## 11. SITE $714^{1}$

## Shipboard Scientific Party ${ }^{2}$

## HOLE 714A

Date occupied: 0850 L, 23 June 1987
Date departed: 0215 L, 24 June 1987
Time on hole: $17 \mathrm{hr}, 25 \mathrm{~min}$
Position: $05^{\circ} 03.6^{\prime} \mathrm{N}, 73^{\circ} 47.2^{\prime} \mathrm{E}$
Water depth (sea level; corrected m, echo-sounding): 2038.3
Water depth (rig floor; corrected m, echo-sounding): 2048.8
Bottom felt (m, drill pipe): 2042.0
Penetration (m): 233.0
Number of cores: 25
Total length of cored section (m): 233.0
Total core recovered (m): 194.6
Core recovery (\%): 83.5
Oldest sediment cored:
Depth (mbsf): 233.0
Nature: foraminifer-bearing nannofossil chalk
Age: late Oligocene
Measured velocity (km/s): 1.779

## HOLE 714B

Date occupied: 0435 L, 24 June 1987
Date departed: 1115 L, 24 June 1987
Time on hole: $6 \mathrm{hr}, 40 \mathrm{~min}$
Position: $05^{\circ} 03.6^{\prime} \mathrm{N}, 73^{\circ} 47.2^{\prime} \mathrm{E}$
Water depth (sea level; corrected m, echo-sounding): 2038.3
Water depth (rig floor; corrected m, echo-sounding): 2048.8
Bottom felt (m, drill pipe): 2042.0
Penetration (m): 122.6
Number of cores: 13
Total length of cored section (m): 122.6
Total core recovered (m): 114.9
Core recovery (\%): 93.7
Oldest sediment cored:
Depth (mbsf): 122.6
Nature: foraminifer-bearing nannofossil ooze
Age: middle Miocene
Measured velocity (km/s): 1.529
Principal results: Site 714 is located in the northern equatorial Indian Ocean at $5^{\circ} 03.6^{\prime} \mathrm{N}$ and $73^{\circ} 47.2^{\prime} \mathrm{E}$ in water depths of 2031.5 m . The site lies on the eastern shoulder of the Maldives Ridge, which forms part of the aseismic ridge extending northward from the Chagos

[^0]Bank to the Laccadive Islands (Fig. 1). Seismic profiles across the Maldives Ridge area reveal massive accumulations from 1 to 1.5 km of sediments and sedimentary rocks on top of the presumed volcanic basement (see "Seismic Stratigraphy" section, this chapter). Our main purpose was to retrieve a complete late Neogene sequence of soft aragonite-bearing (periplatform) oozes.

Originally, we planned to drill the target for Site 714 at a water depth of about 1500 m on a narrow ridge separated from the main bank complex. The occurrence of strong reflectors virtually at the sediment/water interface, presumed (and later confirmed) to be submerged reef limestone, forced us to move the site to a greater water depth ( 2031.5 m ). This locality showed about 0.24 s (two-way traveltime) of transparent reflective layers in the upper part of the sediment sequence, and thus promised to yield a stratigraphic record of at least 200 m of suitable softer oozes. Our primary goal was to study the causal mechanism(s) for the observed connection between climatic changes and periplatform aragonite content.

We cored two holes continuously at Site 714. Hole 714A terminated in foraminifer-nannofossil chalk of late Oligocene age (about 28 Ma ) at 233.0 mbsf , whereas Hole 714B ended in foraminifer-bearing nannofossil ooze at 122.6 mbsf of middle Miocene age. The depositional history contains a single major hiatus, spanning the time interval from 0.5 to about 8.0 Ma (late Pleistocene through late Miocene). Otherwise, the sequence is characterized by continuous sediment accumulation. In fact, this sequence is both more expanded and more complete than any other lower upper Miocene through lower Miocene sequence recovered during Leg 115.

The excellent preservation of foraminifers and calcareous nannofossils in this Miocene section may thus contribute substantially to our understanding of their low-latitude intercorrelation during Miocene times. The middle Miocene also contains moderately well-preserved siliceous faunas and floras, which again emphasizes the potential importance of this tropical Indian Ocean sequence for Miocene biostratigraphy and paleoceanography. To some extent, this balances the disappointing fact that we retrieved only one tenth $(0-0.5 \mathrm{Ma})$ of the anticipated recovery of aragonite-bearing periplatform oozes.

We took 25 cores from Hole 714A, 13 of which were retrieved with the advanced hydraulic piston corer (APC) and the remaining 12 with the extended core barrel (XCB). The total recovery was $83.5 \%$ ( $97.3 \% \mathrm{APC}$ and $69.3 \%$ XCB). From Hole 714B, we retrieved 13 APC cores, with a recovery of $93.7 \%$.

The sequence is comprised of two lithostratigraphic units (see "Lithostratigraphy" section, this chapter), one of which is divided into two subunits.

Unit I ( $0-19.55 \mathrm{mbsf}$ ) consists of late Pleistocene ( $0-0.5 \mathrm{Ma}$ ), dark greenish gray, clay-bearing, foraminifer-bearing nannofossil ooze, grading downcore to a foraminifer-bearing, clayey nannofossil ooze. The average carbonate content is $76 \%$, but the extremes vary from $60 \%$ to $84 \%$.

Unit II (19.55-233.0 mbsf) is divided into two subunits. Subunit IIA (19.55-120.0 mbsf) consists of very light greenish gray, lower upper Miocene through middle Miocene (8.2-15 Ma), clay-bearing, foraminifer-bearing nannofossil ooze. The carbonate content shows a generally progressive increase from lows of about $81 \%-82 \%$ at 9 Ma to peak values of about $89 \%$ at 15 Ma , as estimated from mean values in 1-m.y, time increments. The carbonate values range from $61 \%$ to $95 \%$.

Subunit IIB ( $120.0-233.0 \mathrm{mbsf}$ ) is similar in composition to Subunit IIA, but they are distinguished on the basis of the degree of lithification. The ooze of the latter grades into the clay-bearing, fora-minifer-nannofossil chalk characterizing Subunit IIB. Burrow mottles (moderate) occur throughout Unit II. The carbonate content of Subunit IIB displays more subtle variations than those observed in


Figure 1. Bathymetric map of the central Indian Ocean showing the location of Site 714.

Subunit IIA, and we discerned no obvious trends. Apart from a single extreme value of $61 \%$, all carbonate contents are in the range between $77 \%$ and $93 \%$, with an average of $87 \%$. A 10 -point running mean, however, suggests that the carbonate content changes throughout Unit II as if influenced by a quasicyclical, low-frequency component ( $\sim 0.5 \mathrm{~m} . \mathrm{y} . ?$ ). Figure 2 summarizes the cored stratigraphic sequence.

## Summary of Interpretation

Considering the well-preserved microfossil assemblages, the stratigraphic continuity, and the reasonably high sedimentation rates (10$20 \mathrm{~m} / \mathrm{m} . \mathrm{y}$.) throughout the Miocene sequence of Site 714 , it is clearly disappointing that the intensities of natural remanent magnetization (NRM) are too low to yield a magnetostratigraphic record. We have to rely, therefore, on the calcareous nannofossil correlation to magnetostratigraphy from Site 710 and apply those results to the material recovered from Site 714. This approach, together with the use of well-calibrated planktonic foraminifer datum events from Atlantic and Pacific low-latitude regions (Site 710 was located below the foraminifer lysocline during Miocene times), should, however, provide the necessary means for the reconstruction of a highly resolved chronology at Site 714.

Preliminary attempts to calculate mass accumulation rates of bulk and carbonate sediment are shown in Figure 3. The Pleistocene bulk accumulation rate is approximately twice that of the average of the pre-Pleistocene rates. Nevertheless, the Miocene rates are consistently higher than $1 \mathrm{~g} / \mathrm{cm}^{2} / 1000 \mathrm{yr}$ between 8.2 and 19 Ma , and slightly below that value throughout most of the earliest Miocene and late Oligocene. The accumulation rate of the noncarbonate fraction by and large mimics the changes in bulk and carbonate accumulation,
with the highest Miocene values around the middle/late and early/ middle Miocene boundaries and slightly lower rates during the middle and early Miocene.

At a water depth of 2031.5 m , we assume that carbonate dissolution has had a minor effect on the carbonate accumulation. Moreover, taking into account the consistently high bulk sediment percentage of nannofossils (see "Lithostratigraphy" section, this chapter), we may assume that winnowing has had only a marginal effect on the presented mass accumulation rates. Provided the age-depth model (see "Sedimentation Rates" section, this chapter) for Site 714 is accurate, it follows that the mass accumulation rates illustrated in Figure 3 should reflect the true time-dependent variability in input of carbonate. Clues as to the cause(s) for this variability will emerge as the result of shore-based analyses.

## BACKGROUND AND OBJECTIVES

Hard parts of foraminifers and coccolithophorids are formed of calcite, which represents the chief source material for deepsea carbonates. Aragonite also contributes to the formation of oceanic carbonate sediments, mainly because the pelagic pteropods construct tests of this mineral. Aragonite is less stable than calcite, implying that the former mineral dissolves more easily than the latter. This results in an aragonite-compensation depth which is considerably shallower than that of calcite. Accordingly, the sedimentary lysocline is also shallower for aragonite than for calcite.

Periplatform oozes represent a third source of oceanic carbonate sediment formed within a unique environmental setting.


Figure 2. Stratigraphic summary, Site 714. Black column represents recovered section.

Such oozes are composed of a mixture of the common pelagic input (foraminifers, coccolithophorids, and pteropods) and banktop-precipitated aragonite, and are exclusively deposited in the close vicinity of shallow carbonate banks such as the Bahamas. Aragonite produced on the banks, mostly by benthic algae or by inorganic precipitation, is partly swept away and exported to deeper periplatform areas, where the aragonite becomes preserved if the sites of redeposition are located above the arago-nite-compensation depth.

Analyses of the mineralogic composition of periplatform oozes from the Bahama Bank region have established that the aragonite content of the sediments coincides with the late Pleistocene glacial/interglacial cycles (Kier and Pilkey, 1971; Droxler et al., 1983; Boardman et al., 1986). Increased relative abundance of aragonite is linked with interglacial stages and vice versa. The causal mechanism for this connection between climate and periplatform aragonite content, however, has been the subject of debate (e.g., Boardman and Neumann, 1986; Droxler, 1986). One line of reasoning favors the argument that bankderived sediment is deposited during times of high sea-level stands (e.g., interglacial times) when banktops are flooded and export of banktop-produced aragonite occurs. In this case, the observed variability in aragonite content essentially reflects the primary input signal.

The opposing model argues that, to a major extent, the phase relationship between aragonite content and climatic cycles reflects differences in preservational states of aragonite during glacial and interglacial times. This argument is based on the observation that time-dependent changes in the degree of aragonite preservation are directly linked to changes in ocean chemistry, because the saturation values of different minerals and relative positions of lysoclines or compensation depths are, in turn,
coupled to climatic changes. Thus, the phase relationship between climate and aragonite concentration could represent aragonite dissolution cycles overprinting the primary input signal.

The number of locations producing periplatform oozes is limited. The Maldives Archipelago in the equatorial Indian Ocean is one such area. Our primary objective at Site 714 was to recover a complete sequence of periplatform ooze of Neogene age. This would allow us (1) to derive a highly resolved stable isotope stratigraphy for the Neogene utilizing the high sediment accumulation rates expected in the periplatform sediments; (2) to determine the interaction between sea-level fluctuations and carbonate off-bank transport; (3) to determine the effects of monsoonal circulation in the northern Indian Ocean on carbonate production and preservation as well as on periplatform oozes in general; and (4) to locate and decipher the possible diagenetic imprint on the primary input signal as well as to determine the fate of the metastable shallow-water-derived components (aragonite and magnesian calcite) through burial diagenesis.

## OPERATIONS

## Maldives to Site MLD-2

All who came aboard at Male except for the two Maldivian observers were put ashore. At $1515 \mathrm{hr}, 22$ June 1987, the ship was under way for Site MLD-2. The presite survey at MLD-2 did not look promising to the Co-Chiefs. Surveying continued, therefore, until an alternate site was found approximately 6 nmi south of MLD-2. The Maldivian government had already given its clearance to drill anywhere within a $10-\mathrm{nmi}$ radius of the proposed sites.

We dropped a beacon at $0230 \mathrm{hr}, 23$ June, to establish Site 714. Because of gusting winds to 70 kt , we temporarily lost the beacon as it descended to the bottom. After approximately 1 hr , the ship was positioned over the beacon and was in its dynamic positioning (DP) mode at 0355 hr .

## Hole 714A

We established the mud line at 2031.5 m and commenced APC coring. The hole was advanced to 2149.9 m ( 118.4 mbsf ), with 115.2 m of core recovered for a recovery rate of $97.3 \%$. The pull-out force on Core $115-714 \mathrm{~A}-13 \mathrm{H}$ was $35,000 \mathrm{lb}$, so the decision was made to go to the XCB system. The hole was further advanced to 2264.5 m ( 233 mbsf ), with 12 XCB -coring runs recovering 114.6 m of core for a recovery rate of $69.3 \%$. Although the XCB shoe seals were worn badly by this point in the leg, they were still used. Total penetration was 233 mbsf to 2264.5 m , with 194.6 m of core recovered for a total recovery rate of $83.5 \%$ (Table 1).

## Hole 714B

The mud line was established at 2031.5 m and APC coring commenced. The objective of this hole was to double APC-core the middle/late Miocene sediments. After 13 APC coring runs with 122.6 m of penetration, we met our objective and abandoned the hole. Total penetration was 122.6 mbsf to 2154.1 m with 114.9 m of core recovered for a recovery rate of $93.7 \%$ (Table 1).

## LITHOSTRATIGRAPHY

The sedimentary sequence at Site 714 is dominated by a thick interval of light greenish gray nannofossil ooze and chalk with variable amounts of three secondary components: foraminifers, clay, and, to a lesser extent, biogenic silica. A thin, darker interval of more clay-rich nannofossil ooze overlies this thick interval. The distinctive color change between these two intervals marks an approximately 8 -m.y. hiatus separating Pleistocene sediment above from late Miocene to late Oligocene sediment


Figure 3. Mass accumulation rates of bulk sediment (unfilled area) and biogenic calcium carbonate (diagonal lines) plotted vs. estimated age at Hole 714 A . The gap from 0.5 Ma to 8.2 Ma represents a hiatus. The data represent mean values within 1-m.y. time increments.
below. The color change forms the basis for dividing the sediment into two units (Units I and II). Unit II, the older unit, is further subdivided into two subunits (Subunits IIA and IIB) by the transition from ooze to chalk, which coincides very closely with an increase in foraminifer content. The stratigraphic distribution of the units and subunits is summarized in Figure 4, and each is discussed in detail below.

Unit I: Core 115-714A-1H to Section 115-714A-3H-5, 115 cm ( $0-19.55 \mathrm{mbsf}$ ) and Core $115-714 \mathrm{~B}-1 \mathrm{H}$ to Section 115-714B-3H-1, 135 cm (0-17.75 mbsf); Age: Pleistocene.

Unit I in both holes consists of foraminifer-bearing, claybearing nannofossil ooze alternating with foraminifer-bearing, clayey nannofossil ooze. Color systematically alternates between dark gray and gray ( $5 \mathrm{Y} 4 / 1,5 / 1$ ) and greenish gray (5GY $5 / 1$, $6 / 1)$. Most of the color changes exhibit gradational boundaries. Smear slide analyses reveal a nannofossil content which is consistently between $50 \%$ and $70 \%$, and a foraminifer content between $12 \%$ and $25 \%$. An additional carbonate component is aragonite needles which occur in minor (up to $5 \%$ ) to trace amounts. Also, minor to trace amounts of biogenic silica, quartz, and volcanic glass are found scattered throughout Unit I.

The lowest carbonate content at Site 714 (analyses from Hole 714A) are found in Unit I, averaging about $75 \%$ and dropping to about $60 \%$. There seems to be a general trend toward increasing carbonate content with depth in Unit I. Superimposed upon this general trend are smaller amplitude, higher frequency fluctuations which seem to roughly correlate with the observed color

|  |  |  |  | Length <br> cored <br> (m) | Length <br> recovered <br> (m) | Recovery <br> (\%) |
| :---: | :---: | :---: | :---: | :---: | :---: | ---: |
| no. | Dime <br> (June 1987) | Depth <br> (UTC) | (mbsf) |  |  |  |

cycles in the sequence. Carbonate content in the darker intervals (dark gray and gray) averages below $75 \%$, while in the lighter intervals (greenish gray) carbonate content averages almost $80 \%$. Thus, the cyclic color changes are related possibly to changes in carbonate production and preservation (and/or detrital input).

We identified a single turbidite 50 cm above the $8-\mathrm{m} . \mathrm{y}$. hiatus in both holes, suggesting only minor disturbance in Unit I. The turbidite seems to be a good lithostratigraphic datum in Unit I. This, along with the matching magnetic susceptibility peaks in the upper part of Unit I (see "Paleomagnetism" section, this chapter), demonstrates the good physical correlation between Holes 714A and 714B within the Pleistocene.

Unit II: Section 115-714A-3H-5, 115 cm , through Core 115-714A-25X (19.55-233.00 mbsf) and Section 115-714B-3H-1, 135 cm, through Core 115-714B-13H (17.75-120.9 mbsf); Age: late Miocene to late Oligocene.

All the sediment cored below the hiatus has been grouped into lithologic Unit II. It is composed primarily of greenish gray nannofossil ooze and chalk with higher amounts of foraminifers than Unit I. As would be expected, chalk content increases with depth. Carbonate content is consistently between $80 \%$ and $90 \%$, averaging about $86 \%$, with less than $10 \%$ of the analyses falling outside the $80 \%-90 \%$ range (see Fig. 11) As previously mentioned, the ooze/chalk transition, accompanied by an apparent increase in foraminifer content, divides Unit II into two subunits (see Fig. 4).


Figure 4. Summary of the distribution of lithostratigraphic units and physical correlations for Holes 714A and 714B.

