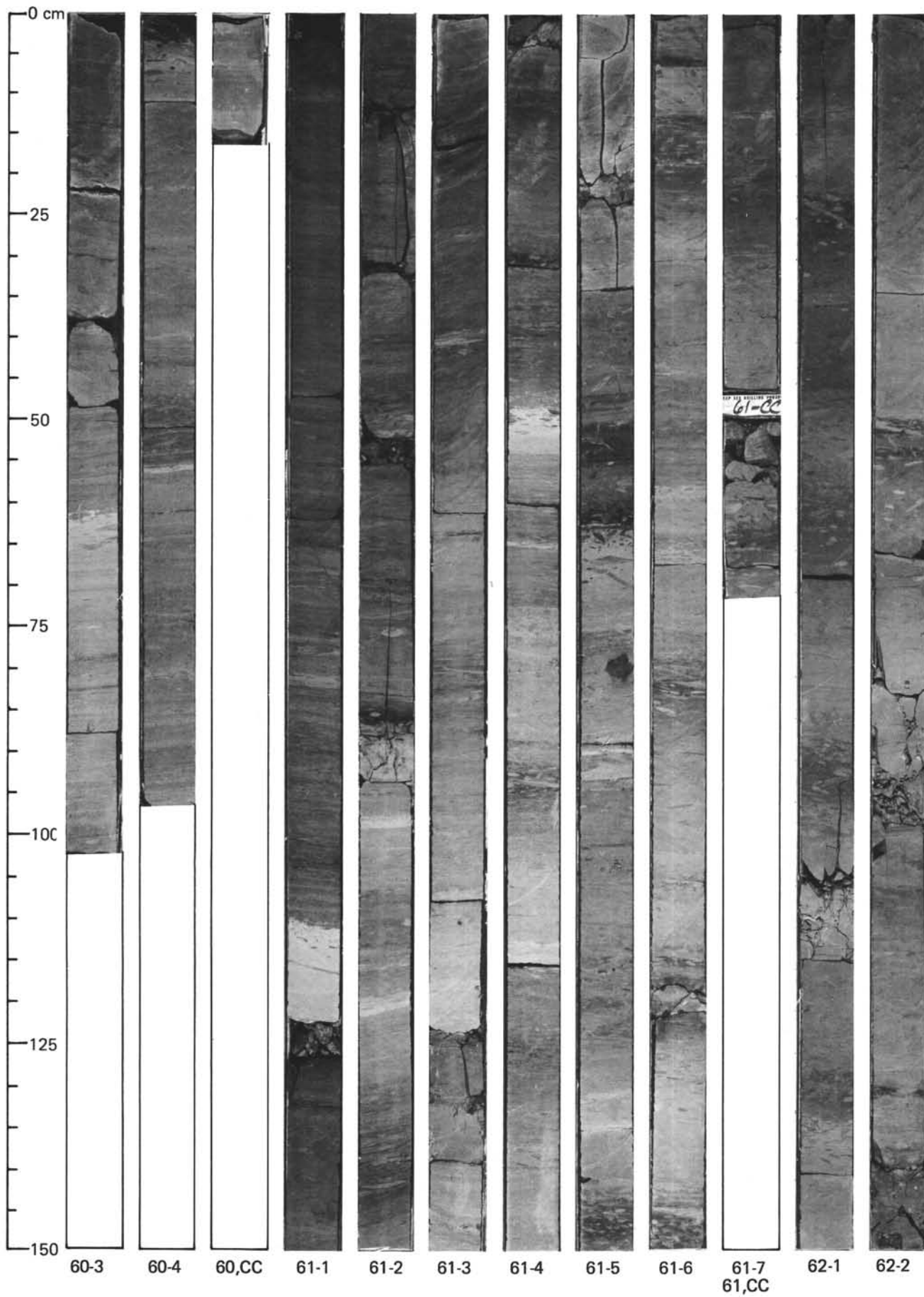
58-4



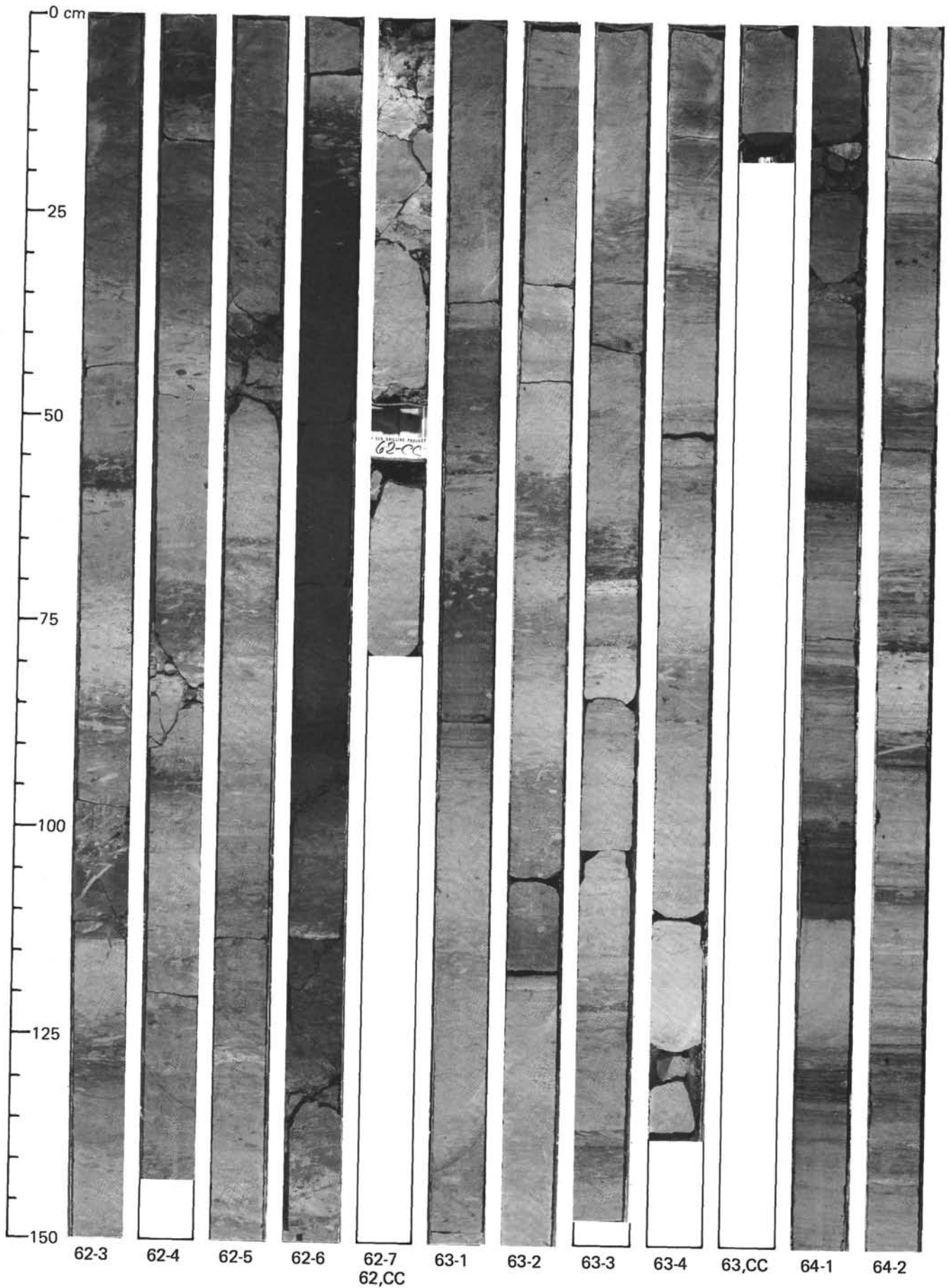
Site 445



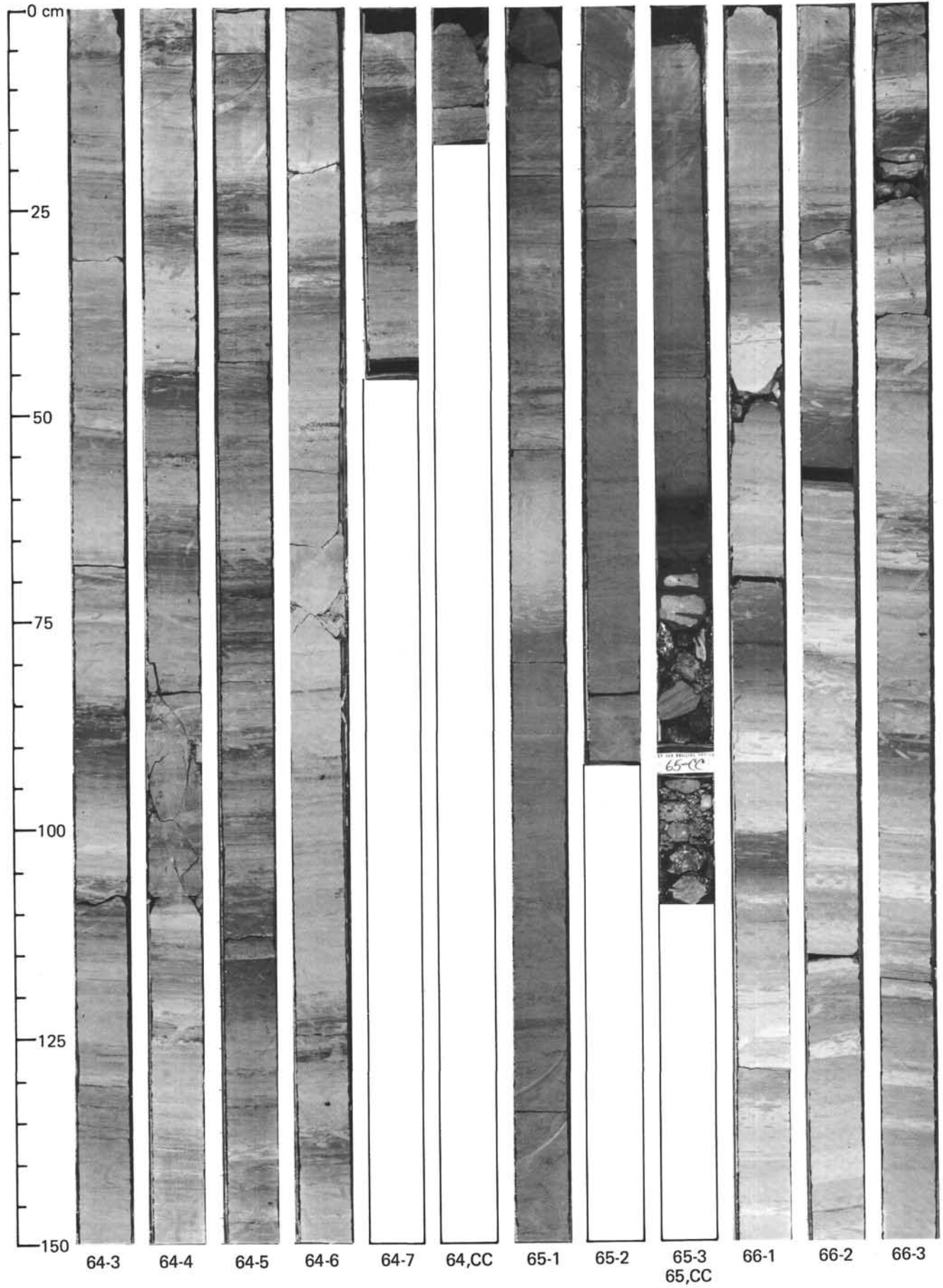
