# 2. SITE 442, SHIKOKU BASIN, DEEP SEA DRILLING PROJECT LEG 58

The Shipboard Scientific Party<sup>1</sup>

# HOLES 442 AND 442A

Date occupied: 12 December 1977

Date departed: 20 December 1977

Time on hole: 7 days, 11 hours

Position: 28° 59.00' N; 136° 03.43' E

Water depth (sea level; corrected m, echo sounding): 4639.0

Water depth (rig floor; corrected m, echo sounding): 4649.0 Bottom felt (m, drill pipe): 4649.0

Penetration (m): 0.5 (Hole 442); 313.5 (Hole 442A)

Number of cores: 1 (Hole 442); 34 (Hole 442A)

Total length of cored section (m): 0.5 (Hole 442); 313.5 (Hole 442A)

Total core recovered (m): 0.10 (Hole 442); 154.26 (Hole 442A)

Core recovery (%): 20 (Hole 442); 49 (Hole 442A)

Oldest sediment cored:

Depth sub-bottom (m): 286.1 Nature: limestone Age: early Miocene (18–21 m.y.) Measured velocity (km/s): 3.75

**Basement:** 

Depth sub-bottom (m): 313.5 Nature: basalt Velocity range (km/s): 3.98-4.66

Principal results: Site 442 is in the west-central part of the Shikoku Basin, on magnetic anomaly 6. The stratigraphic section consists of 164 meters of Pleistocene and Pliocene mud and clay, 45 meters of Pliocene and late-Miocene mud, 68 meters of late- and middle-Miocene mud and volcanic ash, 8.6 meters of middle- and early-Miocene zeolitic clay and claystone, 0.4 meters of early-Miocene limestone, 66 meters of massive basalt flows with normal magnetic polarity, and 92 meters of pillow basalt flows with normal polarity in the upper part and reverse polarity in the lower part. Continuous sedimentation started with pelagic limestone and clay; dominantly hemipelagic sediments were subsequently deposited at or near the CCD. The basement age (early Miocene, 18;-21 m.y.) agrees with the magnetic-anomaly age for anomaly 6. The basalt basement shows higher-thannormal vesicularity and is characterized by the absence of olivine.

#### HOLE 442B

Date occupied: 20 December 1977

Date departed: 27 December 1977

Time on hole: 6 days, 12 hours

Position: 28° 59.04' N; 136° 03.43' E

Water depth (sea level; corrected m, echo sounding): 4634.5

Water depth (rig floor; corrected m, echo sounding): 4644.5

Bottom felt (m, drill pipe): 4644.5

Penetration (m): 455.0

Number of cores: 20

Total length of cored section (m): 187.5

Total core recovered (m): 50.99

Core recovery (%): 27

Oldest sediment cored: Depth sub-bottom (m): 353.0 Nature: clayey nannofossil ooze Age: early Miocene (18-21 m.y.) Measured velocity (km/s): 1.57

#### **Basement:**

Depth sub-bottom (m): 455.0 Nature: basalt Velocity range (km/s): 3.85-5.25

Principal results: See principal results for Holes 442, 442A.

### **BACKGROUND AND OBJECTIVES**

## Background

The marginal basins of the western Pacific owe their origin either to a rifting process analogous to sea-floor spreading (Karig, 1970, 1971; Hayes and Ringis, 1973; among others), to a process analogous to crustal development along a transform fault (Hawkins, 1977), or to entrapment of older oceanic crust through subsequent development of a younger trench (Uyeda and Miyashiro, 1974); Cooper et al., 1976). The north Philippine Sea provides a unique opportunity to test some of these

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ideas about modes of origin of marginal basins, the Shikoku Basin providing an unparalleled testing ground for the rifting model, and the Daito Ridge and Basin region for the crustal-entrapment model.

The marine geology of the Shikoku Basin is summarized best in Karig, Ingle, et al. (1975); Tomoda et al. (1975); Kobayashi and Isezaki (1976); and Watts and Weissel (1975). Tomoda et al. (1975) demonstrated that the magnetic-anomaly pattern of the Shikoku Basin is linear. Later, Watts and Weissel (1975) and Kobayashi and Isezaki (1976) identified the ages of these magnetic anomalies and suggested a symmetrical-spreading history for the Shikoku Basin. This spreading began at a now-extinct spreading center about 28 Ma and ceased about 18 Ma. Watts and Weissel (1975) proposed two episodes of spreading with changing spreading rates, whereas Kobayashi and Isezaki (1976) suggested that spreading was a continuous episode. More recently, a more-detailed map of Shikoku Basin magnetic anomalies was compiled by Kobayashi and Nakata (1977). Deep drilling in the Shikoku Basin, recommended by the Active Margin Panel of JOIDES, provided an unusual opportunity to test the spreading hypothesis and to determine more precisely the magnetic ages suggested by geophysical surveys.

Basement-age determination was attempted during DSDP Leg 31 (Karig, Ingle, et al., 1975) when Site 297 was drilled; however, drilling at that site failed to reach basement. The drilling results did demonstrate that, during the earlier history of rifting, turbidites were deposited in the basin; this was followed by pelagic sedimentation. Presumably, the Nankai Trough, developed later to the northeast, acted as a sediment trap for turbidites derived from the Japanese Islands, although derivation of these turbidites from active and extinct island arcs such as the Kyushu-Palau Ridge and the Iwo Jima Ridge must also be considered.

The nature of the crust underlying marginal basins has also been of interest. Several investigators have demonstrated that marginal basins are floored by oceanic crust (Fischer, Heezen, et al., 1971; Andrews, Packham, et al., 1975). Detailed petrographic and geochemical study of basement rocks from Site 54 in the Parece Vela Basin indicated that these rocks are similar to abyssal tholeiites originating from mid-ocean ridges (Ridley et al., 1974).

Site 442 is on a prominent positive magnetic anomaly, tentatively identified as anomaly 6, on the west side of a hypothetical extinct spreading center in the Shikoku Basin. This site was located along a seismic profile surveyed by the R/V *Kaiyo-Maru* of Japex (IPOD-Japan, 1977) and shown in Figure 1. The seismic-survey line obtained by the D/V *Glomar Challenger* is shown in Figure 2.

## Objectives

The primary drilling objectives at Site 442 were twofold. Of prime importance was paleontological determination of the age of the basalt basement, so as to calibrate the magnetic-anomaly age determinations of previous studies and provide the basis for testing the



Figure 1. Seismic-survey line through Shikoku Basin by *R/V* Kaiyo-Maru (JAPEX).

spreading model proposed for the origin of the Shikoku Basin by concomitant drilling of Site 443. A second major objective was determination of the mineralogy, petrology, and chemical composition of the basalt floor of the back-arc Shikoku Basin, and comparison of these properties with those of oceanic basalts from other marine tectonic domains. In addition, paleomagnetic ages of the basaltic section were to be determined, so as to understand the crustal evolution of the Shikoku Basin in particular and back-arc basins in general.

Additional objectives at Site 442 included determination of changing patterns of sedimentation during the evolution of the Shikoku Basin, and, last, determination of paleocirculation patterns within the basin during its rifting history.

## **OPERATIONS**

Site 442 was approached along a course of 199°, which was established on 11 December 1977 and continued until 1018 hours, 12 December. The course was changed then to 252° to follow a seismic-refraction profile obtained by the R/V *Kaiyo-Maru* (IPOD-Japan, 1977), to reach Site 442 (Figure 2). The ship's speed before the course change was 9 knots, and after the speed was reduced to 6 knots, to obtain a clearer seismic-reflection profile along a line from proposed Site 443 to Site 442 (Figure 3). A standard seismic array of 40- and 80-cubic-inch air guns was towed along with a magnetometer.

On 12 December, at 2116 hours, the 16-kHz beacon was dropped at Site 442, and the site was occupied following a Williamson turn by the ship. Once on site, the crew began a running-in of drill pipe, because it was necessary to remove down-graded drill-pipe on rackers and replace it with aluminum pipe from the hold. This operation was completed at 2145 hours, 16 December, prior to actual spudding-in for a wash-down to determine casing depth.

After the replacement of drill pipe, scientific drilling operations at Site 442 were scheduled according to the following plan:

1. Drilling of Hole 442. Operations included the construction of a bottom-hole assembly and drill string,



Figure 2. Glomar Challenger reflection profile approaching Site 442.



Figure 3. Site location map.

establishing drill-pipe measurement, and, subsequently, obtaining a mudline core. Following the coring of one hole, a wash-in would establish the critical depth for setting the 16-inch casing for a re-entry cone.

2. Hole 442A was to be drilled at a location offset 200 feet from Hole 442. Continuous coring from the surface to near-total penetration was to be undertaken. Total penetration would be limited to 25 meters penetration of basalt below the top of the basalt basement.

3. Hole 442B was to be drilled by multiple re-entry at an offset location of 200 feet from 442A. A re-entry cone would be assembled, keel-hauled and set. Hole 442B would be washed and drilled; spot coring was to cover missing intervals or intervals of low recovery in 442A until reaching the sediment/basement interface established in 442A, when continuous coring would be undertaken.

Drilling operations at Hole 442A began at 2145 hours, 16 December 1977. The drill-pipe measurement was established at 4649.5 meters. At 2130 hours, 17 December, a mudline core was retrieved (Table 1). The wash-in depth for the 16-inch casing was established subsequently at 66 meters (below sea floor). At 0115 hours, 18 December, continuous coring began at Hole 442A (Table 1). Continuous coring proceeded without interruption until 2000 hours, 19 December, when the Bowen power sub leaked oil. Down time for repairs (replacement of a bolt) was 1 hour 30 minutes, and at 2145 hours drilling operations resumed. On December 20, basalt basement was penetrated at 287 meters below sea floor. At 0815 hours, 20 December, the Bowen power sub leaked oil again, and a motor was replaced to prevent further oil leakage. Down time was 2 hours. Drilling resumed at 1015 hours, and two cores of basalt were recovered. At 1500 hours, 20 December, drilling operations at Hole 442A ceased, the drill string was pulled, and assembly of the re-entry cone for Hole 442B was started.

A re-entry cone and 66 meters of casing were assembled prior to coring at Site 442B. The re-entry cone and casing were run into Hole 442B around noon on 21 December. Hole 442B was spudded in on 21 December, at 2310 hours, the drill pipe measurement reaching 4644.5 meters at the sea floor. A hole was washed-in to 267.5 meters below sea floor and continuous coring started around 0800 hours on 22 December at a depth of 267.5 meters below sea floor. Coring was stopped at 1945 hours on 23 December, after nearly 120 meters of penetration, because of torquing on the drill pipe. The drill string was pulled out at 2030 hours. Re-entry was undertaken on 24 December and achieved at 2058 hours. Coring resumed at 0215 hours on 25 December, and four cores were recovered. Because of deteriorating weather and high winds, the drill string was set back in-

TABLE 1 Coring Summary, Site 442

Cores	Date (Dec., 1977)	Time	Depth From Drill Floor (m) Top Bottom	Depth Below Sea Floor (m) Top Bottom	Length Cored (m)	Recovery (m)	Recovery (%)
442-1	17	2246	4649.0-4649.5	0.0-0.5	0.5	0.10	20
	1.4.10			Totals	0.5	0.10	20
442A-1	18	0530	4649.5-4659.0	0.0-9.5	9.5	0.00	0
2	18	0720	4659.0-4668.5	9.5-19.0	9.5	9.15	96
3	18	0845	4668.5-4678.0	19.0-28.5	9.5	8.91	94
4	18	1012	4678.0-4687.5	28.5-38.0	9.5	6.75	71
5	18	1130	4687.5-4697.0	38.0-47.5	9.5	7.94	84
6	18	1315	4697.0-4706.5	47.5-57.0	9.5	8.30	87
7	18	1443	4706.6-4716.0	57.0-66.5	9.5	7.54	79
8	18	1612	4716.0-1725.5	66.5-76.0	9.5	7.51	79
9	18	1746	4725.5-4735.0	76.0-85.5	9.5	3.57	38
10	18	1917	4735.0-4744.5	85.5-95.0	9.5	5.22	55
11	18	2040	4744.5-4754.0	95.0-104.5	9.5	3.40	36
12	18	2223	4754.0-4763.5	104.5-114.0	9.5	3.28	35
13	18	2355	4/63.5-4//3.0	114.0-123.5	9.5	3.41	36
14	19	0129	4773.0-4782.5	123.3-133.0	9.5	9.27	45
13	13	0257	4702.3-4792.0	133.0-142.5	9.5	5.06	33
16	19	0426	4792.0-4801.5	142.5-152.0	9.5	3.30	35
17	19	0554	4801.5-4811.0	152.0-161.5	9.5	2.78	29
18	19	0719	4811.0-1820.5	161.5-171.0	9.5	3.95	42
19	19	0844	4820.5-4830.0	171.0-180.5	9.5	4.72	50
20	19	1009	4030.0-4039.3	180.5-190.0	9.5	1.55	10
21	19	1146	4839.5-4849.0	190.0-199.5	9.5	5.41	57
22	19	1311	4849.0-4858.5	199.5-209.0	9.5	0.73	8
23	19	1443	4858.5-4868.0	209.0-218.5	9.5	6.43	68
24	19	1756	4808.0-4877.5	218.5-228.0	9.5	6.56	69
25	19	1750	40/7.3-4007.0	228.0-237.3	9.5	3.90	05
26	19	1958	4887.0-4896.5	237.5-247.0	9.5	3.10	33
27	19	2308	4896.5-4906.0	247.0-256.5	9.5	3.08	32
28	20	0046	4906.0-4915.5	256.5-266.0	9.5	4.80	51
30	20	0358	4915.5-4925.0	200.0-275.5	9.5	4.05	47
50	20	0550	4925.6-4954.5	275.5-205.0	5.5	4.40	
31	20	0246	4934.5-4939.5	285.0-290.0	5.0	1.18	24
32	20	1217	4939.3-4944.0	290.0-294.5	4.5	2.10	40
33	20	1458	4944.0-4955.5	304 0-313 5	9.5	1.12	12
		1450	4705.5 4705.0	Totals	313.5	154.26	49
				Totais	313.3	134.20	
WASH/DRILL	22		4644.5-4912.0	0.0-267.5			
442B-1	22	0911	4912.0-4921.5	267.5-277.0	9.5	3.97	42
2	22	1056	4921.5-4931.0	277.0-286.5	9.5	1.28	10
3	22	1457	4931.0-4940.3	286.5~296.0	9.5	4.07	49
5	22	1702	4950 0-4959 5	305 5-315 0	9.5	4 35	46
		2012	1000 0 1000 0	216 0 224 6	0.0	2.14	22
0	22	2013	4959.5-4969.0	315.0-324.5	9.5	2.14	23
0	23	0459	4909.0-4978.3	324.5-354.0	9.5	0.90	03
o o	23	0715	4978.3-4988.0	343 5-353 0	9.5	3 37	35
10	23	0921	4997.5-5007.0	353.0-362.5	9.5	1.00	11
	22	1100	50010 5016 5	262.6.272.0	0.0	2.27	
12	23	1247	5007.0-5016.5	302.3-372.0	9.5	0.88	24
12	23	1544	5026.0-5026.0	381 5-391 0	9.5	2 21	22
14	23	1726	5035.5-5045.0	391.0-400.5	95	1.04	11
15	23	1932	5045.0-5054.5	400.5-410.0	9.5	0.57	6
	24	First re	entry				
16	25	0420	5054.5-5061.5	410.0-417.0	7.0	1.70	24
17	25	0631	5061.5-5071.0	417.0-426.5	9.5	0.94	10
18	25	0823	5071.0-5080.5	426.5-436.0	9.5	0.35	4
19	25	1024	5080.5-5090.0	436.0-445.5	9.5	1.40	15
	27	Second	i re-entry				
20	27	1205	5090.0-5099.5	445.5-455.0	9.5	0.72	8
				Totals	187.5	50.99	27

to sediment to wait out the storm. This operation started at 1130 hours and was completed at 1230 hours, 25 December. Around 0420 hours, 26 December, the drill string was removed entirely from the hole. The storm abated during the afternoon of 26 December, and preparations were started for a re-entry attempt around 1830 hours. Re-entry operations began at 0000 hours, 27 December, and were completed successfully at 0445 hours. Coring resumed at 0955 hours, 27 December, but because of bottom hole instability and extreme torquing on the drill string, the drill string was pulled, starting at 1300 hours. Departure from Hole 442B took place at 2242 hours on 27 December.

#### SEDIMENT LITHOLOGY

### Introduction

Three holes were drilled at Site 442. Hole 442 was a pilot hole in which a mudline core was taken, recovering 10 to 15 cm of siliceous silty clay. Hole 442A penetrated 313.5 meters, of which 286.1 meters were sediment, ranging in age from late Pleistocene to early Miocene. Hole 442B was a re-entry hole for deep basalt penetration. Three sediment cores were taken, starting from a sub-bottom depth of 268.5 meters. Basalt was first encountered in Core 442B-3, at a sub-bottom depth of 289.7 meters. Table 2 summarizes the stratigraphy of Site 442.

## **Unit Descriptions**

The descriptions for units I through V (Table 2, Figure 4) are based primarily on analyses of sediments from Hole 442A. Additional descriptions for sub-unit IIIB and unit IV are based on results from Hole 442B.

#### Unit I

Unit I, present in Cores 442A-1 through 442A-18-2, 100 cm, approximately 164 meters thick, is dominantly dark-greenish-gray mud (silty clay) with some clay. The relative homogeneity of the unit precludes further subdivisions into sub-units; however, the following distinctive characteristics are observed:

 The radiolarian content generally exceeds 5 per cent (to a high of 30%) for Cores 442A-1 through 442A Cores 442A-5 through Section 442A-18-2, 100 cm, are barren of radiolarians.

The silt content is (with exceptions) greater than
 per cent for Cores 442A-1 through Section 442-A-18 The clay content generally exceeds 50 per cent for
 Core 442A-1 through Section 442A-18-2.

3. The quartz-feldspar content is (with exceptions) greater than 10 per cent in sediments of Cores 442A-1 through 442A-10, and does not exceed 5 per cent in Core 442A-11 through Section 442A-18-2, 100 cm.

4. Volcanic glass and ash observed in smear slides appear to be present persistently in Core 442A-2 through Section 442A-18-2. Ashy units and (or) pumice fragments generally are present consistently in Cores 442A-2 through 442A-7, 442A-10, 442A-12, 442A-13, 442A-16, and 442A-18.

5. Shear-vane studies indicate that a physical difference exists between sediments of Cores 442A-1 through 442A-9 and those of Cores 442A-10 through 442A-18.

#### Unit II

Unit II is distinguished from Unit I by a change from dark-greenish-gray (5GY 4/1) mud and clay to yellowbrown (10YR 5/4) mud and clay. The unit is 45 meters

TABLE 2								
Lithologic Units	at Site 442							

Interval	Lithologic Unit	Depth and Thickness (m)	General Description	Age
442-1; 442A-1 to 442A-18-2, 100 cm	I	0.0–164.0 (164)	Dark- greenish-gray mud, clay	Quaternary and Pliocene
442A-18-2, 100 cm to 442A-22	п	164.0-209.0 (45)	Yellow-brown mud (164–180.5 m) to brown, dark-brown mud (180.5–209 m)	Pliocene and possibly late Miocene
442A-23 to 442A-30-2, 13 cm	IIIa IIIb	209.0-~277.0 (68.1)	Sub-unit IIIA: yellow-brown mud with siliceous fossils Sub-unit IIIB: gray, dark- greenish intermixed ash, mud (start 259.6 m, Core 28-3, 100 cm)	Pliocene or late Miocene to middle Miocene
442A-30-2, 13 cm to 442A-31-1, 70 cm	IV	277.1-285.7 (8.6)	Dark-brown clay, claystone with bioturbation, zeolites	Middle to early Miocene
442A-31-1, 70 cm to 442A-31-1, 107 cm	v	285.7-286.1 (0.4)	Hard, fine-grained, pink limestone	Early Miocene
		Basalt at 286.	1 meters sub-bottom	
442B-1-1 to 442B-1, CC	IIIb	267.5–277.1 (9.6)	Grayish, olive, greenish- gray clayey ash, clay	Middle Miocene
442B-2 to 442B-3-1, 20 cm	IV	277.1–289.7 (12.6)	Dark-brown zeolitic clays, bioturbated claystones, Mn nodules. Altered ash beds (w/nannofossils in Core 2-1)*	Middle to early Miocene
		Basalt at 289.	7 meters sub-bottom	

thick and present in Section 442A-18-2, 100 cm, through Core 442A-22. It passes into Unit III at 209 meters (Section 442A-23-1) with a change from dark-grayish-brown mud with trace amounts of radiolarians to yellowishbrown (10YR 6/4) sediments with common radiolarians beginning in Core 24 (218.5 m). Foraminifers are generally rare or absent in Unit II, becoming common to abundant in Unit III.

Other lithologic characteristics of the sediments of unit II are: a clay (size) content of 60% per cent or more, a general absence of siliceous fossils, and a volcanic glass content generally less than 3 per cent, but as high as 95 per cent in ash layers.

#### Unit III

Unit III is 68.1 meters thick and present in Core 442A-23 through Section 442A-30-2, 13 cm. The unit is divided into two sub-units. Sub-unit IIIa (50.6 meters thick) contains yellowish-brown mud and clay with radiolarians and is present in Core 442A-23 through Section 442A-28-3, 105 cm. Sub-unit IIIb consists of dark-greenish-gray intermixed volcanic ash, mud, and clay and is present in Sections 442A-28-3, 105 cm, through 442A-30-1, 13 cm. The siliceous-fossil content is considerably lower (trace to 1%) in sub-unit IIIb than in Sub-unit IIIa.

Sub-unit IIIb, in Hole 442A, was cored beginning at 259.6 meters and is 17.5 meters thick. Coring in Hole 442B began at 267.5 meters in sub-unit IIIb, and sediment of the sub-unit continued to a depth of 277 meters. The sub-unit is similar in both holes, consisting of intermixed clay and volcanic ash. However, the lower boundary of the dark-brown, zeolitic clay or claystone of unit IV differs in the two holes. In Hole 442A, the contact with the underlying dark-brown, zeolitic clays or claystones in Section 442A-30-2, 13 cm is clear. However, in Hole 442B the contact is represented by 15 cm of fragments of brecciated, light-gray (5Y7/1) clay within a matrix of nannofossil-bearing clay. Probably this sediment unit was present in Hole 442A, but was washed out during coring and core recovery. Its presence in Hole 442B marks the first significant appearance of carbonate sediments in the upper 275 meters of the sedimentary section.

### Unit IV

Unit IV is present in Sections 442A-30-2, 13 cm through 442A-31-1, 70 cm, and consists of yellowbrown and dark-brown, firm to hard, zeolitic clay and claystone. The sediments are characterized by evidence of extensive bioturbation and clay-size-material and clay-mineral contents greater than 75 per cent (with ex-



Figure 4. Lithology, sonic velocity, and magnetics of sediments from Holes 442A and 442B.

ceptions). In Section 442A-30-3, phillipsite becomes common (20–25%). Micronodules(?) of iron or manganese are common in Sections 442A-30-2 and 442A-30-3.

The darker brown colors appear to be due to oxidation of disseminated manganese and (or) iron. The mottled appearance of the yellow-brown and dark-brown clays is a result of bioturbation. A pink altered-ash zone occurs in Section 442A-30-2.

Sediments of Unit IV were recovered in Hole 442B from 277.1 meters (Core 442B-2) to the top of basalt at 289.7 meters (Core 442B-3). Excellent recovery of the cored section permits good characterization of the unit. In Hole 442B also, the unit contained a brecciated zone consisting of lighter-colored claystone (altered ash) fragments within a matrix of dark-brown, zeolitic clay. It is presumed that the breccia is a drilling artifact, because the general lithologic characteristics do not differ from those of unbrecciated portions of Unit IV.

The unit is a firm, stiff, zeolitic (phillipsite?), dark brown (2.5YR3/3 to 5YR4/4) clay and claystone. Bioturbated, mottled sediment in the upper portion of the unit includes a dark-yellowish-brown (10YR4/4) clay. In general, the unit takes on lighter hues the more extensive the bioturbation; however, bioturbation is present in the lower portions of the unit also, and lighter hues are not evident there.

In many cases, the extensive bioturbation creates a "near breccia" appearance. However, bioturbation evidence does include burrow patterns, such as excellently developed *Zoophycos*.

The upper portion of the unit (Section 442B-2-1 through Core 442B-3) does not contain zeolites; however, it does contain manganese micronodules, evidence of manganese streaking and bioturbation. Zeolites first appear in Section 442B-2-4 and show a maximum content of 30 per cent in the sediment.

Manganese micronodules are ubiquitous, and occasional zones of micronodule concentrations (parallel to bedding) are noted. In the interval, 442B-2-1, 55-57 cm, a nodule-like spheroid, was encountered; it has a diameter of about 2 cm and is composed of altered ash and clay coated with a 0.5- to 2-mm manganese crust. The clay is similar to the surrounding sediment.

