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Locked in the icehouse: evolution of an endemic *Epimeria* (Amphipoda, Crustacea) species flock on the Antarctic shelf

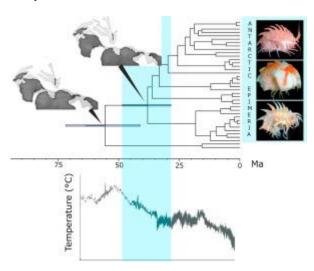
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Abstract:

The Antarctic shelf's marine biodiversity has been greatly influenced by the climatic and glacial history of the region. Extreme temperature changes led to the extinction of some lineages, while others adapted and flourished. The amphipod genus Epimeria is an example of the latter, being particularly diverse in the Antarctic region. By reconstructing a time-calibrated phylogeny based on mitochondrial (COI) and nuclear (28S and H3) markers and including Epimeria species from all oceans, this study provides a temporal and geographical framework for the evolution of Antarctic Epimeria. The monophyly of this genus is not supported by Bayesian Inference, as Antarctic and non-Antarctic Epimeria form two distinct wellsupported clades, with Antarctic Epimeria being a sister clade to two stilipedid species. The monophyly of Antarctic Epimeria suggests that this clade evolved in isolation since its origin. While the precise timing of this origin remains unclear, it is inferred that the Antarctic lineage arose from a late Gondwanan ancestor and hence did not colonize the Antarctic region after the continent broke apart from the other fragments of Gondwanaland. The initial diversification of the clade occurred 38.04 Ma (95% HPD [48.46 Ma; 28.36 Ma]) in a cooling environment. Adaptation to cold waters, along with the extinction of coldintolerant taxa and resulting ecological opportunities, likely led to the successful diversification of Epimeria on the Antarctic shelf. However, there was neither evidence of a rapid lineage diversification early in the clade's history, nor of any shifts in diversification rates induced by glacial cycles. This suggests that a high turnover rate on the repeatedly scoured Antarctic shelf could have masked potential signals of diversification bursts.

Graphical abstract



Highlights

▶ A molecular phylogeny of Epimeria amphipods from every world's oceans is provided. ▶ Antarctic and non-Antarctic Epimeria form 2 distinct clades. ▶ The Antarctic clade arose from a Late Gondwanan ancestor. ▶ The diversification of the Antarctic Epimeria clade appears related to cold-waters.

Keywords: Amphipoda, Southern ocean, Historical biogeography, Phylogeny, Divergence times, Diversification

1. Introduction

The Southern Ocean is traditionally viewed as an isolated ecosystem where the marine fauna has essentially evolved *in situ* (Clarke and Crame, 1989; Dell, 1972; Knox and Lowry, 1977; Lipps and Hickman, 1982). The eastward flowing Antarctic Circumpolar Current (ACC), encircling Antarctica since the Oligocene, is the most powerful sea current on Earth. The Antarctic Polar Front (APF), one of the ACC's jets, marks a sharp change in surface water temperatures and impedes north-south exchange of water. Therefore, it appears to be an important barrier to dispersal (Angel, 1997; Clarke et al., 2005; Dell, 1972). The high degree of species-level endemism of the Antarctic marine fauna is a signature of this long history in isolation (Arntz et al., 1997; Clarke and Crame, 1997; Clarke and Johnston, 2003). Some Antarctic lineages are descendants of Gondwanan ancestors, which arose by vicariance when the supercontinent progressively broke apart (e.g. Brandt, 1992; Clarke and Johnston, 1996; Waters et al., 2007; Williams et al., 2003). While the APF has been

ACC = Antarctic circumpolar current; APF = Antarctic Polar Front; MMCT = Mid-Miocene climate transition; COI = cytochrome b oxidase subunit I; H3 = histone 3; BI = Bayesian Inference; ILD = Incongruence Length Difference test; ML = Maximum Likelihood; BIC = Bayesian Information criterion; PP = posterior probability; BV = bootstrap value; MOTU = molecular taxonomic unit; MCMC = Markov chain Monte-Carlo; UCLD = uncorrelated lognormal distribution; UCED = uncorrelated exponential distribution; BF = Bayes Factor; SD = standard deviation; LTT = lineage through time plot; PB = Pure Birth; BD = Birth-Death; AIC = Akaike Information criterion; DDL = density-dependent linear; DDX = density-dependent exponential; BDS = Birth-Death Shift; dAIC = difference of AIC; wAIC = Akaike weight; MLE = marginal likelihood estimate; MRCA = most recent common ancestor; BAMM = Bayesian analysis of macroevolutionary mixtures; MAP = maximum posterior probability; HPD = higher posterior density; EECO = early Eocene climate optimum; EOCT = Ecocene-Oligocene climate transition