Site 445



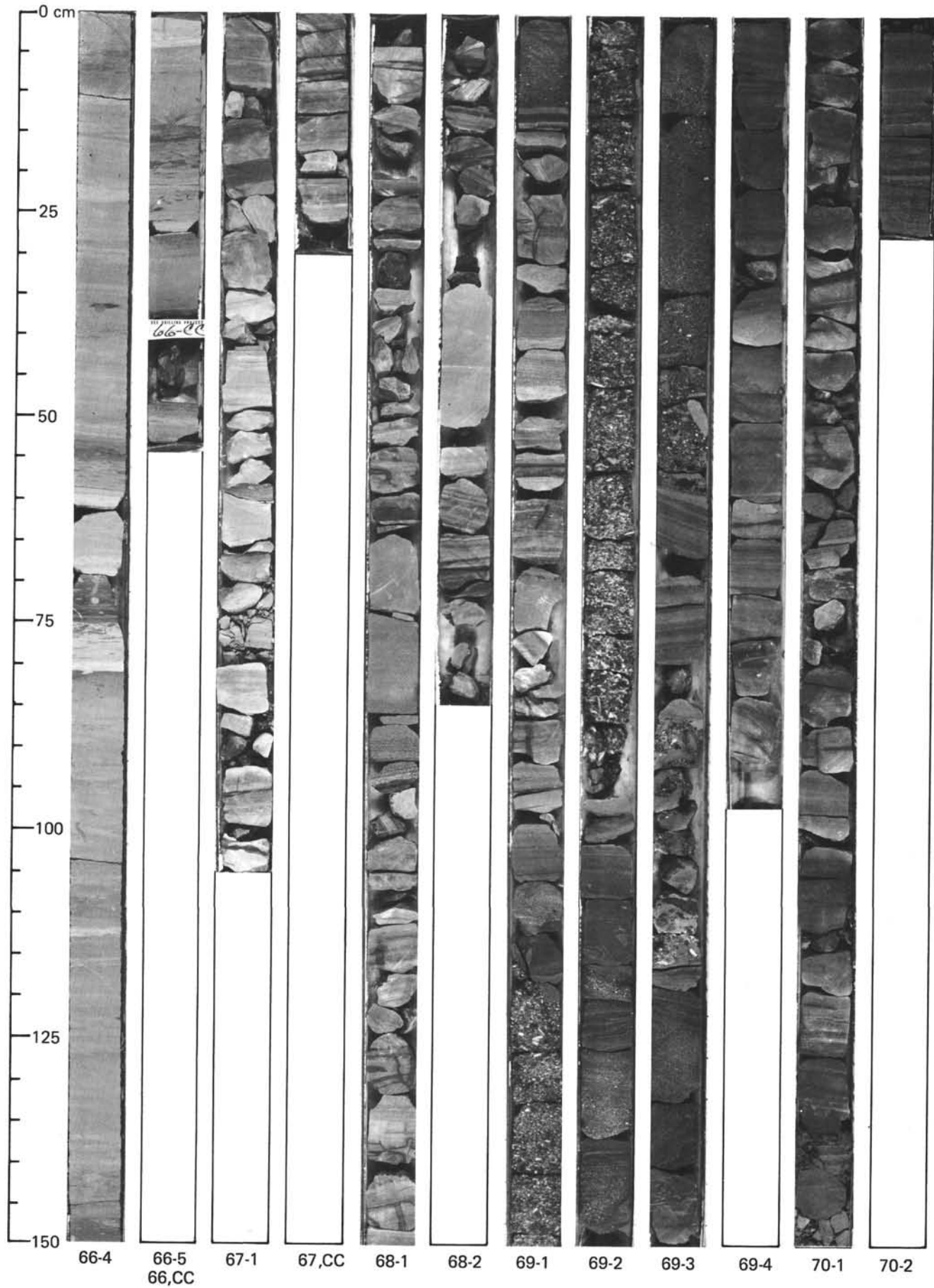
Site 445



Site 445

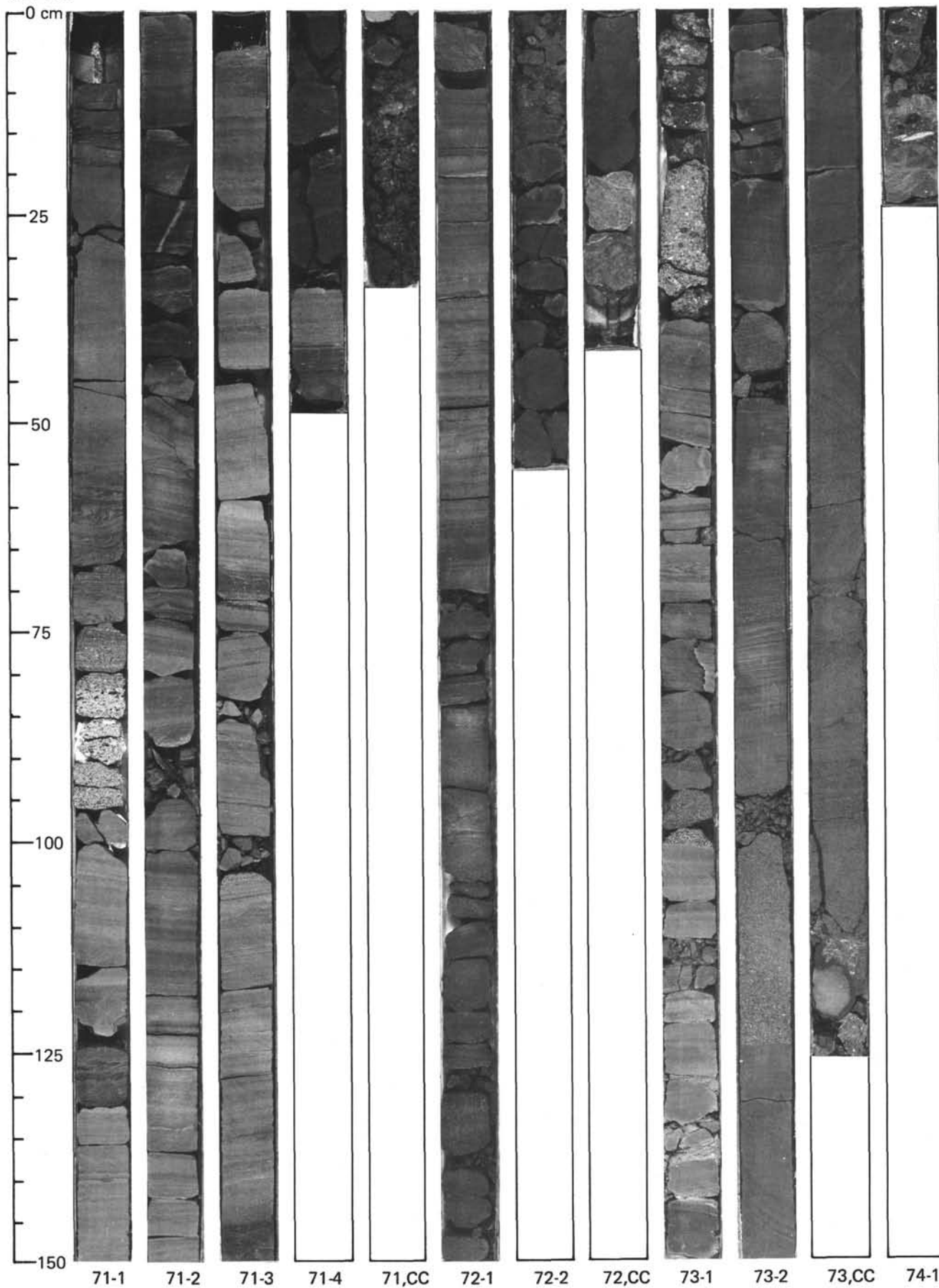