Subunit IIA: Sections 115-714A-3H-5, 115 cm , to 115-714A$14 \mathrm{X}-2,100 \mathrm{~cm}$ (19.55-120.90 mbsf) and Section 115-714B-3H-1, 135 cm , through Core 115-714B-13H (17.75-122.60 mbsf); Age: late Miocene to middle Miocene.

Subunit IIA consists of faintly burrow-mottled, foraminiferbearing, clay-bearing nannofossil ooze and foraminifer-bearing nannofossil ooze with faint, but distinct, alternating color changes from light greenish gray ( $5 \mathrm{GY} 7 / 1$ ) to light gray ( 5 Y $7 / 1$ ) observed in all but the lowest part of the subunit. The amount of clay-bearing ooze is greatly reduced in the lower $60 \%$ of Subunit IIA. Ooze firmness gradually increases with depth.

Examination of the average carbonate content of cores (within Hole 714A) in Subunit IIA reveals a slight decrease in carbonate content (from around $84 \%$ to $81 \%$ ) with depth at the top of the subunit, followed by a steady increase to around $90 \%$ near the base of the subunit. The most carbonate-rich intervals at Site

714 occur within the lower part of this subunit (around 80-110 mbsf; lower part of the middle Miocene). Visual core descriptions reveal no distinguishing characteristic for this interval.

Biogenic silica seems to dominate the noncarbonate component (greater than $15 \%$ estimated relative abundance in smear slide analyses) within the middle third of Subunit IIA (approximately $60-90 \mathrm{mbsf})$. This, along with the increasing carbonate content with depth in this interval, suggests a low clay content. The apparent biogenic silica peak seems to be superimposed upon an interval of rising carbonate content. This indicates, perhaps, that carbonate variations in Subunit IIA are complicated by influxes of biogenic silica, as well as clay, and can be evaluated better when cast in terms of absolute accumulation (see introductory discussion to this chapter).

A minor to trace component in Subunit IIA is pyrite, which is sparsely disseminated in scattered intervals, and also occurs in pyritic burrows and as pyrite staining in thin horizontal bands and vertical streaks. The presence of minor amounts of pyrite, the distinctive greenish gray color, and the odor of hydrogen sulfide gas released during core splitting indicate sulfate-reduced (low oxygen) conditions in Subunit IIA. Other minor to trace components in Subunit IIA include possible aragonite needles near the top and scattered occurrences of volcanic glass, mica, quartz, and dolomite.

Subunit IIB: Section 115-714A-14X-2, 100 cm , through Core 115-714A-25X (120.90-233.00 mbsf); Age: middle Miocene to late Oligocene.

The top of Subunit IIB is marked by the transition from firm nannofossil ooze (with a few scattered chalk nodules) to abundant nannofossil chalk. This transition is clear in Hole 714A (see Fig. 4). Hole 714B terminates just before the transition. Subunit IIB is composed of clay-bearing, foraminifer-bearing nannofossil chalk which grades into foraminifer nannofossil chalk below. Ooze occurs primarily as a minor component throughout the subunit thinly distributed between and around chalk nodules, biscuits, and layers, and is most likely an artifact of drilling (drilling paste). Another indication of drilling disturbance is the relatively common fracturing observed in the chalk throughout the sequence.

The upper part of Subunit IIB (down to 169 mbsf ) is uniformly a very light greenish gray (5G 8/1) and exhibits only faint burrow mottling in scattered intervals. The uniform color is interrupted by an interval of very abundant faint to distinct light greenish gray (5G7/1) laminae occurring in clusters (169180 mbsf ). Within and below this interval, burrow mottling is much more common. Particularly conspicuous are scattered large Zoophycus burrows among other more common and fainter light gray ( $10 \mathrm{YR} 7 / 2$ ) horizontal to subhorizontal burrows.

The average carbonate content for cores in Subunit IIB exhibits a slight decrease from around $88 \%$ at the top of the subunit to a low of $83 \%$ by 146.1 mbsf (Core 115-714A-16X). Below this level values return to near $89 \%$ and remain at this level for the rest of the hole. Only a few analyses fall outside the $80 \%-90 \%$ range. In general, carbonate content in Subunit IIB seems to exhibit smaller low-frequency variations than in Unit I or Subunit IIA.

Biogenic silica occurs in minor amounts throughout most of Subunit IIB, with a slight increase in abundance observed below 175.0 mbsf (Core 115-714A-20X). Other noncarbonate components include clay, volcanic glass, and quartz.

Evidence of natural disturbance and redeposition in Unit II occurs at the very bottom of Subunit IIB in the form of several small turbidites (one in Core 115-714A-22X and six in Core 115-714A-25X), and the occurrence in scattered intervals of large benthic foraminifers identified as shallower water forms (upper
bathyal and carbonate shelf species; see benthic foraminifer discussion, "Biostratigraphy" section, this chapter). The very pale brown ( $10 \mathrm{YR} 7 / 3$ ) turbidite intervals contain sand-size and larger limestone fragments and shallower-water benthic foraminifers, suggesting transport from shallow-shelf environments to intermediate depths.

## BIOSTRATIGRAPHY

## Introduction

Two holes at Site 714, which lie in a water depth of 2031.5 m , penetrated a $233-\mathrm{m}$-thick sedimentary sequence consisting of (1) 19.55 m of an upper Pleistocene nannofossil ooze; (2) a major unconformity spanning a time interval of approximately $8 \mathrm{~m} . \mathrm{y}$.; and (3) a continuous, expanded sequence ( 200 m thick) of lower upper Miocene through upper Oligocene nannofossil ooze and chalk.

Calcareous nannofossils are abundant throughout the section. They show excellent preservation in the Pleistocene and moderately good to good preservation in the Miocene and upper Oligocene. The diversity of the nannofossil assemblages is higher in the Oligocene-Miocene sediments of Site 714 than in any of the previous sites of Leg 115.

Planktonic foraminifers are abundant, being well preserved in the Pleistocene, well to moderately well preserved in the Miocene, and moderately well preserved in the Oligocene.

Benthic foraminifers are common and well preserved in all sediments, except in the lower Miocene interval where preservation deteriorates. Throughout the sequence, benthic foraminiferal assemblages are representative of middle bathyal depths, except in the upper Oligocene where assemblages are associated with redeposited shallower faunas.

Radiolarians, present in the entire stratigraphic interval, are moderately well to well preserved and vary in abundance from rare to common.

Diatoms are restricted to the upper Miocene through upper middle Miocene interval, and to a short interval in the upper Oligocene. Only traces of diatom valves are present in the Pleistocene. Assemblages are diversified and fairly well preserved in the upper Miocene (in contrast to the previous Leg 115 sites) and are poorly preserved in the middle Miocene.

A biostratigraphic summary for Site 714 is presented in Figure 5.

## Calcareous Nannofossils

## Hole 714A

Hole 714A yielded abundant calcareous nannofossil assemblages ranging in age from the late Oligocene to the late Pleistocene. A hiatus spanning approximately $8 \mathrm{~m} . \mathrm{y}$. is present in Section 115-714A-3H-5, in which middle Pleistocene sediments belonging to Zone CN14a (NN19) overlie upper Miocene sediments belonging to Zone CN8 (NN10).

An interesting feature of the calcareous nannofossil assemblages recorded here is the higher taxonomic diversity in comparison with that observed in other Leg 115 sites. Specifically, in the upper Oligocene and the lower Miocene, several species belonging to the Helicosphaeraceae (e.g., the index species Helicosphaera recta and H. ampliaperta) are present at Site 714 but are totally missing in previous sites. In the middle and upper Miocene, the Pontosphaeraceae (Perch-Nielsen, 1985) are also well represented by the genera Scyphosphaera and Pontosphaera. Both Helicosphaeraceae and many species of Pontosphaeraceae are more typical of "nearshore" environments and are often missing or rare in open-ocean sediments. Helicosphaeraceae seem to prefer upwelling areas or nutrient-rich water masses (PerchNielsen, 1985). This finding would suggest marginal marine con-
ditions for this Maldivian sequence, in contrast to more openocean conditions for both the Mascarene and Chagos sites during late Oligocene and early Miocene times.

The high taxonomic diversity and good preservation allow us to consider the sequence at Site 714 as a reference section for the upper Oligocene through the lower-upper Miocene calcareous nannofossil biostratigraphy of the equatorial Indian Ocean.

## Pleistocene

Emiliania huxleyi is present at the top of the Pleistocene sequence, and we placed its first occurrence ( FO ) at the base of Zone CN15 between Samples 115-714A-2H-4, 110 cm , and 115$714 \mathrm{~A}-2 \mathrm{H}-5,40 \mathrm{~cm}$. The interval between Samples 115-714A$2 \mathrm{H}-5,40 \mathrm{~cm}$, and $115-714 \mathrm{~A}-3 \mathrm{H}-4,70 \mathrm{~cm}$, is assigned to the upper Pleistocene Subzone CN14b. The entire Pleistocene sequence is relatively free of reworked forms, and we easily identified the last occurrence (LO) of Pseudoemiliania lacunosa in Sample $115-714 \mathrm{~A}-3 \mathrm{H}-4,150 \mathrm{~cm}$. The interval between the latter sample and the sharp lithologic contact which marks the abovementioned hiatus contains abundant large Gephyrocapsa oceanica, which show morphologic features typical of the upper Pleistocene morphotypes of the species. We assigned this interval, therefore, to the upper part of Subzone CN14a. By using the $P$. lacunosa event and the top of the sequence as control points for the sedimentation rate, an age of 0.5 Ma is obtained for the lowermost Pleistocene sediments just above the hiatus.

## Miocene

The Miocene sequence of Site 714 is represented by a fairly uniform and homogeneous lithology with no evidence of reworking. Moreover, it appears to be continuous and complete from the Oligocene/Miocene boundary to the lower upper Miocene. We recognized all the Miocene zones of Martini (1971) and of Okada and Bukry (1980), which are given in Figure 5.

The youngest Miocene sediments recovered below the hiatus belong to the upper Miocene NN10 (CN8) Zone. The other upper Miocene biozones of both Martini (1971) and Okada and Bukry (1980) are easily recognized by the successive LO of Discoaster hamatus, the FO of D. hamatus, and the FO of Catinaster coalitus (Table 3 and Fig. 5). In this interval the assemblage contains common Pontosphaeraceae (represented by the genera Scyphosphaera and Pontosphaera) and Helicosphaeraceae. The boundary between the middle and upper Miocene occurs at approximately $55-60 \mathrm{mbsf}$ within Zone NN8. The total thickness of the upper Miocene is between 35 and 40 m .

The boundary between middle Miocene Zones NN6 and NN7, and the corresponding boundary between Subzones CN5a and CN5b (FO of Discoaster kugleri), occurs in the upper part of Core $115-714 \mathrm{~A}-9 \mathrm{H}$. The boundary between Zones NN5 and NN6, and the corresponding boundary between Zones CN4 and CN5a, is easily recognized on the basis of the LO of Sphenolithus heteromorphus.

We identified the boundary between Zones NN4 and NN5, and the corresponding boundary between Zones CN3 and CN4, by means of the LO of Helicosphaera ampliaperta. The NN4NN5 zonal boundary, which occurs at ~131-132 mbsf, identifies the boundary between the lower and the middle Miocene.

Sphenolithus belemnos seems to be missing above the Triquetrorhabdulus carinatus LO, and therefore the boundary between Zones NN3 and NN4, originally defined by the LO of S. belemnos, has been drawn by means of the FO of S. heteromorphus. These two events are known to be very close to each other (Perch-Nielsen, 1985). We identified the boundary between Zones NN2 and NN3 on the basis of the LO of T. carinatus, which is abundant in the sequence. In the early Miocene interval, the nannofossil assemblage is quite diversified. Several species of


Figure 5. Biostratigraphic summary of Site 714. Black bars indicate recovery in Holes 714A and 714B.
sphenoliths (Sphenolithus delphix, S. dissimilis, S. capricornutus, and $S$. conicus) are restricted to this interval and to the uppermost Oligocene. Among the Helicosphaeraceae, Helicosphaera euphratis, $H$. perch-nielseniae, and small-sized $H$. carteri are well represented.

The LOs of Sphenolithus ciperoensis and of Helicosphaera recta, which define the base of the CN1 Zone and the NN1 Zone, respectively, and approximate the Oligocene/Miocene boundary, occur simultaneously in Sample 115-714A-22X-2, 130 cm .

## Oligocene

As in the overlying lower Miocene sequence, the assemblages are characterized by the presence of Helicosphaeraceae, including low abundances of $H$. recta. The basal Oligocene sediments recovered in this hole contain abundant S. ciperoensis and, thus, are late Oligocene in age.

The upper Oligocene series (Zone CP19) is represented by about 35 m of sediments, from Section 115-714A-22X-4 to the bottom of the hole. We placed the LO of Sphenolithus distentus (base of Subzone CP19b) between Sample 115-714A-24X-6, 130 cm , and Section 115-714A-24X, CC.

## Hole 714B

We drilled Hole 714B only 20 m away from Hole 714A and only examined the core-catcher samples. Because the sub-bottom depth of Core $115-714 \mathrm{~B}-1 \mathrm{H}$ is 4 m lower than that of Core $115-714 \mathrm{~A}-1 \mathrm{H}$, zonal assignments were slightly different between the corresponding core catchers of these two holes. The zonal assignments for the 12 core-catcher samples recovered from Hole 714B are summarized in Table 2.

Table 2. Zonal assignments for core-catcher samples recovered from Hole 714B.

| Section <br> interval | Zone | Period |
| :---: | :--- | :--- |
| 115-714B- |  |  |
| 1H, CC | Subzone CN14b | late Pleistocene |
| 2H, CC | No recovery |  |
| 3H, CC | Zone CN8 (NN10) | late Miocene |
| 4H, CC | Subzone CN7b (NNN) | late Miocene |
| 5H, CC | Subzone CN7b (NN9) | late Miocene |
| 6H, CC | Zone CN6 (NN8) | middle/late Miocene |
| 7H, CC | Subzone CN5b (NN7) | middle Miocene |
| 8H, CC | Subzone CN5a (NN6) | middle Miocene |
| 9H, CC | Subzone CN5a (NN6) | middle Miocene |
| 10H, CC | Zone CN4 (NN5) | middle Miocene |
| 11H, CC | Zone CN4 (NN5) | middle Miocene |
| 12H, CC | Zone CN4 (NN5) | middle Miocene |
| 13H, CC | Zone CN4 (NN5) | middle Miocene |

## Planktonic Foraminifers

## Neogene

The Neogene planktonic foraminiferal biostratigraphy of Hole 714 A is based on data from core catchers and an additional two samples per core. Planktonic foraminifers are abundant throughout the Neogene sequence, being well preserved in the Pleistocene and well to moderately well preserved in the Miocene.

The upper part of Hole 714A down to Section 115-714A-3H-4, rich in Globorotalia menardii, is assigned to the Pleistocene zonal interval N23-N22 based on the presence of rare Globorotalia truncatulinoides.

An unconformity between sections 4 and 5 of Core 115-714A-3H spans the Pliocene and uppermost Miocene. The sequential order of planktonic foraminiferal events, allowing zonal assignment in the expanded middle to lower Miocene sequence, is as follows:

1. The simultaneous FO of Neogloboquadrina acostaensis and LO of Paragloborotalia siakensis in Section 115-714A- 6H, CC, marks the boundary between Zones N16 and N14. We could not identify Zone N15.
2. The FO of Globigerina nepenthes in Section 115-714A-10H-2 marks the base of N14.
3. The FO of Sphaeroidinellopsis subdehiscens in Section $115-714 \mathrm{~A}-11 \mathrm{H}-2$ marks the base of N13.
4. The FO of Globorotalia fohsi fohsi in Section 115-714A$11 \mathrm{H}-5$ marks the base of N12.
5. The FO of Globorotalia peripheroacuta in Section 115$714 \mathrm{~A}-12 \mathrm{H}-5$ marks the base of N10.
6. The FO of Orbulina in Section 115-714A-12H, CC, marks the base of N9.
7. The FO of Globigerinoides sicanus in Sample 115-714A$15 \mathrm{X}, \mathrm{CC}$, marks the base of N8.

The interval from Sections 115-714A-15X, CC, through 115-714A-20X, CC, often dominated by abundant small-sized faunas, is assigned to the zonal interval N7-N5, whereas samples in Cores 115-714A-21X and -22X containing small-sized Globorotalia kugleri are assigned to the lowermost Miocene Zone N4.

## Paleogene

We recovered Paleogene planktonic foraminifers in core catchers from Sections $115-714 \mathrm{~A}-23 \mathrm{X}, \mathrm{CC}$, to $115-714 \mathrm{~A}-25 \mathrm{X}, \mathrm{CC}$. One additional sample, $115-714 \mathrm{~A}-23 \mathrm{X}-1,50-55 \mathrm{~cm}$, was surveyed in order to specify the Oligocene/Miocene boundary. Paleogene sediments were chalky, and foraminifers were only moderately well preserved. We recognized the following zones:

1. Zone P22 (Sample 115-714A-23X-1, 50-55 cm, and Section 115-714A-23X, CC) was identified by the presence of Globigerina ciperoensis and Paragloborotalia pseudokugleri and by the absence of G. kugleri. Despite only moderate preservation, the fauna contained a full range of species including such solu-tion-susceptible forms as Globigerina angulisuturalis. Paragloborotalia nana and P. siakensis were common. Smaller species and Globoquadrina tripartita and Globigerina binaiensis dominated the coarse fractions.
2. Zone P21A (Sections 115-714A-24X, CC, and 115-714A25X, CC) was determined by the co-occurrence of Paragloborotalia opima and common Streptochilus cubensis. Faunas contained large numbers of the G. tripartita group, G. angulisuturalis, and Catapsydrax unicavus.