shown to be a permeable barrier for a variety of pelagic and deep-sea organisms (Antezana, 1999; Bargelloni et al., 2000; Brandt et al., 2007; Hodgson et al., 1997; Page and Linse, 2002; Pawlowski et al., 2007; Thatje and Fuentes, 2003), the Drake Passage is too deep to allow the dispersal of stenobathic shelf organisms along the benthos (e.g. Hunter and Halanych, 2008; Shaw et al., 2004). However, some benthic taxa lacking a pelagic stage may be able to cross the front by drifting on macroalgae or floating debris (Barber et al., 1959; Coombs and Landis, 1966; Fraser et al., 2009; Helmuth et al., 1994; Waters, 2008). Historical cases of polar emergences — colonizations of the Antarctic shelf from the depths —or submergences — dispersals from the shelf towards the depths — were also inferred for some strictly benthic organisms (Held, 2000; Strugnell et al., 2008). Such movements may be facilitated by the formation of cold and dense water near the continent which sinks northward at a layer below 4000 m depth to become the "Antarctic Bottom Water", thereby forming an isothermal water column between the shelf and the deep-sea (Knox and Lowry, 1977). The Southern Ocean fauna therefore contains a mixture of lineages with different histories, some have evolved *in situ* since before Antarctica existed, others are more recent colonizers (Clarke and Crame, 1989).

All along its geological history, the Antarctic region has faced profound climatic changes, which deeply impacted Southern Ocean marine biodiversity (Clarke and Crame, 1997). The fossil record indicates major changes in the nature of the marine fauna along the cooling trend, which began across the early to middle Eocene boundary (ca. 48-49 Ma). While many taxa went extinct, the availability of previously occupied benthic niches provided ecological opportunities for surviving lineages (Aronson et al., 2007; Brandt, 1999). Ecological opportunity has been inferred to drive initial rapid lineage diversification (Condamine et al., 2013; Losos and Mahler, 2010; Yoder et al., 2010). As lineages rapidly fill unoccupied niches and available habitats run out, the rate of diversification would consequently decrease (Freckleton and Harvey, 2006; Rabosky and Lovette, 2008; Schluter, 2000; Walker and Valentine, 1984).

Alternatively, the simultaneous formation of multiple geographical barriers can be responsible for a burst of allopatric speciation (Rundell and Price, 2009). Hence, by providing ecological opportunities or causing allopatric speciation, several main events in Antarctica's climatic history could have impacted species' origination rates. For instance, the gradual extinction of decapods, possibly beginning with the cooling trend which started in the early Eocene (Thatje et al., 2005), likely triggered the radiation of the peracarids (Aronson et al., 2007). The Eocene-Oligocene boundary (34 Ma) was marked by a sudden drop in temperatures, the first continent-wide glaciations (Lear et al., 2000, 2004), the opening of the Tasmanian and Drake passages (Barker, 2001; De Broyer and Jażdżewska, 2014; Exon et al., 2000; Stickley et al., 2004) and the formation of the ACC (Lawver and Gahagan, 2003; Lyle et al., 2007; Pfuhl and McCave, 2005). The resulting thermal and geographical isolation of the Antarctic region is presumed to have promoted vicariant speciations in Southern

Ocean taxa. Since the Mid-Miocene Climate Transition (MMCT; ~14 Ma), the ice-shelf grounding line periodically extended to the outer shelf break, at least in some places during glacial maxima. These glacial cycles, which intensified in the Late Pliocene-Pleistocene period (Lewis et al., 2008; McKay et al., 2012; Pollard and DeConto, 2009) were inferred to act as a "diversity pump" on the Southern Ocean continental shelf (Clarke and Crame, 1989, 2010; Clarke et al., 1992). The isolation of populations in ice-free refugia during glacial advances would have resulted in allopatric speciations of less dispersive organisms (Thatje et al., 2005).

The Antarctic component of the amphipod genus *Epimeria* (Costa, 1851) was put forward as an example of an Antarctic species flock, i.e. a highly diverse clade of species that originated and diversified in the Antarctic region (Lecointre et al., 2013). However, the assumption of monophyly was based on a previous COI phylogeny of *Epimeria*, comprising 17 Antarctic species, but only two non-Antarctic (New Zealand) species (Lörz et al., 2009). Yet, the genus is cosmopolitan, but particularly well represented in the Southern Ocean, with 26 described species out of a total of 54 worldwide (Coleman, 2007; Lörz, 2009; Lörz et al., 2007, 2009, 2011). Moreover, a recent study of COI and 28S sequence data identified 24 lineages as putative new Antarctic species, showing that the species richness of this genus on the shelf is still greatly underestimated (Verheye et al., 2016). *Epimeria* contains a mixture of regionally-restricted and (almost) circum-Antarctic species (d'Udekem d'Acoz and Verheye, *in press*; Verheye et al., 2016), but no species are found on both sides of the APF. Their mostly benthic (only two pelagic species are known) and brooding ecology, coupled with limited eurybathy, would prevent such long-distance dispersal across deep passages, on ecological timescales (d'Udekem d'Acoz and Verheye, *in press*). In the COI phylogeny of Lörz and Held (2004), comprising six Antarctic and no extralimital species, ages of 15.7 and 34.9 Ma were inferred for the last common ancestor of Antarctic *Epimeria*, using respectively a COI rate of evolution estimated for cirripeds and for alpheid shrimps, and assuming a strict molecular clock. The authors concluded that Southern Ocean *Epimeria*, coupled with a low dispersal potential, makes it a good model to study the connection between the Antarctic shelf and the other oceans, the origin of this component of the shelf benthos and its *in situ* diversification patterns.