Site 445

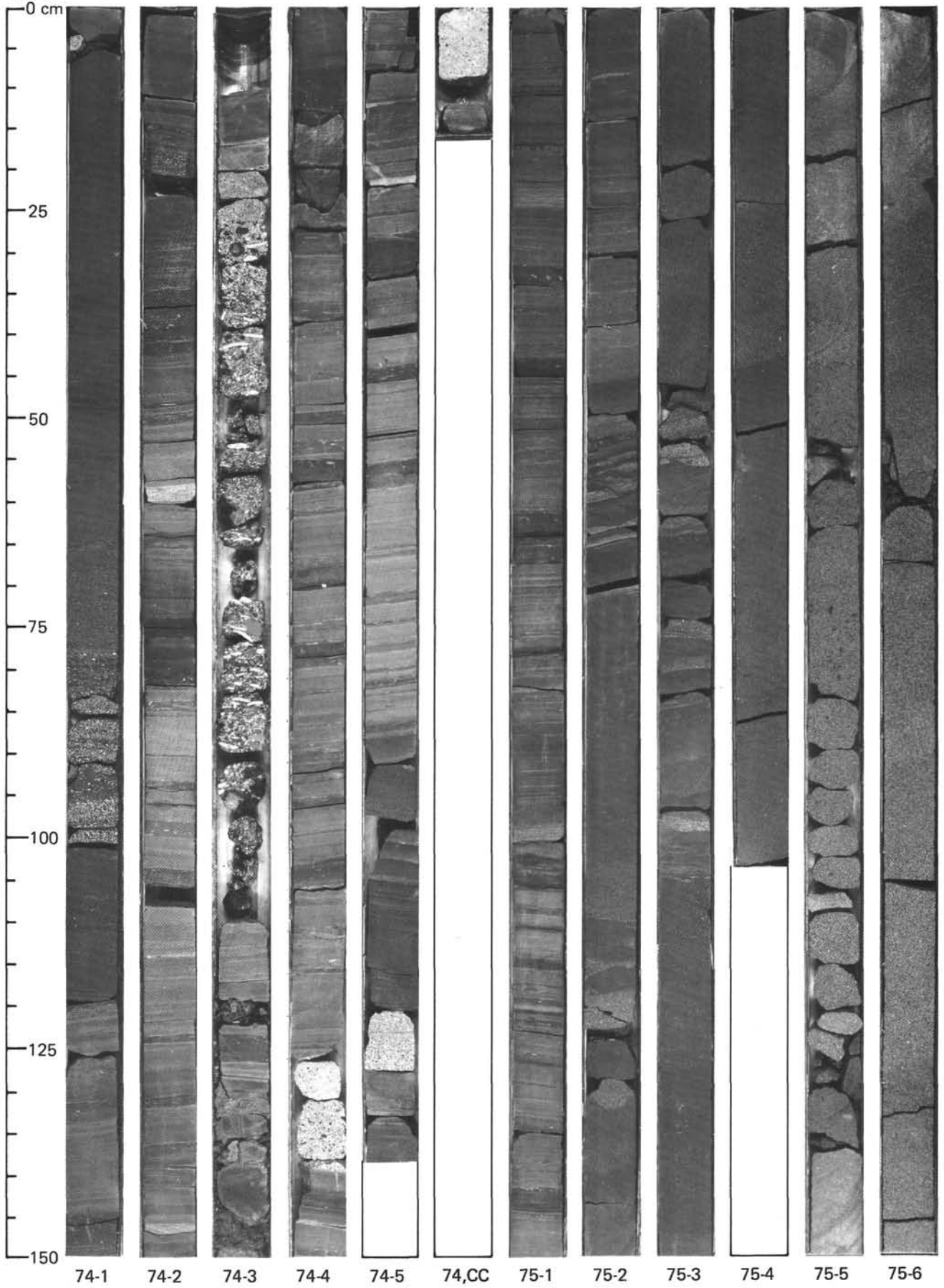




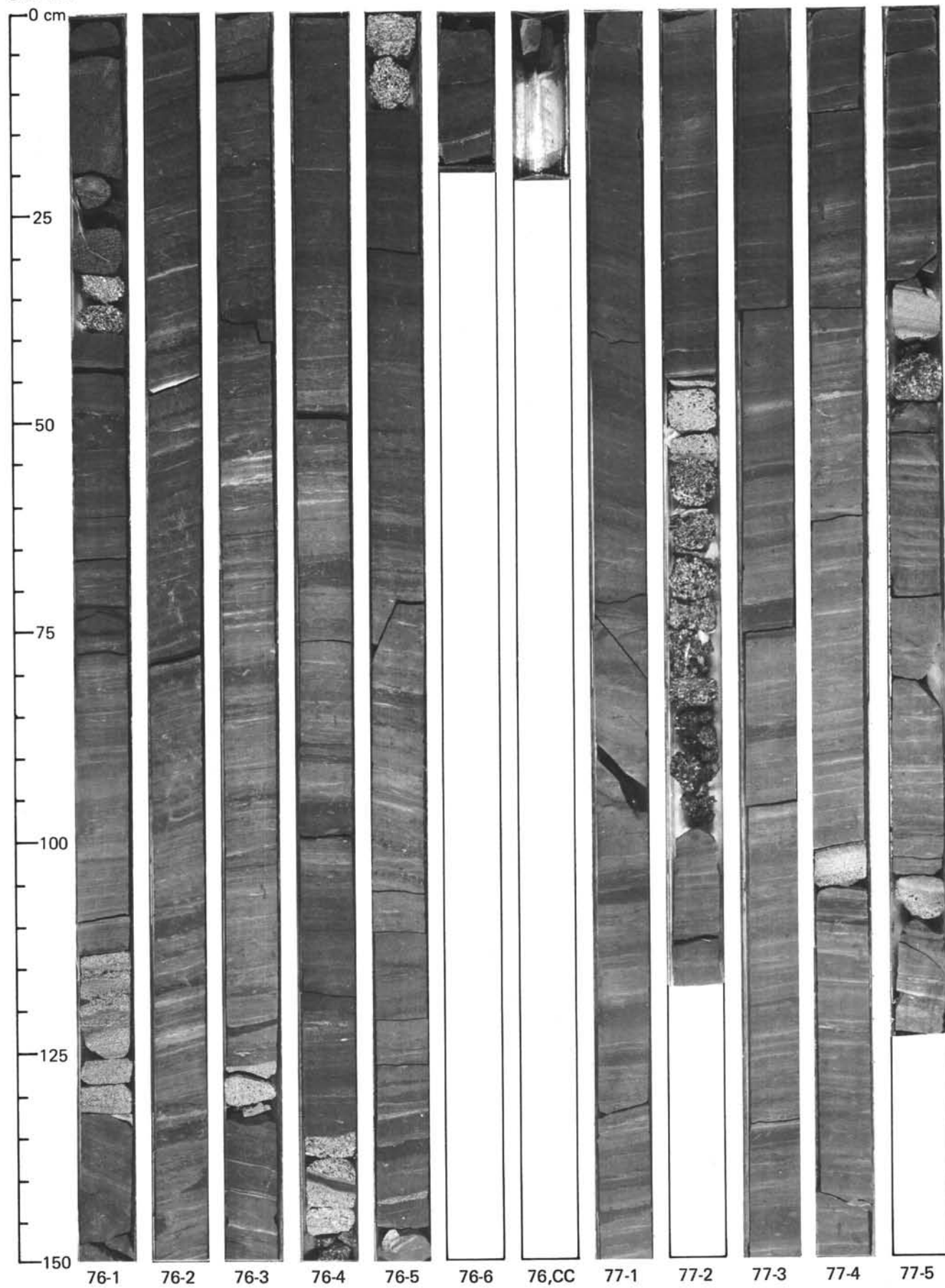
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