## Benthic Foraminifers

Benthic foraminifers were recovered from all cores in Holes 714A and 714B at Site 714. With the exception of the lower Miocene interval, benthic foraminifers were well preserved and numerous throughout the Pleistocene-upper Oligocene sequence. From the Pleistocene through the lower Miocene Zone CN1 (Section 115-714A-20X, CC), benthic foraminifers were in place and representative of middle bathyal depths. In upper Oligocene Zone P21 (Sections 115-714A-24X, CC, and 115-714A-25X, CC), middle bathyal benthic foraminifers were accompanied by upper bathyal and carbonate platform species such as Amphistegina. This implies that the site was still located at middle bathyal depths and that the shallower material was redeposited.

## Pleistocene

The Pleistocene fauna (Sections 115-714A-1H, CC, and 115$714 \mathrm{~A}-2 \mathrm{H}, \mathrm{CC}$ ) were well preserved and more abundant than at Site 707, which was drilled at approximately the same water depth. Nevertheless, the faunal composition was similar at both sites. Middle bathyal depths were indicated by the presence of Osangularia bengalensis, Hoeglundina elegans, and Bulimina striata, with B. rostrata, Uvigerina proboscidea, and Anomalinoides globulosus together with the truly carinate form of Cassidulina laevigata carinata. The low number of uvigerinids in both samples suggests deposition during interglacial episodes.

## Upper Miocene

The upper Miocene (Sections 115-714A-3H, CC, through 115$714 \mathrm{~A}-6 \mathrm{H}, \mathrm{CC}$ ) contains a series of faunas suggesting a more or
less well-developed oxygen minimum at the bottom. These faunas are characteristic of the darker grey-green vs. grey sediment which occurred in cycles throughout this interval. The oxygen minimum was indicated by floods of small bolivinids, the majority of which are ornamented with reticulations, striations, or ridges; these occurred in the grey intervals. In the darker colored sediment, bolivinids may have been equally abundant, but the majority of forms are small and smooth. In Californian offshore basins, where bolivinid abundance correlates with lower oxygen levels in bottom waters, the lowest oxygen levels are characterized by smooth bolivinids. Slightly higher oxygen levels, on the other hand, are characterized by the ornamented, crenulate forms. This suggests that episodic color changes are primary and represent cycles of bottom-water oxygenation.

Upper Miocene faunas contain larger numbers and a greater diversity of uvigerinids than at any other Leg 115 site. Species include Rectuvigerina spinea, Uvigerina auberiana, U. proboscidea, and U. spinulosa. These were most abundant in Cores $115-714 \mathrm{~A}-5 \mathrm{H}$ and -6 H . The presence of uvigerinids suggests higher accumulation rates here than at the other Leg 115 sites, presumably as a result of increased overhead productivity.

Unusually good preservation in these sediments is indicated not only by the diverse benthic faunas, but by the high numbers of miliolids. Pyrgo murrhina and Triloculina lucernula occur together with Quinqueloculina seminulum, Q. venusta, and at least two other triloculinids and a spiroloculinid.

An intermediate water depth is suggested by the presence of the miliolids, and the co-occurrence of Osangularia trans. bengalensis, Bulimina striata, Anomalinoides cf. aragonensis, A. alazanensis, and Rectuvigerina spinea. Both R. spinea and U. spinulosa are restricted to areas above $\sim 1500 \mathrm{~m}$ throughout the late Paleogene and the Neogene. In addition, U. spinulosa is a species characteristic of subtropical intermediate waters under gyre-margin current systems and the equatorial circulation. These ecologic features seem consistent with the location of Site 714 in the late Miocene.

## Middle Miocene

A different bottom environment is indicated by the preservation and character of the faunas in middle Miocene Zones CN5CN4 (Sections 115-714A-8H, CC, through 115-714A-10H, CC). Bolivinids, miliolids, uvigerinids, benthic diversity, and the siliceous component of the coarse fractions decrease dramatically. Typical intermediate faunas include Nuttalides umbonifera, a bottom water and/or corrosivity index. By Zone CN4 (Section $115-714 \mathrm{~A}-10 \mathrm{H}, \mathrm{CC}$ ), only solution-resistant benthic foraminifers occur in the small faunas. Because the samples lie above the chalk/ooze transition and because preservation and abundance improve in the underlying cores, we cannot attribute the faunal change solely to diagenesis. The decrease in bolivinids and consistent occurrence of hispid uvigerinids, together with N. umbonifera, suggest the development of oxidizing, corrosive conditions at the bottom. Future studies of the ostracodes in this interval should test such interpretations because there are several ostracodal tracers for oxygen at the bottom.

During Zone CN4 (Section 115-714A-11H, CC), the benthic fauna becomes more diverse, sediment preservation improves, ornamented bolivinids flood the fine fraction, and a new component is added to the faunas. The austral species Siphonina australis, Stilostomella basicarinata, and a bolivinitid occur together only at this time. The siphoninids become more abundant in general at this time, as Siphonina tenuicarinata is also present. The siphoninids consistently co-occur with glauconite in areas of the North Atlantic, as well as in shallow-water environments, and are tentatively used to suggest micro-reducing (but macro-oxidizing) conditions at the bottom. The uvigerinid is Rectuvigerina spinea. The benthic fauna suggests higher pro-
ductivity surface-water conditions and a possibly cooler, microreducing and macro-oxidizing bottom environment.

## Lower Miocene

Preservation deteriorates in the interval from Zone CN4 through CN1 (Sections 115-714A-12H, CC, through 115-714A$22 \mathrm{X}, \mathrm{CC}$ ), so that benthic faunas are less diverse and probably unrepresentative of the original fauna in the area. As in other intermediate depth areas, benthic faunas are small, preservation is poor, unusual genera and species occur throughout, and bolivinids are only occasionally very abundant. This interval is important because it represents the transition from the Paleogenetype bottom environment to those typical of the Neogene. The majority of Paleogene benthic species become extinct, and typically Neogene species evolve throughout this interval.

## Lower Miocene-Upper Oligocene

From the lowermost Miocene Zone CN1 through the upper Oligocene (Sections 115-714A-21X, CC, through 115-714A-25X, CC ), shallow-water materials were added to the intermediatedepth benthic faunas in ever-increasing amounts. Only one or two shallow-water amphisteginids were found in Sections 115-714A-21X, CC, and 115-714A-22X, CC. By Section 115-714A24X, CC, however, large upper bathyal specimens and carbon-ate-platform benthic foraminifers occur together with the intermediate benthic fauna. This fauna contains Cibicidoides kullenbergi, intermediate-size Globocassidulina subglobosa, Stilostomella subspinosa, and Uvigerina pygmaea, all intermedi-ate- or deeper-water forms of this time. Site 714, therefore, was still situated at intermediate depths in the late Oligocene, but was receiving redeposited shallower-water material throughout the early CN1-CP18 zonal interval represented in the core catchers.

## Radiolarians

Identifiable radiolarians are present in nearly the entire stratigraphic sequence recovered at Site 714. Two stratigraphic intervals are represented, with an unconformity between them: the upper Pleistocene (approximately $0-20.0 \mathrm{mbsf}$ ) and the upper Miocene to upper Oligocene (20.0-233.0 mbsf).

## Pleistocene

Section 115-714A-1H, CC, is of late Pleistocene age, within the Buccinosphaera invaginata Zone. Section 115-714A-2H, CC, is also of late Pleistocene age, but within the Collosphaera tuberosa Zone. Diagnostic taxa include C. tuberosa, Collosphaera orthoconus, Didymocyrtis tetrathalamus, Anthocyrtidium nigriniae, Pterocorys hertwigii, Spongaster tetras, and Amphirhopalum ypsilon. Sample 115-714A-3H-4, 70 cm , corresponds with the A. ypsilon Zone. Radiolarians are common and well preserved at the top of the Pleistocene interval, but preservation deteriorates downward. A sharp unconformity is present in Section $115-714 \mathrm{~A}-3 \mathrm{H}-5$, at a level of about 20.0 mbsf .

## Miocene

Section 115-714A-3H, CC, through Sample 115-714A-5H-4, 70 cm , are assigned to the Didymocyrtis antepenultima Zone of late Miocene age. Radiolarians are common and well preserved in the upper part of this interval, but preservation deteriorates downward. Diagnostic radiolarian taxa include Diartus hughesi, Didymocyrtis antepenultima, Stichocorys delmontensis, Lithopera neotera, Siphostichartus corona, and Calocycletta caepa.

We assigned Section 115-714A-5H, CC, through Sample 115$714 \mathrm{~A}-10 \mathrm{H}-4,70 \mathrm{~cm}$ (41.4-85.0 mbsf), to the Diartus petterssoni Zone of middle Miocene age. Radiolarian abundance varies from rare to common, and preservation varies from good to poor within this interval. Diagnostic taxa include Lithopera thornburgi, L. neotera, Diartus petterssoni, Didymocyrtis laticonus,

Calocycletta caepa, Cyrtocapsella japonica, C. cornuta, Siphostichartus corona, and Stichocorys delmontensis.

Section 115-714A-10H, CC, through Sample 115-714A-14X-4, 70 cm ( $89.5-123.6 \mathrm{mbsf}$ ), are assigned to the Dorcadospyris alata Zone of middle Miocene age. Radiolarians are rare and moderately to poorly preserved in this interval. Diagnostic taxa include Stichocorys wolffii, S. delmontensis, Carpocanopsis bramlettei, Calocycletta virginis, C. costata, Cyrtocapsella cornuta, and Dorcadospyris alata.

The Calocycletta costata Zone of middle Miocene age occurs in Sections 115-714A-14X, CC, through 115-714A-17X, CC ( $126.8-155.7 \mathrm{mbsf}$ ). Radiolarians are common and moderately well preserved. Diagnostic taxa include Calocycletta costata, $C$. virginis, Cyrtocapsella cornuta, C. japonica, Eucyrtidium diaphanes, Didymocyrtis tubaria, D. violina, Stichocorys wolffii, S. delmontensis, and Carpocanopsis bramlettei.

Sample 115-714A-18X-4, 70 cm , through Section 115-714A18X, CC (160.9-165.3 mbsf), are assigned to the Stichocorys wolffii Zone of early Miocene age. Radiolarians are rare and moderately preserved. Diagnostic taxa include Stichocorys wolffii, S. delmontensis, Lychnocanoma elongata, Carpocanopsis cristatum, Eucyrtidium diaphanes, Calocycletta virginis, and Cyrtocapsella cornuta.

We assigned Sample 115-714A-19X-4, 70 cm , through Section 115-714A-20X, CC (170.5-184.7 mbsf), to the Stichocorys delmontensis Zone of early Miocene age. Radiolarians are rare and poorly preserved; the faunal assemblage is dominated by sponge spicules. Poorly preserved radiolarian fragments include Calocycletta virginis, Eucyrtidium cienkowskii, E. diaphanes, Stichocorys delmontensis, Cyrtocapsella elongata, C. cornuta, and Theocorys spongoconus.

Section 115-714A-21X, CC ( 194.4 mbsf ), is assigned to the Cyrtocapsella tetrapera Zone of early Miocene age. Radiolarians are rare and poorly preserved. Diagnostic taxa include Lychnocanoma elongata, Artophormis gracilis, and Calocycletta virginis.

## Oligocene

Sample 115-714A-22X-4, 70 cm , through Section 115-714A23X, CC (199.6-213.8 mbsf), are assigned to the Lychnocanoma elongata Zone of uppermost Oligocene age. Radiolarians are rare and poorly preserved. Diagnostic taxa include Artophormis gracilis, Dorcadospyris ateuchus, Lychnocanoma elongata, L. trifolium, Eucyrtidium cienkowskii, and Theocyrtis annosa.

All samples examined below Section 115-714A-23X, CC ( 213.8 mbsf ), are barren of radiolarians.

## Diatoms

A consistent occurrence of diatoms is restricted to the stratigraphic interval of upper to middle Miocene. Only traces of valve fragments are present in the Pleistocene sections. The upper Miocene assemblages are generally characterized by a fairly high abundance of diatoms showing a moderate to poor preservation. A late Miocene productivity increase may be deduced from the increase in the abundance of Thalassionema spp., a group of species which under present-day conditions occur in high abundances in upwelling areas. Both the middle and upper Miocene diatom assemblages have a fairly high content of neritic species. Diatoms are not present in the early Miocene section of Hole 714A, whereas a rare but fairly well-preserved flora is present in the upper Oligocene.

Site 714 differs significantly from the previous sites of Leg 115 in terms of the stratigraphic occurrence of diatoms. Diatoms have not been observed in a Miocene section older than
late Miocene in any of the previous sites, nor has a true Paleogene flora been recovered in the previous sites.

## Pleistocene and Pliocene

In the recovered Pleistocene core-catcher samples of Holes 714A and 714B, diatoms have a scattered occurrence due to dissolution. In Sections $115-714 \mathrm{~A}-1 \mathrm{H}, \mathrm{CC}$, and 115-714A-2H, CC, valve fragments of Pliocene-Pleistocene diatoms are present. Diatoms are absent in the youngest core catchers of Hole 714B (Sections 115-714B-1H, CC, and 115-714B-2H, CC).

## Miocene

The late Miocene Coscinodiscus yabei Zone is recognized in Holes 714A and 714B in Sections 115-714A-3H, CC, 115-714A$4 \mathrm{H}, \mathrm{CC}$, and $115-714 \mathrm{~B}-4 \mathrm{H}, \mathrm{CC}$. The floras, which are fairly well preserved and diversified, include species such as Coscinodiscus temperei delicata, C. yabei, and Actinocyclus ellipticus.

Sections $115-714 \mathrm{~A}-5 \mathrm{H}, \mathrm{CC}, 115-714 \mathrm{~A}-6 \mathrm{H}, \mathrm{CC}$, and $115-714 \mathrm{~B}-$ $5 \mathrm{H}, \mathrm{CC}$, are placed at the boundary between the Actinocyclus moronensis and C. yabei zones. The biogenic silica content is fairly high in these samples, but strong fragmentation has affected the assemblages. Sections $115-714 \mathrm{~A}-7 \mathrm{H}, \mathrm{CC}$, and $115-$ 714B-6H, CC, are assigned to the $A$. moronensis Zone.

The middle Miocene Craspedodiscus coscinodiscus Zone is recognized in Sections 115-714A-8H, CC, and 115-714A-9H, CC. Sections $115-714 \mathrm{~B}-7 \mathrm{H}, \mathrm{CC}$, and $115-714 \mathrm{~B}-8 \mathrm{H}, \mathrm{CC}$, are referred to the Coscinodiscus gigas var. diorama-Craspedodiscus coscinodiscus Zones. The abundance of diatom valves is low and preservation is poor. No diatoms of early Miocene age are present in Holes 714A and 714B.

## Oligocene

In Sections 115-714A-22X, CC, and 115-714A-23X, CC, the siliceous microfossil assemblages are dominated by radiolarians and sponge spicules. Only a subordinate amount of moderately well-preserved diatoms are present. The diatom flora includes such age-diagnostic species as Rocella vigilans, Coscinodiscus rhombicus, and Bogorovia veniamini, which tentatively place the samples in the $B$. veniamini Zone.

All samples examined below Section 115-714A-23X, CC, are barren of diatoms.

## PALEOMAGNETICS

## Introduction

The results from sediments taken at Site 714 are disappointing; we will discuss them only briefly therefore. Data acquired from pass-through measurements showed no interpretable magnetics except perhaps from the topmost sediments (where the signal is marginal at best).

## Results

The first few meters of Pleistocene sediment recovered proved promisingly stable; however, with increasing depth the intensity of magnetization decreases and the measured directions become erratic. Intensities remain low in all other cores measured (except, perhaps, near core tops), and the directions appear to be largely random.

## Discussion

It appears that the NRM-carrying minerals are not preserved upon burial. Consequently, there seems to be little chance that we can establish an interpretable magnetic stratigraphy at this site. There is no evidence of consistent overprinting, and the directions measured probably reflect the noise of the instrument.

Occasional higher intensity sections (near core tops) most likely result from minor rust contamination.

## Magnetic Susceptibility

Magnetic susceptibility was measured at varying intervals (6, 7 , or 10 cm ), depending on the lithologic homogeneity of the sequence, in all sections of cores from Hole 714A to a depth of 118 mbsf (Cores $115-714 \mathrm{~A}-1 \mathrm{H}$ through -13 H ); and at intervals of 5 cm in two cores from Hole 714B (Cores 115-714B-1H and $-3 \mathrm{H})$ to a depth of 24 mbsf . A limited number of susceptibility measurements were made in Holes 714A and 714B because of the low signal-to-noise ratio exhibited by the sediments recovered. The Pleistocene sediments yielded low susceptibility values ( $5-15 \times 10^{-6} \mathrm{cgs}$ ), and the underlying Miocene sediments were almost totally devoid of magnetizable material, with the single exception of a peak in susceptibility of about $40 \times 10^{-6} \mathrm{cgs}$ toward the top of each hole.

Within the sequence measured at Site 714, there are two main lithologic units (see "Lithostratigraphy" section, this chapter): one of Pleistocene age ( $0-19.54 \mathrm{mbsf}$ ) and one of Miocene age ( 19.54 mbsf to the base of the section measured for susceptibility).

The Pleistocene sequence (Unit I) consists of gray, dark gray, and greenish gray, clay-bearing (sometimes foraminifer-bearing) nannofossil oozes, which produce a malodorous gas (hydrogen sulfide) when the cores are split and which exhibit pyrite staining throughout. These features suggest that the Pleistocene sediments are strongly reduced, possibly as a result of high initial organic carbon concentrations and high rates of deposition. A large flux of organic carbon and a rapid rate of burial favors suboxic processes of organic matter diagenesis, which involve bacterial reduction (dissociation) of nitrates, followed by the precipitation of manganese oxides, iron oxides, oxyhydroxides, and eventually iron sulfides in the sediment (i.e., pyrite; see Froelich et al., 1979; Berner, 1980; Klinkhammer, 1980; Bender and Heggie, 1984).

