

Original paper

Babánekite, $\text{Cu}_3(\text{AsO}_4)_2 \cdot 8\text{H}_2\text{O}$, from Jáchymov, Czech Republic – a new member of the vivianite groupJakub PLÁŠIL^{1*}, Pavel ŠKÁCHA^{2,3}, Jiří SEJKORA², Radek ŠKODA⁴, Milan NOVÁK⁴, František VESELOVSKÝ⁵, Jan HLOUŠEK[†]¹ Institute of Physics, Academy of Sciences of the Czech Republic (ASCR) v.v.i, Na Slovance 2, 182 21 Prague 8, Czech Republic; plasil@fzu.cz² Department of Mineralogy and Petrology, National Museum, Cirkusová 1740, 193 00 Prague 9, Czech Republic³ Mining Museum Příbram, Hynka Kličky Place 293, 261 01 Příbram VI – Březové Hory, Czech Republic⁴ Department of Geological Sciences, Faculty of Science, Masaryk University, Kotlářská 2, 602 00 Brno, Czech Republic⁵ Czech Geological Survey, Geologická 6, 152 00 Prague 5, Czech Republic[†] Jáchymov; deceased on April 27, 2014

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Babánekite, $\text{Cu}_3(\text{AsO}_4)_2 \cdot 8\text{H}_2\text{O}$, a new member of the vivianite group was found in material originating from the Geister vein, Rovnost mine, Jáchymov, Western Bohemia, Czech Republic. It occurs as a supergene alteration mineral in association with members of the lindackerite supergroup (veselovskýite, hloušekite, pradetite and lindackerite), lavendulan, gypsum and an X-ray amorphous Cu–Al–Si–O–H phase. Crystals of babánekite are pinkish to peach-colored, elongated, prismatic and up to 2 mm in length. They exhibit the forms $\{010\}$, $\{100\}$, $\{110\}$, $\{101\}$ and less frequently also $\{001\}$. Crystals are transparent to translucent with a vitreous luster. The mineral has a light pinkish streak. Estimated Mohs hardness is between 1.5 and 2. The cleavage is perfect on $\{010\}$. The calculated density is 3.192 g/cm^3 . Electron-microprobe analysis yielded CoO 8.89, NiO 4.06, CuO 15.31, ZnO 10.87, P_2O_5 0.16, As_2O_5 39.79, SO_3 0.13, H_2O 24.78 (calc.), total 103.99 wt.% giving the empirical formula $(\text{Cu}_{1.12}\text{Zn}_{0.78}\text{Co}_{0.69}\text{Ni}_{0.32}\text{As}_{2.91})[(\text{AsO}_4)_{2.01}(\text{PO}_4)_{0.01}(\text{SO}_4)_{0.01}]_{\Sigma 2.03} \cdot 8\text{H}_2\text{O}$ based on 16 O apfu. The ideal end-member formula of babánekite is $\text{Cu}_3(\text{AsO}_4)_2 \cdot 8\text{H}_2\text{O}$, which requires CuO 38.95, As_2O_5 37.52, H_2O 23.53, total 100.00 wt.%. Babánekite is monoclinic, $C2/m$, with $a = 10.1729(3)$, $b = 13.5088(4)$, $c = 4.7496(1) \text{ \AA}$, $\beta = 105.399(2)^\circ$, $V = 629.28(3) \text{ \AA}^3$ and $Z = 2$. The eight strongest X-ray powder diffraction lines are [d_{obs} $\text{Å}(l)(hkl)$]: 7.936(11)(110), 6.743(100)(020), 3.231(14)(13-1), 2.715(11)(041), 2.3331(10)(15-1), 2.0819(5)(350), 1.6862(16)(080) and 1.6107(4)(55-1). The crystal structure of babánekite, refined to $R_1 = 2.18 \%$ for 864 unique observed reflections, confirmed that the atomic arrangement is similar to other members of the vivianite group of minerals. The mineral is named for Ing. František Babánek (1836–1910), Czech mining and geological expert, who worked in the Jáchymov and Příbram mines.

Keywords: babánekite, new mineral, vivianite group, crystal structure, Jáchymov

Received: 13 September, 2017; accepted: 3 January, 2018; handling editor: F. Laufek

The online version of this article (doi: 10.3190/jgeosci.248) contains supplementary electronic material.

1. Introduction

The vivianite group unifies hydrated arsenates and phosphates with the generalized structural formula $\text{Me}_3(\text{XO}_4)_2(\text{H}_2\text{O})_8$, where *Me* designates sites that are occupied by various divalent cations and *X* the site occupied by As^{5+} or P^{5+} . To date, six As-dominant members of this group are known: annabergite (Ni), erythrite (Co), hörnesite (Mg), manganohörnesite (Mn), köttigite (Zn) and parasymphesite (Fe). In addition, four P-dominant members have been described: arupite (Ni), baričite (Mg, Fe), pakhomovskite (Co) and vivianite (Fe). Broad solid-solutions taking place at the cationic sites have been established especially for As-dominant members.

In this work we report on the new mineral species babánekite, ideally $\text{Cu}_3(\text{AsO}_4)_2 \cdot 8\text{H}_2\text{O}$. It is a Cu- and

As-dominant member of the vivianite group that was found in the Rovnost mine, Jáchymov, Czech Republic. The new mineral honors the Senior Mining Counselor (“Oberbergrath”) František Babánek (born October 10, 1836, in Kamenný Přívoz near Jílové, Bohemia; died February 25, 1910, in Prague). František Babánek was a Czech mining expert, geologist and mineralogist. During his active service, he worked in Příbram and later in Jáchymov, the two most important mining districts in Bohemia, then a part of the Austro-Hungarian Empire.

The new mineral was approved by the CNMNC IMA under the code IMA 2012-007. The type specimen is deposited in the collections of the Department of Mineralogy and Petrology of the National Museum, Prague, Czech Republic, catalog number: PIP 8/2011.