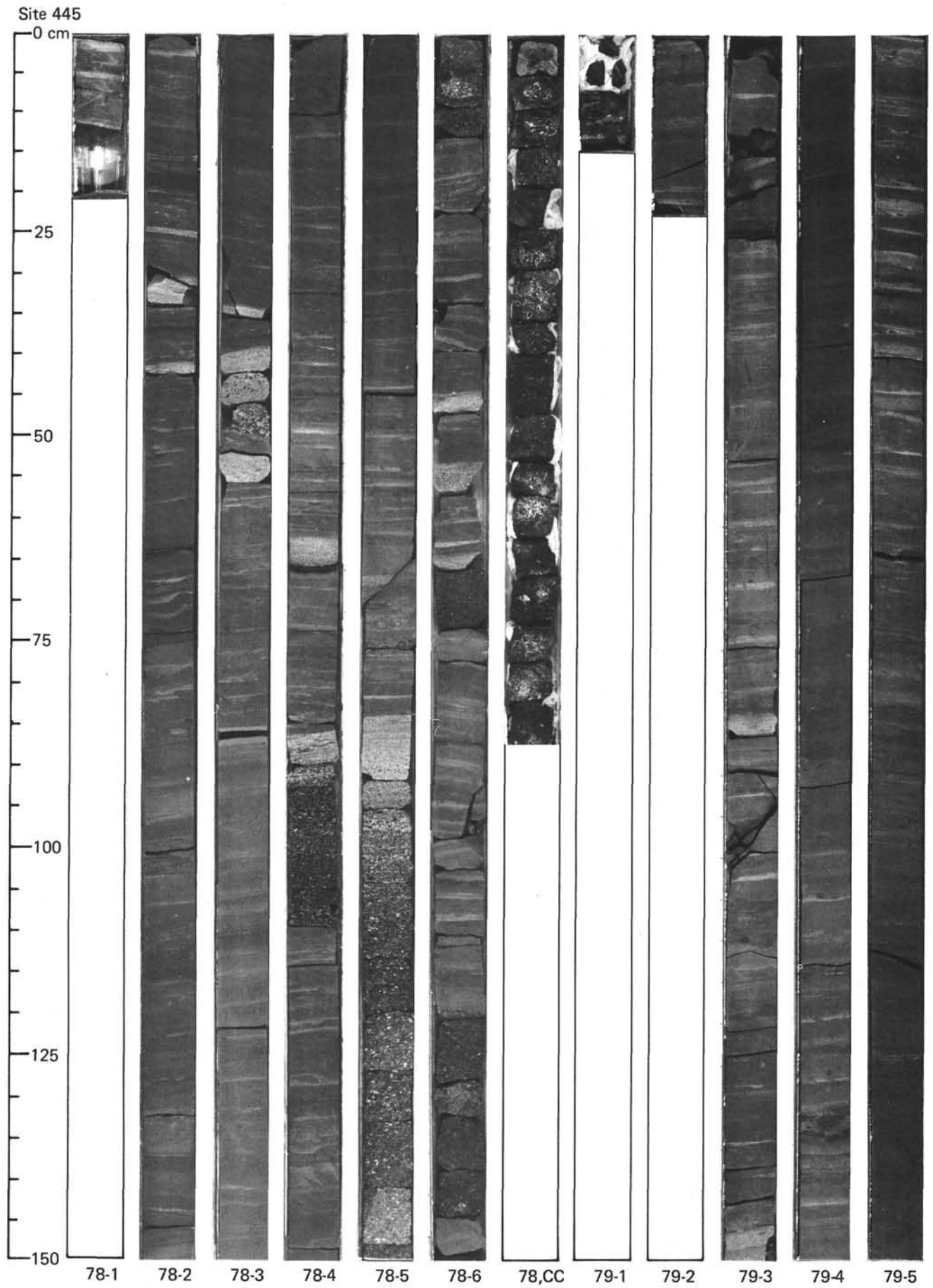


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