The underlying Miocene sequence (a portion of Unit II) consists of foraminifer-bearing nannofossil oozes throughout (with occasional siliceous-bearing horizons), becoming lighter in color downhole: from light brownish gray with occasional pale green laminae at the top of the sequence (e.g., Cores $115-714 \mathrm{~A}-3 \mathrm{H}$, -4 H , and -5 H ), to light greenish gray and pale green toward its base (e.g., Cores 115-714A-18X and -19X). Pyrite staining again was observed throughout the sequence, and the smell of hydrogen sulfide was detected in several of the Miocene cores upon splitting. Therefore, the Miocene sediments at Site 714 are likely to be reduced as strongly as the overlying Pleistocene sediments. The lighter colors and lower susceptibility values of the Miocene sequence are probably the result of their higher carbonate content (see "Geochemistry" section, this chapter) and lower lithogenic clay content. It is likely that most of the magnetic (NRMcarrying) iron oxides and oxyhydroxides originally present in the sediment have been destroyed by bacterial dissociation during suboxic diagenesis of organic matter in both the Pleistocene and Miocene sections. The measured susceptibility values are probably generated mainly by paramagnetic ferrous iron in clay minerals and pyrite.

Variations in magnetic susceptibility within the Pleistocene sequence (though possibly augmented by a significant authigenic component) are sufficiently consistent in a spatial sense to provide a reliable correlation between Holes 714A and 714B, as shown in Figure 6.

## SEDIMENTATION RATES

Site 714 recovered a 233 -m-thick Neogene and late Paleogene sequence, with a long hiatus spanning the interval from late Pleistocene to late Miocene.

Sedimentation rates for Hole 714A are presented in Figure 7 based on biostratigraphic datums from various fossil groups listed in Table 3. The curves are drawn without making corrections for compaction and without considering short-term fluctuations. The reduced sedimentation rates observed in most of the lower and middle upper Miocene sections of the previous sites of Leg 115 are not recorded at Site 714.

## GEOCHEMISTRY

## Interstitial Water Studies

Samples for interstitial water studies were taken from 12 cores at Site 714. The results of the analyses are presented in Table 4 and Figure 8. As a result of the potential for higher amounts of organic material at this site, phosphate determinations were made in addition to the usual analyses.

## Calcium and Magnesium

The downhole gradient of $\mathrm{Ca}^{2+}$ is the lowest observed in any of the holes drilled during Leg 115. The concentration increases from $10.58 \mathrm{mmol} / \mathrm{L}$ at 1.45 mbsf to only $11.65 \mathrm{mmol} / \mathrm{L}$ at 224.75 mbsf . This corresponds to a gradient of $0.004 \mathrm{mmol} / \mathrm{L} / \mathrm{m}$. It must be remembered, however, that the titration technique used (see "Explanatory Notes" chapter, this volume) cannot distinguish between $\mathrm{Ca}^{2+}$ and $\mathrm{Sr}^{2+}$, and therefore it is possible that the apparent increase in $\mathrm{Ca}^{2+}$ is really a result of an increase in $\mathrm{Sr}^{2+}$. Normally, changes in the amount of $\mathrm{Sr}^{2+}$ are substantially less than variations in the concentration of $\mathrm{Ca}^{2+}$. For example, the $\mathrm{Sr}^{2+}$ content of seawater, which is approximately $97 \mu \mathrm{~mol} / \mathrm{L}$, can rise to concentrations of between 700 and $1000 \mu \mathrm{~mol} / \mathrm{L}$ (Baker, 1986; Swart and Guzikowski, 1988) because of the dissolution of organically produced low-magnesium calcite and aragonite, and the precipitation of low-magnesium calcite cements. This process can affect measured $\mathrm{Ca}^{2+}$ concentrations by up to $1 \mathrm{mmol} / \mathrm{L}$. Such a contribution can be lost in the fluctuation of the typical downhole increase in calcium as observed at previous sites drilled during Leg 115.

An increase in pore-water $\mathrm{Sr}^{2+}$ concentration at Site 714 is even more likely since aragonite, which can contain between 7000 and 8000 ppm strontium, is a predominant component in the upper 30 m . As a result of the metastable nature of aragonite, it will tend to dissolve more easily than low-magnesium calcite and hence impart strontium to the pore fluids. The maximum rise of strontium in pore waters of deep-sea sediments appears to be limited by the solubility product of celestite $\left(\mathrm{SrSO}_{4}\right.$; Baker and Bloomer, 1988). In the absence of sulfate reduction, the strontium content of pore fluids cannot exceed approximately $700 \mu \mathrm{~mol} / \mathrm{L}$.

In contrast to most of the other sites examined during Leg 115, no significant changes were observed in the concentration of $\mathrm{Mg}^{2+}$ in Hole 714A (Fig. 9). This type of relationship typically is observed in situations where pore-water chemistry is dominated by carbonate dissolution and precipitation reactions (Gieskes, 1981).

## Alkalinity, Sulfate, Ammonia, and Phosphate

The alkalinity of the pore fluids reaches a maximum of 4.44 $\mathrm{mmol} / \mathrm{L}$ at 39.15 mbsf (Fig. 10). The maximum depletion in sulfate is achieved at 8.75 mbsf and coincides approximately with a rise in phosphate to between 6.99 and $6.22 \mu \mathrm{~mol} / \mathrm{L}$ and in ammonia to $135 \mu \mathrm{~mol} / \mathrm{L}$. Below this depth, phosphate rapidly decreases to seawater values of approximately $2 \mu \mathrm{~mol} / \mathrm{L}$, while sulfate remains depressed and ammonia is enriched throughout the remainder of the core.

## Chlorinity and Salinity

Chlorinity and, to a lesser extent, salinity show small variations: from $526.8 \mathrm{mmol} / \mathrm{L}$ at 1.45 mbsf , to a maximum of 564.7


Figure 6. Correlation between the upper portion of Holes 714A and 714B, based on whole-core magnetic susceptibility profiles of each hole. Lithologic units correspond to those defined in "Lithostratigraphy" section, this chapter.
$\mathrm{mmol} / \mathrm{L}$ at 8.75 mbsf , and finally back to $531.5 \mathrm{mmol} / \mathrm{L}$. This type of change in chlorinity was also seen at Site 709 and may be an inherited signature related to paleosalinity variations of the oceans during Pleistocene times.

## Silica

Concentrations of silica gradually rise from $575 \mu \mathrm{~mol} / \mathrm{L}$ at 1.45 mbsf to $870 \mu \mathrm{~mol} / \mathrm{L}$ at 200.10 mbsf . The gradual increase probably reflects dissolution of biogenic silica components in the hole.

## X-ray Mineralogy, Carbonate, and Organic Analyses

 $X$-ray Mineralogy and Carbonate ContentAll sediment samples squeezed for interstitial waters were analyzed by X-ray diffraction in order to determine their car-
bonate mineralogy. These analyses revealed that aragonite and low-magnesium calcite were the dominant carbonate minerals throughout the upper 25 m of Hole 714A, an interval corresponding to sediments of Pleistocene age. Below the sharp hiatus visible in Core 115-714A-4H, aragonite disappears and lowmagnesium calcite was the only mineral detected.

The carbonate content of Hole 714A varies between 60.05 and $94.74 \mathrm{wt} \%$, with an overall average of $84.8 \mathrm{wt} \%( \pm 5.35$; Table 5 and Fig. 11). The lowest concentrations of aragonite occur in the uppermost Pleistocene interval where carbonate contents average approximately $76 \mathrm{wt} \%$ (Fig. 12). The noncarbonate fraction of the samples is composed primarily of quartz, presumably detrital in nature. The amount of carbonate gradually rises in the lower Pleistocene, and the percentage of quartz decreases. The upper Miocene is characterized by an average carbonate content of $82.77 \mathrm{wt} \%$, which gradually increases to


Figure 7. Sedimentation rates in Hole 714A based on the biostratigraphic events listed in Table 3.
an average of $89.4 \mathrm{wt} \%$ in the middle Miocene. The concentration of carbonate remains at approximately this level throughout the remainder of the core.

## Organic Carbon Analysis

The concentration of organic material in Hole 714A was the highest observed during Leg 115. In the Pleistocene interval, concentrations were all in excess of $0.18 \%$, and in some instances were as high as $0.78 \%$ (Fig. 11). However, there did not appear to be a downhole decrease consistent with the continued oxidation of organic material, as was suggested by the sharp decrease in sulfate and increase in phosphate. Below the hiatus marking the boundary between the Pleistocene and Miocene, concentrations of organic material rapidly decrease to background concentrations (Table 6 and Fig. 11).

In addition, concentrations of methane obtained by headspace analysis (see "Explanatory Notes" chapter, this volume) were detectable above background levels of approximately 2 ppm (Fig. 13 and Table 7). As methane concentrations decrease with depth, the concentrations of methane are related to methanogenesis taking place on an extremely small scale within the sediments themselves. Locally, it is possible that more drastic reductions in the concentration of sulfate have occurred than is observed from squeezing bulk sediment samples.

## PHYSICAL PROPERTIES

## Introduction

The following physical properties were measured at Site 714: index properties including carbonate content, compressionalwave velocities on discrete samples, shear-wave velocities, shear

Table 3. Biostratigraphic datum levels, Hole 714A.

|  | Species event | Depth (mbsf) | Age <br> (Ma) |
| :---: | :---: | :---: | :---: |
| LO | P. lacunosa (N) | 17.3-18.8 | 0.46 |
| LO | C. yabei (D) | 12.4-26.0 | 7.5 |
|  | D. petterssoni-D. hughesi ( R ) | 36.9-41.4 | 8.6 |
| LO | D. hamatus ( N ) | 33.0-33.3 | 8.9 |
| FO | D. hamatus ( N ) | 52.6-54.1 | 10.0 |
| FO | N. acostaensis (F) | $51.0-60.6$ | 10.2 |
| LO | P. siakensis ( F ) | 60.6-70.2 | 10.4 |
| FO | C. coalitus ( N ) | 60.7-62.2 | 10.8 |
| FO | C. coalitus ( N ) | 60.7-62.2 | 10.8 |
| FO | D. petterssoni ( R ) | 79.8-89.5 | 11.5 |
| FO | S. subdehiscens (F) | 79.8-89.5 | 11.8 |
| LO | S. heteromorphus (N) | 87.1-88.0 | 13.2 |
| FO | G. foshi foshi (F) | 96.5-99.1 | 13.5 |
| FO | Orbulina (F) | 108.7-118.4 | 15.2 |
|  | D. dentata-D. alata ( R ) | 123.6-126.8 | 15.3 |
| LO | H. ampliaperta ( N ) | 131.1-132.4 | 16.0 |
| FO | G. sicanus (F) | 136.4-146.1 | 16.6 |
| FO | S. heteromorphus ( N ) | 163.0-163.3 | 18.6 |
| FO | C. costata (R) | 155.7-160.9 | 17.3 |
| FO | S. wolffii (R) | 165.3-170.5 | 18.2 |
| FO | S. delmontensis ( R ) | 184.7-194.4 | 21.1 |
| LO | G. kugleri ( F ) | 184.7-194.4 | 21.8 |
| FO | G. kugleri (F) | 204.1-213.8 | 23.7 |
| LO | S. ciperoensis (N) | 195.7-197.2 | 25.2 |
|  | B. veniamini (D) | 204.1-213.8 | 26.5 |
| LO | S. distentus (N) | 215.3-223.5 | 28.2 |

Note: $\mathrm{FO}=$ first occurrence, $\mathrm{LO}=$ last occurrence, $\mathrm{N}=$ nannofossil, $\mathrm{F}=$ foraminifer, $\mathrm{D}=$ diatom, and $\mathrm{R}=$ radiolarian.
strengths, and thermal conductivities. We completed the transfer of the $P$-wave logger and GRAPE to a vertical tracking system during the drilling of Hole 714A. However, a number of problems arose with implemention of the new system, prohibiting any measurements at this hole.

## Index Properties

The calculated results for wet-bulk density, porosity, water content, and grain density for Site 714 are shown in Figures 14 and 15 and are listed in Table 8. The measured carbonate contents of the same samples are plotted in Figure 16. The wet-bulk density gradually increases with increasing depth, whereas the other index properties generally decrease with depth. However, porosity and water content are constant between 75 and 175 mbsf. There is little scatter in the porosity and water content data, and only sightly more in the bulk- and grain-density results. There are three anomalously low porosity values: at 186 , 216 , and 217 mbsf. They were measured on samples obtained from core catchers.

## Compressional-Wave Velocity and Acoustic Impedance

The compressional-wave velocities $\left(\mathrm{V}_{p}\right)$ obtained from discrete samples from Site 714 are shown in Figure 17 and in Table 9. These data show a fair agreement between Holes 714A and 714 B ; a trend of increasing $\mathrm{V}_{p}$ with increasing depth is also observed. There is a large amount of scatter in the data below 150 mbsf. The discrete acoustic impedances for Site 714 are shown in Figure 18 and in Table 9. Generally, there is a gradual increase of impedance with depth, but no strong impedance contrasts are resolvable.

The GRAPE and $P$-wave logger were mounted on a vertical tracking system for Site 714. The effect of the transfer increased the uncertainty in $\mathrm{V}_{p}$ (determined from the $P$-wave logger) to approximately $2 \%$ for the distilled water standard. More importantly, there was an increase in background noise with the vertical tracking system, which introduced spurious readings. The

Table 4. Interstitial water analyses, Hole 714A.

| Sample interval <br> $(\mathrm{cm})$ | Depth <br> $(\mathrm{mbsf})$ | Ca <br> $(\mathrm{mmol} / \mathrm{L})$ | Mg <br> $(\mathrm{mmol} / \mathrm{L})$ | Cl <br> $(\mathrm{mmol} / \mathrm{L})$ | Al <br> $(\mathrm{mmol} / \mathrm{L})$ | pH | Salinity <br> $(\%)$ | Si <br> $(\mu \mathrm{mol} / \mathrm{L})$ | $\mathrm{SO}_{4}$ <br> $(\mathrm{mmol} / \mathrm{L})$ | $\mathrm{NH}_{4}$ <br> $(\mu \mathrm{~mol} / \mathrm{L})$ | $\mathrm{PO}_{4}$ <br> $(\mu \mathrm{~mol} / \mathrm{L})$ |
| :--- | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Seawater | 0 | 10.58 | 55.21 | 562.9 | 2.47 | 8.5 | 35.4 | 0 | 29.48 | 0 | 1.85 |
|  |  |  |  |  |  |  |  |  |  |  |  |
| 115-714A- |  |  |  |  |  |  |  |  |  |  |  |
|  |  |  |  |  |  |  |  |  |  |  |  |
| 1H-1, 145-150 | 1.45 | 10.58 | 51.21 | 526.8 | 3.30 | 7.6 | 35.2 | 575 | 27.46 | 46.5 | 6.99 |
| 2H-4, 145-150 | 8.75 | 10.72 | 53.95 | 564.7 | 3.88 | 7.5 | 35.2 | 621 | 26.51 | 112.0 | 6.22 |
| 3H-4, 145-150 | 18.75 | 10.68 | 52.23 | 531.5 | 4.32 | 7.6 | 34.8 | 663 | 26.88 | 135.0 | 4.16 |
| 4H-5, 145-150 | 28.08 | 10.83 | 52.16 | 567.5 | 4.08 | 7.5 | 35.5 | 746 | 27.03 | 106.0 | 2.62 |
| 5H-5, 145-150 | 39.15 | 10.81 | 52.10 | 562.0 | 4.44 | 7.6 | 35.2 | 568 | 27.17 | 125.0 | 3.78 |
| 6H-4, 145-150 | 47.35 | 10.94 | 51.89 | 551.8 | 4.23 | 7.6 | 35.4 | 704 | 26.51 | 114.0 | 2.36 |
| 9H-2, 145-150 | 73.15 | 10.92 | 50.71 | 558.3 | 3.99 | 7.7 | 35.0 | 769 | 26.30 | 111.0 | 2.24 |
| 12H-4, 120-125 | 104.80 | 11.01 | 49.98 | 545.3 | 4.19 | 7.6 | 35.2 | 783 | 26.29 | 129.0 | 2.49 |
| 15X-3, 120-125 | 131.00 | 11.27 | 50.04 | 538.9 | 4.00 | 7.5 | 35.2 | 797 | 26.64 | 112.0 | 2.11 |
| 18X-4, 120-125 | 161.40 | 11.07 | 50.64 | 544.4 | 3.78 | 7.5 | 34.8 | 808 | 26.42 | 94.0 | 2.24 |
| 22X-4, 120-125 | 200.10 | 11.43 | 50.77 | 547.2 | 3.63 | 7.6 | 35.2 | 870 | 26.67 | 69.0 | 1.98 |
| 25X-1, 145-150 | 224.75 | 11.65 | 52.54 | 556.4 | 3.36 | 7.8 | 35.8 | 771 | 26.46 | 42.0 | 1.98 |

new tracking system did not affect the GRAPE wet-bulk density results, although the vertical tracking caused slumping and upward migration of gas bubbles for soupy and partially filled core liners.

The $P$-wave logger measurements for $\mathrm{V}_{p}$, GRAPE wet-bulk density results, and acoustic impedances for Hole 714B are shown in Figure 19. The wet-bulk densities, determined by the GRAPE, increase gradually with depth and are in good agreement with the discrete measurements. The $\mathrm{V}_{p}$ measurements are highly variable and contain a number of spurious readings. These uncorrected measurements should be viewed with caution, as should the continuous impedance profile which is calculated from these data.

## Shear Strength and Shear-Wave Velocity

Measurements for shear-wave velocity $\left(\mathrm{V}_{s}\right)$ and shear strength taken at Site 714 are shown in Tables 10 and 11 and Figures 20 and 21. Both $\mathrm{V}_{s}$ and shear strength increase with depth, although there is a large amount of scatter in both sets of data. Shear strength measurements were not taken in the stiffer, more brittle chalks below 127 mbsf .

## Thermal Conductivity

Thermal conductivity measurements for Holes 714A and 714B are shown in Figure 22 and Table 12. Almost all of the measurements were made in Hole 714A. The measurements in Hole 714B were confined to Core 115-714B-1H in order to fill in a small gap in the data at the top of Hole 714A. The thermal conductivity data show a certain amount of small-scale variability in the upper 50 m of this site, but below that depth the magnitude of the variability is reduced in comparison with the other sites of this leg.