Inferences on the historical biogeography of the *Epimeria* Antarctic shelf species are not possible without a robust phylogeny, based on an extensive sampling of both Antarctic and non-Antarctic species. Therefore, we reconstructed a phylogeny of *Epimeria* using three gene fragments (COI, 28S and H3)

and including 12 out of the 26 described Antarctic *Epimeria* species as well as 24 putative new Antarctic *Epimeria* species. The non-Antarctic material is composed of 21 *Epimeria* species collected in every other world's oceans. In order to explore the monophyly and systematics of *Epimeria*, we included representatives of other families which form a clade with *Epimeria*, viz. Acanthonotozomellidae, Dikwidae, Stilipedidae and Vicmusiidae (Verheye et al., 2015). Then, we used a relaxed molecular clock to date the phylogenetic tree and considered the geographical distribution of the specimens to address the following issues: 1. Does the Antarctic component of *Epimeria* form a single clade or does the phylogeny provide evidence of historical dispersal events in and/or out of the shelf (which would make this Antarctic component non-monophyletic)? 2. When and where did the Antarctic component originate i.e. is it a Gondwanan relict or a more recent colonizer? and 3. Does the historical diversification pattern of Antarctic *Epimeria* bear signatures of an early diversification burst, or of shifts in diversification rates which might be associated with climatic events (e.g. glacial cycles)?

2. Material and Methods

2.1. Taxon sampling

We included data of 12 out of the 26 described species of Antarctic *Epimeria*, as well as the 24 putative new Antarctic species of Verheye et al. (2016), i.e. a total of 36 out of 50 known species (72.5 %). Non-Antarctic *Epimeria* included one species from continental Norway (*Epimeria cornigera*), one from the Svalbard Archipelago (*Epimeria loricata*), one from Mexico (*Epimeria morronei*) and 16 putative new species, viz. 14 from the Melanesian subregion (Indonesia, Papua New Guinea, New Caledonia, Solomon, Vanuatu and Fiji islands), one from Taiwan and one from Mozambique. Species from other families that were shown to be related to *Epimeria* in Verheye et al. (2015) (i.e. *Acanthonotozomoides oatesi* and Acanthonotozomellidae n. gen. n. sp. [Acanthonotozomellidae], *Dikwa andresi* [Dikwidae], *Astyra abyssi, Bathypanoploea schellenbergi* and *Alexandrella dentata* [Stilipedidae], and *Acanthonotozomopsis pushkini* [Vicmusiidae]) were included as well, in order to explore the monophyly of the genus. Three iphimediid species were used as outgroup. All specimens used in this study are listed in Table 1.

Samples were collected during several expeditions of the R.V. Polarstern in the Southern Ocean: ANT-XXIV/2, ANTXXIII/8, ANT-XXVII/3 and ANT-XXIX/3 in the Drake Passage, Bransfield Strait, eastern coast of the Antarctic Peninsula and eastern Weddell Sea. Additional specimens were sampled from the Adélie

Coast during the CEAMARC and REVOLTA I, II, III expeditions and in Prydz Bay during the MD42 (SIBEX) expedition of the R.V. Marion Dufresne. One specimen was sampled during the JR144 expedition with the RRS James Clark Ross near Elephant Island. Non-Antarctic specimens were sampled during BIOPAPUA, Papua Niugini (Papua New Guinea), KARUBAR (Indonesia), EXBODI, NORFOLK2, BATHUS3 (New Caledonia), SOLOMON2 and 3 (Solomon Islands), MUSORSTOM8, SANTO (Vanuatu), FIDJI (Fiji Islands), Taiwan 2000 (Taiwan) and MAINBAZA (Mozambique) expeditions. Additional material was obtained by opportunistic collections from Mexico and Norway (including the Svalbard archipelago) (Table 1).

All specimens were preserved in 96–100% ethanol for DNA analysis. Vouchers are deposited at the Royal Belgian Institute of Natural Sciences (RBINS, Brussels, Belgium) and the Muséum National d'Histoire Naturelle (MNHN, Paris, France). For voucher collection ID numbers, see Table 1.

2.2. DNA sequencing

DNA was extracted from the pleopods and abdomen muscles using a NucleoSpin® Tissue kit (Macherey-Nagel) following the manufacturer's protocol for animal tissues. The DNA was eluted in 100 μ l of sterile distilled H₂O (RNase free) and stored at -20 °C.

Partial segments of the mitochondrial cytochrome c oxidase subunit I (COI) (~550 bp), nuclear 28S rDNA (~1400 bp) and Histone 3 (H3) (~360 bp) were amplified by PCR. Amplifications were performed in a 25 μl reaction mix, which contained 0.15 μl Taq DNA Polymerase (5 U μl⁻¹; Qiagen, Antwerpen, Belgium), 2.5 μl 10x CoralLoad PCR Buffer (Qiagen, Antwerpen, Belgium), 2.5 μl dNTPs mix (250 μM of each), 11–16 μl RNase-free water, 1.25 μl of each primer (2 μM), and 1–6 μl of DNA extract. The COI fragment was amplified using the primers Cp-COIF3 (Pilar Cabezas et al., 2013) and COI2R (Otto and Wilson, 2001). The thermal cycling used for the COI amplification followed Pilar Cabezas et al. (2013), except for the annealing temperature set at 51°C. The 28S rDNA fragment was amplified using the primers 28S-3311F (Witt et al., 2006) and 28R (Hou et al., 2007), modified as follows: 5′-GGGACTACCCGCTGAACTTAAGCAT-3′ and 5′-GTCTTTCGCCCCTATGCCCAACTG-3′. PCR amplification settings for 28S rDNA consisted of an initial denaturation for 3 min at 94 °C, followed by 40 cycles of denaturation at 94 °C for 40 s, annealing at 45 °C for 40 s, extension at 72 °C for 90 s, and a final extension at 72 °C for 10 min. The H3 fragment was amplified using the primers H₃aF and H₃aR (Colgan et al., 1998). PCR amplification settings were as in Colgan et al. (1998), with an annealing temperature of 54–60°C.