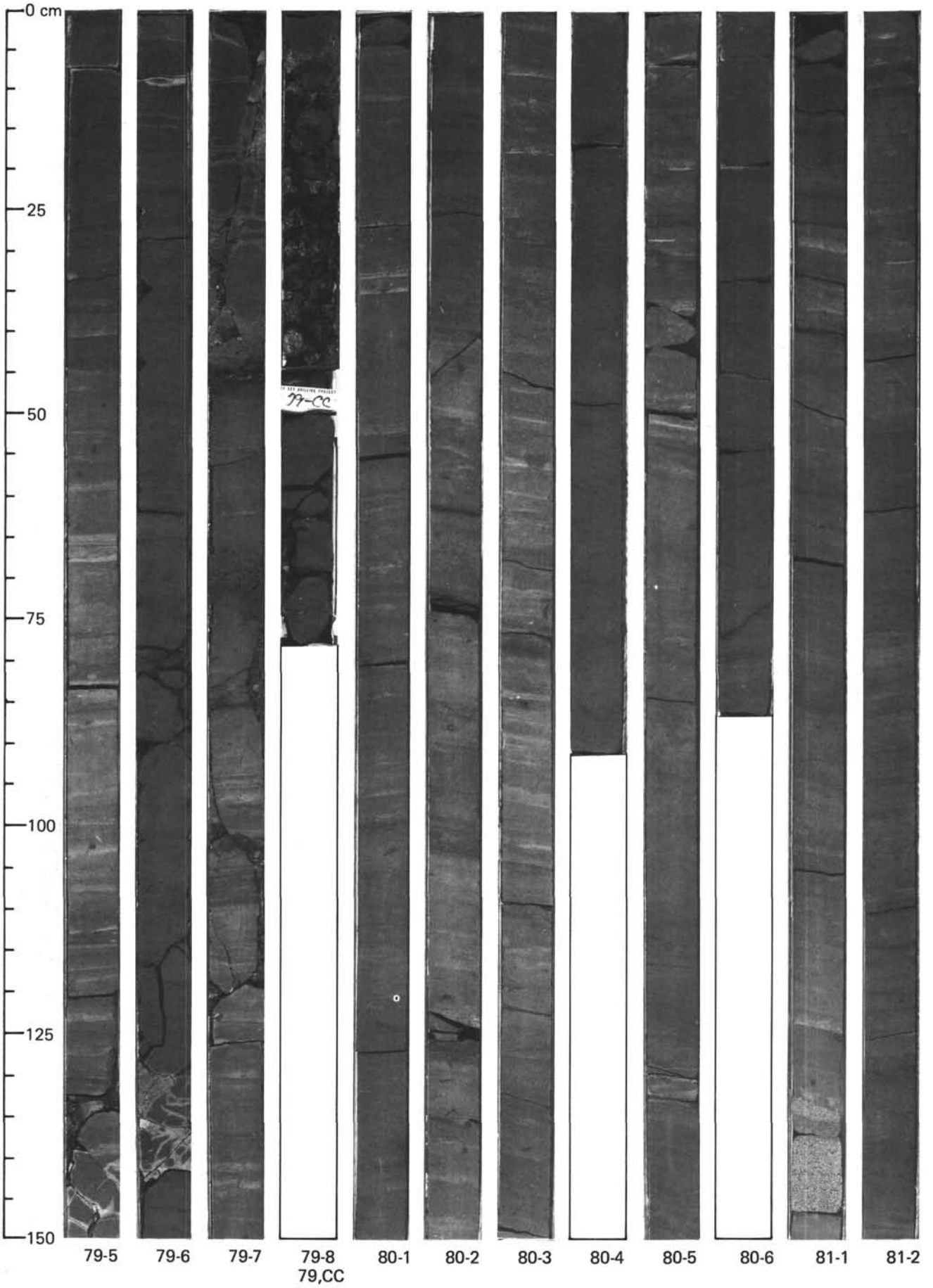


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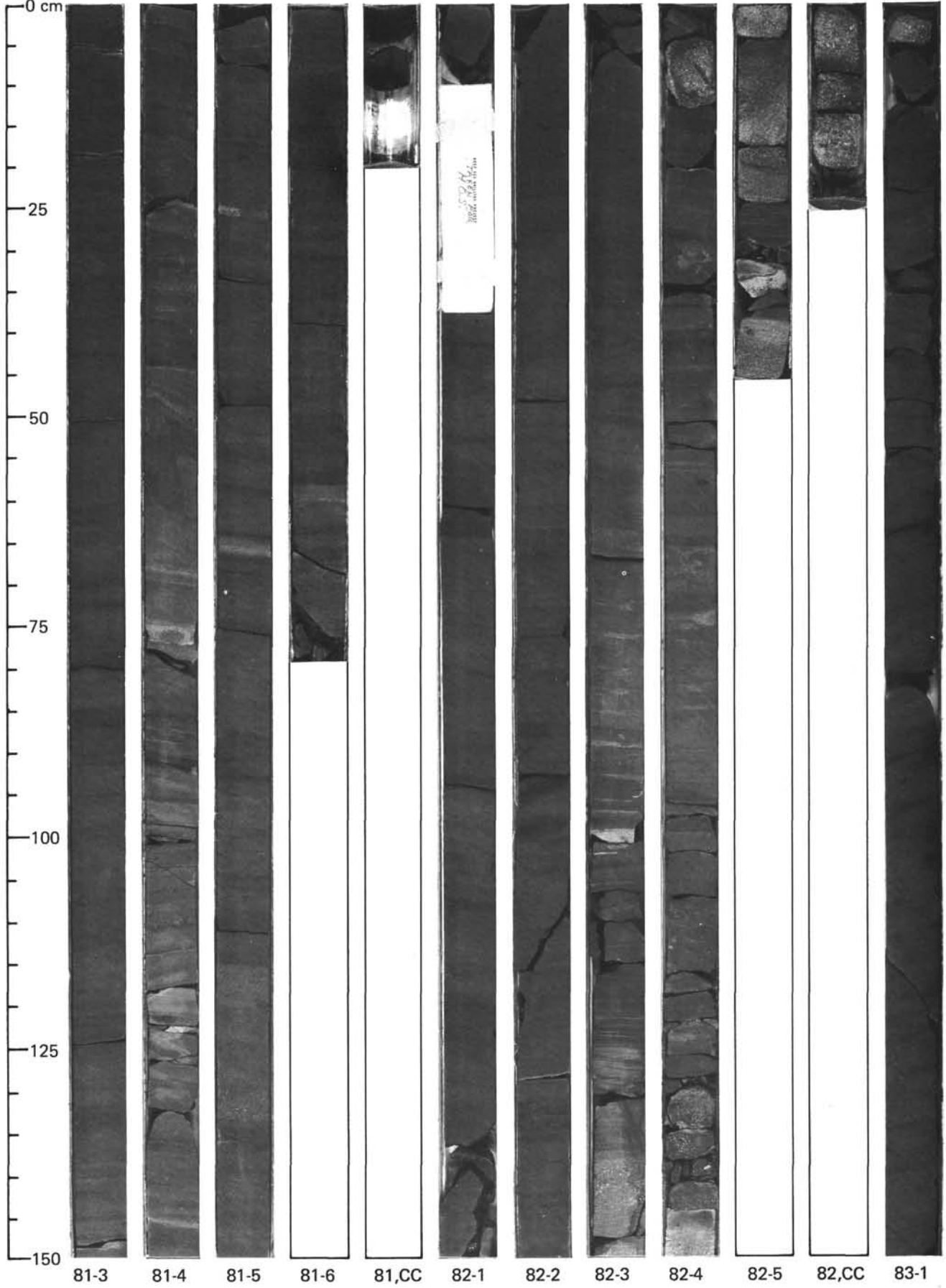