The unit also contains several zones, 10- to 20-cm thick, of lighter-colored [pinkish-gray (7.5YR7/2), light-yellowish-brown (2.5Y6/4), pale-yellow (2.5Y7/4)] altered ash. The ash, now almost completely altered to a clay, contains manganese micronodules and displays bioturbation, particularly at the upper boundary with the dark-brown, zeolitic clay.

#### Unit V

Unit V consists of a 0.4-meter section of hard, dense, fine-grained, pinkish-gray limestones with manganese(?) and smaller-foraminifer tests scattered throughout. This unit overlies basalt in Hole 442A.

The limestone matrix is a fine micrite, with sparite as vein or cavity filling and as a replacement of foraminifer tests. The limestone may be recrystallized nannofossil ooze; however, there is no evidence to support this. Grain size of the micrite appears to decrease up-section from the basalt contact. Other notable characteristics of the limestone include:

1. Evidence of bioturbation. The burrows are filled with micrite which is slightly finer grained than the limestone matrix.

2. The presence of rare iron-manganese minerals; partly replaced (by sparite, Fe-Mn oxides, or other carbonates) foraminifer tests; volcanic-glass spherules; and basalt fragments.

3. Black opaque minerals (Fe or Mn ozides) as vein or cavity fillings; cavity filling by volcanic glass or silica; filling of foraminifer tests by sparry calcite (the tests themselves replaced by micrite).

Sediments of Unit V were not recovered from Hole 442B. it appears that the sediment unit overlying basalt in Hole 442B is the dark-brown, zeolitic clay of unit IV.

## ORGANIC GEOCHEMISTRY

Four gas samples were taken from cores of Hole 442A which showed some signs of gas, but light hydrocarbons were found to be absent in all four samples.

Twenty-one samples were analyzed for organic-carbon and nitrogen contents. The data and results are reported elsewhere (Waples and Sloan, this volume). Organic-carbon contents are uniformly low, ranging from about 0.5 per cent near the sediment/water interface to about 0.05 per cent near basalt basement.

A black, granular material was found in several cores and analyzed to determine whether it was organic. It reacted rapidly with  $H_2O_2$ , suggesting that it might be organic, but element analysis and pyrolysis both showed that it consisted of less than 0.1 per cent organic carbon. It was concluded that the material probably represented manganese micronodules.

#### **INORGANIC GEOCHEMISTRY**

Seven interstitial-water samples were taken for inorganic geochemical measurements. The data are listed in Table 3 and presented on Figure 5. Six samples were taken in the sedimentary section of Hole 442A, within units I, II, III, and IV. One sample from Hole 442B is representative of unit IV.

The *p*H averages 7.71 and ranges from 8.19 to 7.46. Except for an increase in Sample 2 relative to Sample 1, *p*H tends to decrease with increasing depth. All sediment *p*H values are below those values for the IAPSO standard and surface sea water at the site.

Alkalinity averages 4.70 meq/kg for the seven samples, all values exceeding the IAPSO standard and surface sea water values of 2.39 and 2.32, respectively. Alkalinity decreases regularly with increasing depth, and no deviations from this trend were observed. The range of values is 9.74 to 2.97 meq/kg.

Salinity averages 35.3 per mill and chlorinity 19.29 per mill for the seven samples. Only one deviation exists in the expected matching trends for the two variables. For Samples 3 and 4, salinity remains constant at 35.2 per mill; however, the chlorinity drops from 19.38 to 18.67 per mill. However, this may be a result of laboratory error in the chlorinity and (or) salinity measurements.

	Samp (interval i	le Sa n cm) Nu	mple mber	Sub-Botto Depth (m)	m pH	Alkalinity (meq/kg)	Salinity (º/ <sub>00</sub> )	Ca <sup>++</sup> mmol/l	Mg <sup>++</sup> mmol/l	CI- (º/ <sub>00</sub> )	-
	442A-2-3, 1 7-4, 1 13-2, 19-3, 23-2, 27-3,	IA S 43-150 40-150 143-150 0-10 93-100 0-8	PSO SW 1 2 3 4 5 6		8.05 8.27 0 7.78 0 8.19 00 7.87 10 7.59 50 7.52 08 7.57	2.39 2.32 9.74 5.65 4.66 3.75 3.12 2.97	35.2 35.2 35.5 35.2 35.2 35.2 35.2 36.0 35.5	10.55 10.38 10.68 12.52 13.64 14.80 14.41 12.65	53.99 53.81 50.26 46.29 44.70 45.13 47.53 48.29	19.375 19.14 19.41 19.34 19.38 18.67 19.79 19.51	
	442B-2-4, 14	40-150	7	282.90-283.0	00 7.46	2.99	34.4	11.45	51.38	18.94	
	Section	Sub-bottom Depth Interv	al(m)	рH	Salinity ( <sup>0</sup> /00)	C( <sup>-</sup> ( <sup>0</sup> /00)	Alkalinity ( <sup>0</sup> /00)	Ca++ (mmol,	/1)	Mg (mm	++ ol/l)
				78	34 36	18 20	0 10	10 1	15 40 I I	) 45 5	0 55 I
0	Standa Surfac	ard Sea Water e Sea Water	-	•	•	•	•	•			:
	442A=2=3	13.93–14.	00 -	9	Ŷ	P	ſ	9		9	P
	- 442A=7=4	62.90–63.	00 -	þ		0				9	-
om Depth (m)	442A=13=2	116.93–117	2.00 -		0						
Sub-bott	- 442A=19=3	174.00–174	.00 -		R		0				
200	- 442A=23=2	211.43–211	.50 -	4	þ			5		þ	
	- 442A=27=3	250.00-250	.08 -	0	þ					q	
	442A=2=4	282.90-283	9.00 -	6	6	6	6	6			δ

 TABLE 3

 Summary of Shipboard Geochemical Data for Holes 442A and 442B

Figure 5. Interstitial-water geochemistry, Hole 442A.

 $Ca^{++}$  averages 12.88 mmol/l and Mg<sup>++</sup> 47.64 mmol/l for the seven samples.  $Ca^{++}$  values are higher than those recorded for the IAPSO and surface sea water standards, whereas Mg<sup>++</sup> values are lower compared to values for the same standards.

Ca<sup>+</sup> increases regularly to Section 442A-9-3 (unit II) decreasing thereafter.  $Mg^{++}$  decreases to Section 442A-13-2 (unit I), increasing thereafter.

### BIOSTRATIGRAPHY

## Overview

Holes 442, 442A, and 442B are in the Shikoku Basin, approximately 50 km west of a supposedly extinct spreading center. The objectives at this site were to determine the age of basement basalt which is considered to be on magnetic anomaly 6 ( $\sim$  19–20 m.y.), and to obtain a complete record of the sedimentary history of Shikoku Basin.

Core 442-1 recovered only the length of the core catcher of the surface sediment. Although the water depth at this site, 4639 meters, is considered to be slightly below the present CCD, this sample contained late-Quaternary foraminifers, nannofossils, and radiolarians.

Hole 442A was continuously cored through the 286meter sedimentary sequence and approximately 30 meters of basalt. Calcareous nannofossils are the most significant microfossils for age determination, although their occurrence is sporadic in the lower portion of the sedimentary sequence of this hole (Table 4).

Cores 442A-1 to 442A-3 contain rich faunal and floral assemblages of foraminifers, nannofossils and radiolarians. All of these three microfossil groups indicate the age of late Quaternary with mixed tropical and temperate water facies. Due to the closeness to the CCD, calcareous nannofossils suffered slight to moderate dissolution. The foraminifers are fragmented and poorly preserved.

Core 442A-4 yielded well-preserved nannofossils and radiolarians. Both fossil groups indicate an age of middle Pleistocene.

Between Cores 442A-5 and 442A-13, nannofossils are the only significant microfossils preserved; below Core 442A-7, the assemblages indicate an early-Pleistocene age.

Cores 442A-14 through 442A-23 are barren, except for Section 442A-19, CC, which yielded a moderately well-preserved assemblage of Pliocene radiolarians.Cores 442A-21 to 442A-27 contain many pseudomorphs which are casts of foraminifers having no significance for age determination.

Cores 442A-24 and 442A-25 contain only radiolarians; the assemblages indicate ages of late and middle Miocene for these two cores, respectively.

Cores 442A-26 and 442A-27 contain nannofossils and radiolarians. The assemblages of both fossil groups indicate an age of middle Miocene.

In Section 442A-31-1, a dark-brown clay overlies approximately 30 cm of limestone, which in turn directly overlies basalt. Nannofossils in this clay indicate an age of late early Miocene. There was no sign of this limestone being formed under shallow-water conditions.

Limestone directly above basalt was also observed at Site 53, on the western flank of the Iwo Jima Ridge, approximately 700 km south of Site 442. The origin of this limestone was thought to be either inorganic precipitation or recrystallization of nannofossil ooze due either to low-temperature diagenetic reactions or thermal metamorphism associated with igneous activity (Fischer, Heezen, et al., 1971).

At Hole 442B, 22 meters of sediments were cored directly above the basement basalt. Section 442B-1,CC contains poorly preserved radiolarians, whereas Section 442B-2-1 contains a thin layer of calcareous sediment. Nannofossils and foraminifers indicate ages of late early to early middle Miocene for these cores, as does the sediment directly above the limestone at Hole 442A.

Approximately 2 meters of soft sediment (recognized by the drilling rate) were encountered in Core 442B-9, which was drilled through the basalt layers. Although none of this sediment was recovered in the core barrel proper, a small piece of light-yellowish-brown and darkbrown mud was found in the core catcher. The lightcolored material consists almost entirely of nannofossils, with some foraminifers, whereas the dark material contains the same fossil assemblage diluted by non-biogenic material. Both foraminifers and nannofossil assemblages show the age of this sediment to be early early Miocene (17–121 m.y.). This age is near that predicted by magnetic-anomaly study at this site. Although Core 442B-13 contained another fragment of sediment, it does not contain any microfossils.

Although this hole was offset by only 70 meters from Hole 442A, no limestone was recovered; this could be attributed to either patchy distribution of limestone or failure of recovery because of brecciation and subsequent washout of the material by drilling.

The presevation of calcareous microfossils shows the depth of water at this site in relation to the CCD as follows:

1. Slightly above the CCD during the late Quaternary, as indicated by both foraminifers and nannofossils.

2. At about the depth of the CCD during the middle and early Quaternary.

3. Well below the CCD during the Pliocene and late Miocene.

4. Slightly below the CCD during the early and middle Miocene, except during a few short periods slightly above the CCD.

## Foraminifers

The summary for Holes 442, 442A, and 442B is based essentially on core-catcher samples, plus four section samples which were used for controls. The time scale used is that of Berggren and Van Couvering (1974) and Saito (1977).

Foraminifers encountered in Holes 442, 442A, and 442B were sporadic and, where present in some sections, rare and poorly preserved. Therefore, a continuous bio-stratigraphic study of the sequences is not possible.



 TABLE 4

 Biostratigraphic Zones, Site 442 (Holes 442, 442A, 442B)

The only time intervals for which foraminifers could be used with some confidence were the Pleistocene and early Miocene, and even for these intervals foraminifers were rare and showed evidence of dissolution.

The paucity of this group is attributed to the depth of the CCD. Planktonic foraminifers are readily fragmented and dissolved between the lysocline and the carbonate-compensation depth. In samples deposited in this zone, only the most robust and heavily calcified forms survive, and these may be reduced in size because of dissolution of outer shell material. Although benthic foraminifers are usually better preserved, because of morerapid burial and retention of their protective organic sheath, they too show varying degrees of dissolution.

In summary, deposition throughout the sections penetrated in Holes 442, 442A, and 442B was either very close to or below the CCD.

#### Hole 442

One core was taken at Hole 442. The core penetrated Holocene to uppermost-Pleistocene sediments. Dissolution and fragmentation of the planktonic foraminifers indicate deposition close to the CCD during that time. Identifiable planktonic forms were *Globorotalia inflata*, *G. truncatulinoides*, and *Globigerina*; the deep-water benthic assemblage contained *Melonis pompilioides*, *Eponides*, *Dorothia*, and *Lagena*.

#### Hole 442A

Section 442A-2,CC, at a depth of 4668 meters yielded the first material from Hole 442A. The material examined contained a sparse and poorly preserved planktonic and benthic fauna of Pleistocene (N.23) age. There is a possibility of reworking in this sample.

In Core 442A-3, there is an increase in numbers and diversity of the identifiable planktonic and benthic foraminifers. However, the abundance of fragmented tests indicates dissolution and deposition close to the CCD. The assemblage of planktonic forms includes *Globorotalia inflata*, *G. truncatulinoides*, *G. tosaensis*, *Globigerina bulloides*, *Globigerinoides ruber*, *G. diminutus*, *Neogloboquadrina dutertrei*, *Pulleniatina finalis*; this assemblage indicates a Pleistocene (N.22) age for the core. The benthic assemblage contains species of *Uvigerina*, *Cibicides*, and *Globocassidulina*.

Cores 442A-4 through 442A-20 are considered barren of foraminifers, for in all samples only one or two specimens and fragmented chambers were recovered.

In the washed residue of Core 442A-21 is the first appearance of completely replaced casts of foraminifers. It appears that the test of the foraminifer was not replaced, but dissolved after the infilling was diagnetically altered. The crystalline replacement material, identified by X-ray diffraction and atomic absorption, is the rhodochrosite.

In Core 442A-22, there is an increase in the amount of this material, and in Core 442A-23 there is a flood of these completely replaced questionable foraminifers. In this core, the shapes are more diverse and definitive. Down-hole, the last significant occurrence of these forms is in Core 442A-27. Cores 442A-28 through 442A-30 lacked all traces of foraminifers. The brown, silty clay and brecciated material in Core 442A-31 was also devoid of foraminifers, but a thin section of the limestone immediately overlying the basalt did have outlines of recrystallized planktonic forms.

#### Hole 442B

Cores 442B-1 through 442B-3 are barren of foraminifers, except for Section 442B-2-1, 7–9 cm, in which partially dissolved *Sphaeroidinellopsis subdehiscens* or *Prosphaeroidinella* were found. The preservation is so poor that identification is questionable.

S. subdehiscens does not appear in the geologic record until N.13—approximately 12 m.y.—and the 12-m.y. age is at least 2 m.y. less than the nannofossil age, which seems to be more reliable. *Prosphaeroidinel-la* occurs from 14 to 15 m.y.

Core 442B-9 recovered basalt, but the core catcher contained approximately 20 cm<sup>3</sup> of brown, silty clay with some calcium carbonate. Planktonic and benthic foraminifers were found in this material, and the index species *Catapsydrax dissimilis*, although partially dissolved, was recognized.

Based on this species, the age of the sediment, is estimated to be early Miocene (N.6 or below; 17.5–19 m.y.).

#### Nannofossils

Well- to moderately well-preserved Pleistocene and poorly preserved Miocene nannofossil assemblages were observed at this site. The Zonation of cores is summarized in Table 4.

#### Pleistocene

Only one core was recovered in Hole 442. Sample 442-1, CC contains a subtropical assemblage of the *Emiliania huxleyi* Zone (late Pleistocene or Holocene). The preservation of nannofossils is good; they show only slight effects of etching.

Sample 442A-2-1, 50 cm contains a well-diversified assemblage of the *E. huxleyi* Zone, with occasional reworked Pliocene and Miocene forms. The interval between Samples 442A-2-3, 52 cm and 442A-3-2, 105 cm belongs to the *Ceratolithus cristatus* Subzone; the nannofossils, affected by slight etching, are well- to moderately well-preserved, and reworking is moderate. The *Pseudoemiliania lacunosa* Subzone was identified in Samples 442A-3-5, 75 cm to 442A-5, CC; reworking is extensive in the upper part and becomes slight in the lower portion of this zone. The interval between Samples 442A-7-1, 80 cm and 442A-13, CC represents the *Crenalithus dolonicoides* Zone (early Pleistocene). Rare occurrences of Pliocene and Miocene forms are also recognized.

#### Miocene

Nannofossils are rare and badly dissolved in Samples 442A-26-2, 80 cm and 442A-28-1, 95 cm. The coexistence of *Discoaster bollii* and *exilis* suggests middle Miocene (*Discoaster exilis* Zone?). Rare and poorly pre-

#### **SITE 442**

served nannofossils also occur in Sample 442A-31-1, 4 cm. Although Sphenolithus heteromorphus is absent, abundant Discoaster deflandrei, Cyclicargolithus floridanus, and common Discoaster variabilis suggest an age of late early Miocene (S. heteromorphus or Helicosphaera ampliaperta Zones).

Sample 442B-2-1, 10 cm contains an abundant and moderately well-preserved assemblage of either the S. heteromorphus Zone or the H. ampliaperta Zone. A piece of sediment recovered in the core catcher of Core 442B-9 contains abundant but poorly preserved nannofossils. Occurrences of Discoaster druggii, Sphenolithus dissimilis, Triquetrorhabdulus carinatus and T. milowii indicate middle early Miocene (Sphenolithus belemnos or Discoaster druggii Zones).

#### Radiolarians

Of the three holes drilled at this site, only Hole 442A yielded a significant radiolarian fauna.

#### Preservation

There were four preservation zones encountered in Hole 442A. The first four cores contain abundant, wellpreserved radiolarians. Cores 442A-5 to 442A-24 (with the exception of Core 442A-19, which contains a few moderately well-preserved Pliocene radiolarians) are barren. The third zone, from Cores 442A-25 to 442A-27, contains radiolarians which are common and moderately well-preserved. From Cores 442A-28 to 442A-31, the sequence is essentially barren.

### Biostratigraphy

Preservation zone 1 (Cores 442A-1 through 442A-4) is Quaternary to Holocene (*Lamprocyrtis haysi* Zone). Species found in these cores include *L. haysi*, *Spongaster tetra*, *Ommatartus tetrathalamus*, *Theocorythium trachelium*, and *Spongopyle osculosa*.

The large section barren of radiolarians in preservation zone 2 has no biostratigraphic indicator fossils, with the exception of Core 442A-19, which is lower Pliocene.

Preservation zone 3 contains Miocene forms. Because of the presence of common Ommatartus antepenultimus and associated species, Section 442A-24, CC is placed in the late-Miocene Ommatartus antepenultimus Zone. Core Section 442B-25, CC is in the middle-Miocene Cannartus petterssoni Zone, with Cannartus laticonus, Lithopera bacca, Cyrtocapsella japonica, and Stichocorys wolffii. The first appearance of Cyrtocapsella tetrapera is in Section 442A-26, CC, with common Cannartus laticonus and Cyrtocapsella japonica; these species suggest the early middle Miocene (Dorcadospyris alata Zone). The bottom of the preservation zone 3 and the last core to which an age can be given using radiolarians is Core 442A-27. This core is also early middle Miocene, in the Dorcadospyris alata Zone. It contains the same species as Section 442A-26, CC, and also Stichocorys delmontensis.

## SEDIMENTATION RATE

An age-depth plot is shown in Figure 6. The ages of the sediment were obtained from the time scale of Berg-

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Figure 6. Sediment-accumulation-rate curve for Site 442, based on biostratigraphic age determinations.

gren (1972), Berggren and Van Couvering (1974), Bukry (1975), and modifications of the Miocene proposed by Saito (1977). Table 5 shows sediment-accumulation rates for each stratigraphic unit, which also are plotted on Figure 6.

The sediment-accumulation curve shows an upward increase of sediment-accumulation rates; unit V shows the lowest rate of 2.4 m/m.y., and the rate increase to 19.4 m/m.y. and 9.4 m/m.y. for unit IV and unit III respectively. Units IV and V appear to be dominantly pelagic, and low rates are to be expected, whereas unit III appears to represent the first phase of hemipelagic sedimentation, which can show both low and intermediate rates (Klein, 1975). With the increase in hemipelagic deposition of unit II, rates of sediment accumulation increase again, to an average of 10.7 m/m.y.

TABLE 5 Sedimentation Rates, Site 442 (Hole 442A)

Unit	Depth (m)	Interval Thickness (m)	Sedimentation Rate (m/m.y.)		
I	0.0-164.0	164.0	30.9		
II	164.0-209.0	45.0	10.5		
ш	209.0-277.1	68.1	12.6		
IIIa IIIb	209.0-259.6 259.6-277.1	50.6 17.5	11.2 17.5		
IV V*	277.1-285.7 285.7-286.1	8.6 0.4	17.2 0.8		

\*Not recovered in Hole 442B.

Unit I, which is Quaternary and Pliocene age, is characterized by higher rates of sediment accumulation with an average rate of 30.9 m/m.y. Initially, resedimentation of this unit was suspected as a cause for these high rates, but the complete biostratigraphic continuity, lack of evidence of abraded faunal remains, and lack of turbidites clearly point to other causes. An increase in regional volcanism is a second possibility, particularly with the known increase of volcanic ash and glass reported in Quaternary sediments at Site 296 (Donnelly, 1975), and in the circum-Pacific (Kennett et al., 1977). However, the general ratio of volcanic debris to a higher volume of terrigenous sedimentary components in unit I and the older units of Site 442 indicates that, although volcanism may play an important part, it can only explain the high sediment accumulation rate in unit I if combined with an additional sedimentary process. That third process may well be Quaternary fluctuations in sea level over adjoining land areas. In some areas, such as the eastern U.S., rates of sediment supply and accumulation are low in marine-slope settings during high stands of sea level, but increase with a lowering of sea level, when turbidite sedimentation and hemipelagic deposition become dominant (Doyle et al., 1976). The consequence of lowered sea level is to increase the area of erosion on land, thus increasing the volume of sediment supplied to oceanic settings. The distance of lateral transport from shore regions to deeper basins also decreases under such conditions because of shoreline progradation towards the basin center. Under such conditions, it would be expected that high accumulation rates would be dominant. It is suggested that the combination of sea-level fluctuations and the high rate of volcanism during the Quaternary in the region would account for the high sediment accumulation rates in unit Ι.

## **IGNEOUS PETROLOGY**

### Hole 442A

#### Lithology

At Hole 442A, basalt was encountered at 285.7 meters sub-bottom, and was cored continuously to 313.5 meters sub-bottom. Of the 27.8 meters of basalt cored, 8.5 meters were recovered. The entire section consists of aphyric, fine- to medium-grained basalt, with only slight variations in texture, degree of alteration, and vesicularity. Given the short interval cored, and its uniformity, these basalts are treated as a single lithologic unit.

The most remarkable feature of the basalts is their high vesicularity, which averages between 20 and 30 per cent by volume. All the basalts contain from 10 to 30 per cent fine vesicles (<1.0 mm), and many have a second distinct population of 1 to 5 per cent medium-sized vesicles (1.0–5.0 mm) as well.

Based on sometimes subtle differences in the degree of alteration, vesicularity, and texture, the basalts were divided into six sub-units. Despite the lack of intervening glass or chill zones, we feel that each of these subunits represents a separate cooling unit. These units are similar to those found in the upper 30 meters of basement cored nearby at Hole 442B where intervening glass and chill zones are present. Sub-units 1C and 1F differ from the other sub-units in that the basalt has a variolitic texture. Sub-units A and B, and D and E, differ from each other principally in degree of vesicularity and alteration.

#### Petrography

The basalts from Hole 442A are fine- to mediumgrained, and, with the exception of a few scattered plagioclase microphenocrysts, are entirely aphyric. Generally, they consist of randomly oriented plagioclase laths and pyroxene granules in an intersertal groundmass of pyroxene and plagioclase microlites, magnetite, what appears to be devitrified glass, and alteration products, chiefly clays. The intersertal basalts grade into intergranular varieties made up of randomly oriented plagioclase and subophitic pyroxene with intra- and intergranular magnetite. Chromite is a common accessory mineral. No olivine, or evidence of it, was seen in any thin section. Sample 442A-34-1 (Piece 9) differs from the rest of the basalts in that pyroxene is present only as microlites between plagioclase laths, and a variolitic texture is readily discernible in thin section.

Plagioclase is the most abundant constituent of these rocks, amounting to 10 to 30 per cent of the basalt. It is frequently twinned, and carlsbad and pericline twins are present in addition to abundant albite twins. The plagioclase is fairly calcic, generally exceeding An<sub>70</sub>, and larger grains are often zoned, with substantially morecalcic cores.

Pyroxene also occurs in large amounts, generally from 5 to 30 per cent, excluding microlites. It is augite, with a light-brown color, a positive optic sign, and a moderate 2v around 45 to 60°. It ranges from groundmass microlites to granules and subophitic masses around plagioclase and stubby subhedral grains. The pyroxene in Sample 442A-33-4 (Piece 7) differs in that it often occurs as laths intergrown with plagioclase. Hourglass zoning of the pyroxene was observed in a number of thin sections.

Magnetite ranges from less than 1 to 5 per cent of the basalt. It is commonly euhedral and may be enclosed in plagioclase and pyroxene grains; it also occurs intergranularly in the groundmass.