There are a few isolated values at depths greater than 100 mbsf that are much lower than the surrounding values; these may be disturbed areas of the core which have a much higher water content. The higher water content would also account for the lower thermal conductivity. The general pattern shown by the thermal conductivity at this site is a steady increase with depth.

## Summary

The physical properties measurements obtained at Site 714 all varied gradually with increasing depth. There was no major change in physical properties at the hiatus at $\sim 20$ mbsf nor at the deeper transition from soft nannofossil oozes to chalks. Porosity decreased to a depth of 75 mbsf , where it leveled off at $60 \%$. A similar leveling off occurred for the other index properties. The carbonate content, shear strength, $\mathrm{V}_{s}, \mathrm{~V}_{p}$, impedance, and thermal conductivity all increased with depth with varying
amounts of scatter. The quality of the continuous $\mathrm{V}_{p}$ data was reduced because of the new vertical tracking system; hence, one should view the continuous $\mathrm{V}_{p}$ and impedance profiles with caution.

## SEISMIC STRATIGRAPHY

Site 714 is located on the eastern shoulder of the Maldives Ridge, in $2031.5-\mathrm{m}$ water depths (Fig. 23). This elevated plateau extends northward from the Chagos Bank to the Laccadive Islands and is part of the proposed volcanic trace of the Réunion hotspot left on the Indian plate. In this region the ridge is capped by a $100-\mathrm{km}$-wide carbonate bank of perhaps $1-1.5 \mathrm{~km}$ thickness. We intended to position this site on the top of a satellite spur that connects with the main carbonate platform, where pelagic carbonate oozes might be protected from disruption by turbidites. From previous seismic surveys, it appeared as though several hundred meters of soft sediment rested on harder strata (Fig. 24).

The JOIDES Resolution surveys of the site, however, revealed that the ridge was scoured of soft sediment everywhere and only a hard reflector could be seen at or near the sediment-water interface (Fig. 25). We continued to survey down the eastern slope of this ridge in the vicinity of a piston core location (V29-26) known to have recovered 12 m of Pleistocene carbonate ooze. The $3.5-$ and $12-\mathrm{kHz}$ recordings show a $1-\mathrm{km}$-wide basin bounded by sediment hills (Fig. 26). This sediment appears to have accumulated at the base of the ridge by a combination of downslope transport and normal pelagic sedimentation. From the results of the piston coring, we felt there was a reasonable chance that a Quaternary section might be preserved in the small basin.

The single-channel seismic (SCS), water-gun reflection profile run by the JOIDES Resolution over Site 714 (Fig. 25) reveals a series of acoustic reflectors. Any evidence of impedance contrast across the Pleistocene-Miocene unconformity (see "Biostratigraphy" section, this chapter) is masked by reflection from the sediment-water contact. A strong reflector that appears to be continuous with strata near the surface of the nearby ridge occurs at 0.24 s below the seafloor. From drilling we know this is a late Oligocene shallow-water limestone, intersected at about 230 mbsf . In situ $P$-wave velocities of about $1500 \mathrm{~m} / \mathrm{s}$, then, seem appropriate for the overlying Neogene oozes. A second strong reflector appears at 0.36 sbsf and is most likely a basaltic basement.

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Figure 8. Summary of interstitial water analyses, Hole 714A, as a function of sub-bottom depth. Surface seawater is plotted at 0 mbsf.


Figure 9. Relationship between calcium and magnesium concentrations, Hole 714A.


Figure 10. Downhole changes in alkalinity, sulfate, phosphate, and ammonia in Hole 714A. Note the dramatic increase in the concentration of phosphate at 1.45 mbsf and the corresponding decrease in sulfate.

Table 5. Carbonate content of samples from Hole 714A.

| Sample interval (cm) | Depth (mbsf) | Carbonate (wt\%) |
| :---: | :---: | :---: |
| 115-714A- |  |  |
| 1H-1, 30-31 | 0.30 | 81.86 |
| $1 \mathrm{H}-1,100-101$ | 1.00 | 67.47 |
| 1H-2, 30-31 | 1.80 | 72.56 |
| 1H-2, 69-71 | 2.19 | 67.98 |
| $1 \mathrm{H}-2,100-101$ | 2.50 | 60.05 |
| $2 \mathrm{H}-1,30-31$ | 3.10 | 79.41 |
| $2 \mathrm{H}-1,60-62$ | 3.40 | 80.70 |
| $2 \mathrm{H}-1,100-101$ | 3.80 | 80.09 |
| $2 \mathrm{H}-2,30-31$ | 4.60 | 70.81 |
| $2 \mathrm{H}-2,100-101$ | 5.30 | 71.58 |
| $2 \mathrm{H}-3,30-31$ | 6.10 | 71.49 |
| $2 \mathrm{H}-3,100-101$ | 6.80 | 80.78 |
| $2 \mathrm{H}-4,30-31$ | 7.60 | 73.17 |
| $2 \mathrm{H}-4,100-101$ | 8.30 | 76.04 |
| $2 \mathrm{H}-5,30-31$ | 9.10 | 75.81 |
| $2 \mathrm{H}-5,100-101$ | 9.80 | 74.82 |
| 2H-6, 30-31 | 10.60 | 75.22 |
| $2 \mathrm{H}-6,100-101$ | 11.30 | 84.10 |
| $2 \mathrm{H}-7,30-31$ | 12.10 | 84.06 |
| $3 \mathrm{H}-1,30-31$ | 12.70 | 72.32 |
| $3 \mathrm{H}-1,60-62$ | 13.00 | 81.05 |
| $3 \mathrm{H}-1,100-101$ | 13.40 | 80.55 |
| $3 \mathrm{H}-2,30-31$ | 14.20 | 78.40 |
| $3 \mathrm{H}-2,100-101$ | 14.90 | 80.67 |
| $3 \mathrm{H}-3,30-31$ | 15.70 | 66.56 |
| $3 \mathrm{H}-3,60-62$ | 16.00 | 73.65 |
| $3 \mathrm{H}-3,100-101$ | 16.40 | 71.86 |
| $3 \mathrm{H}-4,30-31$ | 17.20 | 74.78 |
| $3 \mathrm{H}-4,60-62$ | 17.50 | 80.50 |
| $3 \mathrm{H}-4,100-101$ | 17.90 | 80.07 |
| $3 \mathrm{H}-5,30-31$ | 18.70 | 81.31 |
| $3 \mathrm{H}-5,60-62$ | 19.00 | 83.64 |
| $3 \mathrm{H}-5,100-101$ | 19.40 | 81.39 |
| $3 \mathrm{H}-6,30-31$ | 20.20 | 81.84 |
| $3 \mathrm{H}-6,60-62$ | 20.50 | 88.56 |
| $3 \mathrm{H}-6,100-101$ | 20.90 | 74.21 |
| $3 \mathrm{H}-7,30-31$ | 21.70 | 88.99 |
| $4 \mathrm{H}-2,30-31$ | 22.48 | 77.19 |
| $4 \mathrm{H}-2,60-62$ | 22.78 | 89.70 |
| $4 \mathrm{H}-2,100-101$ | 23.18 | 83.34 |
| $4 \mathrm{H}-3,30-31$ | 23.93 | 85.69 |
| $4 \mathrm{H}-3,100-101$ | 24.63 | 83.07 |
| $4 \mathrm{H}-4,30-31$ | 25.43 | 86.62 |
| $4 \mathrm{H}-4,86-88$ | 25.99 | 87.60 |
| 4H-4, 100-101 | 26.13 | 87.12 |
| $4 \mathrm{H}-5,30-31$ | 26.93 | 86.14 |
| $4 \mathrm{H}-5,100-101$ | 27.63 | 69.77 |
| $4 \mathrm{H}-6,30-31$ | 28.43 | 86.73 |
| 4H-6, 88-90 | 29.01 | 81.60 |
| 4H-6, 100-101 | 29.13 | 81.86 |
| $4 \mathrm{H}-7,30-31$ | 29.93 | 79.47 |
| $4 \mathrm{H}-7,100-101$ | 30.63 | 82.98 |
| 4H-8, 30-31 | 31.43 | 83.52 |
| $5 \mathrm{H}-1,30-31$ | 32.00 | 83.57 |
| $5 \mathrm{H}-1,100-101$ | 32.70 | 80.71 |
| $5 \mathrm{H}-2,30-31$ | 33.50 | 77.45 |
| $5 \mathrm{H}-2,60-62$ | 33.80 | 78.03 |
| $5 \mathrm{H}-2,100-101$ | 34.20 | 83.38 |
| $5 \mathrm{H}-3,30-31$ | 35.00 | 86.00 |
| 5H-3, 100-101 | 35.70 | 82.54 |
| $5 \mathrm{H}-4,30-31$ | 36.50 | 81.74 |
| $5 \mathrm{H}-4,60-62$ | 36.80 | 79.54 |
| $5 \mathrm{H}-4,100-101$ | 37.20 | 80.30 |
| $5 \mathrm{H}-5$ 30-31 | 38.00 | 84.61 |
| 5H-6, 100-101 | 38.70 | 80.47 |
| $5 \mathrm{H}-6,30-31$ | 39.50 | 81.38 |
| $5 \mathrm{H}-6,60-62$ | 39.80 | 82.67 |
| 5H-6, 100-101 | 40.20 | 75.79 |
| $5 \mathrm{H}-7,30-31$ | 41.00 | 83.57 |
| $6 \mathrm{H}-1,30-31$ | 41.70 | 80.77 |
| $6 \mathrm{H}-1,100-101$ | 42.40 | 82.27 |
| $6 \mathrm{H}-2,30-31$ | 43.20 | 77.45 |
| $6 \mathrm{H}-2,60-62$ | 43.50 | 82.88 |
| $6 \mathrm{H}-2,100-101$ | 43.90 | 73.40 |

Table 5 (continued).

| Sample interval (cm) | $\begin{aligned} & \text { Depth } \\ & \text { (mbsf) } \end{aligned}$ | Carbonate (wt\%) | Sample interval (cm) | $\begin{aligned} & \text { Depth } \\ & \text { (mbsf) } \end{aligned}$ | Carbonate (wt\%) |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 115-714A- |  |  | 115-714A- |  |  |
| $6 \mathrm{H}-3,30-31$ | 44.70 | 84.80 | $11 \mathrm{H}-2,100-101$ | 92.00 | 89.03 |
| $6 \mathrm{H}-3,100-101$ | 45.40 | 72.14 | $11 \mathrm{H}-3,30-31$ | 92.80 | 90.89 |
| $6 \mathrm{H}-4,30-31$ | 46.20 | 84.84 | $11 \mathrm{H}-3,70-72$ | 93.20 | 92.35 |
| $6 \mathrm{H}-4,60-62$ | 46.50 | 86.18 | $11 \mathrm{H}-3,100-101$ | 93.50 | 90.08 |
| $6 \mathrm{H}-4,100-101$ | 46.90 | 69.82 | $11 \mathrm{H}-4,30-31$ | 94.30 | 88.41 |
| $6 \mathrm{H}-5,30-31$ | 47.70 | 83.80 | $11 \mathrm{H}-4,100-101$ | 95.00 | 88.40 |
| $6 \mathrm{H}-5,100-101$ | 48.40 | 86.71 | $11 \mathrm{H}-5,30-31$ | 95.80 | 87.61 |
| $6 \mathrm{H}-6,30-31$ | 49.20 | 80.99 | $11 \mathrm{H}-5,100-101$ | 96.50 | 86.17 |
| $6 \mathrm{H}-6,60-62$ | 49.50 | 82.35 | $11 \mathrm{H}-6,30-31$ | 97.30 | 88.66 |
| $6 \mathrm{H}-6,100-101$ | 49.90 | 83.51 | $11 \mathrm{H}-6,52-54$ | 97.52 | 88.21 |
| $6 \mathrm{H}-7,30-31$ | 50.70 | 80.09 | $11 \mathrm{H}-6,100-101$ | 98.00 | 92.08 |
| $7 \mathrm{H}-1,30-31$ | 51.30 | 85.20 | $12 \mathrm{H}-1,30-31$ | 99.40 | 88.44 |
| $7 \mathrm{H}-1,100-101$ | 52.00 | 84.20 | $12 \mathrm{H}-1,100-101$ | 100.10 | 90.60 |
| $7 \mathrm{H}-2,30-31$ | 52.80 | 79.70 | $12 \mathrm{H}-2,30-31$ | 100.90 | 85.56 |
| $7 \mathrm{H}-2,40-41$ | 52.90 | 74.15 | $12 \mathrm{H}-2,60-62$ | 101.20 | 92.14 |
| 7H-2, 60-62 | 53.10 | 86.18 | $12 \mathrm{H}-2,100-101$ | 101.60 | 89.47 |
| $7 \mathrm{H}-2,100-101$ | 53.50 | 79.73 | $12 \mathrm{H}-3,30-31$ | 102.40 | 90.79 |
| 7H-3, 30-31 | 54.30 | 79.95 | $12 \mathrm{H}-3,100-101$ | 103.10 | 88.96 |
| $7 \mathrm{H}-3,40-41$ | 54.40 | 75.93 | $12 \mathrm{H}-4,30-31$ | 103.90 | 89.42 |
| $7 \mathrm{H}-3,100-101$ | 55.00 | 86.33 | 12H-4, 61-63 | 104.21 | 94.73 |
| $7 \mathrm{H}-4,30-31$ | 55.80 | 79.03 | $12 \mathrm{H}-4,100-101$ | 104.60 | 88.10 |
| 7H-4, 40-41 | 55.90 | 83.10 | $12 \mathrm{H}-5,30-31$ | 105.40 | 88.50 |
| $7 \mathrm{H}-4,60-62$ | 56.10 | 81.43 | 12H-5, 63-65 | 105.73 | 89.54 |
| 7H-4, 100-101 | 56.50 | 75.26 | $12 \mathrm{H}-5,100-101$ | 106.10 | 86.82 |
| 7H-5, 30-31 | 57.30 | 85.53 | $12 \mathrm{H}-6,30-31$ | 106.90 | 89.84 |
| 7H-5, 100-101 | 58.00 | 80.76 | $12 \mathrm{H}-6,100-101$ | 107.60 | 88.41 |
| 7H-6, 30-31 | 58.80 | 76.19 | $12 \mathrm{H}-7,30-31$ | 108.40 | 85.77 |
| 7H-6, 56-58 | 59.06 | 85.32 | $13 \mathrm{H}-1,30-31$ | 109.00 | 90.25 |
| $7 \mathrm{H}-6,100-101$ | 59.50 | 82.22 | $13 \mathrm{H}-1,100-101$ | 109.70 | 87.02 |
| $7 \mathrm{H}-7,30-31$ | 60.30 | 84.55 | $13 \mathrm{H}-2,30-31$ | 110.50 | 90.37 |
| $8 \mathrm{H}-1,30-31$ | 60.90 | 84.21 | $13 \mathrm{H}-2,58-60$ | 110.78 | 88.10 |
| $8 \mathrm{H}-1,100-101$ | 61.60 | 79.19 | $13 \mathrm{H}-2,100-101$ | 111.20 | 92.57 |
| $8 \mathrm{H}-2,30-31$ | 62.40 | 84.67 | $13 \mathrm{H}-3,30-31$ | 112.00 | 90.20 |
| $8 \mathrm{H}-2,60-62$ | 62.70 | 86.18 | $13 \mathrm{H}-3,100-101$ | 112.70 | 90.84 |
| $8 \mathrm{H}-2,100-101$ | 63.10 | 86.43 | $13 \mathrm{H}-4,30-31$ | 113.50 | 90.34 |
| $8 \mathrm{H}-3,30-31$ | 63.90 | 84.41 | $13 \mathrm{H}-4,65-67$ | 113.85 | 87.03 |
| $8 \mathrm{H}-3,100-101$ | 64.60 | 86.04 | $13 \mathrm{H}-4,100-101$ | 114.20 | 88.21 |
| $8 \mathrm{H}-4,30-31$ | 65.40 | 87.55 | $13 \mathrm{H}-5,30-31$ | 115.00 | 91.35 |
| $8 \mathrm{H}-4,60-62$ | 65.70 | 87.74 | $13 \mathrm{H}-5,100-101$ | 115.70 | 85.34 |
| $8 \mathrm{H}-4,100-101$ | 66.10 | 86.46 | $13 \mathrm{H}-6,30-31$ | 116.50 | 87.46 |
| $8 \mathrm{H}-5,30-31$ | 66.90 | 90.60 | $13 \mathrm{H}-6,100-101$ | 117.20 | 87.34 |
| $8 \mathrm{H}-5,100-101$ | 67.60 | 88.14 | $13 \mathrm{H}-7,30-31$ | 118.00 | 86.25 |
| $8 \mathrm{H}-6,30-31$ | 68.40 | 83.29 | 14X-1, 30-31 | 118.70 | 87.73 |
| $8 \mathrm{H}-6,60-62$ | 68.70 | 88.61 | $14 \mathrm{X}-1,100-101$ | 119.40 | 87.36 |
| $8 \mathrm{H}-6,100-101$ | 69.10 | 82.10 | 14X-2, 30-31 | 120.20 | 88.32 |
| $8 \mathrm{H}-7,30-31$ | 69.90 | 87.71 | 14X-2, 63-65 | 120.53 | 88.81 |
| $9 \mathrm{H}-1,30-31$ | 70.50 | 88.65 | $14 \mathrm{X}-2,100-101$ | 120.90 | 86.95 |
| $9 \mathrm{H}-1,100-101$ | 71.20 | 87.96 | $14 \mathrm{X}-3,30-31$ | 121.70 | 88.59 |
| $9 \mathrm{H}-2,30-31$ | 72.00 | 84.35 | $14 \mathrm{X}-3,100-101$ | 122.40 | 85.99 |
| 9H-2, 60-62 | 72.30 | 81.26 | 14X-4, 30-31 | 123.20 | 85.44 |
| $9 \mathrm{H}-2,100-101$ | 72.70 | 84.88 | 14X-4, 63-65 | 123.53 | 87.40 |
| $9 \mathrm{H}-3,30-31$ | 73.50 | 86.01 | $14 \mathrm{X}-4,100-101$ | 123.90 | 82.54 |
| $9 \mathrm{H}-3,75-77$ | 73.95 | 86.58 | 14X-5, 30-31 | 124.70 | 88.80 |
| $9 \mathrm{H}-3,100-101$ | 74.20 | 82.99 | 14X-5, 100-101 | 125.40 | 84.83 |
| $9 \mathrm{H}-4,30-31$ | 75.00 | 86.42 | 14X-6, 30-31 | 126.20 | 88.38 |
| 10H-1, 30-31 | 80.10 | 86.21 | 14X-6, 54-56 | 126.44 | 85.91 |
| $10 \mathrm{H}-1,100-101$ | 80.80 | 86.17 | 14X-6, 65-67 | 126.55 | 89.49 |
| 10H-2, 30-31 | 81.60 | 87.50 | 14X-6, 100-101 | 126.90 | 88.42 |
| 10H-2, 60-62 | 81.90 | 86.31 | $15 \mathrm{X}-1,30-31$ | 127.10 | 79.94 |
| 10H-2, 100-101 | 82.30 | 79.36 | 15X-1, 92-93 | 127.72 | 86.82 |
| $10 \mathrm{H}-3,30-31$ | 83.10 | 85.21 | 15X-2, 30-31 | 128.60 | 85.30 |
| $10 \mathrm{H}-3,60-62$ | 83.40 | 94.74 | 15X-2, 92-93 | 129.22 | 87.26 |
| 10H-3, 100-101 | 83.80 | 86.89 | 15X-2, 100-102 | 129.30 | 87.56 |
| $10 \mathrm{H}-4,30-31$ | 84.60 | 87.08 | 15X-3, 30-31 | 130.10 | 83.56 |
| $10 \mathrm{H}-4,100-101$ | 85.30 | 88.12 | 15X-3, 92-93 | 130.72 | 86.58 |
| 10H-5, 30-31 | 86.10 | 89.11 | 15X-4, 26-28 | 131.56 | 86.78 |
| 10H-5, 60-62 | 86.40 | 86.68 | 15X-4, 30-31 | 131.60 | 85.69 |
| $10 \mathrm{H}-5,100-101$ | 86.80 | 89.97 | 15X-4, 92-93 | 132.22 | 87.18 |
| 10H-6, 30-31 | 87.60 | 90.96 | 15X-5, 30-31 | 133.10 | 87.50 |
| $10 \mathrm{H}-6,100-101$ | 88.30 | 86.52 | 15X-5, 92-93 | 133.72 | 85.20 |
| 10H-7, 30-31 | 89.10 | 91.30 | 15X-6, 22-24 | 134.19 | 89.14 |
| $10 \mathrm{H}-7,60-62$ | 89.40 | 89.99 | 15X-6, 30-31 | 134.27 | 84.59 |
| $11 \mathrm{H}-1,30-31$ | 89.80 | 89.85 | $16 \mathrm{X}-1,30-31$ | 136.70 | 85.44 |
| $11 \mathrm{H}-1,100-101$ | 90.50 | 90.98 | $16 \mathrm{X}-1,100-101$ | 137.40 | 83.25 |
| $11 \mathrm{H}-2,30-31$ | 91.30 | 90.95 | 16X-2, 30-31 | 138.20 | 86.09 |