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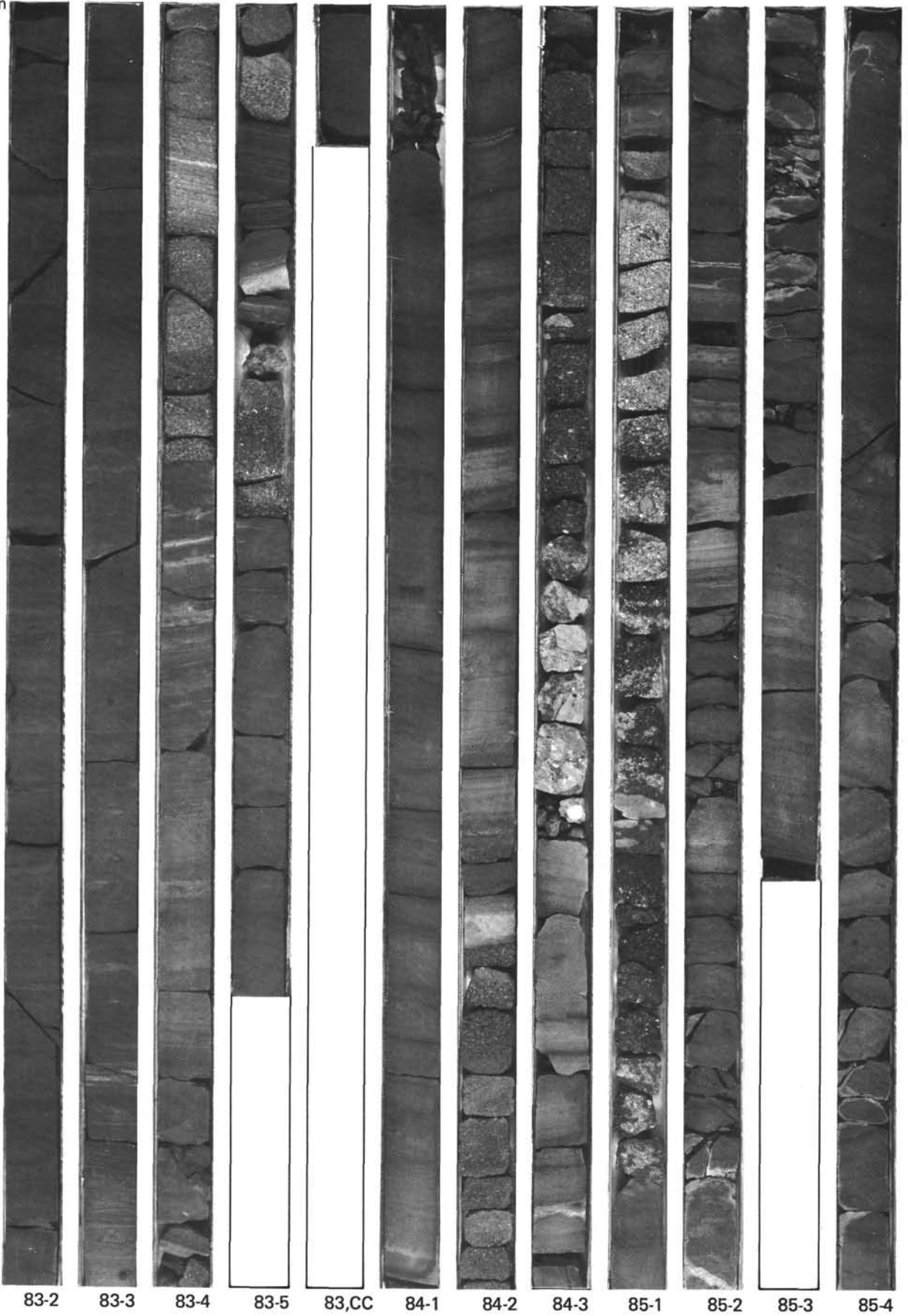




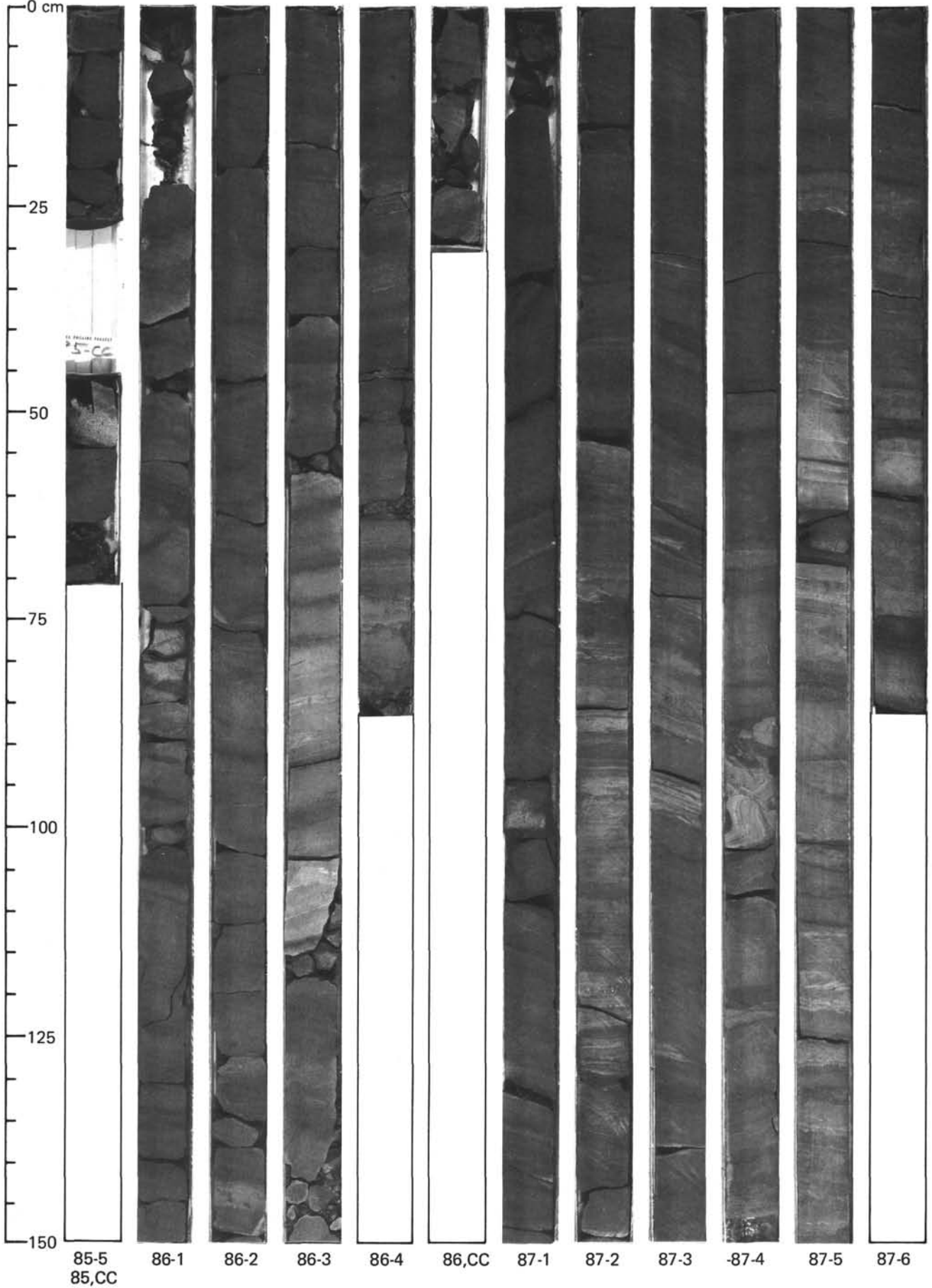