2. Occurrence

The Jáchymov ore district (Western Bohemia, Krušné hory Mountains, Czech Republic) presents a classic example of Ag–As–Bi–Co–Ni–U hydrothermal vein-type mineralization (Ondruš et al. 2003; Hloušek et al. 2014). The ore-bearing veins cut medium-grade metasedimentary rocks of Cambrian to Ordovician age, which surround a Variscan granitic pluton. The majority of the primary ore minerals were deposited from mesothermal fluids during Variscan mineralizing processes. More than 430 minerals have been described from Jáchymov, including an extremely diverse assemblage of supergene minerals (see Ondruš et al. 1997, 2003; Hloušek et al. 2014 for review).

Babánekite was found in an old ore-stope on the Geister vein at the 3rd Geister level of the Rovnost (former Werner) mine. The particular locality, Geister vein, was opened by old mining workings from the surface down to the level of the Daniel drainage adit, at the 1st, 3rd, 6th, Barbora and Daniel levels. This vertical sequence mostly represents a supergene oxidation zone of the vein opened by mining workings, with the very rich zones of the supergene cementation Ag–Hg mineralization at the two last mentioned adit levels (Vogl 1854; Škácha et al. 2014).

The old, near-surface workings (from the 1st to the 6th level), originating partly from the 16th century, were re-opened during the first half of the 19th century. The main aim was mining for U, used for glass-staining, and later mostly for extraction of radium. The re-opening of the 3rd and 6th Geister levels was well documented by the mining official Josef Florian Vogl, after whom the mineral voglite is named (Haidinger 1853; Ondruš et al. 1997). Vogl (1856a, b) described in detail a varied supergene mineral association from the Geister vein constituted by arsenates of the vivianite group and lindackerite su-

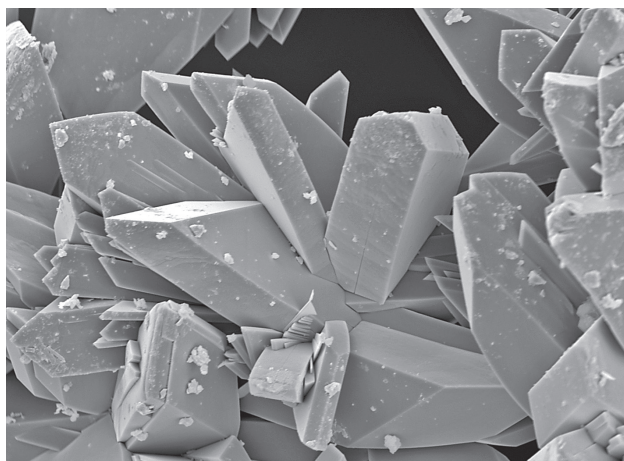


Fig. 1 Detail of babánekite crystals. Secondary electron (SE) image (Hitachi 3700N), 150 μm wide.

pergroup, as well as supergene uranyl-bearing minerals. The richness of the locality results from the occurrence of both the recently/sub-recently formed minerals connected with the post-mining processes and the supergene minerals formed *in-situ* in the oxidation zone (association of uranyl arsenates and vanadates; association of Pb–Cu supergene minerals and minerals containing Y + REE).

Babánekite was found by one of the authors (PŠ) in an old mining field informally called the “lindackerite stope” by mineral collectors. The name is due to frequent occurrences of the lindackerite-supergroup minerals: lindackerite, veselovskýite (Sejkora et al. 2010), hloušekite (Plášil et al. 2014a), pradetite (unpublished data of the authors) and klajite (Plášil et al. 2014b; Hloušek et al. 2014). These minerals crystallize on the strongly altered ore-body consisting mainly of massive tennantite, galena and chalcopyrite with disseminated uraninite in quartz. Babánekite aggregates grow in cavities and on the surface of ore fragments. In close association with babánekite, members of the lindackerite supergroup, lavendulan, gypsum and X-ray amorphous Cu–Al–Si–O–H material were detected, overgrowing relics of the primary minerals partly cemented by supergene mineralization.

3. Physical and optical properties

Babánekite forms elongated, prismatic crystals (Fig. 1) up to 1.5 mm in length, grouped in hemispherical aggregates up to 2 mm (Fig. 2). Crystals exhibit the forms $\{010\}$, $\{100\}$, $\{110\}$, $\{101\}$ and less frequently also $\{001\}$. Crystals, transparent to translucent, are pinkish to peach in color and have a vitreous luster. Babánekite has a light pink streak. The Mohs hardness is estimated between 1.5 and 2.5 (in analogy with other As-dominant members of the vivianite group). Prismatic crystals show perfect cleavage on $\{010\}$. A density of 3.192 g/cm³ was calculated using the empirical formula and unit-cell parameters obtained from single-crystal X-ray diffraction. Direct density measurements were not undertaken due to paucity of pure material. Babánekite is non-fluorescent under short- or long-wave UV radiation. Optical properties could not be determined due to zonation of the babánekite crystals available (Fig. 3); the calculated (Gladstone–Dale rule) average n is 1.6615.

4. Chemical composition

The chemical composition of babánekite was determined at Masaryk University in Brno using a Cameca SX100 electron microprobe (WDS mode, 15 kV, 5 nA, and 10 μm beam diameter). The following X-ray lines and standards were selected: K_{α} lines: P (fluorapatite), Zn (ZnO), Cu (lammerite), Co (metallic Co), Ni (Ni) and S (baryte),



Fig. 2 Babánekite crystal aggregates consisting of elongated prismatic crystals of dominating {010} faces. The bluish material is a PXRD-amorphous Cu–Al–Si–O–H phase; metallic minerals are tennantite and chalcopyrite. The width of the image is 3.5 mm.