Table 5 (continued).
$115-714 \mathrm{~A}$ -
$115-714 \mathrm{~A}$

| 16X-2, 100-101 | 138.90 | 83.35 |
| :---: | :---: | :---: |
| 16X-3, 30-31 | 139.70 | 78.83 |
| 16X-3, 50-52 | 139.90 | 86.42 |
| 16X-3, 100-101 | 140.40 | 83.26 |
| 16X-4, 30-31 | 141.20 | 80.73 |
| 16X-4, 100-101 | 141.90 | 84.59 |
| 16X-5, 30-31 | 142.70 | 84.44 |
| 16X-5, 100-101 | 143.40 | 83.53 |
| 16X-6, 25-27 | 144.60 | 83.75 |
| 17X-1, 30-31 | 146.40 | 84.82 |
| 17X-1, 100-101 | 147.10 | 85.60 |
| 17X-2, 30-31 | 147.90 | 86.17 |
| 17X-2, 92-94 | 148.52 | 85.22 |
| 17X-2, 100-101 | 148.60 | 85.33 |
| 17X-3, 30-31 | 149.40 | 85.45 |
| $17 \mathrm{X}-3,100-101$ | 150.10 | 86.50 |
| 17X-4, 30-31 | 150.90 | 87.95 |
| 17X-4, 100-101 | 151.60 | 82.09 |
| 17X-4, 148-150 | 152.08 | 85.96 |
| 17X-5, 30-31 | 152.40 | 87.07 |
| 18X-1, 30-31 | 156.00 | 87.76 |
| 18X-1, 100-101 | 156.70 | 84.35 |
| 18X-2, 30-31 | 157.50 | 85.69 |
| 18X-2, 100-101 | 158.20 | 78.50 |
| 18X-3, 8-10 | 158.78 | 88.40 |
| 18X-3, 30-31 | 159.00 | 87.42 |
| 18X-3, 100-101 | 159.70 | 81.00 |
| 18X-4, 30-31 | 160.50 | 82.82 |
| 18X-4, 100-101 | 161.20 | 88.14 |
| 18X-5, 10-12 | 161.80 | 82.93 |
| 18X-5, 30-31 | 162.00 | 81.82 |
| 18X-5, 100-101 | 162.70 | 88.80 |
| 18X-6, 30-31 | 163.50 | 88.30 |
| 18X-7, 30-31 | 163.82 | 86.02 |
| 19X-1, 30-31 | 165.60 | 86.51 |
| $19 \mathrm{X}-1,100-101$ | 166.30 | 84.80 |
| 19X-2, 30-31 | 167.10 | 86.54 |
| 19X-2, 52-54 | 167.32 | 88.01 |
| 19X-2, 100-101 | 167.80 | 85.00 |
| 19X-3, 30-31 | 168.60 | 83.17 |
| $19 \mathrm{X}-3,100-101$ | 169.30 | 87.44 |
| 19X-4, 30-31 | 170.10 | 85.61 |
| 19X-4, 100-101 | 170.80 | 88.63 |
| 19X-5, 30-31 | 171.60 | 90.78 |
| 19X-5, 56-58 | 171.86 | 88.77 |
| 19X-5, 100-101 | 172.30 | 89.49 |
| 19X-6, 30-31 | 173.10 | 90.71 |
| 19X-6, 100-101 | 173.80 | 90.94 |
| 20X-1, 30-31 | 175.30 | 86.68 |
| 20X-1, 100-101 | 176.00 | 87.25 |
| 20X-2, 30-31 | 176.80 | 89.53 |
| 20X-2, 60-62 | 177.10 | 92.58 |
| 20X-2, 100-101 | 177.50 | 86.91 |
| 20X-3, 30-31 | 178.30 | 86.71 |
| 20X-3, 100-101 | 179.00 | 88.46 |
| 20X-4, 30-31 | 179.80 | 76.81 |
| $20 \mathrm{X}-4,100-101$ | 180.50 | 87.35 |
| 20X-5, 30-31 | 181.30 | 85.81 |
| 20X-5, 64-66 | 181.64 | 88.14 |
| 20X-5, 100-101 | 182.00 | 87.86 |
| 21X-1, 33-34 | 185.03 | 86.05 |
| 21X, CC, 2-4 | 185.49 | 60.72 |
| 22X-1, 30-31 | 194.70 | 87.36 |
| 22X-1, 100-101 | 195.40 | 87.35 |
| 22X-2, 30-31 | 196.20 | 89.19 |
| 22X-2, 60-62 | 196.50 | 88.21 |
| 22X-2, 100-101 | 196.90 | 88.46 |
| 22X-3, 30-31 | 197.70 | 88.17 |
| 22X-3, 100-101 | 198.40 | 89.09 |
| 22X-4, 14-15 | 199.04 | 91.60 |
| 22X-4, 60-62 | 199.50 | 91.73 |
| 22X-4, 100-101 | 199.90 | 90.05 |
| 22X-5, 30-31 | 200.70 | 90.46 |

Table 5 (continued).

| Sample interval <br> $(\mathrm{cm})$ | Depth <br> $(\mathrm{mbsf})$ | Carbonate <br> $(\mathrm{wt} \%)$ |
| :---: | :---: | :---: | :---: |
| $115-714 \mathrm{~A}-$ |  |  |
| 22X-6, 54-56 | 202.40 | 91.03 |
| 23X-1, 30-31 | 204.40 | 87.77 |
| 23X-1, 100-101 | 205.10 | 87.77 |
| 23X-2, 30-31 | 205.90 | 87.20 |
| 23X-2, 63-65 | 206.23 | 91.16 |
| 23X-2, 100-101 | 206.60 | 87.60 |
| 23X-3, 30-31 | 207.40 | 86.36 |
| 23X-3, 62-65 | 207.72 | 90.46 |
| 23X-3, 100-101 | 208.10 | 89.09 |
| 23X-4, 30-31 | 208.90 | 89.44 |
| 23X-4, 100-101 | 209.60 | 88.76 |
| 23X-5, 30-31 | 210.40 | 89.40 |
| 23X-5, 100-101 | 211.10 | 88.81 |
| 23X-6, 30-31 | 211.90 | 85.65 |
| 23X-6, 63-65 | 212.23 | 88.42 |
| 23X-6, 100-101 | 212.60 | 89.61 |
| 24X-1, 30-31 | 214.10 | 86.94 |
| 24X-1, 100-101 | 214.80 | 88.57 |
| 25X-1, 30-31 | 223.60 | 89.99 |
| 25X-1, 100-101 | 224.30 | 89.34 |
| 25X-, 30-31 | 226.60 | 89.97 |



Figure 11. Carbonate and organic carbon content of sediments from Hole 714A.


Figure 12. Smoothed variations in carbonate content, Hole 714A.

Table 6. Organic carbon analyses, Hole 714A.

| Sample interval <br> (cm) | Depth <br> (mbsf) | Organic <br> carbon <br> $(\%)$ |
| :---: | :---: | :---: |
| 115-714A- |  |  |
| 1H-1, 30-31 | 0.30 | 0.18 |
| 1H-1, 100-101 | 1.00 | 0.78 |
| 1H-2, 100-101 | 2.50 | 0.27 |
| 2H-1, 30-31 | 3.10 | 0.46 |
| 2H-2, 100-101 | 5.30 | 0.35 |
| 2H-3, 100-101 | 6.80 | 0.24 |
| 2H-4, 30-31 | 7.60 | 0.28 |
| 2H-5, 30-31 | 9.10 | 0.71 |
| 2H-6, 30-31 | 10.60 | 0.41 |
| 2H-7, 30-31 | 12.10 | 0.33 |
| 3H-1, 100-101 | 13.40 | 0.39 |
| 3H-2, 30-31 | 14.20 | 0.64 |
| 3H-3, 100-101 | 16.40 | 0.59 |
| 3H-4, 30-31 | 17.20 | 0.30 |
| 3H-5, 30-31 | 18.70 | 0.62 |
| 3H-6, 30-31 | 20.20 | 0.34 |
| 3H-7, 30-31 | 21.70 | 0.06 |
| 4H-3, 30-31 | 23.93 | 0.21 |
| 5H-3, 30-31 | 35.00 | 0 |
| 6H-3, 30-31 | 44.70 | 0.05 |
| 7H-3, 30-31 | 54.30 | 0.16 |
| 8H-3, 100-101 | 64.60 | 0 |
| 9H-3, 100-101 | 74.20 | 0.15 |
| 10H-3, 100-101 | 83.80 | 0 |
| 11H-3, 100-101 | 93.50 | 0 |
| 12H-3, 30-31 | 102.40 | 0 |
| 13H-3, 30-31 | 112.00 | 0 |
| 14X-3, 30-31 | 121.70 | 0 |
| 15X-3, 30-31 | 130.10 | 0 |
| 16X-3, 30-31 | 139.70 | 0.10 |
| 17X-3, 100-101 | 150.10 | 0 |
| 18X-3, 30-31 | 159.00 | 0 |
| 19X-3, 100-101 | 169.30 | 0.19 |
| 20X-3, 100-101 | 179.00 | 0 |
| $21 \mathrm{X}-1,33-34$ | 185.03 | 0.08 |
| 22X-3, 101-101 | 198.41 | 0.17 |
| 23X-3, 100-101 | 208.10 | 0 |
| 24X-1, 100-101 | 214.80 | 0 |
| 25X-1, 100-101 | 224.30 | 0 |
|  |  |  |

Table 7. Concentrations of methane $\left(\mathrm{CH}_{4}\right)$, Hole 714A.

| Sample interval <br> $(\mathrm{cm})$ | Depth <br> $(\mathrm{mbsf})$ | Methane <br> $(\mathrm{ppm})$ |
| :--- | ---: | :--- |
| $115-714 \mathrm{~A}-$ |  |  |
|  |  |  |
| $1 \mathrm{H}-1,145-150$ | 1.45 | 2.0 |
| $3 \mathrm{H}-4,145-150$ | 18.35 | 2.1 |
| $4 \mathrm{H}-5,145-150$ | 28.08 | 2.6 |
| 5H-4, 145-150 | 37.65 | 2.86 |
| $6 \mathrm{H}-4,145-150$ | 47.35 | 4.06 |
| $12 \mathrm{H}-2,120-125$ | 101.80 | 1.91 |
| $15 \mathrm{X}-4,120-125$ | 132.50 | 1.95 |
| $22 \mathrm{X}-4,120-125$ | 200.10 | 2.6 |



Figure 14. Index properties (wet-bulk density, water content, porosity, and grain density) at Hole 714A.


Figure 15. Index properties (wet-bulk density, water content, porosity, and grain density) at Hole 714B.

Table 8. Index-properties data, Site 714.