The PCR products were visualized under blue light on 1.2 % agarose gel stained with SYBR Safe (ThermoFisherScientific, Waltham, MA, USA), with a comigrating 200-bp ladder molecular-weight marker to confirm their correct amplification. Prior to sequencing, PCR products were purified using Exonuclease I (20 U μ I⁻¹) and FastAPTM Thermosensitive alkaline phosphatase (1 U μ I⁻¹) (ThermoFisher Scientific, Waltham, MA, USA), following the manufacturer's protocol. Forward and reverse strands were sequenced with fluorescent-labeled dideoxynucleotide terminators (BigDye v.3.1; Applied Biosystems, Foster City, CA, USA) following the protocol of Sanger et al. (1977) and using an automated ABI 3130xl DNA analyzer (Applied Biosystems, Foster City, CA, USA). Both fragments were sequenced using the PCR primers.

2.3. Phylogenetic analyses

Sequence chromatograms were checked, and forward and reverse sequence fragments were assembled using Codoncode Aligner v.3.7.1. (CodonCode Corporation; http://www.codoncode.com/aligner/). All sequences have been deposited in GenBank (Table 1).

28S sequences were aligned with MAFFT v.7 (Katoh and Standley, 2013) (http://mafft.cbrc.jp/alignment/server/), using the structural alignment strategy Q-INS-i under default settings. As some regions of the 28S sequences were too divergent to be confidently aligned, the software program Aliscore v.2.0. (Misof and Misof, 2009) was used to identify poorly aligned regions for removal with Alicut v.2.3, prior to further analysis. CLUSTALW was used to align the COI and H3 sequences in MEGA6 (Tamura et al., 2013). In order to prevent inclusion of pseudogenes in the analyses, amino acid translations of both fragments were checked for stop codons.

In order to evaluate the congruence between genes and reconstruction methods, preliminary phylogenetic trees were inferred using Bayesian inference (BI) on the separate datasets. The Incongruence Length Difference (ILD) test (Farris et al., 1995) was implemented using Paup*4.0b10 (Swofford, 2003). BI and Maximum Likelihood (ML) were then used to reconstruct phylogenetic relationships based on a dataset concatenated with SequenceMatrix (Vaidya et al., 2011).

The best-fit models of DNA substitution were selected using the Bayesian Information Criterion (BIC) on the concatenated dataset partitioned by gene and by codon position (for COI and H3), in PartitionFinder (Lanfear et al., 2012). This model selection procedure was performed both by assuming a single set of underlying branch lengths for the tree and independent set of branch lengths for each partition and the best scheme was selected based on the BIC value.

BI trees were reconstructed using MrBayes v.3.2 (Ronquist and Huelsenbeck, 2003) on the CIPRES portal (Miller et al., 2010). BI analysis of each alignment included two runs of 10^7 generations. Trees were sampled every 1000 generations using four Markov chains, and default heating values. Convergence was assessed by the standard deviation of split-frequencies (<0.01) and by examining the trace plots of log-likelihood scores in Tracer 1.6 (Rambaut and Drummond, 2005). The first 10 % trees were discarded as burn-in, while the remaining trees were used to construct a 50 % majority rule consensus tree and estimate the posterior probabilities (PP). Nodes with posterior probabilities (PP) \geq 0.95 were considered as significantly supported.

ML trees were estimated using GARLI v.2.0 (Zwickl, 2006). For each dataset, 10 separate ML searches were run independently from different stepwise-reconstructed trees. The best scoring tree across runs was considered for further analyses. Confidence levels of branches were estimated by 1000 bootstrap replicates. Nodes with bootstrap values (BV) \geq 70 were considered meaningful.

2.4. Estimation of divergence times

As intraspecific variation may lead to the overestimation of divergence times (Ho et al., 2008), the time-calibrated reconstruction was based on a reduced multimarker dataset. For the Antarctic *Epimeria* clade, one individual per Molecular Taxonomic Unit (MOTU) identified as a putative species by Verheye et al. (2016) was retained. For other sequences, one individual per clade corresponding to a morphospecies was selected.

BEAST2 (Bouckaert et al., 2014) on the CIPRES portal (Miller et al., 2010) was used to estimate divergence times under a Bayesian approach. A Bayesian Model Averaging method was implemented in BEAST2 with the bModelTest package (Bouckaert and Drummond, 2015) in order to estimate a phylogeny averaged over site models, and not to rely on a likelihood-based method to determine the site model. During the Bayesian analysis, Markov Chain Monte-

Carlo (MCMC) proposals switch between substitution models and estimate the posterior support for gamma-distributed rate heterogeneity, proportion of invariable sites and unequal base frequencies.