L_{α} line: As (lammerite). Other elements, including Al, Bi, Ca, Cl, Fe, K, Mg, Mn, Na, Pb, Si and V, were also sought, but not found (the detection limits for these elements are ~0.05–0.10 wt.% at the analytical conditions used). Peak counting times were 10–20 s and the counting time for the background their half. The measured intensities were processed for matrix effects using the “PAP” correction routine (Pouchou and Pichoir 1985). The H₂O content (8 pfu) was calculated based on the stoichiometry obtained from the structure data and by analogy with other vivianite-group minerals.

An aggregate removed from the holotype specimen – that was also used for single-crystal and powder X-ray diffraction experiments – is characterized by the presence of slightly Cu²⁺-dominated crystals (0.80–1.47 apfu Cu) with sectors/zones enriched by Zn²⁺ (up to 1.03) or Co²⁺ (up to 0.84) or Ni²⁺ (up to 0.50 apfu) (Fig. 3). In any case, Cu²⁺ is significantly prevailing in all zones (Fig. 4). The results of electron microprobe analyses of this aggregate (11 point analyses) are summarized in Tab. 1.

Tab. 1 Chemical composition of babánekite from Jáchymov

Constituent	Mean of 11 spots (wt. %)	Range	SD
CoO	8.89	6.48–10.65	0.04
NiO	4.06	1.20–6.26	0.09
CuO	15.31	10.72–20.01	3.47
ZnO	10.87	8.70–14.51	2.20
P ₂ O ₅	0.16	0.10–0.24	0.04
As ₂ O ₅	39.79	39.01–40.58	0.57
SO ₃	0.13	0.00–0.37	0.11
H ₂ O*	24.78		
Total	103.99		

* H₂O content was calculated from stoichiometry (8 H₂O) derived from the crystal-structure study
SD – standard deviation

The empirical formula of babánekite (calculated on the basis of 16 O apfu) is (Cu_{1.12}Zn_{0.78}Co_{0.69}Ni_{0.32})_{Σ2.91} [(AsO₄)_{2.01}(PO₄)_{0.01}(SO₄)_{0.01}]_{Σ2.03} · 8H₂O. This formula shows a certain Me²⁺-deficiency and thus it is not fully charge balanced (having 0.26 of negative charge excess). The possible presence of other cations (Ca, Pb, Fe, Mg, Mn...) in concentrations under detection limits of the measurement is then the most suitable explanation. The results of crystal structure study do not indicate a presence of OH groups in this mineral. The ideal end-member formula of babánekite is Cu₃(AsO₄)₂ · 8H₂O, which requires CuO 38.95, As₂O₅ 37.52 and H₂O 23.53, total 100.00 wt. %.

Subsequently, we also determined the chemical composition of other crystal aggregates from the holotype sample and also from four additional samples of babánekite from the same locality (Geister vein, Jáchy-

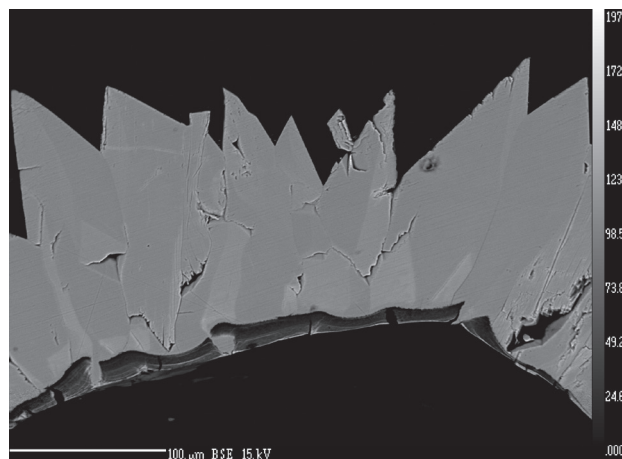


Fig. 3 Zonation of babánekite crystals mainly due to variable Zn contents. Back-scattered electron (BSE) image (Cameca SX100).

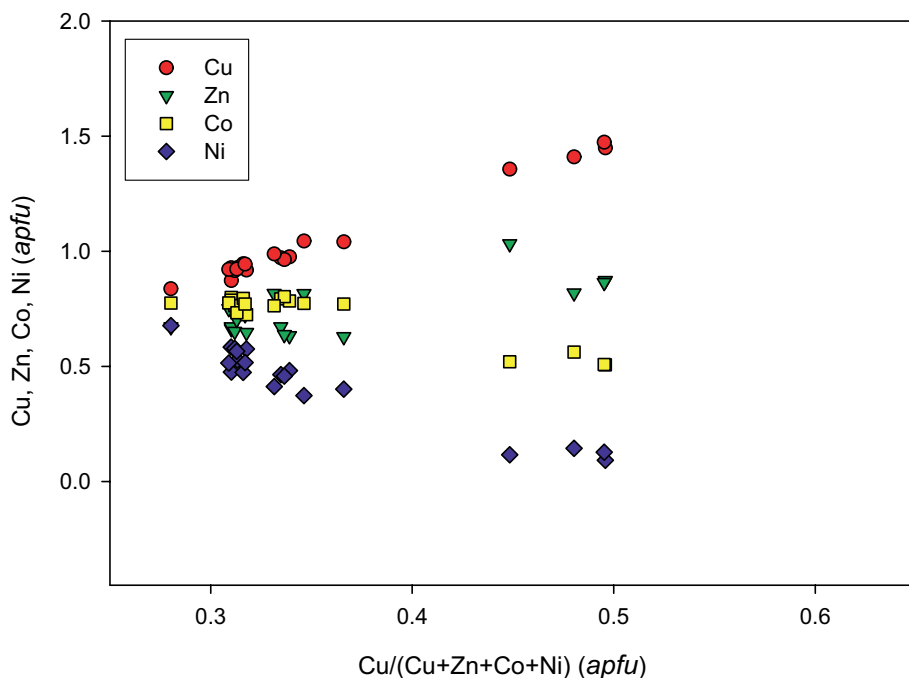


Fig. 4 Plot of $\text{Cu}/(\text{Cu} + \text{Zn} + \text{Co} + \text{Ni})$ vs. individual contents of Cu, Zn, Co and Ni (all in *apfu*) for the holotype sample of babánekite from Jáchymov.

mov), preserved mostly in collections of private mineral collectors. Additional analyses of crystals removed later from the holotype specimen resemble original analyses used for the mineral proposal for CNMNC of IMA (Fig. 4). Apparently, the Cu^{2+} content correlates negatively both with Ni^{2+} and Co^{2+} , but not with Zn^{2+} . Other samples of babánekite from Geister vein investigated by EMPA differ significantly particularly in the Zn^{2+} contents (Fig. 5), which do not exceed 0.14 *Zn apfu*, along with the even more dominating Cu^{2+} than for holotype speci-

men, reaching up to 1.78 *Cu pfu*. The Cu^{2+} contents again correlate negatively with Ni^{2+} (or $\text{Ni} + \text{Co}$).