Small chromian-spinel octahedra ( $\sim 0.01$  mm) were found in eight of the nine thin sections examined. Unlike many oceanic basalts, spinel is often found as an inclusion in pyroxene grains as well as in plagioclase. It is rare in the groundmass. The spinel has a deep rust-red color and is commonly barely transparent. Spinel was present in very small amounts, and commonly only a few grains could be found in an entire thin section.

With the exception of Sample 442A-34-1 (Piece 9), the order of crystallization of the minerals was difficult to determine. Plagioclase, pyroxene, and spinel appear to have crystallized early, and magnetite late. The crystallization of magnetite clearly overlapped that of pyroxene and plagioclase, but may not have overlapped that of spinel. The rarity of chromian spinel in the groundmass probably reflects very early crystallization and a tendency to act as a nucleus for silicate crystallization.

## Alteration

All basalts from Hole 442A apparently have undergone some alteration. The principal evidence of alteration is numerous calcite amygdules (1-5%) and veins. There is a light-brown to yellow discoloration of the basalt, although much of it has little discoloration. In thin section, many basalts are seen to contain small amounts of clay.

# Hole 442B

#### Introduction

Basalt was encountered at approximately 292 meters sub-bottom, and was cored continuously to 455.0 meters. 39.5 meters of basalt was recovered from the 163 meters of basement drilled. A sub-basement lithologic column is shown in Figure 7. Recoveries and intervals have been adjusted for the plastic spacers added to split cores during processing. This changes the recovery from the 39.5 meters actually recovered, to the 51.6 meters shown in Figure 7, but allows ready crossreference between the present positions of rocks in the core liners to their positions in the column.

The upper 59 meters, designated unit 1, consists of aphyric, vesicular cooling units similar to those described for Hole 442A. Beneath unit 1, unusually high drilling rates, similar to those for mud, were encountered for 2 meters. The core catcher contained nannofossil- and foraminifer-bearing mud, containing *Discoaster druggii* (18–21 m.y.), which is of greater age than fossils found in the mud overlying unit 1. Paleontologic evidence suggests an age gap of 1 to 5 m.y. Beneath this horizon, at 353 meters sub-bottom, a sequence of aphyric vesicular pillow basalt, unit 2, was encountered and drilled to the bottom of the hole at 455.0 meters sub-bottom. The lower third of unit 2 is brecciated and contains two short intervals of massive, aphyric basalt intercalated with the pillows.

As at Hole 442A, the most unusual feature of these basalts is their high vesicularity, averaging about 25 per cent by volume.

We believe that the mud found in Section 9, CC represents about a 2-meter-thick horizon at about 351 meters sub-bottom, rather than slumping of mud from higher in the hole. We base this belief on (1) the high drilling rates, (2) the presence of mud in the core catcher with fossils older than those in the sediments overlying basement, (3) the location at the major lithologic break in the basalt sequence, and (4) the absence of any similar recoveries elsewhere in the sub-basement cores.

### Unit 1, Massive Basalts

Unit 1 consists of highly vesicular, massive basalts, similar in most respects to that recovered at Hole 442A. As at Hole 442A, the unit 1 basalts at Hole 442B have been divided into sub-units which we feel represent

separate cooling units. The lithologic differences among these sub-units, as at Hole 442A, are subtle, and include variability in texture, vesicularity, and alteration. Unlike Hole 442A, however, eleven separate chill zones were found, and these separated most of the designated lithologic sub-units shown in Figure 7.

The sub-units in the upper 30 meters of Holes 442A and 442B correlate well, with the exception of the very minor sub-unit 1F at Hole 442B. If 1F at 442B is ignored, there are the same number of sub-units, with similar variations in vesicularity, and the last one at each hole is variolitic.

No actual contacts between subunits were recovered. At some locations adjoining fragments have opposing chill zones, but are not physically connected. Unlike the glass-rich pillow rinds found in unit 2, glass was missing in most of the chill zones and was found only at two locations in Core 442B-6; the largest glass zone was only 3 mm thick. In most cases, chill zones are marked by a sudden large decrease in grain size, discoloration of the rock, pipe vesicles, and, in some instances, a variolitic zone. Where the original orientation was preserved, the chill zones appeared to be nearly horizontal.

As at Hole 442A, all basalts contain 10 to 30 per cent fine vesicles (<1.0 mm) by volume. Many also have a second distinct population of 1 to 5 per cent mediumsized vesicles (1.0 to 5.0 mm). Pipe vesicles are common near chilled margins.

The basalts range from aphanitic to medium grained and have a variety of textures. The majority of the thin sections examined consist of intergrown, fine- to medium-grained augite and plagioclase, with a texture ranging from intersertal through intergranular to diabasic. In a few cases, the intersertal basalts have a groundmass consisting of intergrown plagioclase and clinopyroxene microlites and titanomagnetite grains; in most cases, however, the groundmass consists almost entirely of alteration products, chiefly smectites, but in some cases zeolites, chlorite, and possibly talc. The intergranular basalts grade into the intersertal basalts, consist of fineto medium-grained plagioclase and augite, and are either diabasic (roughly equigranular) or glomeroporphyritic, with a groundmass of finer-grained plagioclase, clinopyroxene, titanomagnetite. Chromite, as in Hole 442A, is found in many of these rocks as an accessory phase, principally enclosed in plagioclase, occasionally in pyroxene, and rarely in the groundmass. The augite is also similar to that in the Hole 442A basalts, but is strongly zoned in the coarser-grained rocks and grades from a pale olive green at the center to brown at the rims. Plagioclase is generally calcic, usually around An<sub>80</sub> or higher.

Variolitic textures were found in sub-units 1G and 1H. These rocks consist of randomly oriented and radiating spherulites of plagioclase in a cryptocrystalline groundmass consisting of feathery plagioclase and clinopyroxene microlites.

Sub-units 1B and 1C contain sparsely phyric basalt with close to 1 per cent plagioclase phenocrysts (0.5-1.0 mm) and very rare clinopyroxene phenocrysts in an intergranular groundmass of 0.1-mm pyroxene granules,



Figure 7. Recovery, lithology, and characteristics of basalts, Hole 442B.

0.1-0.2-mm plagioclase laths, titanomagnetite, and smectite.

Sub-units 1I and 1O are fine-grained, intersertal basalts consisting of plagioclase needles from 0.2 to 3.0 mm long, in a very fine-grained mass of quench plagioclase and pyroxene microlites, and smectites.

Sulfide droplets occur in five out of nine coarsestgrained basalts. These droplets appear to be primarily pyrite and pyrrhotite.

The unit 1 basalts are generally only slightly to very very moderately altered. In a third of the thin sections, however, smectite and other clays are abundant, replacing the groundmass minerals. Alteration is present in a number of the sub-units, notably sub-units 1E and 1K. In these, a green, low-relief mica appears to replace the groundmass, along with a clear, high-birefringence, mica-like mineral (clay or talc?) and a brown, pleochroic, low-relief, high-birefringence, biotite-like mineral which has bird's-eye extinction (talc?). The cavities in the locks where this mineral is found yield a soft, talclike mineral, which we have also tentatively identified as talc. Also present in the same unit are zeolite-filled amygdules. The other principal forms of alteration are nearly ubiquitous calcite-filled amygdules, which may amount to as much as 5 per cent by volume, and numerous calcite-filled veins. Pyrite crystals were also observed, with calcite, in vugs in sub-unit 1K.

### Unit II, Pillow Lavas

Pillow lavas were recovered only in the bottom 100 meters of Hole 442B, and are interrupted by two massive units consisting of coarser-grained material (Figure 7). Twenty per cent of the many basalt fragments recovered in the core have glassy chill margins. The basalts are again very vesicular (30-50%). The vesicles are usually 1 mm or less in size, but can reach 5 to 7 mm. Pipe vesicles are also found, oriented at right angles to the chill zone.

The pillow basalts grade inward from a glass zone, through variolitic basalt, into fine-grained, relatively crystalline, intergranular or intersertal basalt.

In thin section, the glass zones may contain plagioclase, clinopyroxene, or olivine microphenocrysts, although olivine microphenocrysts were found ony in glass from the lower half of the pillow-basalt section (Cores 442B-14 to 442B-20). Olivine generally appears to be absent in the relatively crystalline interiors of the pillows, which consist of plagioclase and clinopyroxene with accessory magnetite. The plagioclase appears to be fairly calcic ( $\sim An_{80}$ ), and the clinopyroxene appears to be augite.

There is little alteration of the pillow lavas. Only a few vesicles contain clay or calcite. In a few intervals, there are caps of limestone on the pillow fragments.

### Geochemistry of Site 442 Basalts

The geochemistry of these basalts is discussed in detail elsewhere in this volume. Major- and traceelement analyses from the shipboard XRF sampling program indicate that these basalts are tholeiites, generally lying within the compositional limits of abyssal tholeiites from both mid-ocean ridges and marginal basins. The aphyric pillow basalts in the lower section of Hole 442B, unit 2, which contain minor olivine microphenocrysts, have anomalously low Cr contents (35–64 ppm). These Cr contents are similar to those of islandarc tholeiites, but other island-arc tholeiite features, such as low Ni contents, extensive range of Fe/Mg ratios, and low abundances of incompatible trace elements, combined with evolved major-element contents, are not present.

### Interpretation

The massive cooling units drilled at Hole 442A and above the pillow basalts at Hole 442B are something of an enigma. The correlation of the massive basalts at the two holes suggests that they are laterally continuous. It is evident from the lack of pillow rinds, the relatively coarse textures, and the nearly-continuous recovery of a single unit in Cores 442A-7 and 442A-8 that these units are not pillow lavas. Even where recovery during drilling of pillow lavas has been poor on previous DSDP legs, glassy margins and pillow rinds have been well represented in the recovery. The presence of glassy margins and fine-grained to aphanitic textures in Core 442A-6 suggests that there may be some intercalated pillows. It is clear, however, that these units must be largely shallow intrusives, or surface flows which for some reason did not develop pillow structures. One possibility is that the basalt was intruded under a sediment cover, some of the sub-units representing sills intruded contemporaneously into the others. This would help explain the paucity of glass in the recovered material (longer cooling intervals).

Chromian spinel is commonly found in oceanic basalts as an inclusion in olivine and plagioclase, and only rarely in pyroxene. Dick (1975) has suggested that this is a result of early crystallization of spinel and the late appearance of pyroxene on the liquidus of most oceanic basalts, and the high solubility of  $Cr_2O_3$  in pyroxene. Thus, spinel is likely to precipitate with pyroxene only where the pyroxene is close to saturation with respect to  $Cr_2O_3$ . Accordingly, one might assume that pyroxene appears on the liquidus of some of the 442A basalts earlier than it does in most abyssal tholeiites.

Vesicular basalts are not unusual in mid-ocean-ridge basalt suites, however, at depths greater than 1000 to 2000 meters, they are rare. Studies of vesicularity of basalts along the Reykjanes Ridge suggest a strong correlation between decreasing vesicularity and increasing water depth. Shipboard studies of the sediment immediately overlying the basalts indicate a water depth near the CCD ( $\sim$  4000 m). Accordingly, unless precipitous and extraordinary subsidence occurred after eruption of the basalts, it is probable that these basalts contained volatiles in excess of those normally found in midocean-ridge basalts.

Compared to basalts of layer 2A drilled on previous DSDP legs, and dredged from many localities throughout the world, the Site 442 basalts are extraordinarily vesicular. In addition, these basalts generally lack phenocrysts. Such basalts are not unusual in other oceanic regions, but these features are noteworthy. The frequency with which chromian spinel occurs as an inclusion in pyroxene is also unusual, as are the low Cr contents of some of the 442B pillow basalts. It is fair to say, then, that these basalts are not entirely typical of what is commonly perceived to be mid-ocean-ridge abyssal tholeiite.

# PALEOMAGNETISM

#### Sedimentary Layers

Paleomagnetism samples for sedimentary cores of Holes 442A and 442B were taken on the average every 1.5 meters in the recovered cores. The positions of samples are listed in Table 6. All samples were cut into cylinders and put in a plastic vial (2.3-cm diameter, 2.3-cm length). Measurements of the natural remanent magnetization (NRM) and the remanent magnetism after alternating-field demagnetization (AFD) were carried out by means of a shipboard flux-gate-type spinner magnetometer (Balanced Fluxgate Rock Magnetometer, Digico Limited).

After measurement of natural remanent magnetization, all the samples were washed magnetically in a unitaxial AF demagnetizer (AC Geophysical Specimen Demagnetizer, Schonstedt) with a peak AF of 200 oe, decreasing to zero at a constant rate of 20 milligauss/cycle. The cylindrical axis of the specimen was aligned paralled to the demagnetizing axis of the demagnetizer. Stationary DC magnetic field, mixture of the geomagnetic field and ship-induced magnetic field, were eliminated by a three-layered, highly permeable metal shield surrounding the AF demagnetizer.

Results of the measurements are listed in Table 6, and a schematic diagram of normal and reversed remanent magnetization of the sedimentary cores is given in Figure 8. Inclination of the original NRM of both the 442A and 442B samples in relation to sub-bottom depth is shown in Figure 9. There is a large scatter in the inclination values which might have been caused by mechanical disturbances of the drill bit. Therefore, a statistical treatment is applied to cores every 50 meters subbottom depth. The results of the statistical calculation are listed in Table 7 and are also plotted in Figure 9, with the standard deviation around mean values.

#### **Basaltic Layers**

Eight cylindrical core samples from Hole 442A, and 42 samples from Hole 442B, were taken for study of paleomagnetism. Of the 50 samples, 42 were AF demagnetized in a stepwise manner (up to maximum 600 oe, with 50 oe step) to find the stability of the NRM. Measurements of remanent magnetizations were carried out on six spinning axes. Demagnetization was made along three mutually perpendicular orientations of each sample. Sampling depths in the cores and the results of measurements are listed in Table 8. Inclination of NRM versus core-section number of basaltic layers for both 442A and 442B are plotted in Figure 10. Maximum demagnetizing field varies from specimen to specimen, depending on their magnetic hardness, as shown in Figure 11. Therefore, the inclination values after AFD were taken at the nearest higher value of MDF. Inclination values of remanent magnetization after AFD are also plotted in Figure 10. Statistical analysis of NRM intensities and AFD inclinations were carried out. Specimens were divided into three groups: (1) all samples of Hole 442A, (2) normally magnetized layer of Hole 442B, and (3) reversely magnetized layer of Hole 442B. Statistical results are given in Table 9.

#### Interpretations

The results of paleomagnetism measurements are summarized as follows:

1. Average intensity of NRM is around  $3 \times 10^{-5}$  gauss/cm<sup>3</sup>, in good agreement with values for typical oceanic basalts.

2. Stability of NRM in both sediments and basalts is sufficiently high to provide data about the paleomagnetic-field direction.

3. Sequences of normal and reversed polarity in the lowermost portion of the sedimentary layers from both Hole 442A and Hole 442B are identical. The bottom of the sedimentary layers has a normal polarity of NRM, which continues to the top of the underlying basaltic layers.

4. Means of the NRM inclination of sediments are close to the mean inclination of the recent geomagnetic dipole field (estimated:  $47.9^{\circ}$ ) around Site 442, although the scatter of inclinations of samples is large (~  $20^{\circ}$ ).

5. NRM inclinations of sediments and basalts are nearly identical.

6. AFD inclinations of the basalts of Hole 442A and the Hole 442B normal group differ by about 9 degrees, and their standard deviations overlap little.

7. AFD inclinations of the Hole 442B normal group and reversed group seem to align in antipodal directions.

Taking into account these facts, we can conclude that:

1. High stability and high  $Q'_n$  values (often exceeding 500) imply that contribution of an induced magnetization to the magnetic anomalies is negligible in the uppermost part of layer 2A of the oceanic crust at this site.

2. The top of the basalts and the lowermost part of the sediments seem to be of the same age. According to shipboard paleontological study of Section 442A-30-3, this age is about 16 to 18 m.y., which includes normal polarity periods with anomaly numbers 5D and 5E.

3. Site 442 has remained at almost the same latitude during the last 17 m.y.

4. Basalts of Holes 442A and 442B were formed at different times, or one of them was relatively tilted by more than 9 degrees. If they were formed at different times, and if the present rate of the secular variation of the geomagnetic field has held for the last 17 m.y., the interval of formation of basalts of Holes 442A and 442B was at least about 1000 years.

5. The time interval of the formation of normal and reversed groups of basalts of Hole 442B was not shorter than 5000 years as indicated by switching time of geomagnetic polarities.

Sample	Depth	J <sub>NRM</sub>	Susceptibility	Incli	ination		
(interval in cm)	(m)	(10 <sup>-5</sup> gauss)	(10-5 gauss/oe)	NRM	AFD	Polarity	Age
442A-2-1, 91-93	10.42	2.09		10.2	39.5	+	
2-2	-	1.57	-	5.0	3.6	+	
2-4, 71-73	14.72	0.54	-	-5.2	0.0	-	
2-5, 15-15	17.37	2.75		-68.3	-75.2	2	
3-2, 30-32	20.81	1.76		-77.5	-32.0	-	0.2 m v
3-3, 46-48	22.47	1.00		71.4	70.9	+	0.5 m.y.
3-4, 53-55	24.04	2.41	-	66.1	32.6	+	
3-5, 100-102	26.01	0.72	-	-60.4	-4.9	1	
3-7, 68-70	28.69	0.45	_	-75.0	-7.5	-	
4-2	-	2.00	_	69.8	37.0	+	
4-3, 20-22	31.71	25.3	-	-10.0	-15.0	-	1.00
4-4, 35-37	33.36	1.64	-	45.7	24.5	+	0.9 m.y.
4-5, 7-9	34.58	0.21		41.9	-2.8	-	
5-2, 106-108	40.57	0.90	-	58.8	18.9	+	
5-3, 73-75	41.74	8.05	-	-0.4	+0.5	±	
5-4, 131-133	43.82	1.29	-	19.3	10.3	+	
5-5, 118-120	45.19	0.12	-	15.3	0.4	t	
6-1, 121-123	48.72	0.07	0.45	71.1	15.4	Ţ	
6-3, 108-110	51.59	0.97	0.49	-37.6	-16.7	2	
6-4, 45-47	52.46	0.57	0.58	-57.4	-9.8	-	
6-5, 10-12	53.61	0.37	0.59	-28.7	-5.3	-	
6-6, 5-7	55.06	0.58	0.55	-16.8	-3.3	2	
7-2, 34-36	58.85	0.97	0.51	-39.9	-11.3	-	
7-3, 34-36	60.35	0.50	0.51	-14.2	-1.8	-	
7-4, 34-36	61.85	0.88	0.54	-62.7	-18.8	-	
7-5, 34-36	63.35	1.16	0.49	-22.3	-9.0	-	
8-4, 45-47	71.46	0.87	0.41	-53.7	-21.9	-	
10-3, 11-13	88.62	0.69	0.71	-64.0	-24.4	-	
10-4, 11-13	90.12	0.34	0.77	-43.4	-10.3	-	
11-3, 6-8	98.07	1.01	0.69	41.4	25.7	+	
13-2, 100-102	116.51	1.59	0.74	71.0	58.3	+	
14-3, 99-101	127.50	0.54	0.84	-38.1	-2.2	-	
15-3, 44-46	136.45	0.83	0.75	-58.6	-17.9	1	
15-4, 9-11	137.60	1.29	0.86	-60.9	-17.1	-	
16-1, 93-95	143.44	1.81	0.67	44.4	48.8	+	
16-2, 33-35	144.34	0.22	0.67	12.9	2.0	+	
18-2 135-137	154.55	1.88	0.73	-+2.1	-25.4	+	
18-3, 7-9	164.59	0.86	0.78	-57.8	-13.5	-	1.6 m.v.
19-1, 130-132	172.31	2.95	0.55	38.1	24.4	+	
19-2, 60-62	173.11	0.13	0.51	47.9	2.5	+	
19-3, 88-90	174.89	2.70	0.59	53.4	36.2	+	
21-1, 16-18	190.17	0.35	0.60	-45.9	-10.6	3	
21-2, 16-18	191.67	0.39	0.82	-37.1	-2.3	30	
21-3, 16-18	193.17	2.22	0.66	-34.1	-13.5	-	
21-4, 16-18	194.67	1.36	0.54	88.3	87.5	+	
21-4, 20-22	194.71	2.14	0.53	46.3	32.2	1	
23-3, 57-59	212.56	3.26	0.69	3.5	6.0	+	
23-4, 143-145	214.94	2.70	0.71	0.6	0.9	±	
23-5, 16-18	215.17	3.19	0.75	-42.0	-41.2	-	
24-1, 128-130	219.79	1.42	0.65	-40.1	-32.8	1	
24-2, 140-142	222.01	0.66	1.02	-75.8	-85.8	2	
24-4, 50-52	223.51	0.22	0.72	-22.0	-24.6	-	
24-5, 50-52	225.01	3.29	1.20	-74.8	-76.1	-	
25-2, 111-113	230.52	0.26	0.92	23.8	38.4	+	
25-3, 126-128	232.27	1.74	1.13	-0.3	-10.1	-	
26-2, 100-102	240.01	4.80	0.92	36.8	39.9	+	13 m.y.
27-2, 100-102	249.51	88.6	1.39	33.5	32.5	+	
28-1, 62-64	257.13	276.1	1.50	46.5	48.8	+	14 m.y.
28-2, 53-55	258.54	2.14	1.01	-10.8	-51.1	-	15 m.y.
29-3. 85-87	269.86	1.19	1.12	-38.4	-29.6	-	
30-1, 51-53	276.02	2.30	1.57	-31.7	30.1	+	
30-2, 22-24 30-3, 76-78	277.23 279.27	1.05 1.22	1.29 1.14	-19.9 70.0	-36.8 67.0	-+	17 m.y.
442B-1-1, 76-78	268.27	0.99	0.48	50.4	-9.9	?	
1-2, 24-26	269.25	0.00	-			?	
1-2, 137-139	270.38	7.16	0.99	44.9	-55.2	+	
2-1, 30-38	278.66	2 46	0.95	-62.0	-64 5	2	
2-3, 9-11	280.10	0.26	0.89	-84.1	-57.7		
2-4, 62-64	282.13	8.64	1.20	-26.4	-48.4	-	
2-5, 47-49	283.49	47.86	1.35	20.5	18.6	+	
3-2, 40-42	288.41	21.12	0.92	11.3	14.0	+	
3-2, 115-117	289.10	14 82	0.97	30.7	73.0	1	

 TABLE 6

 Paleomagnetism Measurements of Sedimentary Cores from Holes 442A and 442B<sup>a</sup>

<sup>a</sup>AFD is obtained by peak alternating demagnetizing field of 200 oe, decreasing to zero at a constant rate of 20 milligauss/cycle; polarity shows whether the inclination of NRM is positive (+) or negative (-); absolute age determined by shipboard paleontological studies.



Figure 8. Remanent-magnetization diagram of sedimentary cores from Holes 442A and 442B. (See Table 6.) Polarity diagrams for some basaltic cores next to the bottom layer of the sediments are also shown.



Figure 9. Absolute values of inclination of NRM versus sub-bottom depth of the sedimentary layers. Data of Holes 442A and 442B are mixed. Mean values of the inclination taken every 50 meters; vertical bars represent standard deviations.

# PHYSICAL PROPERTIES

Physical properties measured on sediments and basalts recovered from Holes 442A and 442B included sonic velocity, density, porosity, water content, shear strength, and thermal conductivity. The analog output from the GRAPE unit provided additional approximations of density and porosity. Sonic velocities were determined at atmospheric pressure and room temperature by measuring the travel time of a sound wave through a known length of material sandwched between two barium-titanate transducers on the Hamilton Frame Velocimeter. Wave-propagation directions were perpendicular to the core axis in sediments, and parallel to the core axis in basalts. All basalt samples were kept water-saturated until their measure-

		Su	b-bottom Dept	th (m)			
	50	)	100	150	200	250	300
Mean Inclination	45	37	43	48	30	37	
Standard Deviation	27	17.5	21	23	24	19	
Dispersion	736	284	395	471	553	316.5	

 TABLE 7

 Statistics of NRM Inclinations of Sedimentary Cores of Holes 442A and 442B<sup>a</sup>

<sup>a</sup>Data are grouped every 50 Meters of Sub-bottom Depth; mean, standard deviation, and dispersion of individual 50-meter intervals are listed.

ment. Wet-bulk densities, porosities, and water content for syringe and chunk samples were calculated from relationships in Boyce (1976). Physical properties are tabulated in Table 10 for sediments and in Table 11 for igneous rocks. Although the data from the analog GRAPE will be referred to, the analog output will not be reproduced here. Special 2-minute GRAPE counts were made on selected basalt samples. A grain density of 2.924 g/cm<sup>3</sup>, the average Site 442 basalt grain density, was assumed to calculate the porosities and wet-bulk densities listed in Table 12. Detailed treatment of the GRAPE technique can be found in shipboard manuals and in Boyce (1976).