We simultaneously inferred the posterior distribution of trees and estimated divergence times assuming a relaxed clock model of evolution, allowing substitution rates to vary among branches. Both uncorrelated lognormal (UCLD) and exponential (UCED) models of rate change were implemented. In order to assess the pertinence of a relaxed estimation, the coefficients of variation of the clock rates were checked in Tracer v.1.6 (Rambaut et al., 2014). The coefficient of variation is the standard deviation of the clock rate distribution divided by its mean, and is used to assess the clock-likeness of the data. Values closer to zero indicate that the data are more clock-like. Therefore, values < 0.1 are generally considered low enough to justify the use of a strict molecular clock (Drummond and Bouckaert, 2015). To identify the best relaxed clock model, the marginal likelihoods of the competing models were estimated and summarized via the path-sampling method (Lartillot and Philippe, 2006), implemented in the MODEL_SELECTION package in BEAST2. Both Yule and Birth-death speciation processes were used as tree priors in combination with the UCLD clock model, and a path-sampling method was again used to select for the best tree model. All the path-sampling analyses were run for 100 steps of 2 x 10^6 generations each. The log Bayes Factors (BF) were calculated as follows: $\log_e BF (M_0, M_1) = \log_e P(X|M_0) - \log_e P(X|M_1)$, where $\log_e P(X|M_1)$ is the marginal \log_e -likelihood estimate for the model M_1 . The strength of support for a given model was based on the interpretation of BF suggested by Kass and Raftery (1995). Values of 2 $\log_e BF$ between 0 and 3 were interpreted as no evidence for the alternative model M_1 over the null model M_0 . When 2 $\log_e BF$ values were above 3, the alternative model M_1 was supported over the null model M_0 , and values over 20 were interpreted as a very strong support for the alternative model.

Likely due to their thin cuticle, peracarid crustaceans do not fossilize well (e.g. Briggs and Kear, 1994; Briggs and Wilby, 1996; Taylor, 1972). Their fossil record is very incomplete and therefore of limited utility for molecular dating. The aquatic fossil amphipods known (Coleman and Myers, 2000; Coleman and Ruffo, 2002; Coleman, 2004, 2006; Jażdżewski and Kulicka, 2000a, 2002; Jażdżewski and Kupryjanowicz, 2010; Jażdżewski et al., 2014; McMenamin et al., 2013; Weitschat et al., 2002) are all from freshwater taxa, phylogenetically distant from *Epimeria* (Verheye et al. 2015). Moreover, the detailed phylogenetic placement of these fossils could generally not be determined, due to their relatively poor preservation (e.g. Coleman, 2004; Coleman and Myers, 2000; Coleman and Ruffo, 2002; Jażdżewski and Kulicka, 2000b). Similarly, there is no unambiguous biogeographical event that could be used to calibrate the tree. We therefore used priors on rates of COI, 28S and H3 evolution based on rates inferred in previous studies. The prior rate of COI was set as a normal

distribution with a mean of 0.018 substitutions/site/My and a standard deviation (SD) of 0.0043. This rate was previously inferred for *Pontogammarus* amphipods (Nahavandi et al., 2013). A normal prior with a mean of 0.003 substitutions/site/My and SD of 0.0007 was used for the 28S gene, a rate inferred for the *Gammarus balcanicus* complex (Mamos et al., 2016). Rates of H3 evolution are, to our knowledge, not available for amphipods. Therefore, the prior rate of H3 was set as a normal distribution with a mean of 0.0019 and SD of 0.0004, a rate inferred for freshwater crabs (Klaus et al., 2010).

Two independent runs were performed with 200 million generations and a sampling frequency of 20000 generations. The first 1000 trees were discarded as burn-in and the results of the two runs were combined using the LogCombiner v1.7.5. Convergence was assessed by trace plots in Tracer v.1.6. and the effective sampling size for all parameters was more than 200 (Rambaut et al., 2014). The maximum clade credibility tree showing the mean nodal height was generated by TreeAnnotator v1.8.0. The final analyses were also run without data to ensure that the prior settings will not bias the results.

2.5. Rate and mode of diversification

The reduced dataset (one individual per species or putative species) was also used for diversification rates analyses because intraspecific polymorphisms can induce a false increase in diversification rates in the most recent history. All the diversification rate analyses were based on the Antarctic *Epimeria* clade. Incomplete taxonomic sampling can result in spurious declines in diversification rates over time (Pybus and Harvey, 2000). Our phylogeny comprises 72.5 % of the known (putative) species. However, as Verheye et al. (2016) discovered at least 24 new *Epimeria* species from the Peninsula, eastern Weddell Sea and Adélie Coast, it is likely that similar studies of material from e.g. the Ross, Amundsen and Bellingshausen Seas will reveal an even higher diversity within the genus. Therefore, the methods used herein to examine diversification patterns tested the effect of different levels of taxonomic sampling on the results: sampling proportions of 72 %, 50 % and 10 % were considered.