5. X-ray crystallography and crystal structure

The X-ray powder diffraction pattern of babánekite (Tab. 2) was obtained from a hand-picked sample utilizing a Bruker D8 Advance diffractometer (National Museum, Prague) with a solid-state 1D LynxEye detector using CuK_α radiation (40 kV, 40 mA).

For minimizing the background of the scan, the powder sample was placed onto the surface of a flat silicon wafer from acetone suspension. Powder data were collected in the Bragg–Brentano geometry covering the range $5\text{--}70^\circ 2\theta$, with the step size of 0.01° and counting time of 30 s per step (total time of experiment about three days). Positions and intensities of diffraction maxima were found and refined using the Pear-

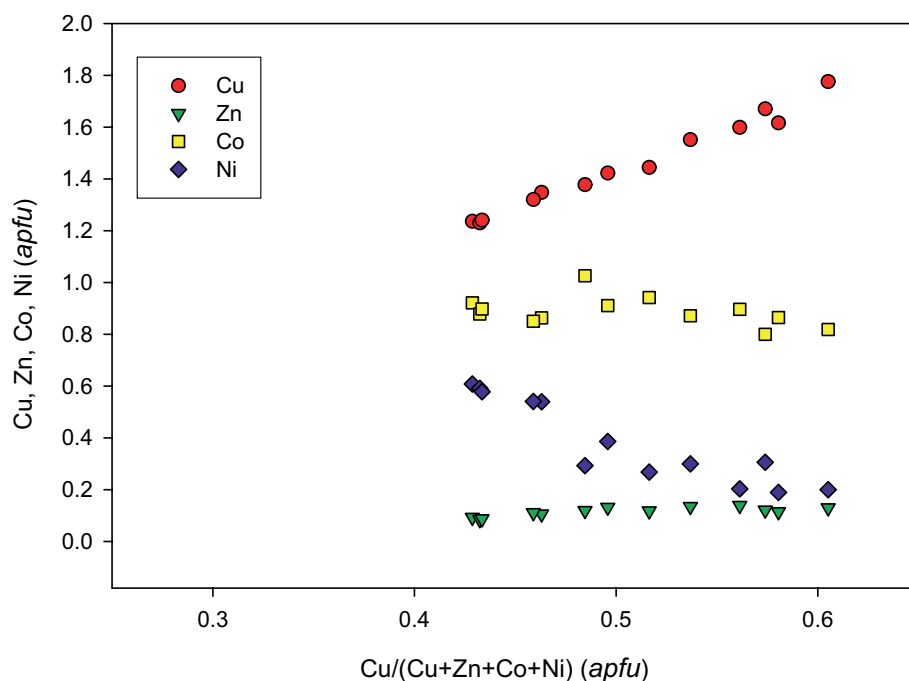


Fig. 5 Plot of $\text{Cu}/(\text{Cu} + \text{Zn} + \text{Co} + \text{Ni})$ vs. individual contents of Cu, Zn, Co and Ni (all in *apfu*) for other studied samples of babánekite from the Geister vein, Jáchymov.

sonVII profile-shape function of the ZDS program package (Ondruš 1995), and the unit-cell parameters were refined by the least-squares algorithm implemented by Burnham (1962).

Refined unit-cell parameters are $a = 10.1850(6)$, $b = 13.4852(6)$, $c = 4.7484(3)$ Å, $\beta = 105.316(5)^\circ$, with $V = 629.01(6)$ Å³. Some observed differences (in the range 0.001–0.02 Å) between parameters obtained from powder and single-crystal data are probably caused by the slightly different chemical composition of the samples (i.e., zonation; see Figs 3 and 4) by these studies.

A $0.097 \times 0.037 \times 0.034$ mm prismatic crystal of babánekite was selected for single-crystal X-ray diffraction experiment using an Oxford diffraction Gemini single-crystal four-circle diffractometer. Graphite-monochromatized MoK_α radiation ($\lambda = 0.71073$ Å) from a conventional sealed X-ray tube was collimated with a fiber-optics Mo-Enhance collimator and detected with an Atlas CCD detector. The unit cell of babánekite was refined from 3645 reflections by the least-squares algorithm of the CrysAlis-Pro package giving a monoclinic cell with: $a = 10.1742(2)$, $b = 13.5104(3)$, $c = 4.7489(1)$ Å, $\beta = 105.416(2)^\circ$ and $V = 629.29(3)$ Å³, $Z = 2$. From the total of 5722 measured reflections, 1005 were unique, and 864 were classified as unique observed with the criterion [$I_{\text{obs}} > 3\sigma(I)$]. An empirical (multi-scan) correction for absorption was applied. The summary of data collection, crystallographic data and refinement are listed in Tab. 3.

The crystal structure of babánekite was refined from the single-crystal X-ray data using the known structure model for erythrite (Wildner et al. 1996) and the full-matrix least-squares algorithm of the Jana2006 software (Petříček et al. 2014) based on F^2 . All non-H atoms were refined anisotropically. In the last cycles of the refinement, the occupancies of the Cu sites were also refined. The refined site occupancy of the Cu sites, giving a sum of 2.92 atoms per unit cell, suggests that Cu is mixed at the sites with lighter atoms such as Co or Ni and also heavier, as Zn, as well.