Table 10 summarizes data from shear-strength measurements on clay samples from Holes 442A and 442B performed with the Soiltest Torvane. Shear-strength values were obtained by reading directly the shear strength from the instrument gauge and calculating the shear strength using the strength-correlation relationship included with the Torvane instructions. The axis of rotation of the vane was perpendicular to the core axis in all determinations. The shear-strength data show no apparent relationship to depth, as recognized on Leg 31 (Bouma and Moore, 1976). When the data are examined in comparison to the lithologic units recognized at Hole 442A, however, definite relationships are evident (Figure 12).

Shear strength increases with depth in the upper 85 meters of unit I, apparently reflecting the transition from silty clays in the upper 38 meters to higher-shearstrength clays below. 'A distinct decrease in shear strength occurs within unit I at 85 meters, a depth at which a considerable amount of drilling breccia was recovered. Shear strength increases again to 164 meters, the bottom of unit I, and decreases to a very low value in unit II. Several sections of drilling breccia were recovered in the upper part of unit II. Unit II is characterized by low shear strength. A third increase in shear strength is evident in unit IIIa. Strength is significantly reduced below 240 meters, where two sections of drilling breccia were recovered. Value of shear strength vary considerably in unit IIIb. The low values were obtained from ash units or intervals with a large ash component. The high shear-strength values are from intervals with a large clay component. The high values for unit IV apparently reflect the high clay content and better lithification of this unit.

The most significant aspects of the shear-strength data are the three intervals with systematic increases of shear strength with depth, followed by large drops in shear strength below. Each shear strength drop is characterized by the recovery of drilling breccia. During drilling a certain shear stress is exerted by the drill bit. If a lithologic unit of low shear strength is penetrated, the yield strength is exceeded, producing drilling breccia. It may be possible to use relationships similar to those presented in figure 12 to predict decreases in shear strength with depth and, therefore, to enhance recovery. This finding may be particulary useful for future drilling near Site 442.

Figure 13 shows sonic velocity as a function of depth. With few exceptions, velocities for sediments deviate little from the average of 1.58 km/s for the entire sediment thickness of 287 meters. Although distinctive lithologic units are present at this site, the sequence is dominated by clay minerals, producing a monotonous acoustic stratigraphy. Wet-bulk densities (Figure 13) also show little deviation from the average of 1.47 g/cm<sup>3</sup>. The estimated average wet-bulk density for the sediments from the analog GRAPE is 1.54 g/cm<sup>3</sup>, in close agreement with the laboratory values.

Thermal-conductivity values for sediments, summarized in Table 10, were measured using the transienthot-wire method on the shipboard Quick Thermal Conductivity Meter manufactured by Showa Denko K. K. The stated accuracy of the meter is 5 per cent. No corrections were made for ocean-bottom pressures and temperatures (Ratcliffe, 1960). Figure 13 shows the variation of thermal conductivity with depth in the hole. Values of thermal conductivity of sediments are well within the range of values observed for deep-sea sediments (e.g., Clark, 1966). The average for Hole 442A sediments is 1.965 mcal/cm-s-°C. No systematic variation of thermal conductivity was observed with depth. A systematic relationship between per cent water content and thermal conductivity reported by previous workers (e.g., Ratcliffe, 1960) was not observed for sediments from Hole 442A. This may be the result of lack of thermal equilibration between the samples and ambient laboratory temperatures; the samples were given 4 hours to equilibrate, but this time may be insufficient.

Porosities in the sediments range from 48 to 85 per cent, average 72.2 per cent, and show no obvrious trend with depth. These exceptionally high porosities result

	TABL	E 8			
Paleomagnetism	of Basalts,	Holes	442A	and	442B <sup>a</sup>

$ \begin{array}{c c c c c c c c c c c c c c c c c c c $		Sub-bottom							
	Sample	Depth	JNRM	Inclination	Inclination	MDF	Xin	0'	-
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	(interval in cm)	(m)	(10 <sup>-5</sup> gauss)	NRM	AFD	(oe)	(10 <sup>-5</sup> gauss/oe)	$Q_n$	Remarks
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	442A-31-1, 90-92	285.91	118.1	35.3	35.8	360	1.13	240	350 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	32-2, 62-64	296.61	104.2	26.2	-		1.08	224	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	33-1, 136-138	305.37	346.5	33.3	36.1	260	1.82	440	250 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	33-2, 101-103	306.52	109.6	33.4	31.9	290	0.92	275	300 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	33-3, 79-81	307.80	134.0	31.4	34.6	240	1.15	269	250 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	33-4, 52-54	309.03	124.0	29.0	30.0	165	1.52	188	200 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	34-1, 51-53	314.02	424.7	41.4	41.9	240	1.59	656	250 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	34-1, 125-127	314.76	300.6	44.2	42.6	180	1.82	381	200 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	442B-3-3, 118-120	290.69	537.7	27.7	-	-	1.68	739	
	3-4, 54-56	291.55	301.1	39.0	45.3	260	1.73	402	300 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	4-1, 117-119	297.18	420.6	50.8	51.8	320	1.87	519	350 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	4-2, 118-120	298.67	119.9	32.0	(50.8)	n.d.	2.07	133	450 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	5-1, 116-118	306.67	224.0	40.6	41.0	220	1.79	289	250 oe
$  \begin{array}{ccccccccccccccccccccccccccccccccccc$	5-2, 1-3	307.02	186.8	45.5	46.5	200	2.17	199	250 oe
	5-3, 37-39	308.86	386.7	44.2	_		2.94	304	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	6-1, 48-50	315.49	272.2	63.4	65.5	320	1.36	461	350 oe
	6-1, 110-112	316.11	552.6	55.6	56.7	270	1.93	660	300 oe
	6-2, 110-112	317.61	568.8	49.2	-		2.71	484	1074047 - 9431 1
7-1, 59-61325.10211.355.0 $   7.37$ 667-1, 106-108325.57164.860.354.6706.5258100 oe8-1, 47-49334.48456.450.951.97010.6899100 oe8-2, 24-26335.75902.747.8 $ -$ 9.162278-3, 55-57337.56465.748.950.9707.14151100 oe8-4, 3-5338.54255.052.250.0908.0673100 oe8-5, 23-25340.24446.848.751.3708.48122100 oe8-6, 84-86342.35381.445.5 $ -$ 9.71918-7, 107-109343.08261.348.048.312011.0355150 oe9-1, 24-26343.75455.8 $-1.2$ $-3.7$ 1002.92360150 oe9-1, 33-35343.84448.9 $-3.1$ $-2.1$ 903.44301100 oe9-2, 73-75345.741940.147.647.8802.801599100 oe9-3, 105-107347.56437.643.249.91302.34432150 oe9-3, 107-107347.56437.643.249.91302.34432150 oe11-1, 60-62363.11108.233.328.93601.69148400 oe12-1, 12-129365.26	6-3, 28-30	318 29	105.3	67.7	477	80	7.87	31	100 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	7-1, 59-61	325 10	211.3	55.0	-	-	7 37	66	
$k_{11}$ $k_{12}$ $k_{12}$ $k_{13}$ $k_{14}$ $k_{15}$ $k_{16}$ $k_{16$	7-1 106-108	325 57	164.8	60.3	54.6	70	6.52	58	100 06
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	8-1 47-49	334 48	456.4	50.0	51.0	70	10.68	99	100 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	8-2 24-26	335 75	902.7	17.9	51.9	70	9.16	227	100 00
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	8-3 55-57	337 56	165 7	47.0	50.0	70	7 14	151	100 08
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	8-3, 33-37	229 54	403.7	40.9	50.9	00	2.06	73	100 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	8 5 32 35	240.24	233.0	32.2	51.2	20	0.00	122	100 00
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	9 6 94 96	242.25	201 4	40.1	51.5	70	0.40	01	100 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	87 107 100	242.55	361.4	45.5	10.2	120	9.71	55	150 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	0.1.24.26	343.08	201.5	48.0	48.3	120	11.05	260	150 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	9-1, 24-26	343.75	455.8	-1.2	-3.7	100	2.92	201	100 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	9-1, 33-35	343.84	448.9	-3.1	-2.1	90	3.44	301	100 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	9-1, 103-105	344.54	544.8	53.8	56.3	100	3.38	3/1	150 00
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	9-2, 73-75	345.74	1940.1	47.6	47.8	80	2.80	1599	100 oe
9-3, 105-107 $347,56$ $437.6$ $43.2$ $49.9$ $130$ $2.34$ $432$ $150$ oe11-1, 60-62 $363.11$ $108.2$ $33.3$ $28.9$ $360$ $1.69$ $148$ $400$ oe11-2, 127-129 $365.28$ $57.30$ $48.8$ $  1.27$ $104$ 11-3, 18-20 $365.69$ $112.0$ $43.7$ $42.1$ $190$ $2.16$ $120$ $200$ oe12-1, 40-42 $372.41$ $90.96$ $31.7$ $27.3$ $380$ $1.21$ $173$ $400$ oe12-1, 114-116 $373.15$ $89.80$ $32.6$ $29.8$ $490$ $1.45$ $143$ $500$ oe13-1, 60-62 $382.11$ $102.3$ $52.6$ $38.1$ $340$ $1.80$ $131$ $350$ oe13-2, 100-102 $384.01$ $93.97$ $37.7$ $34.9$ $380$ $1.44$ $151$ $400$ oe $13-3, 2-4$ $384.53$ $156.4$ $22.7$ $20.4$ $460$ $1.64$ $220$ $500$ oe $14-1, 38-40$ $391.39$ $      16-1, 29-31$ $410.30$ $504.8$ $-47.4$ $   1.82$ $640$ $16-1, 35-37$ $410.36$ $154.6$ $-42.2$ $-43.5$ $320$ $2.07$ $172$ $350$ oe $17-1, 104-106$ $420.55$ $491.2$ $-59.9$ $-60.4$ $480$ $1.66$ $682$ $500$ oe $17-1, 104-106$ $420.55$ $491.2$ $-59.9$ $-60.4$ $480$ <td>9-3, 32-34</td> <td>346.83</td> <td>768.1</td> <td>45.7</td> <td>45.8</td> <td>260</td> <td>2.00</td> <td>886</td> <td>300 oe</td>	9-3, 32-34	346.83	768.1	45.7	45.8	260	2.00	886	300 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	9-3, 105-107	347.56	437.6	43.2	49.9	130	2.34	432	150 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	11-1,60-62	363.11	108.2	33.3	28.9	360	1.69	148	400 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	11-2, 127-129	365.28	57.30	48.8	1.500		1.27	104	
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	11-3, 18-20	365.69	112.0	43.7	42.1	190	2.16	120	200 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	12-1, 40-42	372.41	90.96	31.7	27.3	380	1.21	173	400 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	12-1, 114-116	373.15	89.80	32.6	29.8	490	1.45	143	500 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	13-1, 60-62	382.11	102.3	52.6	38.1	340	1.80	131	350 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	13-2, 100-102	384.01	93.97	37.7	34.9	380	1.44	151	400 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	13-3, 2-4	384.53	156.4	22.7	20.4	460	1.64	220	500 oe
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	14-1, 38-40	391.39	111.6	47.0	45.4	520	1.15	224	500 oe
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	14-1, 38-40	391.39	_		44.8	_			
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	16-1, 29-31	410.30	504.8	-47.4		-	1.82	640	
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	16-1, 35-37	410.36	154.6	-42.2	-43.5	320	2.07	172	350 oe
$ \begin{array}{cccccccccccccccccccccccccccccccccccc$	16-1, 105-107	411.06	349.2	-45.2	-46.6	310	1.25	644	350 oe
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	17-1, 104-106	420.55	491.2	-59.9	-60.4	480	1.66	682	500 oe
19-2, 62-64       440.63       182.8       -27.8       -36.4       260       2.62       161       300 oe         20-1, 27-29       448.28       387.9       -22.1       -21.6       460       1.88       476       500 oe	19-1, 43-45	438.92	217.2	-31.0	-34.3	180	2.34	214	200 oe
20-1, 27-29 448.28 387.9 -22.1 -21.6 460 1.88 476 500 oe	19-2, 62-64	440.63	182.8	-27.8	-36.4	260	2.62	161	300 oe
	20-1, 27-29	448.28	387.9	-22.1	-21.6	460	1.88	476	500 oe
Orientatio	, -,			I	2110		1100		(Orientation

<sup>a</sup>MDF is the median destructive field (of AF demagnetization) at which the remanent magnetism of a specimen decreases to 50% of its initial value;  $X_{in}$  is the initial susceptibility of a specimen;  $Q'_{in}$  is the Konigsberger ratio of NRM; peak field strength listed in remarks column is that at which inclination of AFD remanent magnetization was taken; for other notations, refer to Table 6.

from the loss of adsorbed water during heating to 110°C for water-content determinations.

The sonic velocity through the limestone recovered at 285.4 meters is 3.75 km/s, producing a major velocity discontinuity (Figure 13). The limestone is directly underlain by basalts at Hole 442A, which yield sonic velocities ranging from 3.85 to 5.25 km/s (Sample 442B-19-2, 12 cm). Velocities are variable in the upper

25 meters of basalt (Figure 14) and increase between 324 and 334 meters, a zone of gray, glomeroporphyritic basalt. Below this zone, velocities drop to 3.9 to 4.2 km/s. The variability of sonic velocities in basalt increases below 410 meters, with a maximum of 5.25 km/s and a minimum of 4.19 km/s. A major discontinuity in the basalt stratigraphy may occur at the 410-meter level (see section on magnetics). Wet-bulk densities follow the



Figure 10. Inclination of remanent magnetization and median destructive field (MDF) of basalts, plotted against sample positions in the cores. (Solid circles and squares represent positive inclination, normal polarity; open circles and squares represent negative inclination, reversed polarity.)

same trend with depth (Figure 14) as do the sonic velocities. This result is not surprising, because sonic velocity is linearly related to density (Figure 15).

The average wet-bulk density of basalts recovered from Holes 442A and 442B is 2.50 g/cm<sup>3</sup>, lower than

the average oceanic-basalt density (Christensen and Salisbury, 1975). The variation with depth of wet-bulk density determined from laboratory measurements and two-minute GRAPE counts is shown in Figure 15. The low densities are caused by the relatively high porosities



Figure 11. Stability of NRM with respect to AF demagnetization. The horizontal axis represents the peak alternating magnetic field, decreasing to zero at a constant rate.  $J_{NRM}$  and  $J_{AFD}$  are intensities of NRM and AFD remanent magnetization of a specimen.

TABLE 9 Statistical Treatment of Paleomagnetism Results of Basalt Samples, Holes 442A and 442B<sup>a</sup>

	11	nclination (degrees)	1	Intensity (10 <sup>-5</sup> gauss/cm <sup>3</sup> )			
Sample Set	Number of Samples	Mean	Standard Deviation	Number of Samples	Mean	Standard Deviation	
Hole 442A, all	7	36.2	4.2	8	208.5	129.5	
Hole 442B, normal	26	45.3	10.0	35	364.5	354.0	
Hole 442B, reversed	6	40.7	13.2	7	326.9	144.5	

<sup>a</sup>Samples from Section 442B-9-1 showed anomalous orientations of NRM which are not included here.

of the vesicular basalts. Figure 15 shows the relationship between basalt porosity and wet-bulk density. The lines on Figure 15 correspond to the predicted influence of porosity on wet-bulk density for basalts with near-ideal grain densities of 2.9 and 3.0 g/cm<sup>3</sup>. Because the data for Site 442 basalts plot between the theoretical lines, it is apparent that mineral alteration does not, in this case, significantly influence wet-bulk density, and that porosity is the dominant control of wet-bulk density. This observation is confirmed by the high grain densities of the Site 442 basalts (Table 11). Christensen and Salisbury (1972) attributed low densities of oceanic basalts to the effects of submarine alteration. In this case, however, the low wet-bulk densities are directly related to the unusually high vesicularity of the basalts from Holes 442A and 442B.

Sonic velocity and wet-bulk density for Site 442 basalts are linearly related (Figure 16), with a correlation coefficient of 0.94. The slope is 3.06, and the intercept is -3.32; these values are significantly different from the parameters of the velocity-density relationship established by Christensen and Salisbury (1972) for

deep-sea basalts. This discrepancy results from importance of porosity of Site 442 basalts. Figure 17 shows an excellent inverse relationship between porosity and sonic velocity, with a correlation coefficient of -0.97. This is the type of relationship expected when waterfilled pore space influences sonic velocity of rocks. Clearly, wet-bulk densities and sonic velocities of Site 442 basalts are controlled by the porosity, and mineral alteration has little or no effect.

Thermal conductivity of basalts is variable (Table 11) and averages 3.40 mcal/cm-s-°C. The variability is probably related to the variable porosity, which determines the per cent water saturation. In general, increased water saturation will increase the thermal conductivity of rocks (e.g., Clark, 1966).

The average sonic velocity for the recovered basalts is much higher than the velocity for layer 2A (2.8 km/s) determined in a sonobuoy survey near this site. This discrepancy requires that unrecovered lower-velocity material must be interlayered with the basalt, or that the basalt must be severely fractured. Using the porosityvelocity relationship discussed above, the large-scale formation porosity would need to be considerably in excess of 30 per cent to explain the observed layer 2A velocities. This situation was encountered on Leg 37 while drilling the FAMOUS area (e.g., Hyndman, 1977). Alternatively, the sonobuoy profile for the nearby area may not be appropriate to this particular site.

### CORRELATION OF GEOPHYSICAL DATA WITH DRILLING RESULTS

## Introduction

Site 442 was chosen to be located on shot point 3550, line 2-2 of the S/S *Kaiyo-Maru* multichannel seismicreflection profile (shot-point spacing 50 m, standard processing with 24-fold stack, deconvolution and timevariant filter) (Figure 1). The site is also on the western shoulder of a positive magnetic anomaly with appreciably large amplitude (about 300 gammas peak to peak) and wavelength (about 30 km) which was identified as anomaly 6 (19–20 m.y.), based upon its characteristic shape, with the aid of age information from DSDP Site 297 (Karig, Ingle, et al., 1975; Watts and Weissel, 1975; Kobayashi and Nakata, 1977).

A sonobuoy measurement was made along a NNW-SSE line about 20 km from the present site (Murauchi and Asanuma, per. comm.).

We attempt to correlate these site-survey data, and underway geophysical observations recorded when approaching and leaving the site, with the shipboard data about lithologies, paleontological ages, physical properties, and paleomagnetism direction and intensity of the cores recovered from various sub-bottom depths.

## Sonic Velocity and Sub-Bottom Depth

Both previous and underway seismic-reflection profiles show a layering of semi-transparent sediment with two-way normal time of about 0.37 second overlying the acoustic basement. Shipboard measurement of the sonic velocity indicated that the  $V_{\rm P}$  of sediment recovered is

Sample (interval in cm)	Lithology	Sonic Velocity (km/s)	Thermal Conductivity (mcal/cm-s-°C)	Shear Strength (× 10 <sup>-5</sup> dynes/cm <sup>2</sup> )	Wet-Bulk Density (g/cm <sup>3</sup> )	Porosity (%)	Water Content (%)
442A-2-1, 96-111	silty clay	1.499	2.267	0.24	-	-	-
2-3, 143-145	"	-		-	1.41	77.13	56.17
2-4, 77-79	**		-		1.44	75.91	54.10
2-6, 30-33	**	1.519		0.24			-
3-2, 88-98	**	1.484	2.622		-	—	<u>~_</u> ~
3-3, 20-22	**		-	and the second sec	1.54	70.45	46.98
3-4, 36-39	55	-	_	0.72	1.000	-	—
3-4, 81-91	**	1.503	2.753	475	3 <del>11</del>		-
4-2, 107-123	**	-	2.522	0.29	-		-11 may
4-2, 135-137	55	34 <u>11</u>	-	-	1.46	76.34	53.55
4-3, 68-78	**	1.505	2.553	0.29	1.49	74.06	50.98
4-5, 55-58	**	-		0.72		-	—
5-1, 95-98	gray clay			0.53	—		-
5-3, 69-85	**	1.484	2.558	0.29	1.50	75.26	51.33
5-4, 95-98	**		1. The second	1.24	100		_
5-5, 73-83	33	1.212	1.836	1.68	-		—
6-1, 95-98	**	-		0.53	1.000	100	-
6-2, 69-72	"		-	1.15			—
6-3, 83-93	55	1.385	1.969	0.72	1.49	74.89	51.56
6-5, 115-129	**	8.555	2.525	0.38	-	—	-
7-1, 99-102	**	-	-		1.50	48.82	33.42
7-2, 48-51	**	844 - C	-	1.15		-	-
7-4, 17-27	**	1.214	2.461	2.15	1.54	67.37	44.91
7-4, 140-150	**	-	-		1.53	66.89	44.76
7-5, 90-93	59			0.53	-	-	—
8-4, 74-84	**	1.472	1.983	1.66	1.57	85.55	55.70
8-5, 80-83	**	3	-	1.63		-	_
10-3, 86-98	**	1.518	1.836	0.48	1.48	80.16	55.31
13-2, 134-144	55	1.502	1.397	1.15	1.41	74.04	53.76
13-2, 143-145	**	1. <del>-</del>	—	1.00	1.47	76.46	53.35
14-3, 89-99	**	1.536	2.192	1.63	1.47	75.43	52.74
15-2, 37-52	**	1.531	3.153	2.15	1.48	75.21	51.99
15-3, 33-36	**	-	-	2.59	-		
15-4, 33-43	**	1.544	0.987	4.21	1.49	75.56	51.90
16-2, 41-57	yellow-brown clay	2.091	1.731	2.30	1.55	71.51	47.41
18-3, 60-70	**	1.538	2.286	3.54	1.59	68.56	44.15
19-2, 70-82	**	1.515	2.083	3.93	1.53	71.62	47.92
19-3, 0-10	23	—	-		1.48	68.95	47.79
20-1,88-98	brown clay	1.889	2.361	1.63	1.58	79.92	51.71
21-1, 34-37	**	-	_	1.20		-	—
21-2, 89-92	**	1.516		1.68	1.35	61.82	46.97
21-2, 116-119	**	1.543	-	. <del></del>	0.00	-	2000
21-3, 8-11	**	5 <b>—</b>	-	1.20	2	-	$(-)_{i}$
23-2,93-95	yellow-brown clay	2 <b>-</b>		6	1.53	72.69	48.66
23-4, 5-15	**	1.528	1.339	a de terrera	-	-	
24-4, 18-28	**	1.518	3.139	0.77	1.44	76.35	54.15
24-5, 25-35	**	1.535	2.142	4.21		-	-
25-3, 25-39	39	1.528	2.031	1.82	1.50	70.65	48.31
25-4, 73-76	**	1.982	-	3.16			
26-2, 117-127		1.931	1.381		1.32	63.93	49.52
26-2, 120-121	**	-		5.08	-	-	$\sim - \sim$
26-2, 124-125	**	-	222	5.75		-	
27-3, 0-8	**			1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-1-	1.43	72.65	52.16
28-3, 78-88	**	2.753	2.325	1.15		-	-
28-3, 118-128	gray-green clay	1.580	2.072	1.44	1.28	50.98	40.86
29-1, 44-47	"			4.50	_		—
29-1, 84-99	**	1.554	1.303		1.32	79.26	61.61
29-2, 31-42	"	1.613	2.161	0.72	_	-	
29-2, 44-47	**	0.000	and and a second se	6.70	-	-	—
29-2, 99-109	.55	1.510	1.969		_	-	-
29-3, 44-47	**	3	2	9.00	_	-	. (H)
29-3, 94-104	**	1.511	1.228		1.50	75.30	51.58
30-2, 81-91	brown claystone	1.557	1.525	$\sim 12.4$	10050000 1000		9 <b>-</b> 9
30-3, 85-97	**	1.568	2.356	$\sim 12.0$	1.47	68.86	48.10
31-1, 57-60	limestone	3.747		-	-	—	_

 TABLE 10

 Summary of Physical Properties of Sediments, Site 442

Wet-Bulk Water Sonic Thermal Shear Strength Density Porosity Content Sample Velocity Conductivity  $(g/cm^3)$  $(\times 10^{-5} \text{ dynes/cm}^3)$ (interval in cm) Lithology (km/s)(mcal/cm-s-°C) (%) (%) 442B-1-1, 70-73 gray-green clav 0.48 1.550 \_ 1-2, 13-17 1.15 ----,, 1-2, 88-93 2.49 \_\_\_\_ -1.565 1-2, 123-126 yellow-brown clay 1.598 1.15 --1-3, 34-37  $\sim 11.5$ ., 1-3, 44-47 1.556 6.70 2-1, 115-118 brown claystone 79.14 1.35 2-2, 130-133 1.605 6.70 .,, 2-3, 54-57 5.84 -2-5, 12-15 ,, 1.607  $\sim 12.0$ \_ \_ ,, 3-2, 70-73 1.568

TABLE 10 - Continued

TABLE 11 Summary of Physical Properties of Igneous Rocks, Site 442

Sample (interval in cm)	Piece No.	Sonic Velocity (km/s)	Thermal Conductivity (mcal/cm-s-°C)	Wet-Bulk Density (g/cm <sup>3</sup> )	Grain Density (g/cm <sup>3</sup> )	Porosity (%)
442A-31-1, 89-96	Ib	4.535	_	2.52	2.92	21.13
32-1, 104-114	12b		3,542		_	
32-2, 54-59	2a	4.050	1020200	2.50	2.97	23.60
32-2, 81-91	2d	-	3.728		-	-
33-2, 22-32	3a	-	3.336	-	-	2
33-2, 97-102	4d	4.663		2.60	2.91	16.28
33-3, 72-82	91	-	3.328			10.60
33-3, 78-88	93		3 158			
34-1,73-85	10b	3.984	3.022	2.43	2.91	24.78
442B-3-3, 49-55	3	3.952	-	2.37	2.94	29.50
3-4.0-10	1a	-	1.625	-		
4-1, 57-59	9	4.145	-	2.40	2.93	27.29
4-2, 109-119	16		4.014	-		
5-2, 106-108	2h	4.155		1.1		-
5-3, 96-106	8a	-		2.40	2.89	25.69
5-3, 96-106	8u		3.319	-	-	-
6-2, 40-42	6	3.846	-	2.33	2.88	29.15
7-1, 49-65	7	4.133	3.839	2.50	2.98	23.96
8-1, 80-90	11	-	4.197	1000		_
8-2, 92-96	1k	4.918	-	2.70	2.95	12.50
8-5, 127-137	4		3.744	1000		_
9-1, 67-77	6b	-	3.658	-		-
9-1, 110-116	6d	4.159		2.41	2.96	25.02
11-2.121-123	16	4.130		- C - C - C - C - C - C - C - C - C - C		-
13-1, 32-40	6	3.928	24	2.42	2.95	27.04
15-1.77-80	13	4.141	-	2.42	2.88	24 46
16-1, 120-123	18	4.741		2.54	2.87	17.65
17-1, 104-106	20	4.187		_		
19-2. 12-14	3	5.250	200	2.80	2.97	9.11
20-1, 27-29	5	4.077	-	-	-	-

TABLE 12 Wet-Bulk Density and Porosity from 2-minute GRAPE Counts for Igneous Rocks, Site 442

Sample (interval in cm)	Piece No.	Wet-Bulk Density (g/cm <sup>3</sup> )	Porosity (%)
442A-31-1, 86-96	1b	2.48	23.39
32-2, 46-48	1f	2.48	23.20
33-2, 14-16	2	2.61	16.73
34-1, 50-52	9	2.38	28.85
442B-3-3, 49-51	3	2.24	35.93
4-1, 57-59	9	2.40	27.39
5-2, 73-75	3e	2.50	22.56
8-3, 135-137	4c	2.59	17.76
8-6, 12-14	1	2.43	26.03
9-1, 116-118	6e	2.49	22.94
11-2, 123-125	16	2.47	24.08
12-1, 55-57	8	2.39	28.21
13-1, 26-28	5	2.43	26.07
16-1, 124-126	18	2.71	11.41
17-1, 104-106	20	2.54	20.42
19-1, 46-48	6	2.41	27.17
19-2, 15-17	3	2.74	9.58

about 1.55 km/s on the average throughout the cores from this site. Thickness of sediment thus estimated is 287 meters, which is consistent with the depth of the first recovery of lithified rock (286 meters; limestone).