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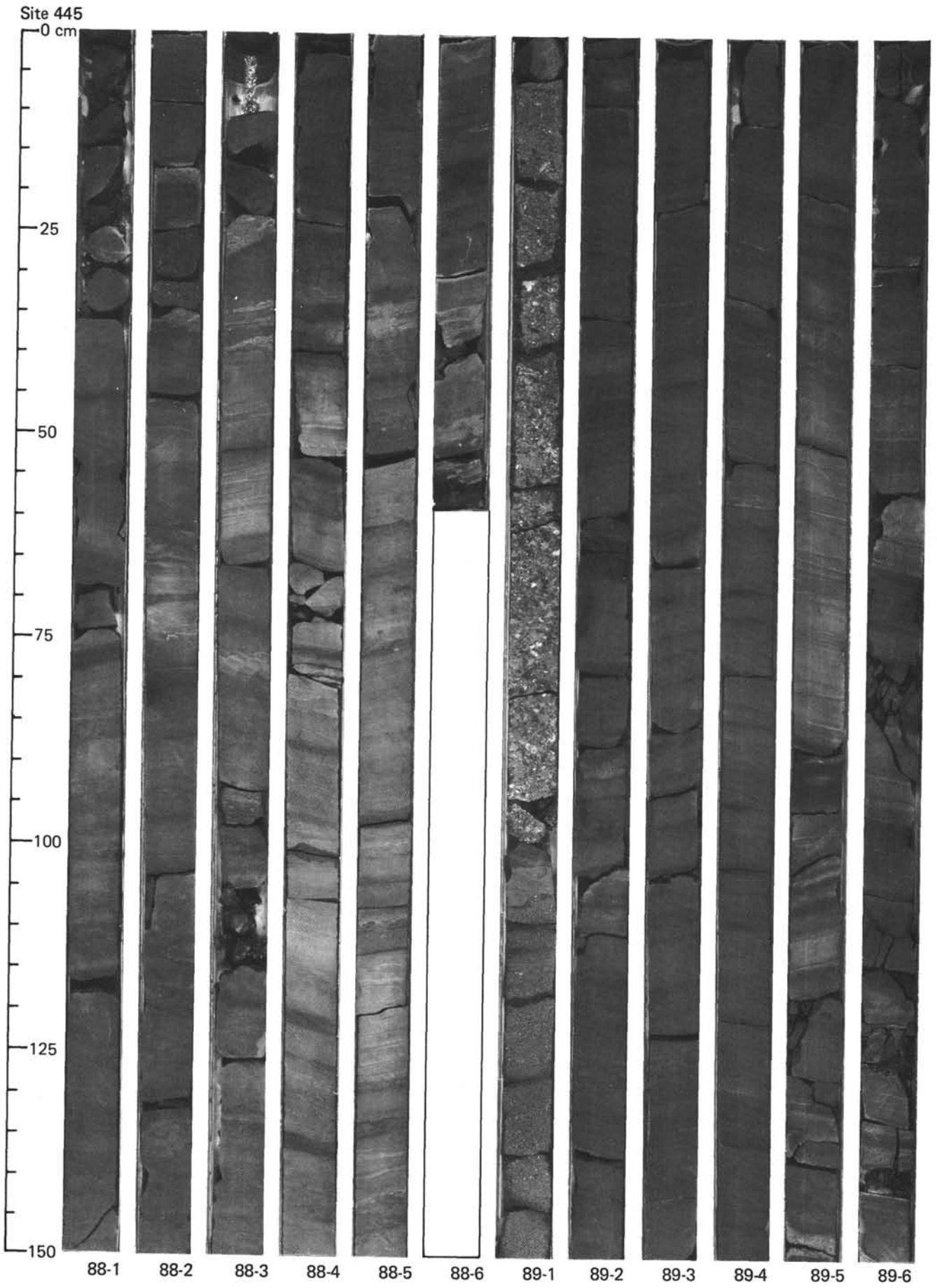