A mean semilogarithmic lineage through time (LTT) plot was constructed using the R package Ape to visualize the temporal pattern of lineage diversification. A straight line is expected under constant diversification rate. A departure from this straight line in the distant past may indicate a diversification rate change: (1) a concave plot either indicates a decelerating diversification rate or incomplete taxon sampling, (2) a convex plot may indicate accelerating diversification or a non-zero background extinction rate. The command sim.bd.taxa.age from the R package TreeSim (Stadler, 2011) was used to simulate

100 trees under Pure Birth (PB) and Birth-Death (BD), with speciation and extinction rates estimated with bd.shift.optim from the R package TreePar (Stadler, 2011). Both functions were used assuming a sampling fraction of 72 %. The LTT plot of the empirical phylogeny was compared to the 95 % confidence intervals of the expected pattern under PB and BD, to detect eventual deviations from the null hypothesis of constant rates.

In order to account for incomplete taxonomic sampling in methods that do not include a correction for missing species, we used the CorSim function from the R package TreeSim to simulate the missing splits on the empirical phylogeny (Cusimano et al., 2012). Missing speciation events were simulated 200 times under the assumption that evolution followed a constant BD model. Simulated branching times are added to the empirical branching times to obtain 200 completed (semi-empirical) datasets. Speciation and extinction rates used for the simulation were those estimated under BD with the bd.shifts.optim function of the TreePar package. Missing taxa were assumed to be located randomly across the tree, as both deep and shallower nodes are likely missing from the phylogeny. Three semi-empirical datasets were obtained for 72 %, 50 % and 10 % sampling.

The γ statistic was calculated on the semi-empirical datasets, using the GamStat function of the R package LASER (Rabosky, 2006a). This statistic tests for departure from a constant-rate pure birth model. Negative γ values indicate a prevalence of nodes closer to the root than expected under a Pure Birth process, therefore suggesting a decreasing rate of diversification through time. Positive γ values indicate either an increasing rate or non-zero extinction rate (Pybus and Harvey 2000).

We compared the fit of the branching times to various models of lineage accumulation, using the Akaike Information Criterion (AIC), for the three semi-empirical datasets. AIC is calculated as $-2\iota + 2k$, where ι is the log-likelihood value and k is the number of free parameters of the model (Burnham and Anderson, 2002). All model- fitting analyses were conducted with the R packages LASER (Rabosky, 2006a) and TreePar (Stadler, 2011). The constant-rate models included PB (constant speciation rate λ and no extinction) and BD (constant speciation rate λ and extinction rate μ). Two density-dependent models were included, which assume that the diversification rate decreases as the lineage population reaches some threshold density. The density-dependent linear (DDL) model assumes that λ decreases linearly and there is no extinction. The density-dependent exponential (DDX) model assumes that λ decreases exponentially and there is no extinction (Rabosky and Lovette, 2008). λ and μ may also change through time in response to external factors. Therefore, we included Yule-2-rate (Rabosky, 2006b) and Birth-death-shift (BDS) (Stadler, 2011) models in order to test whether and when discrete shifts in diversification

rate occurred during the clade's history. In between shifts, these models simplify to the constant-rate PB or BD. The mean and SD of the –Log Likelihood value, models' parameters and AIC scores were computed for each of the semi-empirical datasets. In order to compare the relative fit of the models, the 95 % confidence interval of the AIC values was computed. The difference in AIC (dAIC) between each model and the best-fitting model (with the lowest AIC) was computed. Minimal and maximal dAIC were calculated considering the AIC values comprised in the 95 % confidence intervals. The amount of statistical confidence for each model is represented by the Akaike weights (wAIC).

In order to test rate-variable models that allow for increasing or decreasing rates of speciation, extinctions and declining diversity, we also used the method of (Morlon et al., 2011), implemented in the R package RPANDA. The fit of the following models was compared using the AIC: (1) Bcst: constant speciation rate, no extinction (PB); (2) Bvar: exponential variation of speciation rate, no extinction; (3) BvarDcst: exponential variation of speciation rate, constant extinction; (4) BcstDcst: constant speciation and extinction rates (BD); (5) BcstDvar: constant speciation and exponential variation in extinction rate; and (6) BvarDvar: exponential variation in both speciation and extinction rates. These models were tested assuming sampling fractions of 0.1, 0.5 and 0.72.

The BAMM (Bayesian Analysis of Macroevolutionary Mixtures) 2.4.0 (Rabosky, 2014) software was used to explore eventual shifts in regimes across the branches of a phylogenetic tree, a regime being a constant or time-varying process of speciation and extinction. These heterogeneous mixtures of macroevolutionary rate regimes are sampled with reversible-jump MCMC. Priors were estimated using the setBAMMpriors command. MCMC chains were run for 10 million generations, and sampled every thousand generations. We checked for convergence of the MCMC chains and ESS (at least more than 200) using the R package coda (Plummer et al., 2005). The first 10% of samples were discarded as burn-in. The R package BAMMtools (Rabosky et al., 2014) was used to calculate the BF and the 95% credibility set for the shift configurations and to plot diversification rates through time (Rabosky, 2014). BAMM analyses were computed assuming different levels of taxon sampling: 10, 50 and 72 %.