Sonobuoy observation by Murauchi and Asanuma provided information on  $V_{\rm P}$  of layer 2A (thickness of about 570 meters) underlying the sediments as low as 2.8 km/s. One of several possible explanations of this discrepancy is interlayering of both sediments and basalts showing lower velocity with successions of pillow-lava flows. If the proportion of lava flows with  $V_{\rm P} = 4.6$  is only 40 per cent, and the remaining 60 per cent is sediments with  $V_{\rm P} = 1.55$ , an estimated bulk sonic velocity of 2.8 km/s is in agreement with the observed value. This percentage of lava flows seems reasonable if we consider the poor recovery (<30%) of rocks below Core 442B-9. Fragmentation and brecciation of rocks which prevented deeper penetration of this hole possibly may be another cause of low bulk velocity.

According to the sonobuoy data, layer 2A is underlain here by a 450-meter-thick layer with  $V_{\rm p} = 4.0$  and a 270-meter-thick layer with  $V_{\rm P} = 4.7$ , overlying layer 3 with  $V_{\rm P} = 6.9$ . Composition and structure of these layers, however, were not determined from the present drilling because of insufficient penetration depth.

#### Magnetic Anomaly, Paleomagnetism, and **Paleontological Age**

Occurrence of normal and reversed polarities of natural remanent magnetization in rocks indicates that normally magnetized layers of rocks responsible for magnetic anomaly 6 observed at the site should exist below the reversely magnetized layers recovered at the bottom of Hole 442B. It also implies that ages of formation of rocks recovered by the present drilling are slightly younger than the postulated age of anomaly 6 (19-20 m.y.).

Paleontological age of sediment overlying the uppermost massive basalt flows is middle early Miocene (middle Burdigalian). The most up-to-date magnetic-polarity time scale (LaBrecque et al., 1977) shows that two intervals of normal magnetic polarity with time spans of 0.3



Figure 12. Shear strength versus sub-bottom depth for Site 442 sediments.



Figure 13. Sonic velocity, wet-bulk density, and thermal conductivity versus sub-bottom depth for Site 442 sediments.



Figure 14. Sonic velocity and wet-bulk density versus sub-bottom depth for Site 442 basalts. Crosses on wet-bulk-density graph correspond to values determined from 2-minute GRAPE counts, and dots correspond to values measured in the laboratory.



Figure 15. Wet-bulk density of Site 442 basalts as a function of porosity. The lines are theoretical predictions of wet-bulk density for a rock of a given grain density  $(\varrho_g)$  with varying porosity.



Figure 16. Sonic velocity of Site 442 basalts as a function of wet-bulk density.



Figure 17. Sonic velocity of Site 442 basalts as a function of porosity.

to 0.5 m.y., corresponding to anomalies 5D and 5E, are included (together with succeeding reversed intervals) in the middle Burdigalian stage. A sediment layer intercalated between the massive and pillow basalt flows provides an age of 18 to 21 m.y. (Core 442B-9). Thus, it seems likely that both massive and pillow flows were formed during the period of anomaly 5E or, otherwise, the pillow flows were erupted during 5D, and massive lavas were erupted during 5D after a pause of volcanic activity for about 0.5 m.y. of reversed polarity.

Intensity of natural remanent magnetization of pillow-lava flows averaging about  $3 \times 10^{-3}$  emu/cm<sup>3</sup> is consistent with intensities of usual basalts collected by drilling at other sites and by dredge hauls (e.g., Lowrie, 1974). Approximate estimation of thickness of subbottom magnetic layers responsible for the observed magnetic anomaly (Talwani et al., 1971) indicates that magnetic layers only about 1000 meters thick and at a sub-bottom depth of 5 km, would be sufficient to cause an anomaly of 300 gammas in peak-to-peak amplitude, if the layers are *uniformly* magnetized.

However, layer 2A beneath this site does not seem to be the layer causing the observed anomaly because:

1. The layer contains alternating sequences of normal and reversed polarities which apparently cancel at the surface.

2. The layer consists of interlayered basalts and sediments. Bulk intensity of natural remanent magnetization is reduced with percentage of basalts; if basalts are only 40 per cent, the bulk intensity would be  $0.8 \times 10^{-3}$  emu/cm<sup>3</sup>. Therefore, it seems more likely that layers 2B and 2C (total estimated thickness at this site 720 meters) are primarily responsible for the observed magnetic anomaly. This implication is also consistent with the age relationship showing that layer 2 is slightly younger than the anomaly.

## SUMMARY AND CONCLUSIONS

### Summary

The stratigraphic succession at Site 442 consists of seven lithologic units, five of which are sedimentary and range in age from early Miocene to Quaternary, and two of which are basalt. From the top downward, the units are:

Unit I (0-164 m): dark-greenish-gray mud and clay (Quaternary).

Unit II (164–209 m): yellow-brown mud (early Pleistocene and Pliocene) and (180.5–209 m) dark-brown mud with siliceous microfossils (Pliocene and possibly late Miocene).

Unit III (209-277.1 m):

Sub-unit IIIa (209-259.6 m): yellowish-brown mud and clay with volcanic ash (middle and late Miocene).

Sub-unit IIIb (259.6–277.1 m): interbedded gray and dark-greenish clay and volcanic ash (middle Miocene).

Unit IV (277.1-285.7 m): dark-brown zeolitic clay and claystone (early Miocene).

Unit V (285.7–286.1 m): pink limestone (early Miocene).

Basalt Unit 1 (286.1–353 m): aphyric, massive vesicular basalt.

Basalt Unit 2 (353–445 m): aphyric, vesicular, pillowbasalt flows.

The complete sedimentary section was recovered at Hole 442A. Hole 442B was drilled 70 meters north of Hole 442A and recovered sediments equivalent to Subunit IIIb and unit IV between 267.5 meters below the sea floor and the top of basalt unit I. The absence of unit V in Hole 442B suggests that its distribution is patchy or that is represents a local sediment pond.

The depth of deposition of the sedimentary units at Site 442 is shown in Figure 18. The lower part of unit III and units IV and V were deposited at a depth slightly above the CCD, whereas the lower part of unit II and the upper part of unit III were deposited below the CCD. Unit I and the upper part of unit II were deposited above the CCD.



Figure 18. A. General curve showing estimated water depth of CCD in Pacific Ocean (after van Andel et al., 1975, fig. 29, p. 47). B. Relative depth of deposition at Site 442 compared to CCD curve for Pacific Ocean.

The sediment-accumulation rate increases upward through the section. Deposition of sediment appears to have been uninterrupted since the early Miocene, the high rates of sediment accumulation recorded during the Quaternary being a direct consequence of both increased regional volcanism and increased sediment yield during periods of lower sea level.

Most of the clays at this site are hemipelagic, authigenic minerals composing less than 10 per cent of the total sediment. Pelagic components are dominant only in units IV and V.

The organic chemical components in the sediments consist ony of kerogen and bitumen; light-hydrocarbon gases are absent. Organic-carbon content is low in units I and IV, very low in unit II and III, and virtually absent in unit V.

The pH of the sediment averages 7.71, the alkalinity averages 4.7 meq/kg, the salinity averages 35 per mill, and the chlorinity averages 19.29 per mill.

Examination of the physical properties shows several major changes. The sonic velocity of the sediments of units I through IV averages 1.58 km/s, whereas for the limestone of unit V it averages 3.75 km/s. Sonic velocities for the two basalt units average 4.2 km/s, with a range from 3.85 to 5.25 km/s. Average density for the sediments is 1.47 g/cm<sup>3</sup>, and for the basalt 2.5 g/cm<sup>3</sup>.

Porosity averages 72.2 per cent for sediments (range 48 to 80%), and ranges from 10 to 29 per cent for basalts. The density and porosity of the basalts is lower than the average for oceanic basalts.

A systematic change was recorded in the shear strength of the sediments. Sediment shear strength increased with depth, but at certain horizons shear strength suddenly decreased. These horizons occurred at intervals where drilling breccia increased and recovery decreased. Shear strength of the sediments ranged from  $0.2 \times 10^{-5}$  to  $12 \times 10^{-5}$  dynes/cm<sup>2</sup>.

Paleomagnetism data show that Site 442 has been approximately at the present latitude since the early Miocene. The average intensity of natural remanent magnetization is  $3 \times 10^{-3}$  emu/cm<sup>3</sup>. Analysis of magnetic reversals in the lower part of basalt unit 2 indicates that the source of the magnetic-anomaly is below the present level of penetration in layer 2A and most probably occurs in layer 2B.

Visual examination and microscopic petrographic analysis of the basalts at Site 442 show that they are tholeiitic and that they are characterized by an absence of olivine, by high vesicularity, and by a cooling history dominated by an early crystallization of pyroxene. Both extrusive flows and intrusive sills are present.

## Conclusions

Our data from Site 442 permit the following conclusions:

1. The depositional surface of sedimentation at Site 442 was slightly above or slightly below the carbonatecompensation depth (CCD). Deposition began above the CCD, but appears to have been below the lysocline, as indicated by the poor preservation of foraminifers during the early and early middle Miocene. From the late middle Miocene through the early Pliocene, sediment accumulation occurred below the CCD. During the late Pliocene and Quarternary, deposition occurred slightly above the CCD (Figure 18). The exact depth of deposition cannot be determined exactly, because no data exist concerning present or past elevations of the CCD in the Shikoku Basin. However, if one assumes that the general Pacific Ocean CCD curve of Van Andel et al. (1975) represents a maximum CCD curve for the Shikoku Basin, the depositional surface at Site 442 probably ranged from a water depth of no less than 4000 meters to no more than 4300 meters.

2. The clays and claystones of units I, II, and III are hemipelagic and were deposited in a distal or basinal facies. The terrigenous components account for an average of 80 per cent of total mineralogical components. The exact source of the sediments was not determined. However, Site 442 is on the eastern edge of a westward-thickening clastic wedge, recognized from seismic surveys by Murauchi and Asanuma (1974, 1977); this clastic wedge thickens against the Kyushu-Palau Ridge. Although it has not been proven, the Kyushu-Palau Ridge appears to be the most likely source for these sediments. 3. The age of the basaltic basement was dated from the sediments recovered in 442B-9, CC as earliest Miocene (18–21 m.y.; *Discoaster druggii* Zone of nannofossils, foraminifer Zone N.6). This biostratigraphic age is in agreement with the age of magnetic anomaly 6 postulated by Kobayashi and Nakata (1977), although the anomaly itself probably comes from layer 2B, which was not reached. This biostratigraphic age determination provides a key confirmation of the magnetic-anomaly age determination in back-arc and marginal basins and also provides a critical baseline for testing the idea of a spreading origin for the Shikoku Basin, an objective requiring drilling at Sites 443 and 444.

4. The basalts recovered at Site 442 are unique in several ways when compared to mid-ocean basalts, particularly those of the Mid-Atlantic Ridge. Both the massive lava flows of basalt unit 1 and the pillow lavas of basalt unit 2 are anomalously vesicular. High vesicularity is expected from basalts extruded in shallow water. A strong inverse correlation exists between vesicularity and water depth of basalt extrusion; vesicle-rich basalts almost disappear at a depth of 2000 meters. Because the extrapolated depth of sediment deposition immediately above basement exceeds 2000 meters, it would appear that the Site 442 basalts may have been characterized by a volatile content higher than normal for mid-oceanridge basalts.

The basalts at Site 442 also differ from Atlantic Ocean basalts in that they contain no olivine, whereas many basalts associated with mid-ocean ridges are olivine-bearing. The site 442 basalts contain many chromian spinels which are enclosed by pyroxene, indicating that the pyroxene in these basalts may have crystallized very early in the cooling of the melt (Dick, per. comm.). This early crystallization of pyroxene is also at variance with observations from other oceanic basalts, where pyroxene appears to crystallize later in the cooling history (Dick, 1975). The high volatile content, the lack of olivine, and the evidence of early pyroxene crystallization of the Site 442 basalts suggest that these basalts differ megascopically and petrographically from many mid-ocean-ridge basalts. However, the megascopic and petrographic characters of the Site 442 basalts are similar to those of some ophiolites, particularly in the western U.S. Chemical analysis of the glassy rinds of the Site 442 basalts shows similarity of the Site 442 basalts with those of mid-ocean ridges, however. (Dick et al., this volume)

5. The rate of sediment accumulation was low during the Miocene and Pliocene and became high during the Quarternary. The preservation of organic carbon in the sediments appears to reflect changes in the rate of sediment accumulation.

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Information on core description sheets, for ALL sites, represents field notes taken aboard ship under time pressure. Some of this information has been refined in accord with postcruise findings, but production schedules prohibit definitive correlation of these sheets with subsequent findings. Thus the reader should be alerted to the occasional ambiguity or discrepancy.





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SITE 442 HOL	LEA	COF	RE	5 CORED I	NTER	VAL:	38.0-47.5 m		SITE	442	HO	DLE	A		CORE	6	CORED I	NTERVAL	47,5-57.0 m	
TIME-ROCK UNIT UNIT BIOSTRAT ZONE FORAMS T H MANNOS	SOSSIL SOVE	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE	STRUCTURES LITHOLOGIC SAMPLE		LITHOLOGIC DESCRIPTION	TIME-ROCK UNIT	BIOSTRAT	FORAMS	SONNAN	SOL	R	SECTION	MEIEKS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES		LITHOLOGIC DESCRIPTION
(V) pper Pleistocene B Beudoemiliania Incunosa Subzone (N) B Beudoemiliania Incunosa Subzone (N) B B B Beudoemiliania Incunosa Subzone (N) B B B Beudoemiliania Incunosa Subzone (N)	B	2 3 4 5 6 CC				•	5GY 4/1- 5G 4/1 with 5Y 4/1 pumice fragments 5GY 4/1, N4 5GY 4/1, N4 5GY 4/1, N4 5GY 4/1 N4 5GY 4/1 N4 5GY 4/1	Mud Generally intense deformation throughout; cocasional clay pellets (nodules), pumice fragments. Colors are dark gray green (5GV 4/1), dark gray (144). Pumice ab zones are gray (5Y 5/1). SME ATS: 1-99 Vitrie mud (Minor) (20% volcanic glass, 51% clay minerals) 2-76 Mud Sint 2-5% Muca TR Clay >60% Heavies 2-5% Clay minerals 67% Volcanic glass, 55% (day minerals) 5-108 Vitrie Mud (Minor) (25% volcanic glass, 55% clay minerals) 5-108 Vitrie mud (Minor) (15% volcanic glass glass clay minerals) 5-108 Vitrie mud (Minor) (15% volcanic glass glass clay minerals) 5-108 Vitrie mud (Minor) (15% volcanic glass glass cla	Pleistocene		RP B	B 3 3	B B		4 5 6		() () () () () () () () () () () () () (	0 0 0	5Y 4/1 5GY 4/1 5GY 4/1	Mud Intense deformation throughout, increased brockiton, soap, Pumile Iapilli noted, Colors predominate dark greenid-gray (5Y 4/1 to 5GY 4/1), Motting is noticable, Chire colors include dark gray (5Y 4/1), Brockita or chunky clay layers noted in Section 3, blend in with softer massive sity clay. SMEARS: 3-88 Vitric Mud (Minor) (30% volcanic glans, 52% clay minerats) 3-103 Mud Sand 5% Outritz, Feltdgar 15% Sit 20% Havies, Oraques 4% Clay 75% Clay minerats 2-84 (0.3, 44.8, 54.9) 4-70 (0.4, 659, 43.7) 6-15 (0.3, 42.9, 56.3) CARBON CARBONATE: 2-24 (0.3, 0.2, 1) 4-70 (0.4, 0.2, 1) 4-70 (0.3, 0.2, 1) 6-10 (8)

CC

LUND IN FORMULA LUND IN FOSSIL LUND IN FOSSI	SECTION METERS	GRAPHIC ON THE ALL STORES	LITHOLOGIC DESCRIPTION	TIME-ROCK UNIT	BIOSTRAT	FORAMS A	SOSSIL SARACTER	SECTION	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE	66,5-76,0 m	LITHOLOGIC DESCRIPTION
B AM B	2		5Y 4/1- 5GY 4/1     Mud       5GY 4/1     Interne deformation throughout yet firm with scattered zones of clay chunks (breeclated zones),       SGY 5/1     1-38 Mud       Sand < 1%			всм	8	0.5 1 1.0 2		0.00.0000	5G 4/1	Mud Intense deformation, brecciated and soupy in upper part to moderate in lower. Colors - gray (SY 57) are dominant. SMEARS: 180, 4-90 Mud Sand < 1% Quartz, Feldspar 15%, Silt >40% Citay minerals 80% Citay >50% Volcanic glass 1- 2% Quartas 1% Spicules 1- 2% GRAIN SIZE: 4-28 (03, 420, 57.7) CARBON-CARBONATE:
(N) auoz sapioziun	3		- 2-100 Mud (Clay) (<5% quartz, feldspar; 20% silt) 5GY 5/1 4-52 Volcanic ah (Minor) 4-100 Mud (20% quartz, feldspar; 5-7% volcanic glass) 10R 4/1 5-68 Auh-Rich Mud (Minor) GRAIN SIZE: 2-60 (0,1, 432, 56,4) 4-60 (1,1, 422, 56,9)	Lower Pieistocene	Crenalithus doronicoldes Zone (N	B FM	в	3	Void	9 0 9		4-30 (0.3, 0.2, 0) CARBONATE BOMB: 1-68 (8)
Crentilithus dore	4		CARBON-CARBONATE: 2494 (0.4, 0.3, 1) 5Y 6/1-5G 4/2 4-84 (0.5, 0.3, 2)		0			4			- 5YR 3/2 5Y 5/1	
BB	5		5GY 5/1			B B	B	5			5G 4/1	



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5GY 5/1

5GY 5/1

TIME-ROCK UNIT	11	0	FOSSIL						ICE	ARY	J					
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		. 1							a		11			Heavies	1%	
15							1.12		6					Clay minerals	88%	
DC DC			1			-			11					Volcanic glass	1%	
ö		č					1.2		0			ECV A/A		Opaques	1%	
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				- 1		2	1.2		10					Clay minerals	85%	
							1.5		0					Vorcanic glass	5%	
							12		11					Zastite	1- 2%	
							1.1		6					Carbonate upprecidiert	1. 2%	
							1.2		11					Nannafoesila	TE	
		в	R	в		3	1		10	2		-		Siliceous fossils	TR	
		в	в	в		CC	1				•					

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VOID

3

4 CC

BBB

0

5Y 5/1

5Y 5/1

CARBON-CARBONATE: 3-83 (0.3, 0.2, 1)

CARBONATE BOMB:

3-35 (6)

56
SITE 442 HOL	ΕA	CORE 15	CORED I	NTERV	AL:	133.0-142.5 m		SITE 4	42	HOL	ΕA	CC	RE	17 CORED I	NTERVAL	152.0-161.5 m	
TIME-ROCK UNIT BIOSTRAT ZONE FORAMS AMANNOS	SOVE	SECTION METERS	GRAPHIC ITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	LITHOLOGIC		LITHOLOGIC DESCRIPTION	TIME-ROCK UNIT	ZONE	NANNOS P	SOF	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE STRUMENTARY STRUCTURES LITHOLOGIC SAMPLE		LITHOLOGIC DESCRIPTION
B RP	в	0.5		0000		5Y 5/1	Mud Drilling breccia to 120 cm - Section 1 then change to moderate deformation. Color change from gay (5Y 5/1) to greenish-gray (5Y 5/1) noted at 120 cm - Section 1. SMEARS:		1	B RM	8	1	0.5		• • • • •	5Y 5/1	Mud Drilling breecia (soupp) of mud (SY 5/1) with pebbles (drilling) of blackish dark green gray clay (SBG 4/1). SME ARS: 2-7, CC, 1-30 Mud Sand < 1% Ouartz, Feldspar 2- 3%
		2				5GY 5/1	1-122, 3-100, CC Mud Sard < 1% Ouartz, Feldspar 5% Silt >35% Heavies, Mica TR Clay >60% Clay minrais 90% Volcanic glass 2% Oneques 1% Carbonate unspecified 2% Songerspindles TR					2					Siti ⇒ 35%   Clay minerais   91%     Clay ⇒ 60%   Volcanic glass   5%     Opaques   1%   5%     CC = Mud with 8% glass   1%   1%     1:30 = darker mud with up to 25% volcanic glass   (three occur as fragments).   5%     GRAIN SIZE:   2.51 (0.4.53 (3.6.3)   3
вв	в			1			2-70 Mud (Minor?) lwith sericite 20%) 4-6 Med (Minor) (with reolite 2%)				B			-00		101 5 171 0	CARBON-CARBONATE: 2-81 (0.2, 0.2, 0)
		3				5Y 5/1 5GY 5/1 5Y 4/2	GRAIN SIZE: 2-11 (0.4, 368, 62.8) 4-24 (0.1, 34.7, 65.2) CARBON-CARBONATE: 2-16 (0.4, 0.1, 2) 4-16 (0.2, 0.2, 0)	TIME - ROCK	ZONE	NANNOS H	SQRA	ER NOILOBS	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY SEDIMENTARY SEDIMENTARY SEDIMENTARY SEMPLE	161.5-1/1.0 m	LITHOLOGIC DESCRIPTION
В	в	4		ĺ		5Y 5/1-5GY 5/1 to 5Y	6/1			B RM	в	1	0.5		00000	5Y 5/1	Mud (Sitty Clay) Highly brecciated in Section 1 with pabble-like (drill) fragments of (SY 5/1) mud and (N4) clay (with ash), Yetlowih brown (10/N 5/4) aitry clay begins a 105-140 cm in Section 2. Moderate deformation and darker zones at
	E A OSSIL RACTER		GRAPHIC	TARY AND	AL:	142.5-152.0 m							1.0		000		base of Section 3, 70 cm. SMEARS: 2-110, CC Mud (yellow brown)
TIME - R UNI BIOSTI ZONI FORAMS NANNOS	RADS	AETE	ITHOLO GY	DRILLING DISTURBJ SEDIMEN	SAMPLE		LITHOLOGIC DESCRIPTION								000		Clay >60% Clauric, reitopar Clay >60% Clay minerals 97% Opaques 1% Volcanic glass 1%
6	в	0.5		0000		5Y 5/1	Mud Variable deformation from breccia to moderate, firm, Colors are gray (SY 5/1) to some olive gray (SY 4/2) in Section 2, Scattered pumice fragments noted.									Change to yellow brown (10YR 5/4)	3-70 Clay (Altered Ash) Silt 15% Quartz, Feldspar 3-53 Clay 85% Heavies 19 Clay (altered?) 889 Volcanic glass 255 Opaques 19
8		1.0					SMEARS:   17-5, 275, CC     Sand   <1%			B 8	в	3	-			10YR 5/4	GRAIN SIZE: 2-120 (0.5, 38.1, 61.5) CARBON-CARBONATE: 2-129 (0.2, 0.2, 1)
		2				5Y 5/1- 5Y 4/2	GRAIN SIZE: 2-61 (0.2, 32.2, 67.6) CARBON-CARBONATE:										
ВВ	в	3	3			pumice 5Y 5/1	2-22 (0.2, 0.2, 0) CARBONATE BOMB: 2-30 (6)										