Tab. 2 Powder X-ray diffraction data for babánekite

I_{rel}	d_{obs}	d_{calc}	I_{calc}	h	k	ℓ	I_{rel}	d_{obs}	d_{calc}	I_{calc}	h	k	ℓ
11	7.936	7.940	37	1	1	0	3	2.1965	2.1966	4	1	5	1
100	6.743	6.743	100	0	2	0	3	2.1965	2.1969	6	2	2	-2
2	4.909	4.912	10	2	0	0	5	2.0819	2.0818	12	3	5	0
1	4.581	4.580	5	0	0	1	1	2.0437	2.0437	1	2	6	0
3	4.394	4.394	15	1	1	-1	1	2.0171	2.0177	2	0	6	1
4	4.087	4.087	5	1	3	0	1	1.9822	1.9822	3	4	4	-1
1	3.969	3.970	2	2	2	0	1	1.9503	1.9499	5	3	3	-2
2	3.903	3.903	14	2	0	-1	1	1.9443	1.9441	4	5	1	0
1	3.644	3.645	8	1	1	1	2	1.9105	1.9102	11	1	3	2
4	3.370	3.371	2	0	4	0	1	1.8331	1.8333	3	5	3	-1
14	3.231	3.231	41	1	3	-1	<1	1.7942	1.7944	1	2	6	1
2	3.181	3.182	8	3	1	0	1	1.7854	1.7856	2	3	5	1
5	2.999	2.999	19	3	1	-1	2	1.7795	1.7796	3	1	7	-1
5	2.980	2.980	26	2	0	1	16	1.6862	1.6856	8	0	8	0
4	2.779	2.780	7	2	4	0	3	1.6618	1.6620	10	1	5	2
5	2.725	2.726	24	2	2	1	4	1.6107	1.6106	4	5	5	-1
11	2.715	2.715	19	0	4	1	1	1.5909	1.5910	3	6	2	0
5	2.646	2.647	17	3	3	0	1	1.5524	1.5524	4	5	3	1
1	2.601	2.601	1	1	5	0	<1	1.5376	1.5378	4	6	0	-2
3	2.551	2.551	6	2	4	-1	1	1.5105	1.5105	4	6	4	-1
3	2.4510	2.4506	17	4	0	-1	2	1.4900	1.4899	5	4	0	2
10	2.3331	2.3327	17	1	5	-1	2	1.4672	1.4672	5	2	8	1
1	2.3040	2.3032	7	4	2	-1	1	1.4103	1.4101	3	5	5	1
2	2.2474	2.2475	1	0	6	0	2	1.3888	1.3888	3	4	8	-1
1	2.2325	2.2327	2	2	4	1							

d values quoted in Å

Tab. 3 Crystallographic data and refinement details for babánekite

Structural formula	$\text{Cu}_3(\text{AsO}_4)_2(\text{H}_2\text{O})_8$
a, b, c [Å]	10.1729(3), 13.5088(4), 4.7496(1)
β [°]	105.399(2)
V [Å ³]	629.28(3)
Z	2
D_{calc} [g/cm ³]	3.206
Space group	$C2/m$
Temperature	298 K
Detector; wavelength	Atlas CCD; MoK_α (0.71073 Å)
Crystal dimensions [mm]	$0.097 \times 0.037 \times 0.034$
Collection mode	ω rotational scans
Limiting θ angles [°]	$3.02\text{--}31.13^\circ$
Limiting Miller indices	$-13 < h < 14, -18 < k < 19, -6 < l < 6$
No. of reflections	5722
No. of unique reflections	1005
No. of observed reflections (criterion)	864 [$I_{\text{obs}} > 3\sigma(I)$]
R_{int} , coverage	0.033, 95%
Absorption correction (mm ⁻¹), method	10.20, multi-scan
Transmission (min/max)	0.514/0.730
F_{000}	589
Refinement by Jana2006 on F^2	
Parameters refined, constraints, restraints	69, 4, 6
R_1, wR_2 (obs)	0.0218, 0.0508
R_1, wR_2 (all)	0.0277, 0.0527
GOF obs/all	1.47/1.40
$\Delta\rho_{\text{min}}, \Delta\rho_{\text{max}}$ (eÅ ⁻³)	-0.70, 0.65
Weighting scheme, details	$\sigma, w = 1/(\sigma^2(I) + 0.0004I^2)$

Tab. 4 Atom positions, occupancies, and displacement parameters (in Å²) for babánekite

	<i>x/a</i>	<i>y/b</i>	<i>z/c</i>	<i>U_{eq}</i>	<i>U¹¹</i>	<i>U²²</i>	<i>U³³</i>	<i>U¹²</i>	<i>U¹³</i>	<i>U²³</i>	<i>BV</i>
As	0.18474(4)	0	0.13164(7)	0.00955(11)	0.00547(17)	0.01346(19)	0.00880(17)	0	0.00029(12)	0	4.96(1)
Cu1 [#]	0.5	0	0.5	0.01041(19)	0.0053(3)	0.0152(3)	0.0092(3)	0	-0.0006(2)	0	1.87(1)
Cu2 [#]	0	0.11392(3)	0.5	0.01112(14)	0.0089(2)	0.0129(2)	0.0116(2)	0	0.00284(16)	0	1.90(1)
O1	0.3535(3)	0	0.1357(5)	0.0149(7)	0.0068(11)	0.0257(14)	0.0114(11)	0	0.0009(9)	0	1.86(1)
O2	0.15612(17)	0.10610(12)	0.2920(4)	0.0129(5)	0.0113(8)	0.0129(8)	0.0144(8)	0.0003(6)	0.0032(7)	-0.0027(7)	1.80(1)
O3	0.0978(2)	0	-0.2216(5)	0.0136(7)	0.0111(12)	0.0164(12)	0.0107(11)	0	-0.0018(9)	0	1.91(1)
O4	0.4019(2)	0.11629(16)	0.6978(4)	0.0251(7)	0.0113(90)	0.0418(13)	0.0192(10)	-0.0038(9)	-0.0014(7)	0.0077(9)	2.13(6)
O5	0.1012(2)	0.22550(14)	0.7839(4)	0.0215(6)	0.0236(10)	0.0209(10)	0.0203(9)	-0.0002(8)	0.0062(8)	-0.0008(8)	2.12(6)
H5a	0.044(3)	0.2831(15)	0.778(5)	0.0258*							1.02(5)
H4a	0.381(3)	0.097(2)	0.879(4)	0.0302*							1.01(5)
H5b	0.128(3)	0.2060(19)	0.990(3)	0.0258*							1.00(3)
H4b	0.3132(17)	0.126(2)	0.560(5)	0.0302*							1.05(3)