3. Results

3.1. Data overview

We obtained 132 COI, 139 28S and 159 H3 sequences of *Epimeria* species and related taxa, and three iphimediid sequences used as outgroup. The aligned COI sequences contained 613 bp, with 379 variable sites (among ingroup taxa). The 28S alignment was 1797 bp long (after removal of ambiguous regions with Aliscore), with 1154 variable sites. The length of the H3 alignment was 369 bp, with 119 variable sites. The datasets were concatenated, resulting in an alignment of 2779 bp. The best models and partitioning scheme suggested by PartitionFinder are indicated in Table 2. A single set of underlying branch lengths was assumed for the tree.

| Subset Partitions | Best Model |
|--------------------------|-------------|
| 28S | GTR+I+G |
| COI_pos1 | SYM + I + G |
| COI_pos2 | GTR+I+G |
| COI_pos3 | GTR +I+G |
| H3_pos1, H3_pos2 | JC + I |
| H3_pos3 | GTR + G |

Table 2. Best partitioning scheme and models of DNA substitution, inferred by PartitionFinder.

3.2. Phylogenetic analyses

Congruence between gene trees and methods – The ILD test rejected the null hypothesis of congruence between all tree comparisons (p = 0.001). Upon examination of the tree topologies, incongruences between the three gene trees affected mostly unsupported nodes. Only one supported phylogenetic relationship within the Papuasian clade of *Epimeria* sp. 3 differs between the COI and 28S gene trees. As this unique incongruence does not affect the conclusions of this study, the datasets were concatenated. Differences between the topologies of the two reconstruction methods (ML and BI) were minimal. In all cases, these ambiguities affected only unsupported nodes.

Epimeria and related taxa – The Antarctic component of *Epimeria* is monophyletic with maximal support. All non-Antarctic *Epimeria* species also form a strongly supported clade (PP=1.00, BV=92). However, the monophyly of *Epimeria* is not supported by BI, as Antarctic *Epimeria* species form a sister clade to

two stilipedid species (PP=0.99), while this sister relationship remains unresolved by ML (BV=68). Deeper relationships between the Antarctic and non-Antarctic *Epimeria* clades, *Astyra abyssi*, the Acanthonotozomellid species and *Acanthonotozomopsis pushkini* are not supported. *Acanthonotozomoides oatesi* and *A. pushkini* form a maximally-supported clade (Fig. 1).

Antarctic *Epimeria* clade – The molecular systematics of the Antarctic *Epimeria* clade has been studied in detail in Verheye et al. (2016) and the morphological taxonomy in d'Udekem d'Acoz and Verheye (*in press*).

Non-Antarctic *Epimeria* clade – Twenty monophyletic morphospecies representing putatively new species are observed among Indo-Pacific *Epimeria* (Fig. 1). An indepth systematics study of these non-Antarctic *Epimeria* species is out of the scope of this paper and should be dealt with elsewhere.

Divergence times

The coefficients of variation were much higher than 0.1 for the three genes (COI: 1.1, 28S: 0.961, H3: 0.963), indicating that the sequences analyzed did not evolve at a constant rate along the branches. Therefore, we proceeded to use a relaxed molecular clock. Results of the path-sampling analysis and calculation of the BF are presented in Table 3. The UCED relaxed clock was strongly favored over the UCLD model. The BF strongly supports the Birth-Death over the Yule model (Table 3).

| | MLE | 2 lnBF |
|-------------|--------|--------|
| UCED | -36461 | 10278 |
| UCLD | -41600 | - |
| Birth-Death | -35922 | 34 |
| Yule | -35939 | - |

Table 3. Marginal likelihood estimation (MLE) values recovered by path-sampling. Bayes Factors (2 InBF) were estimated from the MLE to compare two relaxed clock models (UCED and UCLD) and two tree models (Yule and Birth-Death).

Bayesian posterior divergence times recovered by BEAST under an exponential relaxed clock model and Birth-Death tree model were consistent across the two runs. A mean age of 55.58 Ma (95 % HPD: 71.85–41.02 Ma) was estimated for the most recent common ancestor (MRCA) of the clade comprising the stilipedids and Antarctic *Epimeria*. The MRCA of the Antarctic *Epimeria* clade was given a mean age of 38.04 Ma (95 % HPD: 48.46–28.36 Ma). The MRCA of the non-Antarctic *Epimeria* clade was dated at 35.81 Ma (95 % HPD: 47.54–25.24 Ma), which separates into a Melanesian and an American-European-African lineage (Fig. 2).

3.3. Rates of diversification

Based on the γ statistic, we found no evidence of a decelerating lineage accumulation rate towards present time in the origination pattern of the Antarctic *Epimeria* clade. When incomplete sampling is taken into account, assuming 72 % or 50 % sampling, the γ values are positive, but the null hypothesis of constant rate of speciation is not rejected (p>0.01). When a 10 % sampling is assumed, the γ value is significantly positive (p<0.01), indicating either an increasing rate of speciation, or non-zero extinction (Table 4).

| Dataset | γ | p-value |
|---------------|-------------|-------------|
| | • | 2-tailed |
| Semi-emp 10 % | 11.02 (0.8) | 0 (0)* |
| Semi-emp 50 % | 1.419 (0.6) | 0.14 (0.18) |
| Semi-emp 72 % | 0.37 (0.47) | 0.63 (0.21) |
| Empirical | -0.27 | 0.79 |

Table 4. Results of the CR test applied on the empirical and semi-empirical datasets.