85

SITE 442	2	HOL	E A	CC	DRE	23	CORED	NTERV	AL:	209.0-218,5 m		SITE	442	HC	LE A		COR	E 2	4 CORED I	NTER	AL:	218.5-228.0 m	
TIME-ROCK UNIT BIOSTRAT	FORAMS	CHA SONNAN	SOF SUPERIOR	SECTION	METERS	GR	RAPHIC	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	LITHOLOGIC		LITHOLOGIC DESCRIPTION	TIME-ROCK UNIT	BIOSTRAT	FORAMS	FOSSIL ARACI	ER	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE SEDIMENTARY	LITHOLOGIC SAMPLE		LITHOLOGIC DESCRIPTION
		В	B B	1 2 3 4 5 CC	0.5-	2   2				10YR 8/4 dark gray brown (10YR 4/2) plant debrii(?)	Mud Firm, hard, bracciated especially in Section 1.   Color dominately light yellowish-brown (10YR 6/4)   mottid with dark grayish-brown (10YR 6/2).   1, 33 35 cm. 3, 33 35 cm. & CO common and shumdant   SMA Statum 100 (10 (10 (10 (10 (10 (10 (10 (10 (10	Upper Milocene	Ommatartus antepenultimus (R)	E	B CM		3 4 5 CC		VOID		•	10YR 6/4	Mud   Firm clay, breeclated, Dominant yellowish- brown (10YR 64) with bluish-gray (5B 5/1) fragments, Scattered in Section 2, dark grayibh-brown sporting in Section 4, 1, 33-35 cm, 3, 31-33 cm: abundant foram casts SME ARS: 176 Mud   Sand < Six

SITE 442	но	LE A	 COR	E 2	25 CORED IN	NTERV	AL:	228.0-237,5 m		SITE	442	н	OLE	A	c	ORE	26 CORED I	NTERVAL:	237.5-247.0 m	
TIME-ROCK UNIT BIOSTRAT ZONE	FORAMS C	SOSSI	SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBANCE SEDIMENTARY STRUCTURES	LITHOLOGIC		LITHOLOGIC DESCRIPTION	TIME-ROCK UNIT	BIOSTRAT	FORAMS	FOR SOUNAN	ACTE	SECTION	METERS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES LITHOLOGIC SAMPLE		LITHOLOGIC DESCRIPTION
Middle Miocene Cannartus pettessoni (R)	в	СМ	2	3.5				10YR 5/4 with 10YR 4/3-3/3- 10YR 4/2-10YR 3/2	Mud   Firm, brecciated clay: dark yellowish-brown (10/R 5/4) to brown (10/R 4/3), Mottiling with gravibh-brown (10/R 5/2) dark gravibh-brown (10/R 4/2); pyritic zones noted in Sections 2 and 3. Dark layers may mark relist bedding: colors wery dark gray (10/R 3/1) to very dark gravish- brown (10/R 3/2), 3. 41-43 cm: Rare foram casts SMFARS     2.14 3. dm: Rare foram casts SMFARS     2.10 5 Siliceous Mud (5-5% distorn, radiolarian, sponge spicules, 70-75% clay minerals)     2.14 3. Mud (pyritic)     Sand < 1% Ouartz, Feldspar 10-15% Silit >40% Clary minerals 68-70% Clay >50% Pyrite 74-0% Siliceous fragments 1%     3.40 Mud (with pyritic, siliceous) (5% pyrite, 3-5% distorm, radiolarian)     4.6 Mud (with siliceous fragment) (3-5% pyrite, distorm, radiolarian, siliceous?)     4.2 Mud (with siliceous) (7-2% volcancig quas 3-5% distorm, radiolarian, 10-5% pyrite, 50% distorm, radiolarian, 10-2% volcancig quas 3-5% distorm, radiolarian,	Middle Miccene	Dorcadospyris alata (R)	IAM SUIDT SUIVA assentation B	RP CI B CI	P	2	1.0			10YR 5/4 10YR 5/4	Mud     Breccia in Section 1 with burrows noted in the lower part. Interne deformation continues to 55 cm, Section 2 with moderate below, Scattered diffuse areas with pyrite accurate blow, Scattered Burrows with pyrite accurate in Section 2.     Burrows with pyrite accurate blow, Scattered SMEARS:     2, 68-70 cm: Rare foram cents SMEARS:     27.6 Mud Sand < 2% Quartz, Feldoar
			4			1	•	10YR 5/3	6-8% quark; feldbari     4-92   Mud     (6-11% illiceous fragments, 5-7% pyrite)     GRAIN SIZE:     2-41 (15, 48.3, 50.2)     4-41 (1,1, 42.6, 56.2)	TIME-ROCK	BIOSTRAT TA	FORAMS O I	HAR SONNAN		SECTION	WETERS	GRAPHIC	DISILING SEDIMBANCE SEDIMBANCE STRUCTURES LITHOLOGIC SAMPLE	247.0-256.5 m	LITHOLOGIC DESCRIPTION
	В	СМ	CC					10YR 3/3 with 10YR 2/1-10YR 3/1	CARBON-CARBONATE: 2-60 (0.2, 0.1, 1) 4-60 (0.1, 0.1, 0)	Middle Miocene	Dorcadospyris alata (R)		вс	M	1	0.5		00000000	10YR 5/4 10YR 5/4 mottlad with 10YR 4/2	Mud Section 1 = drilling breccia, Moderate to intense disturbance in Section 2 (50-150) with yellow horown (10YR 4/2), Pumice pebble noted in Section 2, 2, 90-92 cm: Rare foram casts SMEARS: 2-135 Mud (dark color) (Minor) Send 10% Mica TR Clay 80% Hawies 1% Clay minerals 84% Volcanic glass 1% Purite 7% Radiolarlans 2% Sonome circles 12%

3% TR 1% 84% 1% 2% 1%

Sponge spicules 2-140 Mud (light color) (TR radiolarians, 1% sponge spicules, 5% volcanic glass)

GRAIN SIZE: 2-103 (3.6, 47.0, 49.4) CARBON-CARBONATE: 2-107 (0.1, 0.1, 0)

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LW. SAMPLE

UN   CONSISTING (1) (1) (1) (1) (1) (1) (1) (1) (1) (1)	SITE	442	HC	LE	A	cc	DRE	28	CORED	IN1	TERVA	LI	256.5-266.0 m		SITE	44	H	OLI	ΕA		ORE	2	9 CORED I	NTERV	AL:	266.0-275.5 m				
Note   Note <th< th=""><th>TIME-ROCK UNIT</th><th>BIOSTRAT</th><th>FORAMS</th><th>FOSS ARA SQV8</th><th>CTER</th><th>SECTION</th><th>METERS</th><th>GI</th><th>RAPHIC</th><th>DRITTING</th><th>DISTURBANCE SEDIMENTARY STRUCTURES</th><th>SAMPLE</th><th></th><th>LITHOLOGIC DESCRIPTION</th><th>TIME-ROCK</th><th>BIOSTRAT</th><th>FORAMS</th><th>FC HA SONNAN</th><th>SOLA</th><th>R</th><th>SECTION</th><th>MEIEKS</th><th>GRAPHIC LITHOLOGY</th><th>DRILLING DISTURBANCE SEDIMENTARY STRUCTURES</th><th>SAMPLE</th><th></th><th>LITHOLOGIC DESCRIPTION</th><th></th><th></th><th></th></th<>	TIME-ROCK UNIT	BIOSTRAT	FORAMS	FOSS ARA SQV8	CTER	SECTION	METERS	GI	RAPHIC	DRITTING	DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLE		LITHOLOGIC DESCRIPTION	TIME-ROCK	BIOSTRAT	FORAMS	FC HA SONNAN	SOLA	R	SECTION	MEIEKS	GRAPHIC LITHOLOGY	DRILLING DISTURBANCE SEDIMENTARY STRUCTURES	SAMPLE		LITHOLOGIC DESCRIPTION			
CC Clay (92% clay, 2% volcanic glass, 3% pyrite with GRAIN SIZE:	Middle Miccene	Discoaster exilis Zone (N)?	B F B E	1P B 1 B		1 2 3 4 CC	0.5						10YR 5/4 5G 5/1 5G 5/1	Mud Section 1 m 100 cm is near total dilling breecia with irregular dark (pyrite) (10YR 4/1) zones, Section 3: 35-45 cm - darkening from 10YR 5/4 to 10YR 4/1 - day with pyrite; 60-95 cm - (10YR 8/4) (day, 88-105 cm - zone of color change from 10YR 8/4 to 5GY 4/1 including dark gras brown (10YR 8/4 to 5GY 4/1 including dark gras brown (10YR 8/4 to 5GY 4/1 including vith graving-reen (56 4/2) band (2 cm) at 135 cm. 1, 37-38 cm: very rare foram casts SMEARS: 34-6 Mud (5% pryrite) 340 Mud (30% volcanic glass) 3-102 Mud (30% volcanic glass) 3-103 Volcenic Ash (39% clay minerals) 3-110 Volcanic Ash (39% clay minerals) 3-135 Volcanic Ash (39% clay minerals) 3-145 Volcanic Ash (39			8 8	B B	B B	-	1 1.0 2 2 3 <u>CC</u>			a superior states and a second se		10YR 5/4 5Y 7/1 2.5Y 4/4 2.5Y 4/4	Mud Section 1 – 0-10 cm - interbedde continued from Core 28; 10-80 cr thrown 1074 54/ clay; rim, mo with black specking, 80-130 cm - specking/strasking; organicat?): 1 light gay ash. Section 2 – 0-50 cm - with interb L2,5Y 0.42 - light brownish-gray / (1074 5/4); 50-80 cm - light yelf (L2,5Y 0.42) and to dark gray br ah, 80-160 cm clay (1074 5/4); yard to olive brown clu; 12,5Y 4/4 punice fragments, motifed = burs SMEARS: 1-90 Clay Sard < 2% Duartz, Feld Sit > 40% Haavie Clay > 50% Clay micreal Volcanic glas Dopaces 1-145 Clayvey/Sitty Volcanic Ah (10% volcanic glass, 39% clay miner 2-55 Volcanic glass, 39% clay miner 2-110 Mud (19% clay minerals, 2% quartz, feld 3-80 Pyrite-bearing Clay 3-100 Clay (85% clay, 5% volcanic glass) CC Clay (92% clay, 2% volcanic glass, 3% p GRAIN SIZE:	I ash/clay - vellowit irrately di irrately di irrately di irrately di irrately as irrately di irrately di irratel	at http://www.interfeature.com/ //2/ ///////////////////////////////	

CARBON-CARBONATE: 2-103 (0.1, 0.0, 0)

CARBONATE BOMB: 3-4 (6) 3-137 (6)



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BIOSTRI	FORAMS	NANNOS	RADS		SECTION	METERS	GRAPHIC LITHOLOGY	DISTURBAN	SAMPLE	LITHOLOGIC DESCRIPTION
Lower or middle Miocene Helicosphaera ampliaperta Zone (N) or Schandlibus haterconschaer Zone (N)	B B	RP	В		2	0.5	A constructions	000		Section 1 — Q-85 cm — drilling bescris of light olive gray (5Y 6/2) virit zoolitic mudstone (85% clay, 12-15% vectaric) glas); dark reddish brown (5YR 3/4) sity clay (with 3-55 vectaric glas, 7% operating site (32 vector); dark modified brown (5YR 3/4) sity clay (with 3-55 vectaric glas, 7% operating site (32 vector); dark modified brown (5YR 3/2) site (32 vector); dense fine limestone with Mn specks, 76-81 cm — Limestone, 81-101 cm — Limestone, 101-107 cm — Weathered basalt, limestore and dark reddish brown (5YR 3/2) mud fragments. Basalt et 107 cm. GRAIN SIZE 1-22 (0.5, 26,1, 73,3) CARBONATE BOMB: 1-30 (6) 1-30 (8,7)





Depth 286.5 to 286.9 m

Visual Description

0-41 cm: aphyric basalt, gray to dark gray; vesicular. Fine vesicles (40%); <1 mm; medium vesicles (2%); >1 up to 4 mm.





ISUAL	CORE DESCRIPTION
FOR	IGNEOUS BOCKS

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Depth 290.0 to 291.5 m

#### Visual Description

0-150 cm: aphyric basalt; dark gray; vesicular; slightly altered. Fine vesicles 40% (<1 mm); medium vesicles ~ 2% (> 1 mm and partly up to 5-7 mm).

98-150 cm: vesicles (~50%) contain some secondary minerals.



			58442	2 A 3 2
			Depth 291.5 to 2	292.8 m
Visual Descriptio	n			
0-128 cm: aphyr medium vesicles secondary miner 98-128 cm: vesic	ic basalt; dar (>1 mm) oc als. Ies are 20-30	k gray, vesicul cur in the 0-51 %	ar, lightly altered. Fine ve cm interval - 2%, 50% of	sicles 40% (<1 mm); the vesicles contain
128-138 cm: vari (30-40%) up to 2	olitic basalt 0 mm in an	gray to dark g aphanitic grou	ray; varioles are an aggrega ndmass.	ite of plagioclase
Thin Section Des	cription – 6	2 cm		
Groundmass: pla	gioclase 30%	, 0.6 x 0.1 mr	n, An <sub>54</sub> , subhedral; clinop	yroxene 7%, 0.4 x 0.
mm, augite, anhe	dral; glass 30	0%, interstitial		
Vesicles: 30%, 0.	1-2.0 mm, re	ound irregular.		
Texture: intergra	inular-interse	ertal.		
Alteration: 3% c	arbonate in	vesicles.		
Thin Section Des Groundmass: pla glass+py+pe+mt	eription — 1 gioclase 10% 50%, microli	26 cm 5, 0.1-2 mm, 7 ite (half crysta	An <sub>62</sub> , euhedral, skeletal, li illine).	ght zonation;
Vesicles: 40%, 0.	05-1.5 mm,	round irregula	r, unfilled.	
Texture: variolit	ic - interserta	al.		
Shipboard Data				
Bulk Analysis:	0 cm	80 cm	Magnetic Data:	62 cm
SiO <sub>2</sub>	48.80	49.60	Intensity (emu/cc)	104.2
Al202	14.82	15.17	Inclination before	
Fe <sub>2</sub> O <sub>2</sub>	1.12	1.03	demag.	26.2
FeÔ	7.39	6.80		
MgO	5.64	7.72	<b>Physical Properties:</b>	54 cm
CaO	15.37	13.93	Vp (km/s)	4.05
Na <sub>2</sub> O	3.07	3.20	Porosity (%)	23.60
K20	0,47	0.30	Wet Bulk Density	2.50
TiO2	1.15	1.19	Grain Density	2,97
P205	0.14	0.12		
MnO	0.15	0.11	Other Data:	81 cm
LOI			Therm, cond.	
H20 <sup>+</sup>			(mcal/cm-s-°C)	3.73
H20-				
CO2	1000			
Cr	234	240		
Ni	57	130		
Sr	183	182		

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86

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



# Depth 291.5 to 292.8 m

### Visual De

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14.02	10.17	inclination before	
1.12	1.03	demag.	26.2
7.39	6.80		
5.64	7.72	<b>Physical Properties:</b>	54 cm
15.37	13.93	Vp (km/s)	4.05
3.07	3.20	Porosity (%)	23,60
0,47	0.30	Wet Bulk Density	2.50
1.15	1.19	Grain Density	2.97
0.14	0.12		
0.15	0.11	Other Data:	81 cm
		Therm, cond.	
		(mcal/cm-s-°C)	3.73





# Depth 294.5 to 296.0 m

#### Visual Description

0-10 cm: light gray aphyric basalt 40% vesicles (<1 mm), 5% vesicles (>1 mm) calcite infilling of larger vesicles

10-150 cm: light gray aphyric basalt 30% vesicles (<1 mm), up to 5% vesicles (>1 mm) occasional calcite infilling of larger vesicles. Largest vesicles are approximately 2 to 3 mm across.

Alteration - light.

#### Thin Section Description - 9 cm

VISUAL CORE DESCRIPTION

FOR IGNEOUS ROCKS

Groundmass: plagioclase 25%, 0.5-1.0 mm, An<sub>59</sub>, laths; clinopyroxene 20%, 0.05-0.2 mm, anhedral; magnetite 3%; other 25%, cryptocrystalline matrix. Vesicles: 25%, 0.1-1.0 mm,

Texture: intersertal.

Alteration: 1% carbonate in vesicles; 1% clays in vesicles.

#### and Date Shipb

Zr

Shipboard Data	
<b>Bulk Analysis:</b>	124 cm
SiO <sub>2</sub>	47.97
Al203	14.16
Fe <sub>2</sub> O <sub>2</sub>	1.08
FeÔ	7.16
MgO	6.80
CaO	16.93
Na <sub>2</sub> O	2.75
K20	0.29
TiO	1.11
P205	0.14
MnO	0.13
LOI	
H20+	
H20-	and the second
cô <sub>2</sub>	
Cr	216
Ni	53
Sr	186

82

Magnetic Data:	136 cm
Intensity (emu/cc)	346.5
Inclination before	
demag.	33.3
Stable Inclination	36.1



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



101 cm

109.6

33.4

31.9

97 cm

4.66

16.28

2.60

2.91

3.34

22 cm

# Depth 296.0 to 297.5 m

#### Visual Description

Sr

Zr

Largest vesicles are up to 5 mm across. Infilling of vesicles by brown and light greenish-gray alteration products. 20-33 cm: (3A) - fracture with weathered surfaces. 129-137 cm: (41) - at 132 cm large vesicle occurs with well-developed crystals of calcite(?). Thin Section Description - 102 cm Groundmass: plagioclase 35%, 0.1-3.0 mm, An<sub>60</sub>, laths; clinopyroxene 25%, 0.1-0.5 mm, anhedral; magnetite 4%, 0.05-0.1 mm, granular; other 20%, cryptocrystalline matrix. Vesicles: 15%, 0.3-1.0 mm, Texture: intersertal.

Light gray, vesicular, variolitic basalt, Vesicle approximately 30% (<1 mm) and 5% (>1 mm).

# Alteration: 1% clays lining vesicles.

189

86

Shipboard Data		
Bulk Analysis:	57 cm	Magnetic Data:
SiO2	50.17	Intensity (emu/cc)
Al203	15.02	Inclination before
Fe2O3	1.22	demag.
FeO	8.08	Stable Inclination
MgO	7.36	
CaO	11.54	Physical Properties:
Na <sub>2</sub> O	3.18	Vp (km/s)
K20	0.44	Porosity (%)
TiO2	1.25	Wet Bulk Density
P205	0.13	Grain Density
MnO	0.13	
LOI		Other Data:
H20 <sup>+</sup>		Therm, cond.
H-0-		(mcal/cm-s-°C)
cô,		
Cr T	233	
Ni	49	



L	EG		SIT	Ē	HOLE	0	OF	E	SE	ст.
5	8	4	4	2	A		3	3		3

# Depth 297.5 to 299.0 m

# Visual Description

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ial.

Light gray, vesicular, variolitic basalt. Vesicles approximately 20% to 30% (<1 mm) and 2% to 10% (>1 mm). Largest approximately 3 mm across. Some infilling of vesicles by brown alteration products. 24-32 cm: (4) - pieces split by heavily weathered fracture.

66-79 cm: (9a) - two distinct vesicle-rich zones at 66-68 cm and 71-73 cm.

130-141 cm: (15) - fracture at lower end with weathered surfaces.

#### Thin Section Description - 80 cm

Groundmass: plagioclase 30%, 0.1-1.0 mm, An<sub>59</sub>, laths; clinopyroxene 20%, 0.1-0.5 mm, anhedral; magnetite 3%, <0.02 mm, granular; other 20%, cryptocrystalline matrix. Vesicles: 25%, 0.05-1.0 mm. Texture: intersertal.

Alteration: 1% carbonate filling vesicles; 1% clays lining vesicles,

#### Shipboard D

SiO<sub>2</sub> AJ<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO

H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O<sup>+</sup> CO<sub>2</sub> Cr

Ni

Sr

Zr

Shipboard Data		
Bulk Analysis:	68 cm	
SiO2	50.04	
A1203	14.93	
Fe2O3	1.18	
FeO	7.80	
MgO	8.31	
CaO	11.68	
Na <sub>2</sub> O	3.25	
K20	0.38	
TiO2	1.24	
P205	0.12	
MnO	0.12	
LOI		
H20 <sup>+</sup>		
H20-	-	

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245

77

182

89

79 cm	
134.0	
31.4	
34,6	
72 cm	78 cm
3.33	3.16
	79 cm 134.0 31.4 34.6 72 cm 3.33



Graphic Represen

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VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



# Depth 299.0 to 300.2 m

#### Visual Description

Light gray, vesicular, variolitic basalt. Vesicles approximately 20% to 30% (<1 mm) and 2% (>1 mm). Some vesicles infilled by calcite. Alteration internally to olive green products (approximately 20%).

0-15 cm: (1) - 2 large vesicles at the base of this piece (14-15 cm) are lined by crystals of calcite and yellow brown palagonite(?).

24-52 cm: (2B) - large inclusion approximately 4 cm across at top (25-26 cm). 89-97 cm: (5) - drilling breccia.

#### Thin Section Description - 53 cm

Groundmass: plagioclase 30%, 0,2-1,0 mm, An<sub>69</sub>, laths; clinopyroxene 20%, 0,1-1,0 mm, anhedral; magnetite 2%, 0.02-0.1 mm, granular; other 25%, cryptocrystalline matrix. Vesicles: 19%, 0.1-1.0 mm, anhedral. Texture: intersertal.

Alteration: 2% carbonate infilling vesicles; 2% clays lining vesicles.

#### Thin Section Description - 109 cm

Groundmass: plagioclase 25%, 0.3-1.0 mm, An<sub>60</sub>, laths, some in spherulitic groups; clinopyroxene 15%, 0.01-1.0 mm, anhedral, often fine-grained, not acicular; magnetite 2%, 0.01-0.2 mm, granules; other 8%, cryptocrystalline matrix. Vesicles: 50%, 0.05-2.0 mm. Texture: intersertal.

# Shipboard Data Bulk Analysis: SiO, Al2<sup>0</sup>3 Fe2<sup>0</sup>3 Fe0 MgO CaO Na20 K20 TiO2 P205 Mn0 LOI

H20<sup>+</sup> H20<sup>+</sup> CO2

Cr

Ni

Sr

71

3.27

0.60

1.23

0.12

0.13

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214

35

178

111 cm	Magnetic Data:	52 cm
50.09	Intensity (emu/cc)	124.0
13.90	Inclination before	
1.26	demag.	29.0
8.32	Stable Inclination	30.0
7.39		
11.81		



L	EG		SIT	E	HOLE	С	OF	E	SE	CT.
5	8	4	4	2	A		3	4		1

125 cm

300.6

44.2

42.6

#### Depth 304.0 to 305.4 m

Light gray, vesicular, variolitic basalt. Vesicles approximately 10% to 30% <1 mm and occasionally >1 mm. 59-69 cm: (10A) - fracture with weathered zone on either side (mainly palagonite). 93-103 cm: (11) - drilling breccia.

131-141 cm: (13) - drilling breccia.