U_{eq} is defined as a third of the trace of the orthogonalized *U^{ij}* tensor

* – refined with isotropic displacement parameters. [#] – refined occupancies for Cu1 and Cu2 sites are 0.973(3) and 0.974(2), respectively

BV – bond-valence sums calculated using Dist option in Jana2006; bond-valence parameters taken from Gagné and Hawthorne (2015)

Regarding the fact that Co, Ni, Cu and Zn (confirmed by the microprobe analysis) have very similar scattering curves, only Cu was used in the later refinement as a proxy. Both the constrained and unconstrained refinement, which took into account the presence of all atoms mentioned above (and also allowing some degree of the ordering) in the cationic octahedral sites, did not lead either to better or more reasonable results. The constrained refinement fixing the composition, expressed by the empirical formula obtained from the electron-microprobe study, converged to the much worse indices of agreement ($R = 0.0290$, $wR = 4.01$, $GOF = 2.41$ for 864 unique observed reflections) compared to the fit using just the Cu at the sites. Positions of the hydrogen atoms were localized from difference-Fourier maps. The distance between the H atom and the corresponding donor O atom was constrained during the refinement (with a soft constraint of 0.90 Å with a certain weight, that produced *D–H* lengths in the range 0.74(3)–0.86(3) Å). The displacement parameters of the H atoms were set to be $1.2 \times U_{iso}$ of the corresponding donor O atom. The final cycles converged with $R = 0.0218$, $wR = 0.0508$ and $GOF = 1.47$ for 864 observed unique reflections. The correctness of the refined structure model was confirmed by an independent structure solution employing the charge-flipping method (Palatinus and Chapuis 2007). This independent solution obtained differs from the model based on Wildner et al. (1996) only in the order of the estimated errors. Atom positions, displacement parameters and results of the bond valence analysis are listed in Tabs 4 to 5. Some of the geometrical characteristics (also given in Tab. 4) were calculated using the program Vesta (Momma and Izumi 2008). The bond-valence analysis was carried out following the procedure of Brown (2002). The CIF file, also containing a block with the reflections, is deposited at the Journal's web page www.jgeosci.org.

5.1. Crystal structure

The refined structure of babánekite shows characteristic features of the vivianite-type minerals and synthetic compounds, characterized by the general formula $Me^{2+}_3(XO_4)_2 \cdot 8H_2O$ ($Me = Mg, Fe, Co, Ni, Zn; X = P, As$) (Mori and Ito 1950; Cesbron et al. 1977; Hill 1979; Fejdi et al. 1980; Giuseppetti and Tadini 1982; Dormann et al. 1982; Takagi et al. 1986; Riou et al. 1989; Bartl 1989; Wildner et al. 1996; Yakubovich et al. 2001; Capitelli et al. 2007, 2012; Assani et al. 2010; Yoshiasa et al. 2016; Antao and Dhaliwal 2017). Vivianite-type structures consist of $MeO_2(H_2O)_4$ octahedra and dimers of $Me_2O_6(H_2O)_4$ octahedra that are linked via XO_4 tetrahedra (where $X = P, As$ and probably can be also V) and hydrogen bonds to form complex layers parallel to (010). Adjacent layers are linked by hydrogen bonds only (Fig. 6).