| Section interval (cm) | Depth (mbsf) | Water content (\%) | Porosity (\%) | Wet-bulk density ( $\mathrm{g} / \mathrm{cm}^{3}$ ) | $\begin{aligned} & \text { Dry-bulk } \\ & \text { density } \\ & \left(\mathrm{g} / \mathrm{cm}^{3}\right) \end{aligned}$ | Grain density ( $\mathrm{g} / \mathrm{cm}^{3}$ ) | $\begin{aligned} & \text { Carbonate } \\ & \text { content } \\ & (\mathrm{wt} \%) \end{aligned}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 115-714A- |  |  |  |  |  |  |  |
| 1H-2, 71 | 2.21 | 47.98 | 72.47 | 1.56 | 0.81 | 2.89 | 67.98 |
| 2H-1, 60 | 3.40 | 45.53 | 71.86 | 1.62 | 0.88 | 2.90 | 80.70 |
| $2 \mathrm{H}-3,100$ | 6.80 | 45.99 | 71.27 | 1.59 | 0.86 | 2.95 | 83.42 |
| $2 \mathrm{H}-5,100$ | 9.80 | 49.94 | 73.58 | 1.54 | 0.77 | 2.82 | 72.44 |
| $2 \mathrm{H}-6,100$ | 11.30 | 46.55 | 71.76 | 1.62 | 0.86 | 2.95 | 83.24 |
| $3 \mathrm{H}-1,60$ | 13.00 | 47.98 | 72.77 | 1.60 | 0.83 | 2.93 | 81.05 |
| $3 \mathrm{H}-3,60$ | 16.00 | 48.79 | 73.35 | 1.54 | 0.79 | 2.92 | 73.65 |
| $3 \mathrm{H}-4,60$ | 17.50 | 45.82 | 70.76 | 1.59 | 0.86 | 2.89 | 80.50 |
| $3 \mathrm{H}-5,60$ | 19.00 | 52.46 | 75.66 | 1.47 | 0.70 | 2.84 | 83.64 |
| 3H-6, 60 | 20.50 | 48.71 | 72.95 | 1.56 | 0.80 | 2.87 | 88.56 |
| 4H-2, 60 | 22.78 | 47.76 | 72.15 | 1.58 | 0.83 | 2.86 | 89.70 |
| 4H-4, 86 | 25.99 | 48.74 | 73.57 | 1.57 | 0.81 | 2.96 | 87.60 |
| 4H-6, 88 | 29.01 | 45.76 | 71.00 | 1.62 | 0.88 | 2.93 | 81.60 |
| $5 \mathrm{H}-2,60$ | 33.80 | 47.33 | 72.54 | 1.56 | 0.82 | 2.97 | 78.03 |
| 5H-4, 60 | 36.80 | 46.32 | 71.59 | 1.57 | 0.84 | 2.95 | 79.54 |
| 5H-6, 60 | 39.80 | 46.61 | 70.88 | 1.57 | 0.84 | 2.82 | 82.67 |
| 6H-2, 60 | 43.50 | 45.33 | 71.84 | 1.64 | 0.89 | 2.85 | 82.88 |
| $6 \mathrm{H}-4,60$ | 46.50 | 40.73 | 66.32 | 1.68 | 1.00 | 2.90 | 86.18 |
| 6H-6, 60 | 49.50 | 42.68 | 66.02 | 1.65 | 0.95 | 2.64 | 82.35 |
| 7H-2, 60 | 53.10 | 44.24 | 68.93 | 1.63 | 0.91 | 2.83 | 85.93 |
| 7H-4, 60 | 56.10 | 48.02 | 72.12 | 1.55 | 0.81 | 2.83 | 81.43 |
| 7H-6, 56 | 59.06 | 40.69 | 66.75 | 1.67 | 0.99 | 2.96 | 85.32 |
| 8H-2, 60 | 62.70 | 44.79 | 70.07 | 1.60 | 0.88 | 2.92 | 86.18 |
| 8H-4, 60 | 65.70 | 41.31 | 67.89 | 1.48 | 0.87 | 2.86 | 87.74 |
| 8H-6, 60 | 68.70 | 41.85 | 67.96 | 1.71 | 0.99 | 2.98 | 88.61 |
| $9 \mathrm{H}-2,60$ | 72.30 | 45.19 | 69.26 | 1.63 | 0.89 | 2.76 | 81.26 |
| $9 \mathrm{H}-3,75$ | 73.95 | 41.11 | 66.15 | 1.66 | 0.98 | 2.83 | 86.58 |
| $10 \mathrm{H}-2,60$ | 81.90 | 38.08 | 63.91 | 1.73 | 1.07 | 2.92 | 86.31 |
| $10 \mathrm{H}-3,60$ | 83.40 | 41.03 | 66.87 | 1.68 | 0.99 | 2.94 | 84.74 |
| $10 \mathrm{H}-5,60$ | 86.40 | 39.82 | 65.83 | 1.80 | 1.08 | 2.95 | 86.68 |
| $10 \mathrm{H}-7,60$ | 89.40 | 42.07 | 66.36 | 1.72 | 1.00 | 2.75 | 89.99 |
| $11 \mathrm{H}-3,70$ | 93.20 | 43.32 | 68.47 | 1.71 | 0.97 | 2.87 | 92.35 |
| $11 \mathrm{H}-6,52$ | 97.52 | 46.41 | 70.03 | 1.67 | 0.89 | 2.73 | 88.21 |
| $12 \mathrm{H}-2,60$ | 101.20 | 42.96 | 66.56 | 1.68 | 0.96 | 2.67 | 92.14 |
| 12H-4, 61 | 104.21 | 42.67 | 67.86 | 1.71 | 0.98 | 2.87 | 94.73 |
| $12 \mathrm{H}-6,63$ | 107.23 | 43.73 | 68.24 | 1.70 | 0.96 | 2.80 | 89.54 |
| $13 \mathrm{H}-2,58$ | 110.78 | 42.48 | 67.23 | 1.69 | 0.97 | 2.81 | 88.10 |
| $13 \mathrm{H}-4,65$ | 113.85 | 40.69 | 65.45 | 1.73 | 1.03 | 2.80 | 87.03 |
| 13H-6, 65 | 116.85 | 43.37 | 66.60 | 1.66 | 0.94 | 2.63 | 89.49 |
| 14X-2, 63 | 120.53 | 47.01 | 70.55 | 1.62 | 0.86 | 2.73 | 88.81 |
| $14 \mathrm{X}-4,63$ | 123.53 | 45.23 | 69.47 | 1.65 | 0.91 | 2.79 | 87.40 |
| 14X-6, 54 | 126.44 | 44.11 | 68.25 | 1.65 | 0.92 | 2.75 | 85.91 |
| 15X-2, 100 | 129.30 | 45.60 | 71.36 | 1.67 | 0.91 | 2.73 | 87.56 |
| 15X-4, 26 | 131.56 | 42.82 | 66.78 | 1.70 | 0.97 | 2.72 | 86.78 |
| 15X-6, 22 | 134.19 | 42.17 | 65.96 | 1.70 | 0.98 | 2.69 | 89.14 |
| $16 \mathrm{X}-3,58$ | 139.98 | 44.30 | 68.16 | 1.66 | 0.92 | 2.72 | 86.42 |
| 16X-5, 25 | 142.65 | 43.56 | 67.78 | 1.68 | 0.95 | 2.76 | 83.75 |
| $17 \mathrm{X}-2,92$ | 148.52 | 45.61 | 68.69 | 1.60 | 0.87 | 2.64 | 85.22 |
| 17X-4, 148 | 152.08 | 43.35 | 68.09 | 1.74 | 0.99 | 2.82 | 85.96 |
| 18X-3, 8 | 158.78 | 44.61 | 69.90 | 1.69 | 0.94 | 2.92 | 88.40 |
| 18X-5, 10 | 161.80 | 43.91 | 68.36 | 1.73 | 0.97 | 2.79 | 82.93 |
| 19X-2, 52 | 167.32 | 44.64 | 69.64 | 1.78 | 0.98 | 2.88 | 88.01 |
| 19X-5, 56 | 171.86 | 43.38 | 66.56 | 1.90 | 1.08 | 2.63 | 88.77 |
| 20X-2, 60 | 177.10 | 45.38 | 71.03 | 1.78 | 0.97 | 2.99 | 92.58 |
| 20X-5, 64 | 181.64 | 42.18 | 66.84 | 1.78 | 1.03 | 2.80 | 88.14 |
| $21 \mathrm{X}, \mathrm{CC}, 2$ | 185.49 | 12.32 | 20.56 | 1.83 | 1.61 | 1.87 | 60.72 |
| 22X-2, 60 | 196.50 | 45.28 | 69.24 | 1.66 | 0.91 | 2.75 | 88.21 |
| $22 \mathrm{X}-4,60$ | 199.50 | 41.20 | 66.13 | 1.74 | 1.03 | 2.82 | 91.73 |
| 22X-5, 54 | 201.44 | 39.48 | 64.75 | 1.83 | 1.11 | 2.85 | 91.03 |
| $23 \mathrm{X}-2,63$ | 206.23 | 39.24 | 64.82 | 1.82 | 1.11 | 2.89 | 91.16 |
| $23 \mathrm{X}-3,63$ | 207.73 | 38.75 | 63.26 | 1.83 | 1.12 | 2.76 | 90.46 |
| 23X-6, 63 | 212.23 | 40.90 | 68.33 | 2.01 | 1.19 | 2.80 | 88.82 |
| 24X-1, 60 | 214.40 | 42.40 | 67.69 | 1.73 | 0.99 | 2.88 | 86.94 |
| 24X, CC, 21 | 215.51 | 34.50 | 59.33 | 1.95 | 1.27 | 2.81 | 88.57 |
| 24X, CC, 24 | 215.54 | 8.53 | 20.45 | 2.79 | 2.55 | 2.80 | 89.99 |
| 25X-1, 60 | 223.90 | 40.04 | 65.56 | 1.87 | 1.12 | 2.89 | 89.34 |
| $25 \mathrm{X}, \mathrm{CC}, 20$ | 226.50 | 37.92 | 62.69 | 2.07 | 1.28 | 2.78 | 89.97 |
| $115-714 \mathrm{~B}$ - |  |  |  |  |  |  |  |
| 1H-2, 60 | 2.10 | 55.52 | 76.83 | 1.49 | 0.66 | 2.68 | 71.06 |
| 1H-6, 60 | 8.10 | 50.85 | 72.85 | 1.60 | 0.79 | 2.62 | 79.97 |
| $3 \mathrm{H}-3,60$ | 20.00 | 51.70 | 74.04 | 1.52 | 0.74 | 2.69 | 79.04 |
| 3H-6, 60 | 24.50 | 50.33 | 74.03 | 1.59 | 0.79 | 2.84 | 83.85 |
| 4H-3, 60 | 29.60 | 49.91 | 73.22 | 1.51 | 0.76 | 2.77 | 80.70 |
| 4H-6, 60 | 34.10 | 49.22 | 71.12 | 1.67 | 0.85 | 2.56 | 82.07 |

Table 8 (continued).

| Section interval (cm) | Depth <br> (mbsf) | Water content (\%) | Porosity (\%) | Wet-bulk density ( $\mathrm{g} / \mathrm{cm}^{3}$ ) | Dry-bulk density $\left(\mathrm{g} / \mathrm{cm}^{3}\right)$ | Grain density ( $\mathrm{g} / \mathrm{cm}^{3}$ ) | Carbonate content (wt\%) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 115-714B- |  |  |  |  |  |  |  |
| 5H-3, 60 | 39.30 | 50.27 | 72.88 | 1.59 | 0.79 | 2.68 | 80.67 |
| $5 \mathrm{H}-6,32$ | 43.52 | 47.20 | 71.46 | 1.61 | 0.85 | 2.83 | 83.04 |
| $6 \mathrm{H}-3,60$ | 49.00 | 47.38 | 71.05 | 1.62 | 0.85 | 2.76 | 77.69 |
| 6H-6, 60 | 53.50 | 47.44 | 70.46 | 1.61 | 0.85 | 2.67 | 82.48 |
| 7H-3, 60 | 58.60 | 48.13 | 71.50 | 1.55 | 0.80 | 2.73 | 79.07 |
| 7H-6, 60 | 63.10 | 40.78 | 65.60 | 1.65 | 0.98 | 2.80 | 85.41 |
| $8 \mathrm{H}-3,60$ | 68.20 | 43.31 | 67.70 | 1.65 | 0.93 | 2.78 | 85.11 |
| 8H-6, 60 | 72.70 | 41.94 | 65.44 | 1.64 | 0.95 | 2.65 | 84.93 |
| $9 \mathrm{H}-3,71$ | 78.01 | 39.40 | 63.79 | 1.65 | 1.00 | 2.74 | 84.50 |
| 9H-6, 64 | 82.44 | 41.26 | 66.20 | 1.63 | 0.96 | 2.82 | 84.41 |
| $10 \mathrm{H}-3,85$ | 87.75 | 42.12 | 65.77 | 1.60 | 0.93 | 2.67 | 84.45 |
| 10H-6, 67 | 92.07 | 40.01 | 64.07 | 1.66 | 0.99 | 2.71 | 89.01 |
| $11 \mathrm{H}-3,68$ | 97.28 | 43.23 | 67.76 | 1.59 | 0.90 | 2.79 | 82.21 |
| $11 \mathrm{H}-6,76$ | 101.86 | 43.29 | 67.40 | 1.59 | 0.90 | 2.74 | 85.64 |
| $12 \mathrm{H}-3,71$ | 106.91 | 43.94 | 68.00 | 1.61 | 0.90 | 2.74 | 87.36 |
| $12 \mathrm{H}-6,92$ | 111.62 | 41.55 | 65.16 | 1.65 | 0.96 | 2.66 | 88.52 |
| $13 \mathrm{H}-3,76$ | 116.66 | 43.02 | 67.35 | 1.60 | 0.91 | 2.76 | 82.34 |
| 13H-6, 97 | 121.37 | 45.03 | 68.65 | 1.56 | 0.86 | 2.70 | 83.96 |



Figure 16. Carbonate content of samples for which the index properties were measured at Site 714.


Figure 17. Compressional-wave velocities measured from discrete samples at Site 714.

Table 9. Compressional-wave velocity and acoustic impedance data, Site 714.

| Section interval (cm) | $\begin{aligned} & \text { Depth } \\ & \text { (mbsf) } \end{aligned}$ | $\begin{gathered} \mathrm{V}_{p} \\ (\mathrm{~m} / \mathrm{s}) \end{gathered}$ | $\begin{gathered} \text { Impedance } \\ \left(\mathrm{g} / \mathrm{cm}^{2} \cdot \mathrm{~s} \cdot 10^{4}\right) \end{gathered}$ |
| :---: | :---: | :---: | :---: |
| 115-714A- |  |  |  |
| 2H-1, 60 | 3.40 | 1507 | 24.41 |
| $2 \mathrm{H}-3,100$ | 6.80 | 1505 | 23.93 |
| $2 \mathrm{H}-5,100$ | 9.80 | 1508 | 23.22 |
| $2 \mathrm{H}-6,100$ | 11.30 | 1485 | 24.06 |
| $3 \mathrm{H}-1,60$ | 13.00 | 1500 | 24.00 |
| $3 \mathrm{H}-3,60$ | 16.00 | 1499 | 23.08 |
| 3H-5, 60 | 19.00 | 1541 | 22.62 |
| 3H-6, 60 | 20.50 | 1562 | 24.36 |
| 4H-2, 60 | 22.78 | 1524 | 24.08 |
| 4H-4, 86 | 25.99 | 1532 | 24.06 |
| 4H-6, 88 | 29.01 | 1525 | 24.40 |
| 5H-2, 60 | 33.80 | 1521 | 23.72 |
| $5 \mathrm{H}-4,60$ | 36.80 | 1549 | 24.32 |
| 6H-2, 60 | 43.50 | 1539 | 25.24 |
| $6 \mathrm{H}-4,60$ | 46.50 | 1582 | 26.58 |
| 6H-6, 60 | 49.50 | 1579 | 26.05 |
| 7H-2, 60 | 53.10 | 1578 | 25.72 |
| 7H-4, 60 | 56.10 | 1568 | 24.30 |
| 7H-6, 56 | 59.06 | 1565 | 26.13 |
| $8 \mathrm{H}-2,60$ | 62.70 | 1582 | 25.31 |
| 8H-4, 60 | 65.70 | 1548 | 22.91 |
| 8H-6, 60 | 68.70 | 1574 | 26.92 |
| $9 \mathrm{H}-1,60$ | 70.80 | 1521 | 24.79 |
| $10 \mathrm{H}-2,60$ | 81.90 | 1593 | 27.56 |
| $10 \mathrm{H}-4,60$ | 84.90 | 1549 | 26.03 |
| $10 \mathrm{H}-5,60$ | 86.40 | 1552 | 27.94 |
| $12 \mathrm{H}-2,60$ | 101.20 | 1554 | 26.11 |
| $12 \mathrm{H}-4,61$ | 104.21 | 1535 | 26.25 |
| $13 \mathrm{H}-2,58$ | 110.78 | 1541 | 26.04 |
| $13 \mathrm{H}-6,65$ | 116.85 | 1569 | 26.04 |
| 14X-3, 54 | 121.94 | 1618 | 26.68 |
| 15X-2, 100 | 129.30 | 1654 | 27.62 |
| 15X-4, 26 | 131.56 | 1649 | 28.03 |
| 15X-6, 22 | 134.19 | 1642 | 27.91 |
| $16 \mathrm{X}-3,58$ | 139.98 | 1681 | 27.95 |
| $17 \mathrm{X}-2,92$ | 148.52 | 1584 | 25.34 |
| $17 \mathrm{X}-4,148$ | 152.08 | 1744 | 30.34 |
| 18X-3, 8 | 158.78 | 1767 | 29.86 |
| 18X-5, 12 | 161.82 | 1682 | 29.10 |
| 19X-2, 50 | 167.30 | 1660 | 29.55 |
| 19X-5, 56 | 171.86 | 1559 | 27.75 |
| 20X-2, 60 | 177.10 | 1680 | 29.90 |
| 21X, CC, 2 | 185.49 | 1705 | 31.20 |
| 22X-2, 60 | 196.50 | 1633 | 27.11 |
| 23X-2, 63 | 206.23 | 1619 | 29.47 |
| 23X-3, 63 | 207.73 | 1663 | 30.43 |
| 24X, CC, 21 | 215.51 | 1779 | 34.69 |
| 24X, CC, 23 | 215.53 | 4787 | 133.50 |
| 115-714B- |  |  |  |
| 1H-2, 60 | 2.10 | 1516 | 22.59 |
| 1H-6, 60 | 8.10 | 1532 | 24.51 |
| $3 \mathrm{H}-3,60$ | 20.00 | 1544 | 23.47 |
| 3H-6, 60 | 24.50 | 1520 | 24.17 |
| 4H-3, 60 | 29.60 | 1539 | 23.24 |
| 4H-6, 60 | 34.10 | 1533 | 25.60 |
| $5 \mathrm{H}-3,60$ | 39.30 | 1520 | 24.17 |
| 5H-6, 32 | 43.52 | 1540 | 24.79 |
| $6 \mathrm{H}-3,60$ | 49.00 | 1584 | 25.66 |
| 6H-6, 60 | 53.50 | 1567 | 25.23 |
| 7H-3, 60 | 58.60 | 1563 | 24.22 |
| 7H-6, 60 | 63.10 | 1594 | 26.30 |
| $8 \mathrm{H}-3,60$ | 68.20 | 1560 | 25.74 |
| 8H-6, 60 | 72.70 | 1560 | 25.58 |
| 9H-6, 64 | 82.44 | 1529 | 24.92 |
| $10 \mathrm{H}-3,85$ | 87.75 | 1551 | 24.82 |
| 10H-6, 67 | 92.07 | 1589 | 26.38 |
| $11 \mathrm{H}-3,68$ | 97.28 | 1529 | 24.31 |



Figure 18. Acoustic impedance calculated from wet-bulk density and compressional-wave velocity of discrete samples at Site 714.


Figure 19. Continuous $P$-wave logger, computed acoustic impedance, and GRAPE wet-bulk density records for Hole 714B.

Table 10. Shear-wave velocity data, Hole 714A.


Figure 20. Shear-wave velocity at Hole 714A.