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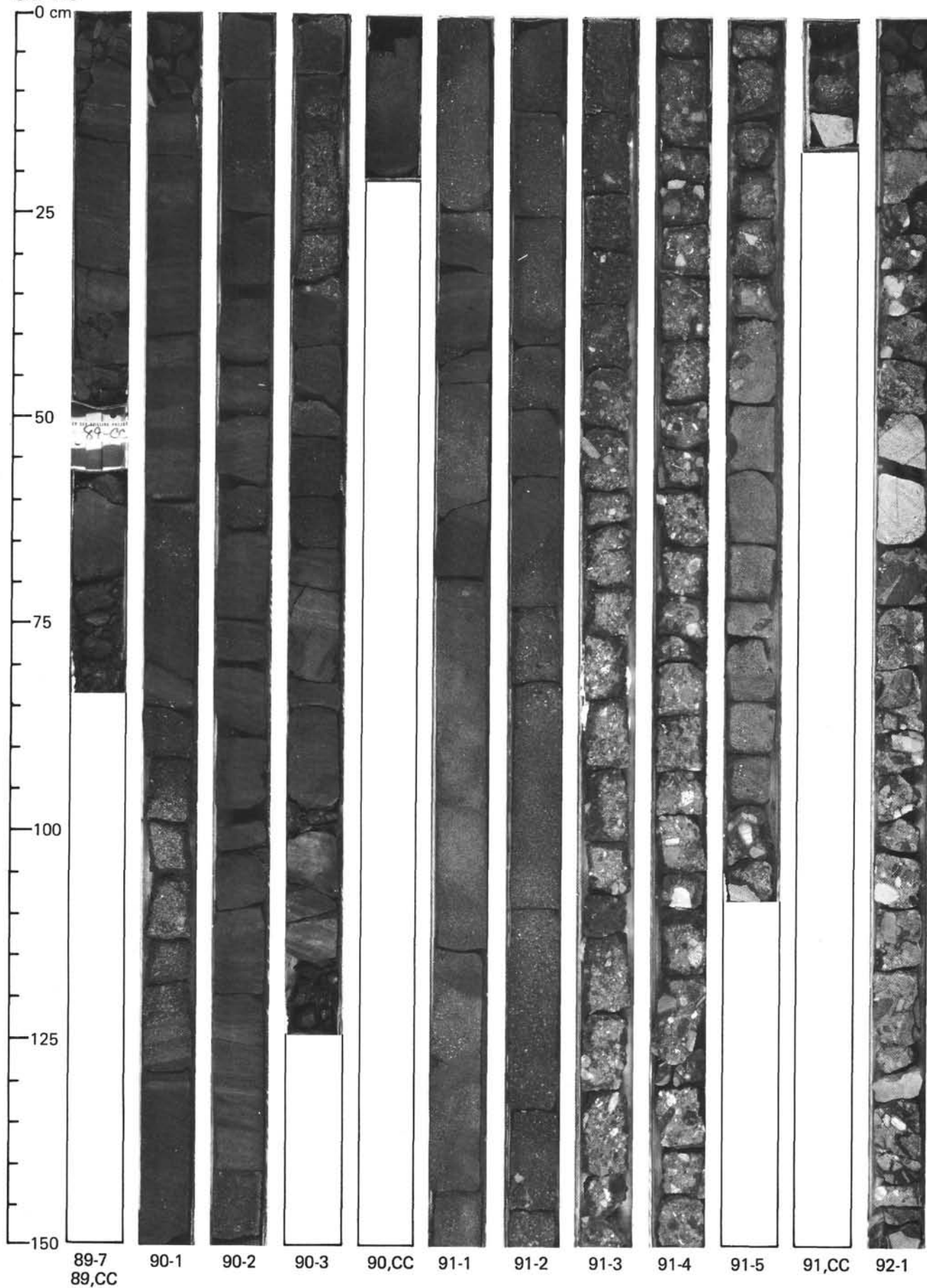


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