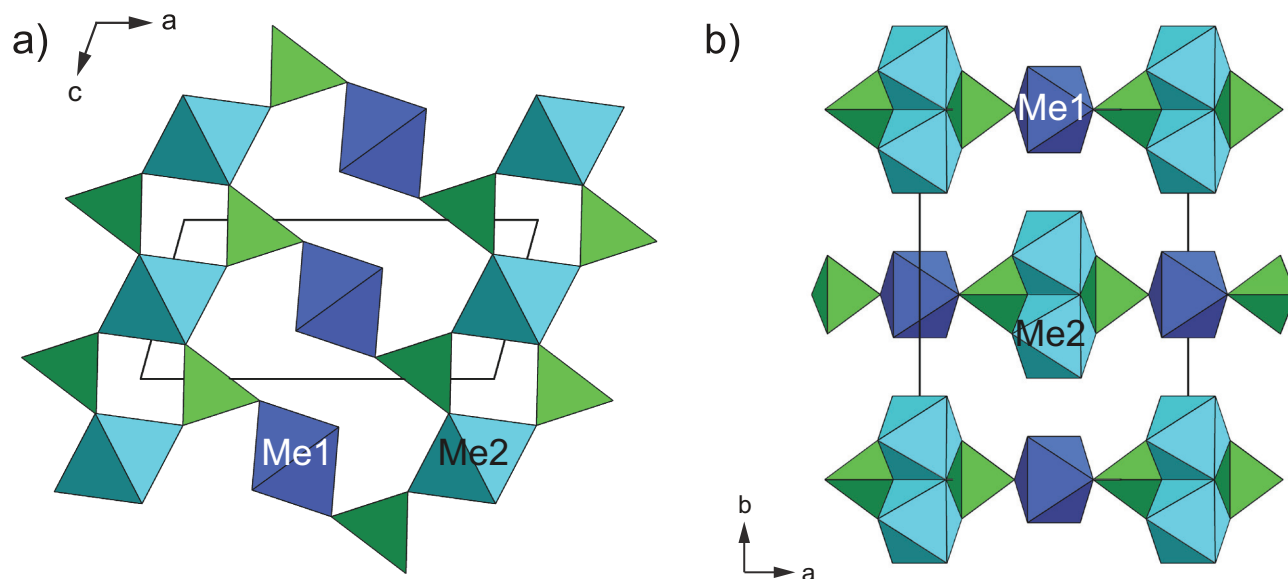


Fig. 6 The crystal structure of babánekite. **a** – Part of infinite sheet viewed along [010] consisting of vertex-sharing *Me1* and *Me2* distorted octahedra and AsO_4 tetrahedra (in green). **b** – The stacking of infinite sheets perpendicular to [010]; adjacent sheets are linked by H-bonds (H atoms and bonds omitted for clarity). Unit-cell edges are outlined in black solid lines.

The Me^{2+} cations, in babánekite dominantly Cu^{2+} , occupy two symmetrically distinct sites with the site-symmetries $2/m$ and 2 (Wyckoff sites $2d$ and $4h$ of the space group $C2/m$). The Cu1-centred (*Me1*) polyhedron forms a monomeric entity, which is linked *via* arsenate group to the dimer of Cu2-centred (*Me2*) polyhedra. The Cu1-centred polyhedron is considerably more irregular (in the terms of “effective coordination number”, $\text{ECoN} = 5.07$, bond-length distortion index $\Delta = 0.05$) than the Cu2-centred polyhedron ($\text{ECoN} = 6.00$, $\Delta = 0.003$). This is a general feature also observed for other members of the vivianite-group of minerals (e.g., Giuseppetti and Tadini 1982, Yakubovich et al. 2001). It is probably caused due to less-firmly bonded ligands to the *Me1* site: only two atoms are O^{2-} cations, the rest of the ligands represent the H_2O groups. The bonding in dimers of the *Me2*-centered polyhedra is much more rigid: only four of the ten ligands belong to the H_2O groups.

6. Discussion: the vivianite group of minerals – nomenclature issues

The present description of babánekite enlarges the family of As-dominant

members in the vivianite group to seven. In Tab. 6, a general comparison of these members is given based on

Tab. 5 Selected interatomic distances (in Å) in the crystal structure of babánekite

Cu1–O1	1.961(2)	Cu2–O2	2.084(2)
Cu1–O1 ^{iv}	1.961(2)	Cu2–O2 ^{vi}	2.084(2)
Cu1–O4	2.201(2)	Cu2–O3 ^{vii}	2.099(1)
Cu1–O4 ^{iv}	2.201(2)	Cu2–O3 ^{viii}	2.099(1)
Cu1–O4 ^v	2.201(2)	Cu2–O5	2.101(2)
Cu1–O4 ⁱ	2.201(2)	Cu2–O5 ^{vi}	2.101(2)
<Cu1–O>	2.122	<Cu1–O>	2.094
V_p	12.59	V_p	12.19
σ^2	7.18	σ^2	6.95
Δ	0.051	Δ	0.003
ECoN	5.067	ECoN	5.998
	As–O1	1.712(3)	
	As–O2	1.684(2)	
	As–O2 ⁱ	1.684(2)	
	As–O3	1.676(2)	
	<As1–O>	1.690	
H-bonds	$D\cdots H$	$D-A$	< $D-H\cdots A$ >
	O4–H4a \cdots O1 ^{vii}	2.753(3)	151(3)
	O4–H4b \cdots O2	2.725(2)	164(3)
	O5–H5a \cdots O4 ^{xv}	2.899(3)	167.7(18)
	O5–H5b \cdots O2 ^{vii}	2.833(3)	151(2)

Δ , Bond-length distortion after Brown and Shannon (1973); σ^2 , bond-angle distortion after Robinson et al. (1971); ECoN, effective coordination number after Hoppe (1979); V_p , polyhedral volume (in Å³)

Calculations by Vesta (Momma and Izumi 2008)

Symmetry codes: (i) $x, -y, z$; (ii) $x, y, z-1$; (iii) $x, 2y, z-1$; (iv) $-x+1, y, -z+1$; (v) $-x+1, -y, -z+1$; (vi) $-x, y, -z+1$; (vii) $x, y, z+1$; (viii) $-x, y, -z$; (ix) $-x, y, -z-1$; (x) $x+1/2, -y+1/2, z$; (xi) $-x+1/2, -y+1/2, -z+1$; (xii) $-x+1/2, -y+1/2, -z+2$; (xiii) $-x+1, y, -z+2$; (xiv) $-x, y, -z+2$; (xv) $x-1/2, -y+1/2, z$

Tab. 6 Comparison of crystallographic and optical data for As-dominant members of vivianite group of minerals