Table 11. Motorized shear strength data, Site 714.

| Section interval (cm) | Depth <br> (mbsf) | Peak $(\mathrm{kPa})$ |
| :---: | :---: | :---: |
| 115-714A- |  |  |
| $2 \mathrm{H}-2,102$ | 5.32 | 4.7 |
| 2H-4, 96 | 8.26 | 2.3 |
| 2H-6, 61 | 10.91 | 5.2 |
| $3 \mathrm{H}-2,80$ | 14.70 | 4.1 |
| $3 \mathrm{H}-4,20$ | 17.10 | 6.4 |
| 3H-6, 93 | 20.83 | 12.2 |
| $4 \mathrm{H}-3,64$ | 24.27 | 10.5 |
| $4 \mathrm{H}-5,62$ | 27.25 | 17.5 |
| 4H-7, 63 | 30.26 | 16.9 |
| 5H-2, 56 | 33.76 | 16.3 |
| $5 \mathrm{H}-4,67$ | 36.87 | 19.8 |
| 5H-6, 57 | 39.77 | 20.9 |
| $6 \mathrm{H}-2,56$ | 43.46 | 11.1 |
| 6H-4, 58 | 46.48 | 2.9 |
| 6H-6, 54 | 49.44 | 27.9 |
| $7 \mathrm{H}-2,56$ | 53.06 | 14.0 |
| $7 \mathrm{H}-2,66$ | 53.16 | 23.3 |
| 7H-4, 60 | 56.10 | 20.9 |
| $8 \mathrm{H}-2,57$ | 62.67 | 12.3 |
| $8 \mathrm{H}-4,60$ | 65.70 | 35.5 |
| 8H-6, 59 | 68.69 | 38.4 |
| $9 \mathrm{H}-1,66$ | 70.86 | 20.4 |
| 9H-2, 66 | 72.36 | 11.1 |
| $9 \mathrm{H}-3,66$ | 73.86 | 19.2 |
| $10 \mathrm{H}-6,63$ | 87.93 | 26.8 |
| $10 \mathrm{H}-4,59$ | 84.89 | 18.6 |
| $10 \mathrm{H}-2,59$ | 81.89 | 21.5 |
| $11 \mathrm{H}-6,45$ | 97.45 | 40.7 |
| $11 \mathrm{H}-3,54$ | 93.04 | 29.7 |
| $11 \mathrm{H}-3,60$ | 93.10 | 30.3 |
| $12 \mathrm{H}-2,54$ | 101.14 | 39.6 |
| $12 \mathrm{H}-4,57$ | 104.17 | 33.7 |
| 12H-6, 59 | 107.19 | 69.8 |
| $13 \mathrm{H}-2,47$ | 110.67 | 42.5 |
| $13 \mathrm{H}-4,55$ | 113.75 | 33.7 |
| $13 \mathrm{H}-4,61$ | 113.81 | 42.5 |
| $13 \mathrm{H}-6,54$ | 116.74 | 64.0 |
| 14X-2, 57 | 120.47 | 44.8 |
| 115-714B- |  |  |
| 1H-2, 84 | 2.34 | 3.5 |
| 1H-6, 84 | 8.34 | 3.5 |
| $3 \mathrm{H}-1,128$ | 17.68 | 23.3 |
| $3 \mathrm{H}-1,140$ | 17.80 | 14.0 |
| 3H-4, 86 | 21.76 | 13.4 |
| 4H-2, 68 | 28.18 | 18.6 |
| 4H-6, 68 | 34.18 | 17.8 |
| $5 \mathrm{H}-3,68$ | 39.38 | 11.6 |
| 5H-5, 68 | 42.38 | 11.9 |
| $6 \mathrm{H}-3,72$ | 49.12 | 11.1 |
| $6 \mathrm{H}-6,72$ | 53.62 | 28.2 |
| 7H-3, 73 | 58.73 | 9.3 |
| 7H-6, 72 | 63.22 | 25.0 |
| $8 \mathrm{H}-3,72$ | 68.32 | 22.7 |
| 8H-6, 72 | 72.82 | 32.6 |
| $9 \mathrm{H}-3,67$ | 77.97 | 37.2 |
| 9H-6, 71 | 82.51 | 27.3 |
| $10 \mathrm{H}-3,82$ | 87.72 | 160.0 |
| 10H-6, 63 | 92.03 | 26.8 |
| $11 \mathrm{H}-3,73$ | 97.33 | 36.7 |
| $11 \mathrm{H}-6,72$ | 101.82 | 51.8 |
| $12 \mathrm{H}-3,66$ | 106.86 | 39.6 |
| $12 \mathrm{H}-6,68$ | 111.38 | 42.5 |
| $12 \mathrm{H}-6,72$ | 111.42 | 21.5 |
| $13 \mathrm{H}-4,71$ | 118.11 | 46.0 |
| $13 \mathrm{H}-6,94$ | 121.34 | 35.2 |



Figure 21. Shear strength at Site 714.


Figure 22. Thermal conductivity at Hole 714A.

Table 12. Thermal conductivity data, Site 714.

| Section <br> interval <br> (cm) | Depth <br> (mbsf) | Thermal <br> (conductivity <br> (W-m |
| :---: | :---: | :---: |
| $115-714 \mathrm{~K}$ |  |  | $\mathrm{~K}^{-1}$ )

Table 12 (continued).

| Section <br> interval <br> $(\mathrm{cm})$ | Depth <br> $(\mathrm{mbsf})$ | Thermal <br> conductivity <br> $\left(\mathrm{W} \cdot \mathrm{m}^{-1} \cdot \mathrm{~K}^{-1}\right)$ |
| :---: | :---: | :---: |

115-714A-(cont.)

| 6H-5, 80 |  |  |
| :--- | ---: | :--- |
| 68.20 | 1.217 |  |
| $6 \mathrm{H}-6,80$ | 49.70 | 1.158 |
| $6 \mathrm{H}-70$ | 50.80 | 1.247 |
| $7 \mathrm{H}-2,80$ | 53.30 | 1.189 |
| $7 \mathrm{H}-4,80$ | 56.30 | 1.113 |
| $7 \mathrm{H}-5,80$ | 57.80 | 1.108 |
| $7 \mathrm{H}-6,80$ | 59.30 | 0.974 |
| $7 \mathrm{H}-7,40$ | 60.40 | 1.134 |
| $8 \mathrm{H}-1,80$ | 61.40 | 1.168 |
| $8 \mathrm{H}-2,80$ | 62.90 | 1.192 |
| $8 \mathrm{H}-3,80$ | 64.40 | 1.253 |
| $8 \mathrm{H}-5,80$ | 67.40 | 1.254 |
| $8 \mathrm{H}-6,80$ | 68.90 | 1.260 |
| $8 \mathrm{H}-7,40$ | 70.00 | 1.145 |
| $9 \mathrm{H}-1,80$ | 71.00 | 1.224 |
| $9 \mathrm{H}-4,18$ | 74.88 | 1.268 |
| $10 \mathrm{H}-1,80$ | 80.60 | 1.155 |
| $10 \mathrm{H}-2,80$ | 82.10 | 1.257 |
| $10 \mathrm{H}-4,80$ | 85.10 | 1.138 |
| $10 \mathrm{H}-5,80$ | 86.60 | 1.256 |
| $10 \mathrm{H}-6,80$ | 88.10 | 1.212 |
| $10 \mathrm{H}-7,40$ | 89.20 | 1.267 |
| $11 \mathrm{H}-2,80$ | 91.80 | 1.288 |
| $11 \mathrm{H}-3,80$ | 93.30 | 1.298 |
| $11 \mathrm{H}-4,80$ | 94.80 | 1.306 |
| $11 \mathrm{H}-5,10$ | 95.60 | 1.166 |
| $11 \mathrm{H}-5,20$ | 95.70 | 1.237 |
| $11 \mathrm{H}-5,30$ | 95.80 | 1.279 |
| $11 \mathrm{H}-5,40$ | 95.90 | 1.314 |
| $11 \mathrm{H}-5,60$ | 96.10 | 1.277 |
| $11 \mathrm{H}-5,70$ | 96.20 | 1.246 |
| $11 \mathrm{H}-5,80$ | 96.30 | 1.200 |
| $11 \mathrm{H}-5,90$ | 96.40 | 1.292 |
| $11 \mathrm{H}-5,100$ | 96.50 | 1.360 |
| $11 \mathrm{H}-5,120$ | 96.70 | 1.231 |
| $11 \mathrm{H}-5,130$ | 96.80 | 1.202 |
| $11 \mathrm{H}-5,140$ | 96.90 | 1.180 |
| $11 \mathrm{H}-7,15$ | 98.65 | 1.301 |
| $12 \mathrm{H}-2,80$ | 101.40 | 1.292 |
| $12 \mathrm{H}-3,80$ | 102.90 | 1.297 |
| $12 \mathrm{H}-4,80$ | 104.40 | 1.162 |
| $12 \mathrm{H}-5,80$ | 105.90 | 1.287 |
| $12 \mathrm{H}-7,40$ | 108.50 | 1.237 |
| $13 \mathrm{H}-1,80$ | 109.50 | 1.314 |
| $13 \mathrm{H}-2,80$ | 111.00 | 1.457 |
| $13 \mathrm{H}-3,80$ | 112.50 | 1.360 |
| $13 \mathrm{H}-5,80$ | 115.50 | 1.308 |
| $13 \mathrm{H}-6,80$ | 117.00 | 1.232 |
| $13 \mathrm{H}-7,40$ | 118.10 | 1.013 |
| $14 \mathrm{X}-1,80$ | 119.20 | 1.225 |
| $14 \mathrm{X}-3,80$ | 122.20 | 1.208 |
| $14 \mathrm{X}-4,80$ | 123.70 | 1.177 |
| $14 \mathrm{X}-5,95$ | 125.35 | 1.104 |
| $14 \mathrm{X}-6,80$ | 126.70 | 1.233 |
| $15 \mathrm{X}-2,80$ | 129.10 | 1.254 |
| 10 |  |  |

Table 12 (continued).

| Section interval (cm) | $\begin{aligned} & \text { Depth } \\ & \text { (mbsf) } \end{aligned}$ | $\begin{gathered} \text { Thermal } \\ \text { conductivity } \\ \left(\mathrm{W} \cdot \mathrm{~m}^{-1} \cdot \mathrm{~K}^{-1}\right. \text { ) } \end{gathered}$ |
| :---: | :---: | :---: |
| 115-714A-(cont.) |  |  |
| 15X-3, 80 | 130.60 | 1.223 |
| 15X-5, 80 | 133.60 | 1.239 |
| $16 \mathrm{X}-2,80$ | 138.70 | 1.126 |
| 16X-3, 80 | 140.20 | 1.282 |
| 16X-5, 80 | 143.20 | 1.231 |
| 17X-2, 80 | 148.40 | 1.244 |
| 17X-3, 80 | 149.90 | 1.247 |
| $17 \mathrm{X}-4,80$ | 151.40 | 1.376 |
| 17X-5, 30 | 152.40 | 1.179 |
| 18X-2, 80 | 158.00 | 1.257 |
| 18X-3, 10 | 158.80 | 1.154 |
| 18X-3, 40 | 159.10 | 1.192 |
| 18X-3, 60 | 159.30 | 1.214 |
| 18X-3, 70 | 159.40 | 1.255 |
| 18X-3, 80 | 159.50 | 1.264 |
| 18X-3, 90 | 159.60 | 1.025 |
| 18X-3, 100 | 159.70 | 1.243 |
| 18X-3, 120 | 159.90 | 1.185 |
| 18X-3, 130 | 160.00 | 1.162 |
| 18X-5, 50 | 162.20 | 1.257 |
| 19X-1, 80 | 166.10 | 1.208 |
| 19X-3, 80 | 169.10 | 1.232 |
| 19X-4, 80 | 170.60 | 1.276 |
| 19X-5, 80 | 172.10 | 0.978 |
| 19X-6, 80 | 173.60 | 1.306 |
| 20X-2, 80 | 177.30 | 1.315 |
| 20X-3, 80 | 178.80 | 1.266 |
| 20X-4, 80 | 180.30 | 0.770 |
| 20X-5, 80 | 181.80 | 1.393 |
| 21X-1, 56 | 185.26 | 1.275 |
| 22X-2, 80 | 196.70 | 1.309 |
| 22X-3, 80 | 198.20 | 1.009 |
| $22 \mathrm{X}-4,80$ | 199.70 | 1.391 |
| 23X-1, 80 | 204.90 | 1.290 |
| 23X-2, 80 | 206.40 | 0.861 |
| 23X-4, 80 | 209.40 | 1.426 |
| 23X-6, 14 | 211.74 | 1.375 |
| 24X-1, 80 | 214.60 | 1.288 |
| 25X-1, 80 | 224.10 | 1.361 |
| 25X-2, 80 | 225.60 | 1.482 |
| 115-715B- |  |  |
| 1H-1, 80 | 0.80 | 1.082 |
| 1H-2, 40 | 1.90 | 1.052 |
| 1H-2, 80 | 2.30 | 0.979 |
| $1 \mathrm{H}-3,40$ | 3.40 | 0.985 |
| 1H-3, 120 | 4.20 | 1.070 |
| 1H-4, 40 | 4.90 | 0.917 |
| 1H-4, 80 | 5.30 | 1.108 |
| 1H-4, 120 | 5.70 | 1.052 |
| 1H-5, 80 | 6.80 | 1.132 |
| 1H-5, 120 | 7.20 | 0.974 |
| 1H-7, 20 | 9.20 | 1.152 |



Figure 23. Bathymetry in the region of Site 714, eastern margin of the Maldives Ridge (after unpublished data from NAVOCEANO, U.S. Navy surveys).


Figure 24. Wilkes 827 (17 September 1976) single-channel seismic (SCS) reflection line near Site 714. A narrow ridge effectively protects the site from turbidites flowing down from carbonate banks to the west.


Figure 25. The JOIDES Resolution single-channel seismic (SCS), water-gun reflection profile over Site 714. Strong reflectors occur at 0.24 and 0.36 s (two-way traveltime) beneath the seafloor, which correspond to the top of an Oligocene shallow-water limestone reef and basaltic basement, respectively. The hard reflector seen at the ocean floor to the left (west) of Site 714 is probably limestone. Above the reef lies a sequence of Neogene carbonate oozes.


Figure 26. The JOIDES Resolution $12-\mathrm{kHz}$ profile over Site 714. This site is located in a small basin bounded by sediment hills (slump structures?). See Figure 20 in "Underway Geophysics" chapter, this volume, for location of Site 714 track line.






|  |  |  |  |  |  |  | ORE 4H CORED |  | RVAL 2053．5－2063．2 mbsil $22.0-31.7 \mathrm{mbsf}$ |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 㲴 菦 岂 岂 |  |  | － |  |  |  |  |  | LTHMLOGIC OESCRIPTION |
|  |  |  |  |  |  |  |  |  | FORAMIIIFER．BEARING CLAY．EEARRING NANNOFOSSIL OOZE <br> Major lithology：Foraminifer－bearing clay－bearing nannofossil ooze， mildly burrow－mottled，with alternating intervals of light greenish gray （5GY 7／1）and light gray（ $5 Y$ 7／1）；mild pyrite staining throughout the core． <br> Minor lithology：Sponge spicule－bearing foraminifer－bearing clay－ bearing nannofossil ooze，light brownish gray（2．5Y 6／2），in distinct layer at Section 7，115－135 cm． |










| SITE |  | 714 |  | HOL |  | A |  | COR |  | 9 CO | RED | INTER | VAL 2101.7-2111.3 mbsi: 70.2-79.8 mbsf |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |  |  |  |  | $\underset{\substack{\text { GRAPHIC } \\ \text { LITMOLOGY }}}{ }$ |  |  | LITHOLOGIC DESCRIPTION |
|  |  |  |  | C. coscinodiscus |  |  |  | 2 <br> 3 <br> 3 |  | (1) |  |  | SPONGE SPICULE-BEARING FORAMINIFER-BEARING NANNOFOSSIL OOZE <br> Major lithology: Sponge spicule-bearing foraminifer-bearing nannofossil ooze, burrow-mottled; light greenish gray (5YR 7/1) alternates with light gray (5Y 7/1) and very light greenish gray (5G 8/1); moderately pyrite-stained. <br> SMEAR SLIDE SUMMARY (\%): <br> COMPOSITION: |




CORED INTERVAL $2121.0-2130.6 \mathrm{mbsl}: 89.5-99.1 \mathrm{mbs} f$
Lthologic description
BEARING CLAY-BEARING NANNOFOSSIL OOZE
Major lithologies: Foraminifer-bearing nannofossil ooze, and Moraminifer-bearing clay-bearing nannofossil ooze, very light greenish
gray ( 5 G 811 ) alternates with greenish gray ( $5 \mathrm{G} 7 / 1$ ): burrow-motling and pyritic burrows in Section 1, thin subhorizontal horizons of gray
(N6) bands in Sections 1 and 2 . -

|  | ${ }_{\mathrm{D}}^{4,16}$ | ${ }_{\text {D }}{ }^{16}$ |
| :---: | :---: | :---: |
| TEXTURE: |  |  |
| Sand | 15 | 10 |
| Silt | 15 | 15 |
| Clay | 70 | 75 |
| COMPOSITION: |  |  |
| Clay | 5 | 5 |
| Volcanic glass |  | Tr |
| Dolomite | $\pi$ | $\pi$ |
| Foraminiters | 20 | 13 |
| Nannotossils | 74 | 80 |
| Radiolarians | - | Tr |
| Sponge spicules | 1 | 2 |
| Bioclasts (bivalve?) | Tr | - |

## 

## 

## 3



















Major lithology: Foraminiter-nannotossil chalk, faintly burrow-mottled,
very very light greenish gray ( $5 \mathrm{G} 8 / 1$ ), occurs as series of long interva
faint thin horizons of light greenish gray ( $5 \mathrm{GY} 7 / 1$ ) distributed throughout core; zone of a bundant bioturbation, Section $1,60-70 \mathrm{~cm} ; 6$
light greenish gray $(5 \mathrm{G} 7 / 1)$ Zoophycos burrows silicified limestone fragments along inside of liner, Section $1,0-15 \mathrm{~cm}$ ight greenish gray ( $5 \mathrm{G} 81 / 1$ ), occurs around and between chalk intervals

SMEAR SLIDE SUMMARY (\%):

|  | 1,80 |
| :--- | :--- |
| TEXTURE: | 0 |
| Sand | 30 |
| Silt | 20 |
| Clay | 50 |

Quart
Clay
Volcanic glass
Volcanic glass
Foraminifers
Nannotossils
Foraminiters
Nannotossils
Radiolarians
Radiolarians
Sponge spicule


UPPER OLIGOCENE
CP 19b (NP 25)

$$
\sum_{0} \sum_{4} \sum_{4}^{n} \frac{a}{\alpha}
$$

$$
\begin{aligned}
& 4
\end{aligned}
$$

$$
\begin{aligned}
& \text { cc }
\end{aligned}
$$




SITE 714 HOLE


















[^0]:    ${ }^{1}$ Backman, J., Duncan, R. A., et al., 1988. Proc. ODP, Init. Repts., 115: College Station, TX (Ocean Drilling Program).
    ${ }^{2}$ Shipboard Scientific Party is as given in the list of Participants preceding the contents with the addition of Isabella Premoli Silva, Dipartimento di Scienze della Terra, Universitá di Milano, Via Mangiagalli 34, I-20129 Milano, Italy.