#### Thin Section Description - 52 cm

Groundmass: plagioclase 5%, 0.05-1.0 mm, An<sub>56</sub>, acicular; other 55%, cryptocrystalline matrix. Vesicles: 40%, 0.02-3.0 mm.

# Thin Section Description - 126 cm

245 64

186

87

Groundmass: plagioclase 21%, 0.05-1.5 mm, An<sub>67</sub>, laths; clinopyroxene 21%, 0.05-0.5 mm, anhedral; magnetite 2%, 0.01-0.1 mm, granular; other 21%, cryptocrystalline matrix. Vesicles: 30%, 0.1-2.0 mm.

Alteration: 1% carbonate infilling vesicles; 4% clays lining vesicles.

78 cm	Magnetic Data:	51 cm
49.96	Intensity (emu/cc)	424.7
15.44	Inclination before	
1.14	demag.	41.4
7.52	Stable Inclination	41.9
8.12		
12.46	<b>Physical Properties:</b>	73 cm
3.13	Vp (km/s)	3.98
0.28	Porosity (%)	24.78
1.17	Wet Bulk Density	2.43
0.12	Grain Density	2.91
0.12		
	Other Data:	73 cm
	Therm cond.	
<u></u>	(mcal/cm-s-° C)	3.02
245		

SITE 442	HO	LE	в	c	ORE	1		CO	ED I	NTE	RVA	L:	2	67.5-2	77.0	n									SIT	44	2 1	IOL	E B		co	RE	2	co	RED I	NTERV	AL:	277.	0-286.5 m							_
TIME-ROCK UNIT BIOSTRAT ZONE	FORAMS D	FOS	CTE	R		MEIEKS	GR	AP)	HIC DGY	DISTURBANCE	STRUCTURES	SAMPLE					LIT	HOLO	GIC D	DESCRI	IPTION	N			TIME-ROCK	BIOSTRAT	FORAMS	FC SONNAN	RAC	L	SECTION	METERS	GLIT	RAP	HIC DGY	DISTURBANCE	LITHOLOGIC SAMPLE			1	ITHOL	OGIC DES	CRIPTION			
	RE	3 FP			0. 1) 1) 1) 1) 1) 1) 1) 1) 1) 1)			SAM					11 5 5 5 5 5 5 5 7 11 11	9YR 5/1 5 Y 4/3 G 5/1 Y 5/2 S 5/1 Y 5/2 S 5/1 S 5/7 6//	6 4 3 3		G Set 11 11 12 2 5 5 14 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5 5	ay and A cition 1: 271 cm: 0745 5/21 cm: 1732 cm way ash 1732 cm way ash 1732 cm way ash 1732 cm 1732 cm 1732 cm 1745 cm 174	Ash is stoff, br is stoff, br is stoff, br is gray (1) is gray (1) is gray (1) is gray (2) is gray (5) is gray (5	reaciated (SY 5/1) (N5), a reliow b o clay b m stiff c ray (SY m stiff c ray (SY ray (S) ray (SY ray (SY	d, light 1, , massik rown (11 rown (12 sob bed (12 s	yellow ve, soft (oYR 5 down 1 d vitric d vitric d vitric from 1 s), firm s), firm s)	vish-brown t sticky locand), S(4) clay; S(4) clay; S(5) clay with to greenial mod. s zeolits, s zeo	n ih- hce (, par) rrz,	I access to Middle Alfanana		BFF (N) hitem envelopmentation of the anti- subdividue ensemblementati	B	B		1 2 3 4 5 CC	0.5						5Y 7/ 7.5YI 7.5YI 10YR 5YR / 10YR 5YR / 10YR 10YR 10YR	1-10YR 6/4 8 3/2 8 3/2 8 3/2 8 3/2 8 3/2 4/4 8 3/2 4/4 8 3/2 4/4 8 3/2 4/4 8 3/2 4/4 8 3/2 8 3/3 3/3		Clays, Z Section 0-13 cm 13-150 (2011) 13-150 (2011) 13-15	eolitic Clays, 1: Encodential of the second secon	Altered Ash light gray (5) 'ft 5/4) nan wn (7,5YR 3/4) dark brown n (7,5YR 3/2) wy: moderast n (11-30, 73 wy: moderast n (11-30, 73 (7,5YR 3/2) titled by biot inikish-gray i h-brown (11 -5,5YR 3/2) titled by biot wm (10YR 4/4) dary to belo dary to belo bation, Mn resolitic clay n m (7,5YR 4/4) bioturbatio/	(7 711) clay (7 711) clay (7 514)-beac (7 514)-beac (7 514) classification (7 514) class	with micronod code, micronod participation code at the second own fin micro own fin fin dicro own fin dicro fin di	Sules/
	B	BR	1		.c	-	-	1.	201	1			1	111 111			-	_					_																							

89





150



#### Depth 291.0 to 291.9 m

0-86 cm: basalt, aphyric, dark gray, vesicular, lightly altered. Vesicles (1 mm or less) about

26-35 cm: a few vesicles, but some have sizes up to 7-10 mm with yellow secondary minerals. 26-31 cm: probable chill zone.

Phenocrysts: plagioclase < 1%, 0.4-1.5 mm, euhedral.

Groundmass: plagioclase 30%, 0.05-1.0 mm, An<sub>60<sup>e</sup></sub> subhedral; clinopyroxene 20%, 0.05-0.1 mm, augite, anhedral; glass + mt 20%.

Alteration: 1% carbonate in vesicles and 1% clays in vesicles,

Bulk Analysis:	13 cm	Magnetic Data:	54 cm
SiO2	50.33	Intensity (emu/cc)	301.1
41203	15.06	Inclination before	
Fe203	1.25	demag.	39.0
FeÔ	8.24		
AgO	7.14	Other Data:	0 cm
CaO	11.77	Therm. cond.	
Va <sub>2</sub> O	3.35	(mcai/cm-s-* C)	1.63
\$20	0.39		
río,	1.30		
205	0.13		
InO	0.14		

70

Z Piece

cm

0 -

50-

Graph

60

100

20

6

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60

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0 0 0

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6 0 S

13 D . 00 00 ~ °

150

100-10

5



LI	EG		SIT	Ē	HOLE	с	OR	E	SE	
5	8	4	4	2	в			4		[

#### Depth: 296.0 to 297.5 m

Visual Description

0-150 cm: basalt - aphyric, gray to dark gray, vesicular, lightly altered. Vesicles (<1 mm) 40%; in some places. Vesicles are 1-5 mm, 1-2%, partly filled with calcite(?). 88-102 cm: chill zone.

#### Thin Section Description - 116 cm

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

Groundmass: plagioclase 30%, 0.1-0.5 mm, An<sub>60</sub>, subhedral; clinopyroxene 20%, 0.03-0.1 mm, augite, anhedral; glass + mt 20%. Vesicles: 28%, 0.01-0.1 mm Texture: intersertal. Alteration: 1% carbonate in vesicles.

# Shipboard Data

Bulk Analysis: 57 cm SiO2 49.72 15.99 AJ203 Fe203 Fe0 1.16 7.66 MgO 6.57 CaO 12.93 Na20 K20 TiO2 P205 Mn0 3,32 0.37 1.21 0,15 0.15 LOI \_ H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>-</sup> CO<sub>2</sub> -\_ 258.00 60.00 189.00 89.00

Intensity (emu/cc)	420.6
Inclination before	
demag.	50.8
Stable inclination	51.8
Physical Properties:	57 cm
Vp (km/s)	4.15
Porosity (%)	27.29
Wet Bulk Density	2.40
Grain Density	2.93

117 cm

Magnetic Data:



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



Depth 297,5 to 299.0 m

#### Visual Description

0-149 cm: basalt-aphyric, dark gray, vesicular, lightly altered.

0-80 cm: vesicles (<1 m) 40%.

110 cm

50.55

15,19

1.16

7.69

7.25

12.07

3.49

0.39

1.22

0.12

0.11

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222.00

42.00

187.00

85.00

80-149 cm: vesicles 40%, including ones more than 1 mm across (15%). Vesicles partly contain a white mineral (calcite?).

#### Shipboard Data

Bulk Analysis: SiO2 Al2<sup>0</sup>3 Fe2<sup>0</sup>3 Fe0 MgO CaO Na<sub>2</sub>O K20 TiO2 P205 Mn0 LOI H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O<sup>-</sup> CO<sub>2</sub> Cr

Ni

Sr

Zr

Magnetic Data:	118 cm
Intensity (emu/cc)	119.9
Inclination before	
demag.	32.0
Stable inclination	50.8
Other Data:	109 cm
Therm, cond,	
(mcal/cm-s-"C)	4.01



Visual Descriptio	n aburia dark aray, yat	ici
0-27 cm: basarea	phyric, dark gray, ves	ince
Shipboard Data		
Bulk Analysis:	21 cm	
SiO2	49,63	
Fa 0.	1 18	
E-03	7.82	
MaO	7.52	
CaO	12.52	
NagO	3.79	
K-0	0.35	
TiO2	1.21	
P205	0.12	
MnO	0.13	
LOI		
H20'		
H20		
CO2	010 00	
Cr Ni	210.00	
041 C-	194.00	
7.	90.00	
2.1	55.55	

L	EG		SIT	ΓE	HOLE	со	RE	SE	ст.
5	8	4	4	2	в	T	4		3

## Depth: 299.0 to 299.9 m

vesicular, lightly altered. Vesicles (1 mm and less) 40%.



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

L	EG		SIT	TE	HOLE	c	OR	E	SE	ст.
5	8	4	4	2	в			5		1

#### Depth 305.5 to 307.0 m

#### Visual Description

Light gray, vesicular, aphyric basalt, Green tinge increasing with weathering. Vesicles approximately 30 to 40% < 1 mm and 2 to 10% > 1 mm, some infilling by smectite (light olive green) and brown (stain?), clay, occasional calcite, 34-40 cm (6A): weathered surface on angled side, olive yellow alteration products.

64 cm (Top 7C): weathered surface, brown alteration products,

108 cm (Top 10A): yellow brown alteration products.

3.33 0.30 1.31 0.12 0.12 --------\_ 222.00 84.00

188.00

89.00

120-134 cm (10B through 10C): cross cutting fracture with weathered surfaces; light olive brown alteration products and calcite.

120 cm	Magnetic Data:	116 cm
50.27	Intensity (emu/cc)	224.0
15.25	Inclination before	
1.15	demag.	40.6
7.58	Stable inclination	41.0
8.79		
11.68		



LE	G		SIT	re	HOLE	cc	RE	SE	ст.
5	8	4	4	2	в		5	Γ	2

109 cm

25.69

2.40

2.89

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### Depth 307.0 to 308.4 m

#### Visual Description

Zr

- 0-37 cm: green-gray, weathered vesicular, aphyric basalt. Smectite and calcite infilling of vesicles, Fracture from 18-28 cm is lined by calcite and brown alteration zone. White calcite vein at 34-36 cm, cuts into lower fresh zone.
- 37-77 cm: gray, vesicular, variolitic basalt cut by white calcite veins. Some infilling of vesicles mainly by calcite.
- 77-140 cm: green-gray, weathered, vesicular variolitic basalt. Smectite and calcite infilling of vesicles, At 126-132 cm between 2I and 2J - fracture with weathered surface and 1-2 mm wide weathered zone.
- Sharp boundary between fresh and weathered basalt . Thin section at 75 cm and covers lower boundary. Vesicles throughout approximately 30-40% < 1 mm and up to 5% > 1 mm.



84.00

91.00 91.00



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

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\_\_\_\_

226.00

74.00

179.00

90.00

L	EG		SIT	E	HOLE	c	OR	E	SE	CT.
5	8	4	4	2	в			5		3

# Depth 308.4 to 309.9 m

### Visual Description

Shipboard Data

Bulk Analysis:

SiO.

Al-O-

Fe<sub>2</sub>O<sub>3</sub> FeO

MgO

CaO

Na20 K20 TiO2

P205 MnO

LOI

H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>+</sup> CO<sub>2</sub>

Cr Ni

Sr

Zr

Green-gray, vesicular, aphyric basalt. Smectite and/or calcite lining of some vesicles. Vesicles approximately 30 to 40% < 1 mm and approximately 2% > 1 mm. Fracture in Pieces 1D through 1E (29-51 cm) along zone of vesicles. Brown weathered surface,



Magnetic Data:	37 cm
Intensity (emu/cc)	386.7
Inclination before	
demag.	44.2
Other Data:	96 cm
Therm. cond.	
(mcal/cm-s-°C)	3.319







Gray, vesicular, aphyric basalt, green tinge. Vesicles approximately 30 to 40% <1 mm up to 5% >1 mm. Some calcite and smectite infilling of vesicles.



150



53 cm

49.95

14,46

1.17

7.71

7.89

13.05

3.20

0.30

1.21

0.13

0.12

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\_\_\_\_

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-

237.00

64.00

130.00

88.00



Depth: 315.0 to 316.5 m

#### Visual Descriptions

Shipboard Data

Bulk Analysis:

SiO2

Al2<sup>0</sup>3 Fe2<sup>0</sup>3 Fe0

MgO

CaO

Na20 K20 TiO2

P205 Mn0

LOI H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O CO<sub>2</sub>

Cr

Ni

Sr

Zr

0-100 cm: light to moderately weathered aphyric fine-grained basalt, 30% fine vesicles (<1.0 mm), 3% medium and coarse vesicles and vugs up to 1 cm across, 1% calcite filled amygdules. Gray, large vug at center of 2 1/2 x 4 cm dark patch.

101-105 cm: chilled margin (no orientation), 2 mm thick glassy zone,

106-150 cm: same basalt as 0-100 cm interval, larger vugs surrounded by 1 cm path of dark discoloration - probably a reaction zone with contents of vug during cooling; 1-2% calcite amygdules.

Magnetic Data:	48 cm	110 cm
Intensity (emu/cc)	272.2	552.6
Inclination before		
demag.	63.4	55.6
Stable Inclination	6.5	56.7





Depth: 316.5 to 318.0 m

# Visual Description

- 0-24 cm: gray aphyric vesicular basalt. Fine grained, 20-30% fine vesicles, 1-2% medium vesicles, 3% calcite filled amygdules.
- 24-31 cm: quench or chill zone with thin glass rind, aphyric vesicular basalt similar to 0-24 cm interval. Flow structures parallel to rind.
- 30-36 cm: fine-grained aphyric basalt with flow structures. Basalt similar to above.
- 38-44 cm: chill zone with 3 mm glassy zone and variolitic zone grading into aphanitic basalt. ~5-10% very fine vesicles in basalt. Basalt has distinct orange red cast to gray color.
- 46-56 cm: loose basalt fragments.
- 56-68 cm: variolitic basalt, fine-grained greenish gray ~15% fine vesicles.
- 70-73 cm: loose basalt fragments.
- 74-150 cm: aphyric vesicular basalt, gray colored with slight greenish cast. Fine-grained -40% fine vesicles, 1-2% medium vesicles, Ca ide staining where calcite is absent and at marg

#### Shipboard Data

Cr

Ni

Sr

Zr

Bulk Analysis: 32 cm 147 cm SiO2 50.04 50.38 Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO 14.88 15.66 1.13 1.09 7.45 2.20 MgO 8.02 9.84 CaO 12.99 11.02 Na20 K20 TiO2 P205 Mn0 3.14 3.13 0.31 0.24 1.22 1.26 0.11 0.13 0.11 0.13 LOI -----\_ H<sub>2</sub>0<sup>4</sup> H<sub>2</sub>0 CO<sub>2</sub> ---\_ -------277.00 247.00 71.00 82.00 152.00 183.00 86.00 83.00

Magnetic Data	110 cm
Intensity (emu/cc)	568.8
Inclination before	
demag.	49.2
Physical Properties:	40 cm
Vp (km/s)	3.85
Porosity (%)	29.15
Wet Bulk Density	2.33
Grain Density	2.88



т

Piece

cm

0

50-

100-

150 -

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



### Depth: 318.0 to 318.36 m

#### Visual Description

0-16 cm: dark gray to brownish gray basalt, light to moderate weathering, fine vesicles. 18-24 cm: drilling breccia cemented with rock flour paste.

27-30 cm: splotchy variolitic basalt grading upper 1/2 mottled dark and light gray. White calcite vein 1.0-2.0 mm. Vesicles. Fresh.

#### Shipboard Data Bulk Analysis: 28 cm 50.38 SiO2 Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO 15.66 1.09 7.2 MgO 8.2 CaO 11.98 Na20 K20 TiO2 P205 Mn0 3,45 0.34 1.26 0.12 0.13 LOI \_ H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>+</sup> CO<sub>2</sub> ----Cr 232.00 Ni 49.00 Sr 185.00

89.00

Zr

Intensity (emu/cc)	105.
Inclination before	
demag.	67.
Stable Inclination	47.





106 cm

164.8

60.3

54.6

# Depth: 324.5 to 375.7 m

### Visual Description

0-25 cm: gray aphyric basalt. Orangish cast, grades downwards from aphanitic to finegrained. Vesicular: 10-20% fine, 2% medium. Apparently chill margin for underlying basalt.

26-34 cm: basalt fragments.

36-86 cm: fresh basalt - fine-grained, approaching diabase, 5-10% medium vesicles, 5-10% fine vesicles.

86-93 cm: drilling debris and basalt fragments and rock powder.

94-108 cm: fine-grained gray aphyric basalt. Upper 1/3 has 4% medium vesicles. Vesicles scarce to absent in lower 2/3. Otherwise is similar to 36-86 cm interval.

100-107 cm: drilling debris and rock powder.

VISUAL CORE DESCRIPTION

FOR IGNEOUS ROCKS

45-52 cm: same as rest in 36-86 cm.

# Shipboard Data

Shipboard Data	46	110	Magnetic Date:	50 cm
Bulk Analysis:	40 CM	110 cm	magnetic Data:	00 cm
SiO2	50.73	50.57	Intensity (emu/cc)	211.3
AI203	14.77	10.48	Inclination before	
Fe <sub>2</sub> O <sub>3</sub>	1.25	1.26	demag.	55.0
FeO	8.24	8.32	Stable Inclination	
MgO	9.47	8.92		
CaO	10.23	10.55	<b>Physical Properties:</b>	49 cm
Na <sub>2</sub> O	3.24	3.20	Vp (km/s)	4.13
K20	0.24	0.20	Porosity (%)	23.96
TIO	1.22	1.38	Wet Bulk Density	2.50
PoOs	0.14	0.11	Grain Density	2.98
MnO	0.16	0.16		
LOI			Other Data:	
H-0+		191143	Therm. cond.	
H_0-		10000	(mcal/cm-s-°C)	3.84
cô,				
Cr	209.00	211.00		
Ni	59.00	201.00		
Sr	161.00	135.00		
Zr	93.00	67.00		



150

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



#### Depth 334.0 to 335.2 m

#### **Visual Description**

Sr

Zr

0-117 cm: gray aphyric diabase, 1% vesicles at top and less than that through rest of Section. Larger vesicles have free growing glomerocrysts of pyrite and calcite (Piece 2C). At 36 cm, white calcite vein and at several other breaks in core. Pyrites found along vein surfaces where exposed. Fresh-little evidence of any weathering. Grain size fairly uniform down section. Between 23-39 cm ~1-3 mm phenocrysts which could be slightly altered plagioclase which stands out rather than phenocrysts.



201.00

135.00

67.00

Magnetic Data:	47 cm
Intensity (emu/cc)	456.4
Inclindation before	
demag.	50.9
Stable Inclination	51.9
Other Data:	80 cm
Therm. cond.	
(mcal/cm-s-°C)	4.20

Graph Shinh Ö cm n 1A 50-T,M A T.M 1K 100-Continuous piece in next section

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

ц	EG		SIT	ΓE	HOLE	c	OR	E	SE	ст.
5	8	4	4	2	в			8		2

# Depth 335.2 to 336.6 m

# Visual Description

Gray aphyric diabse - < 1% vesicles. Calcite veins common and core has fractured along these (particularly between 0 and 45 cm). Pyrites also found on vein surfaces. Rock is fresh, but much of plagioclase has a greenish cast which may indicate alteration medium grained.

#### Shipboard Data

Magnetic Data:	24 cm
Intensity (emu/cc)	902.7
Inclination before	
demag.	47.8
Physical Properties:	92 cm
Vp (km/s)	4.92
Porosity (%)	12.50
Net Bulk Density	2.70
Grain Density	2.95



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



#### Depth: 336.6 to 338.0 m

# Visual Description

0-140 cm: diabase identical to that described for Sections 1 and 2. Diabase, dense, dark greenish-gray, vesicular, moderately altered. Vesicles 5%, 1-3 mm, all filled with calcite and chlorite.

22-80 cm: vesicles, up to 5 mm in diameter. Diabase-aphyric, fine-grained.

Shipboard Data

Magnetic Data:	55 cm
Intensity (emu/cc)	465.7
Inclination before	
demag.	48.9
Stable Inclination	50.9

150 -

SECT.







È

B

23 cm

48.7

51.3

127 cm

3.744

446.8

CORE

8 5

#### Visual Description

SiO<sub>2</sub> Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO

MgO

CaO

Na20 K20 TiO2 P205 Mn0

LOI

H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>-</sup> CO<sub>2</sub>

Cr Ni

Sr

Zr



0-50 cm and 120-133 cm: vugs filled by calcite with chlorite and at 50-120 cm by chlorite. Vesicles unfilled.

Magnetic Data:

demag.

Other Data:

Therm. cond.

Intensity (emu/cc)

Inclination before

Stable Inclination

(mcal/cm-s-°C)

#### Shipboard Data Bulk Analysis: 74 cm 49.57 14.39 1.21 8.00 10.87 10.38 3.03 0.15 1.18 0.11 0.15 \_ \_\_\_\_ --------255.00

100.00

159,00

84.00



LE	EG		SIT	E	HOLE	co	DRE	SE	ст.
5	8	4	4	2	в	Τ	8		6

#### Depth: 340.7 to 342.1 m

# Visual Description

al Sto

The rocks in this sections is identical to the diabase of Sections 1-5. Diabase, aphyric, vesicular. There is difference between the two parts of the section.

0-65 cm: is identical to the diabase of Section 5. Vesicles 5-10°, <1 mm and 3-5%, 1-3 mm, filled with calcite and chlorite.

65-145 cm: vesicles are open, unfilled. It is the freshest part of the core.

#### Thin Section Description - 84 cm

Groundmass: plagioclase 40%, 0.5-5 mm, An<sub>60</sub>, euhedral, subhedral; clinopyroxene 25%, 1-4 mm, augite, anhedral; magnetite 1%, 0.1-0.6 mm

Vesicles: 5%, 0.7-6.0 mm, round, on the walls rims chlorites and zeolites.

Texture: subophitic.

Alteration: 30% zeolites (in groundmass replacing vesicles) and chlorite (in vesicles replacing groundmass).

Shipboard Data	
Bulk Analysis:	141 cm
SiO2	49.64
AlaDa	13.82
Fe <sub>2</sub> O <sub>2</sub>	1.20
FeO	7.92
MgO	11.47
CaO	10.17
Na <sub>2</sub> O	3.11
K20	0.18
TiO	1.26
P205	0.14
MnO	0.13
LOI	
H-0+	
H_0-	
cô,	
Cr 1	258.00
Ni	105.00
Sr	169.00

91.00

Zr

Magnetic Data:	84 cn
Intensity (emu/cc)	381.4
Inclination before	
demag.	45.5



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



# Depth: 342.1 to 343.6 m

### Visual Description

0-50 cm: upper 50 cm appears identical to diabase described in this core, Sections 1-6. In comparison with the lower part of Section 6, the vesicles are filled with calcite and chlorite.

50-146 cm: another type of alteration of the same diabase. Color returns to a vellow-grav. Vesicles are filled with clay minerals. Calcite vein observed in Pieces 12A and 12B (thickness 1-3 mm).

146-147 cm: reddish color. In this part and above (4-5 cm) the diabase is dense. Probably it is the top of a lava flow.

#### Thin Section Description - 107 cm

50.57

15.26

1.16

7.68

8.70

10.78

3.74

0.43

1.23

0.12

0.13

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228.00

66.00

154.00

94.00

Groundmass: plagioclass 45%, 14 mm, Ango, subhedral, skeletal; clinopyroxene 35%, 1-3 mm, augite, anhedral; magnetite 1%, 0.1-0.5 mm, skeletal, rectangle Vesicles: 10%, 1-3 mm, round.

Texture: ophytic.

Alteration: 10% zeolites (in vesicles) and chlorite (in vesicles, groundmass replacing groundmass), partly colored by hydro iron oxide.