mineral	babānekite	erythrite	annabergite	höresite	manganohöresite	köttigite	parasymplesite
type locality	Jáchymov, CZ	Schneeberg, Germany	Annaberg, Germany	Banat, Hungary	Långban, Sweden	Schneeberg, Germany	Kiura Mine, Japan
reference	this paper	Wildner et al. (1996)	Wildner et al. (1996)	Palache et al. (1951)	Gabrielson (1954)	Dana (1850)	Ito et al. (1954)
Me^{2+} (ideal.)	Cu	Co	Ni	Mg	(Mn^{2+}, Mg) ₃	Zn	Fe^{2+}
Me^{2+} (meas.)	($Cu_{1.12}Zn_{0.78}Co_{0.69}Ni_{0.32}$) _{Σ2.91}	($Co_{2.01}Fe_{0.74}Ni_{1.02}$) _{Σ3.00}	($Ni_{2.48}Mg_{0.50}Fe_{0.02}$) _{Σ3.00}	–	($Mn_{1.65}Mg_{1.32}$) _{Σ2.97}	($Zn_{2.44}Co_{0.42}Ni_{1.14}$) _{Σ3.00} ²⁾	–
space group	<i>C2/m</i>	<i>C2/m</i>	<i>C2/m</i>	<i>C2/m</i>	<i>P2₁/c</i>	<i>C2/m</i>	<i>C2/m</i>
<i>a</i>	10.1729(3)	10.251(3)	10.179(2)	10.137(2) ¹⁾	10.38	10.241(3)	10.276(4) ²⁾
<i>b</i>	13.5088(4)	13.447(4)	13.309(3)	13.445(2)	28.09	13.405(3)	13.480(5)
<i>c</i>	4.7496(1)	4.764(1)	4.725(1)	4.754(1)	4.774	4.757(2)	4.771(2)
β	105.399(2)	104.98(1)	105.00(1)	101.73(2)	105.7	105.21(2)	105.02(5)
<i>V</i>	629.28(3)	634.4	618.2	634.17	1340.04	630.2(2)	–
<i>Z</i>	2	2	2	2	4	2	2
strongest lines	6.743/100	6.65/100 [#]	6.58/100 [#]	8.55/10 [#]	8.19 (80)	7.95 (37) ³⁾	7.91 (70)
in XRD	7.936/11	7.89/6	7.82/25	5.11/5	7.01 (100)	6.70 (100)	6.68 (100)
powder pattern	3.23/14	3.34/8	4.33/20	4.30/6	3.25 (60)	3.22 (39)	4.41 (40)
	2.715/11	3.22/12	3.18/26	3.59/9	3.09 (60)	2.99 (28)	3.91 (30)
	1.686/16	2.70/8	2.98/30	3.45/4	3.02 (70)	2.73 (26)	3.24 (50)
optical data		Biaxial (+) [#]	Biaxial (+) [#]	Biaxial (+) [#]	Biaxial (+) [#]	Biaxial (+)	Biaxial (+) [#]
		$\alpha = 1.622-1.629$	$\alpha = 1.622$	$\alpha = 1.563$	$\alpha = 1.579$	$\alpha = 1.619-1.622$	$\alpha = 1.620-1.628$
	$N_{calc} = 1.662$	$\beta = 1.660-1.663$	$\beta = 1.658$	$\beta = 1.571$	$\beta = 1.589$	$\beta = 1.638-1.645$	$\beta = 1.648-1.660$
		$\gamma = 1.681-1.701$	$\gamma = 1.687$	$\gamma = 1.596$	$\gamma = 1.609$	$\gamma = 1.671-1.681$	$\gamma = 1.685-1.705$
		$2V = \text{up to } 90^\circ$	$2V = 84^\circ$	$2V = 65^\circ$	$2V = 65-70^\circ$	$2V = 74-85^\circ$	$2V = 86^\circ$

¹⁾ unit-cell parameters from Jambor and Dutrizac (1995)

²⁾ from Sturman (1976)

³⁾ calculated, Hill (1979)

[#] in Anthony et al. (1990)

the available data from the literature and the current study. For all these minerals, a broad cationic isomorphy is observed (Jambor and Dutrizac 1995, and others). The Me^{2+} cations occupy two symmetrically distinct sites in their crystal structure. On the one hand, partial ordering of cations at the $Me1/Me2$ sites was documented by the structure studies for some of the Mg-containing members of the vivianite group (Giuseppetti and Tadini 1982; Rojo et al. 1996; Wildner et al. 1996; Yakubovich et al. 2001). On the other hand, for the members containing Co, Zn, Ni or Cu (Hill 1979; Wildner et al. 1996; Capitelli et al. 2007; this study) no such a preference for a certain site was found. Theoretically this can be an artefact, caused by the problematic refinement due to the very similar scattering curves of Co, Ni, Cu and Zn, and therefore impossibility to discern them successfully using X-rays at the same site. Based on current study, the preference of Cu^{2+} for the Cu1 site, which is far more irregular than the Cu2 site (stronger effect of the Jahn-Teller distortion), even though highly probable, also cannot be reliably documented from current X-ray diffraction data. However, the possible preferential entering of the cations was discussed in the study of Antao and Dhaliwal (2017) who documented dominance of the heavier cations (Co, Ni, Zn) at $Me1$ site and lighter cations (Fe^{2+}) at the $Me2$ site. Nowadays, the classification of the particular mineral member of the vivianite group is based on the dominant cation at both sites (i.e. $\Sigma(Me1+Me2)$: Sturman and Mandarino 1976; Giuseppetti and Tadini 1982; Yakubovich et al. 2001).

There have been several descriptions of the increased Cu^{2+} contents in mineral members of the vivianite group. A significant Cu^{2+} content in erythrite from Silberberg, near Rattenberg, Tyrol (Austria) was reported by Putz et al. (2012) and from Leogang, Salzburg by Auer (2017). Babānekite has been also found in Germany in the Clara mine (Blaß and Draxler 2015). Recently, we studied other samples from the Geister

vein, Jáchymov (Czech Republic) and also determined Cu-enriched köttigite (up to 0.99 Cu *pfu*), erythrite (up to 0.98 Cu *pfu*) and annabergite (up to 0.41 Cu *pfu*). The research on these mineral varieties is still ongoing and the results will be the subject of a forthcoming paper.

Acknowledgements. We thank Karla Fejfarová (BIOCEV, ASCR, v.v.i. and Charles University in Prague) and Michal Dušek (Institute of Physics, ASCR, v.v.i.) for help with analytical works. The manuscript benefited from thorough reviews of an anonymous reviewer and Uwe Kolitsch. Handling editor František Laufek is acknowledged for taking care of the manuscript. This research was supported by the project of GAČR No. 17-09161S to J.P. and J.S.

Electronic supplementary material. Supplementary crystallographic data for this paper (including the selected interatomic distances, results of bond-valence analysis and a CIF file) are available online at the Journal website (<http://dx.doi.org/10.3190/jgeosci.248>).

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