#### Shipboard Data Bulk Analysis:

SiO2

Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO

MgO

CaO

Na20 K20 TiO2

P205 Mn0

LOI

H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O<sup>+</sup> CO<sub>2</sub>

Cr

Ni

Sr

Zr

Magnetic Data:	107 cm
Intensity (emu/cc)	261.3
Inclination before	
demag.	48.0
Stable Inclination	48.3



08



- Depth: 343.5 to 345.0 m
- 0-49 cm: basalt, glomerophyric yellow-gray, vesicular, altered. Vesicles 15% (1-2 mm and less) are open and contain a few iron oxides. Aggregates of plagioclase (diameter 1-1.5 mm)
- 49-150 cm: basalt, aphyric, dense, yellow-gray, vesicular, altered. Vesicles less than 0.5 mm (5-10%) partly filled with calcite and iron oxide.
- 65-80 cm: two calcite veins with oxidized ore minerals (probably pyrite).

Phenocrysts: plagioclase 50%, 0.3-3 mm, An58, subhedral, sometimes light zonation; clinopyroxene 20%, 0.5-2 mm, augite, euhedral.

Groundmass: groundmass substitute by secondary minerals.

Texture: glomerophyric, intersertal, subophitic.

Alteration: clays, zeolites, chlorite - 25% in groundmass, groundmass partly filled vesicles.

 $\label{eq:Groundmass: plagioclase 20\%, 0.2-1 mm, An_{60}, subhedral, needles; clinopyroxene 25\%, 0.2-3 mm, augite, anhedral; magnetite < 1, 0.04-0.08; glass 50\%, microlite, Pl+Py+Mt.$ Vesicles: 5%, 0.2-1.0 mm, partly filled carbonate and clay.

Alteration: 3% carbonate in vesicles and 2% clays in vesicles.

84.00

4 cm	107 cm	Magnetic Data:	24 cm	33 cm	103 cm
9.46	49.31	Intensity (emu/cc)	455.8	448.9	544.8
5.42	13.18	Inclination before			
1.10	1.19	demag.	- 1.2	- 3.1	53.8
7.35	7.82	Stable Inclination	- 3.7	- 2.1	56.3
9.22	9.02				
1.03	12.60				
2.97	3.05	<b>Physical Properties:</b>	110	cm	
0.30	0.59	Vp (km/s)	4.	16	
1.16	1.11	Porosity (%)	28.	.02	
0.11	0.11	Wet Bulk Density	2.	41	
0.14	0.13	Grain Density	2.	96	
		Other Data:	67 c	m	
		Therm, cond.			
		(mcal/cm-s-°C)	3.	.66	
0.00	282.00				
8.00	100.00				
4.00	173.00				



0 0

150

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



#### Depth: 345.0 to 346.5 m

#### Visual Description

0-150 cm: basalt is identical to that described at the base of Section 1. Aphyric, dark gray,

vesicular, altered. Vesicles, 0.5 mm and less than 10%.

114-150 cm: basalt more dense and has 1% vesicles (1-3 mm).

114-126 cm: vesicles partly filled with calcite.

0-10 cm: a calcite vein occurs with small iron oxides (Piece 1).

#### Thin Section Description - 73 cm

Phenocrysts: plagioclase 10%, 0.4-1 mm, An<sub>60</sub>?, needles. Groundmass: glass 50%, microlites, Py+PI+glass. Vesicles: 40%, 0.1-1.0 mm, unfilled, round Texture: intersertal.

Data		
sis:	37 cm	Magnetic Data:
	49.32	Intensity (emu/
	14.04	Inclination befo
	1.12	demag.

7.42

7.66

14.12

3.13

0.39 1.16

0.12

0.12

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270.00

100.00

173.00

84.00

Intensity (emu/cc) 1940.1 Inclination before demag. 47.6 Stable Inclination 47.8

73 cm



LE	EG		SIT	E	HOLE	cc	ORE	SE	CT.
5	8	4	4	2	В	Т	9		3

# Depth: 346.5 to 347.6 m

# Visual Description

0-110 cm: basait, aphyric, dark gray, vesicular, altered. More vesicles than in the basait of Section 2 (10-15%, <1 mm). Partly filled with calcite and iron oxides.

In Pieces 1B, 1C, and 2E: a calcite vein occurs with an altered zone next to it (iron oxide alteration).

#### Thin Section Description - 32 cm

Groundmass: plagioclase 30%, 0.5-2 mm, An<sub>68</sub>, needles, partly laths; clinopyroxene 20%, 0.2-1.5 mm, augite, subhedral, euhedral; glass 45%, mixrolites, PI+Py+Mt.

Vesicles: 3%, 0.2-2.0 mm, partly filled by calcite.

Texture: subophitic - intersertal.

Alteration: 1% carbonate, at 0.2-1.5 mr

#### Shipboard Data

Bulk Analysis:	54 cm
SiO <sub>2</sub>	50.61
Al202	13.44
FeoOs	1.34
FeÔ	8.87
MgO	9.33
CaO	10.06
Na <sub>2</sub> O	3.42
K20	0.35
TIO	1.42
P205	0.15
MnO	0.14
LOI	-
H20 <sup>+</sup>	-
H20-	
cõ,	
Cr	251.00
Ni	118.00
Sr	177.00
Zr	106.00

m re	placing vesicles.		
	Magnetic Data:	32 cm	105 cm
	Intensity (emu/cc)	768.1	437.6
	Inclination before		
	demag.	45.7	43.2
	Stable Inclination	45.9	49.0



cm

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VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



# Depth: 353.0 to 354.5 m

# Visual Description

Light gray, vesicular, aphyric, aphanitic basalt (approximately 10% vesicles). Glassy margins at 30 cm (6), 41 cm (8), 60 cm (11), 87 cm (16), 92 cm (17), 97 cm (18), and 116 cm (21). Vesicles generally <1 mm across and collect in zones approximately 1 cm below glassy margins, some infilling by clay minerals and occasional calcite. 97-101 cm (18): several fractures infilled by alteration products (clays?).

#### Thin Section Description - 94 cm

Groundmass: plagioclase 10%, 0.02-0.2 mm, laths or acicular, microlites; glass 70% and cryptocrystalline material 20%. Texture: glassy.

#### Thin Section Description - 94 cm

Phenocrysts: plagioclase 0.5%, 0.2-1 mm, An58, laths. Groundmass: cryptocrystalline matrix 68.5%. Vesicles: 20%, 0.02-1 mm Texture: sparsely phyric. Alteration: 1% clays in vesicles.

110 cm

50.16

15.21

1.33

8,79

5.47

12.27

3.33

0.40

1.58

0.19

0.17

-

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150.00

50.00

162.00

105.00

# Shipboard Data

Bulk Analysis: SiOn Al203 Fe<sub>2</sub>O<sub>3</sub> FeO MgO CaO Na20 K20 TiO2 P205 Mn0 LOI H<sub>2</sub>0<sup>4</sup> H<sub>2</sub>0 CO<sub>2</sub> Cr Ni Sr

Zr





82

SITE

5 8 4 4 2 B

Depth 362.5 to 364.0 m

CORE SECT.

111

LEG

- 2) 14-128 cm: light gray-green, vesicular, aphyric basalt except for dark gray zones next to glassy margins (approximately 5 mm thick).
- 86-95 cm and 113-121 cm (Piece 16): abundant large (1-3 mm) infilled vesicles possibly formed by accretion of small vesicles (<1 mm).
- 3) 128-150 cm: light gray, vesicular, aphyric basalt. Vesicles generally <1 mm (approximately

143-150 (Piece 20): 2 linear groups of larger vesicles line by alteration products.

Phenocrysts: spinel 0.1 mm, rounded, triangle, 1 piece red-brown.

Groundmass: plagioclase 20%, 0.2-1 mm, acicular, some with plumose texture (pyroxene

Alteration: 1% clays lining fracture and replacing some crytocrystalline material.

Magnetic Data:	60 cm
Intensity (emu/cc)	108.2
Inclination before	
demag.	33.3
Stable Inclination	28.9



LI	EG		SIT	E	HOLE	c	OR	E	SE	SECT.	
5	8	4	4	2	в	1	1	1		2	

#### Depth 364.0 to 365.4 m

#### Visual Description

Light gray, vesicular, aphyric basalt when lightly to moderately weathered green-gray where moderately weathered.

0-2 cm (Piece 1): glassy zone showing alteration to yellow palagonite.

87-88 cm (Piece 12A): glassy zone completely replaced by yellow palagonite and chilled margin, altered.

Some vesicles lined by brown alteration product others by light olive green smectite. 37-47 cm (Piece 6): drilling breccia.

# Thin Section Description - 28 cm

Groundmass: plagioclase 25%, 0.02·1 mm, An<sub>62</sub>, elongate laths; clinopyroxene 10%, 0.02· 0.5 mm, acicular; magnetite 2%, <0.02 mm, granular; other 28%. Vesicles: 35%, 0.05-0.5 mm, rounded.

Texture: intersertal.

# Shipboard Data

Cr

Ni

Sr

Zr

70 cm Bulk Analysis: SiO<sub>2</sub> Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO 49.99 15.38 1.28 8.48 MgO 6.41 CaO 12.05 Na20 K20 TiO2 P205 Mn0 3.25 0.39 1.50 0.21 0.16 LOI -----H<sub>2</sub>0<sup>4</sup> H<sub>2</sub>)<sup>-</sup> CO<sub>2</sub> --192.00 58.00 185.00

114.00

Magnetic Data:	127 cm
Intensity (emu/cc)	57.3
Inclination before	
demag.	48.8
Physical Properties:	121 cm
Vp (km/s)	4.13



VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



# Depth: 365.4 to 366.1 m

#### Visual Description

Light gray, vesicular, aphyric basalt gray-green when moderately weathered. Vesicles approximately 20% <1 mm. 10-20 cm: several large vesicles 3 mm to 10 mm across, lined by light gray or brown material. 55-67 cm: alteration zone (brown coloration) through Piece 7. 46-54 cm (Piece 6): drilling breccia. Thin Section Description - 19 cm Phenocrysts: plagioclase 1%, 0.3-1.20 mm, laths, euhedral.

Groundmass: plagioclase 30%, 0.2-0.5 mm, An<sub>69</sub>, acicular laths, microlites; magnetite 0.5%, 0.01-0.02%, granular, other, cryptocrystalline matrix. Vesicles: 35%, 0.02-1 mm. Texture: very sparsely phyric (aphyric) Alteration: 1% clays lining vesicles.

#### Shipboard Data

H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O<sup>+</sup> CO<sub>2</sub>

Cr

Ni

Sr

Zr

Bulk Analysis:	24 cm	Magnetic Data:	18 cm
SiO <sub>2</sub>	49.98	Intensity (emu/cc)	112.0
AloOa	15.74	Inclination before	
Fe <sub>2</sub> O <sub>2</sub>	1.28	demag.	43.7
FeÔ	8.48	Stable Inclination	42.1
MgO	5.19		
CaO	12.26		
Na <sub>2</sub> O	3.30		
K20	0.38		
TiO2	1.54		
P205	0.21		
MnO	0.18		
LOI	2.00		

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267.00

71.00

178.00

115.00





VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

L	EG		SIT	Ē	HOLL	c	OR	E	SE	ст.
5	8	4	4	2	в	1	1	2		1

27.3

114 cm

89.80

32.6

29.8

#### Depth: 372.0 to 373.3 m

#### **Visual Description**

Shipboard Data

Bulk Analysis:

SiO<sub>2</sub> Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO

MgO

CaO

Na20 K20 TiO2 P205 Mn0

LOI

H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>+</sup> CO<sub>2</sub>

Cr

Ni

Sr

Zr

Light gray, vesicular, aphyric basalt, gray-green when moderately weathered, chill zones below glass (relatively fresh).

0-5 cm (Piece 1), 33-39 cm (Top of Piece 6), 96-99 cm (Bottom of Piece 13), and 113-125 cm (Piece 15 below chilled margin): all regions with large vesicles 1 mm to 5 mm across. In above regions 30-40% vesicles, otherareas approximately 10% vesicles generally <1 mm. Vesicles are lined by smectite, calcite or brown material.

71-77 cm and 102-110 cm: drilling breccia. 52 cm

49.27

15.66

1.38

9.11

4.58

12.21

2.97

0.58 1.55 0.31 0.19

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-

194,00

67.00

181.00

120.00

Magnetic Date:	40 cm
wagnetic Data.	40 611
Intensity (emu/cc)	90.96
Inclination before	
demag.	31.7

Stable Inclination



150

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

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201.00

66.00

171.00

124.00



Depth: 381.5 to 383.0 m

#### Visual Description

vesicles mainly	<1 mm up to 10%	>1 mm.	
0-130 cm: light	to moderately weat	hered, green tinged, light gray, ve	sicular, aphyric basalt.
No glass margi	ns or chilled zones i	n this section, uniform texture.	
30-139 cm: drill	ling breccia, include	s glass fragments.	
39-150 cm: mod	derately weathered,	green-gray vesicular, aphyric, aph	nanitic basalt,
10% vesicles of	f which 7.5% are be	tween 1 mm and 5 mm across.	
hin Section Des	cription – 61 cm		
roundmass: play	gioclase 25%, 0.2-1	nm, Ango, laths; clinopyroxene	25%, 0.1-0.5 mm, anhedral;
agnetite 0.5%, -	< 0.02 mm, granular		
ther: 34.5%, cry	ptocrystalline matr	ix.	
esicles: 15%, 0.2	2-0.5 mm.		
exture: intersert	tal.		
hipboard Data			
ulk Analysis:	34 cm	Magnetic Data:	60 cm
02	50.12	Intensity (emu/cc)	102.3
1203	15.38	Inclination before	
e203	1.33	demag.	52.6
eÕ	8.79	Stable Inclination	38.1
Ogl	5.97		
Os	12.09	Physical Properties:	32 cm
a-0	3.16	Vp (km/s)	3.93
20	0.42	Porosity (%)	27.04
0,	1.68	Wet Bulk Density	2.42
05	0.24	Grain Density	2.95
InO	0.19		



5	8	4	4	2	в	1	1	3		2
LE	G		SIT			OR	E	SE	CT.	

# Depth: 381.5 to 383.0 m

# Visual Description

Moderately weathered pillow basalt. Aphyric, fine-grained to aphanitic. Plagioclase laths just visible, 0-15% vesicles, mostly fine. Calcite veins in center of Piece 17, side of Pieces 1, 2, 3, 4, 5, 8, 10, and 11. May be additional mineral in veins - zeolite?.

Piece 17, 127-130 cm: has glassy chill margin.

Pieces 8, 16, and 17: have composite vesicles containing highly vesicular darker basalt. <1% amygdules; calcite filled.



Cr

Ni

Sr

Zr

Magnetic Data:	100 cm
Intensity (emu/cc)	93.97
Inclination before	
demag.	37.7
Stable Inclination	34.9



100-

150 -

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T,N

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS

LEG SITE CORE SECT 3 5 8 4 4 2 B 1 3

#### Depth: 383.0 to 383.3 m

#### Visual Description

0-10 cm: aphyric fine-grained basalt. Cracks contain some calcite, < 1% calcite amygdules, <1% vesicles.

10-15 cm: vesicular aphyric basalt with glass rind. Composite vesicles with highly vesicular dark basalt in their interiors.

18-25 cm: fine-grained basalt, 3% fine vesicles. A few amygdules - (possibly iddingsitized olivine, <1% small 3x1 mm plagioclase phenocrysts.

#### Thin Section Description - 3 cm

Groundmass: plagioclase 33%, 0.1-1 mm, An<sub>60</sub>, acicular, microlites; magnetite, trace, very finegrained; other 33%, cryptocrystalline matrix, Vesicles: 33%, 0.1-0.4 mm.

Texture: intersertal,

Alteration: 1% clays replacing cryptocrystalline matrix.

51.00

182.00

115.00

#### Shipboard Data

Ni

Sr

Zr

Bulk Analysis:	19 cm	Magnetic Data:	2 cm
SiO <sub>2</sub>	49.63	Intensity (emu/cc)	156.4
Al203	15.39	Inclination before	
Fe2O3	1.32	demag.	22.7
FeO	8.71	Stable Inclination	20.4
MgO	5.38		
CaO	12.36		
Na <sub>2</sub> O	3.21		
K20	0.46		
TIO2	1.53		
P205	0.35		
MnO	0.14		
LOI			
H20 <sup>+</sup>			
H20-			
cô <sub>2</sub>			
Cr	195.00		

CORE SECT.

2





L	EG		SIT	E	HOLE	0	CORE		SECT.	
5	8	4	4	2	в	-j	1	5		1

### Depth:

#### Visual Description

Aphyric pillow basalt, Pillow rinds common and appear to zone from glass (nonvesicular) to dense brown (moderately weathered) basalt which is aphanitic to fine-grained into dark gray lightly weathered highly vesicular basalt. Vesicular basalt has ~30% vesicles, brown to moderately weathered zone has ~5% vesicles (generally medium sized) and glass generally has no vesicles. Gray basalt has brown vesicular inclusions ~1/2 cm across. Four glassy chill margins. Exterior of pillow rind of Piece 2 is coated with 1 mm thick coating of crystalline calcite.

# Thin Section Description - 69 cm

Groundmass: plagioclase 20%, 0.05-0.5 mm, An<sub>55</sub>, acicular or elongate laths; clinopyroxene 15%, 0.05-0.4 mm; magnetite < 0.02 mm, trace; other 30%. Vesicles: 35%, 0.1-1.5 mm .

# Texture: intersertal. Shipboard Data

Bulk Analysis:	40 cm
SiO2	49.91
AlaÕa	15.76
Fe2O3	1.25
FeO	8.24
MgO	5.85
CaO	12.76
Na <sub>2</sub> O	2.80
K20	0.45
TIO2	1.26
P205	0.40
MnO	0.16
LOI	
H20+	
H20-	
cõ,	
Cr	49.00
Ni	32.00
Sr	219.00
7.	00 00

Zr

hysical Properties:	77 cm
/p (km/s)	4.14
orosity (%)	24.46
Vet Bulk Density	2.42
Grain Density	2.88



150

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



#### Depth: 410.0 to 411.5 m

# Visual Description

10-15 cm: pillow rind, oriented vertically, only glass in core.

0-10 cm: lightly weathered aphyric vesicular basalt. Composite vesicles of highly vesicular brown basalt in gray basalt. Fine vesicles ~20% coarse ~3%. Fine-grained.

10-45 cm: dark gray aphyric vesicular basalt, pillow rind on upper fragment. Fine vesicles ~10%, medium ~7%. Fine-grained.

46-40 cm: gray brown aphyric vesicular basalt. Fine vesicles 20%, no medium vesicles. 51-56 cm: reddish gray vesicular basalt with chill margin. Composite vesicles. Medium vesicles ~ 10% fine ~2%.

57-146 cm: fine-grained aphyric basalt, <1% fine vesicles at top, none at base of section.

#### Shipboard Data

**Bulk Analys** SiO<sub>2</sub> Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO MgO CaO Na20 K20 TiO2 P205 Mn0 LOI H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>+</sup> CO<sub>2</sub> Cr

sis:	24 cm	47 cm	111 cm	Magnetic Data:	29 cm	35 cm	105 cm
	50.21	49.85	50.55	Intensity (emu/cc)	504.8	154.6	349.2
	16.35	16.20	15.67	Inclination before			
	1.31	1.24	1.14	demag.	-47.4	-42.2	-45.2
	8.63	8.15	7.58	Stable Inclination		-43.5	-46.6
	5.95	5.84	7.77				
	11.78	12.78	12.24	<b>Physical Properties:</b>	120 cm		
	3.26	2.73	2.96	Vp (km/s)	4.47		
	0.55	0.44	0.35	Porosity (%)	17.65		
	1.34	1.25	1.16	Wet Bulk Density	2.54		
	0.18	0.21	0.12	Grain Density	2.87		
	0.16	0.17	0.12				
	35.00	52.00	50.00				
	39.00	42.00	59.00				
	208.00	210.00	200.00				
	95.00	88.00	81.00				





LEG SITE SECT. CORE 5 8 4 4 2 B 1 17

# Depth: 417.0 to 418.5 m

#### Visual Description

SiO2

Al20

Fe<sub>2</sub>O<sub>3</sub> FeO

MgO

CaO

Na20 K20 Ti02

P205 Mn0

LOI

H<sub>2</sub>0<sup>+</sup> H<sub>2</sub>0<sup>+</sup> CO<sub>2</sub>

Cr

Ni

Sr

Zr

Aphyric fine-grained to aphanitic, gray basalt, 10-25% vesicles, <1% very small plagioclase phenocrysts. Some samples have composite vesicles, lightly weathered. Glass at Pieces 1, 16, 17, 18, 19, 20, 21, 22, 23, 28.

Piece 1 has crystalline calcite covering glass rind, and rind has a thick palagonite zone 2-3 mm thick

Piece 17 has indurated sediment with palagonite fragments adhering to fracture surface of the basalt.









LEG		SITE			HOLE	CORE		SECT.		
5	8	4	4	2	в	Т	1	9		1

43 cm

217.2

-31.0

-34.3

Depth: 436.0 to 437.5 m

# **Visual Description**

Alter

0-145 cm: basalt representing top of lava flow. Aphyric, fine-grained. Pieces 1 and 2 have glass on the one side on the surface. The glass is dark, the rock - dark-gray, the border is gray (1-2 mm). Vesicular, lightly altered. Vesicles about 5-10% (<1 mm). Vesicles 1-2 mm, 2%. 14-150 cm: basalt, aphyric, dark gray, vesicular, lightly altered. Vesicles 10-15% (<1 mm), unfilled.

100-150 cm: dense basalt.

#### Thin Section Description - 42 cm

Groundmass: plagioclase 25%, 0.2-1.5 mm, An<sub>64</sub>, euhedral, subhedral; clinopyroxene 30%, 0.03-0.5 mm, augite, anhedral; glass + Mt 20%, in groundmass and around vesicles. Vesicles: 25%, 0.1-2.0 mm, round irregular.

Texture: intergranular-intersertal. Alteration: very few zeolites.

1.60

0.16 0.15

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58.00

59.00

173.00

123.00

# Shipboard Data Bulk Analysis:

SiO<sub>2</sub> Al<sub>2</sub>O<sub>3</sub> Fe<sub>2</sub>O<sub>3</sub> FeO

MgO CaO

Na<sub>2</sub>O K20 TiO2 P205 MnO

LOI H<sub>2</sub>O<sup>+</sup> H<sub>2</sub>O<sup>+</sup> CO<sub>2</sub>

Cr

Ni

Sr

Zr

45 cm	Magnetic Data:
51.65	Intensity (emu/cc)
14.87	Inclination before
1.28	demag.
8.48	Stable Inclination
6.64	
11.48	
3.75	
0.37	



150 -

VISUAL CORE DESCRIPTION FOR IGNEOUS ROCKS



#### Depth: 437.5 to 438.2 m

#### Visual Description

Orientation

D,S,P

Graphic Represen

6

6

0

0

6.0

0

0

0

0-28 cm: upper 28 cm appears identical to that described at the base of Section 1. 28-60 cm: similar basalt but more vesicular. Aphyric, dark gray, lightly altered. Vesicles <0.5 mm, 15%.

60-70 cm: pillow lava. Pieces 9 and 10 with glass on the one side of surface. Basalt is dense (fine-grained) dark gray, lightly altered. Vesicles < 2 mm, 5-10%. Basalt close to that of top of lava flow 0-14 cm, Section 1), but fresher.

Thin Section Description - 62 cm

Phenocrysts: plagioclase 3%, 0.7-0.1 mm, An<sub>60</sub>, euhedral; clinopyroxene 1%, 0.5-0.1 mm, augite, subhedral.

Groundmass: plagioclase mictolite; clinopyroxene mictolite.

Vesicles: 30%, 0-1.5 mm, round, irregular, unfilled.

Texture: variolitic.

#### Shipboard Data

Sr

Zr

Bulk Analysis:	36 cm	51 cm	Magnetic Data:	62 cm
SiO2	50.93	50.65	Intensity (emu/cc)	182.8
Al203	15.78	15.46	Inclination before	
Fe <sub>2</sub> O <sub>2</sub>	1.26	1.28	demag.	-27.8
FeO	8.32	8.48	Stable Inclination	-36.4
MgO	5.64	5.76		
CaO	12.35	12.22	Physical Properties:	12 cm
Na <sub>2</sub> O	3.56	3.62	Vp (km/s)	5.25
K20	0.38	0.44	Porosity (%)	9.11
TiO	1.59	1.53	Wet Bulk Density	2.80
POOF	0.20	0.18	Grain Density	2.9
MnO	0.16	0.19		
LOI	1			
H_0+	-			
H_0-				
cóa				
Cr	59.00	60.00		
NI	57.00	62.00		

189.00 188.00 123.00 129.00



# SITE F CORE SECT. 5 8 4 4 2 B 2 0 1

27 cm

387.9

-22.1

-21.6

4.08

# Depth: 445.5 to 446.2 m

Pieces at 0-4 cm, 18-24 cm, 55-60 cm: have glass veins. Vesicles 30-40% (<5-7 mm).

S
Ξ
H
N
4
N

89



Hole 442A (no photograph available for Cores 442-1 and 442A-1)












SITE 442



Hole 442A



99





SITE 442







104





106