No evident deviation from the null hypotheses of constant-rate models (BD or PB) is visible on the LTT plots (Fig. 3).

The BDL analyses indicate PB as the best supported model of lineage diversification, closely followed by Yule2rate. The estimated probability that PB is the best model for our data among the evaluated models (wAIC) is only 31 %, which demonstrates that the method does not strongly support one model over the others. Table 5 shows the results for the best-scoring models only and Table A1 for all models evaluated. When incomplete sampling is taken into

account, PB is the best model for the 72 % semi-empirical dataset (wAIC = 33 %), Yule2rate is the best model for the 50 % semi-empirical dataset (wAIC=27 %) and BD is the best model for the 10 % semi-empirical dataset (wAIC=62 %) (Table A1).

| Models | LH | shift times | λ | AIC | dAIC | wAIC |
|-----------|--------|----------------|-------|-------|------|------|
| PB | -28.47 | - | 0.078 | 58.94 | 0.00 | 0.31 |
| Yule2rate | -26.62 | 2.09 | 0.088 | 59.24 | 0.30 | 0.27 |
| Tuicziate | -20.02 | 2.03 | 0.027 | 33.24 | 0.30 | 0.27 |

Table 5. Parameter estimates and comparison of the fit of different lineage diversification models to the empirical dataset. Results are shown for the two best-scoring models. LH is the log-likelihood value of the model. The shift times (Ma) are indicated for the models implying discrete shifts in diversification rates. λ is the speciation rate. For each model, the Akaike Information criterion (AIC) was computed. The best-fitting model with the lowest AIC score is indicated in bold. dAIc is the difference between the AIC score of the evaluated model and the AIC score of the best-fitting model. wAIC are the Akaike weights.

The evaluation of Morlon et al.'s (2011) models of lineage diversification with RPANDA yielded the following results. When a 0.72 sampling fraction was assumed, Bcst (PB) was the best fit among the evaluated models (wAIC = 0.45), followed by Bvar (wAIC = 0.22), which implies an increasing speciation rate and no extinction. The Bvar model (again implying an increasing speciation rate and no extinction), was the best-fitting model when assuming a sampling fraction of 0.5. However, the latter had almost the same probability than Bcst (PB) of being the best fit among the evaluated models (wAIC = 0.32 and 0.30, respectively). For the 0.1 sampling fraction, Bvar (with an increasing speciation rate and no extinction) was also the best-fitting model with a wAIC of 0.54. Table 6 shows the results for the best-scoring models only and Table A2 for all models evaluated.

| Dataset | Model | AIC | dAIC | wAIC | λ | α |
|---------------|-------|--------|------|------|-------|--------|
| Sampling 72 % | Bcst | 243.44 | 0.00 | 0.45 | 0.093 | - |
| | Bvar | 244.88 | 1.44 | 0.22 | 0.11 | -0.015 |

| Sampling FO 9/ | Bcst | 244.34 | 0.16 | 0.30 | 0.11 | - |
|----------------|------|--------|------|------|------|-------|
| Sampling 50 % | Bvar | 244.18 | 0.00 | 0.32 | 0.14 | -0.03 |
| Sampling 10 % | Bvar | 242.72 | 0.00 | 0.54 | 0.37 | -0.07 |

Table 6. Parameter estimates and comparison of the fit of Morlon et al.'s (2011) lineage diversification models with RPANDA. Results are shown for the best-scoring model and the second best fit whenever the wAIC of the best-scoring model was < 0.5. For each model, the Akaike Information criterion (AIC) was computed. The best-fitting model with the lowest AIC score is indicated in shaded grey. dAIC is the difference between the AIC score of the evaluated model and the AIC score of the best-fitting model. wAIC are the Akaike weights. λ is the speciation rate and α is the parameter controlling the exponential variation of the speciation rate with time.

For all considered sampling fractions (0.72, 0.5 and 0.1), the maximum posterior probability (MAP) shift configuration returned by BAMM — which is the distinct shift configuration with the highest posterior probability— is a single macroevolutionary regime across the whole tree and, i.e. no discrete rate shift. This is also the only regime contained in the 95 % credible shift set. Bayes Factor analyses also favour a model with no rate shifts, for all sampling scheme (Table 7). The rate through time plots show a constant diversification rate for the 0.72 sampling scheme, while it is increasing for the 0.1 and 0.5 sampling schemes (Fig. 4). However, the current version of BAMM assumes that all regimes are time-variable. In order to investigate the rate variation observed for the 0.1 and 0.5 sampling schemes, we computed the 95 % Higher Posterior Density (HPD) interval of the post-burnin samples of the lambda shift parameter. This parameter is the standard deviation of the normal distribution of speciation rates. A value of zero therefore indicates rate constancy. As the 95 % HPD of the lambda shift parameter includes 0 in both cases (95 % HPD for 10 % sampling [-0.005, 0.057] and for 50 % sampling [-0.022; 0.057]), there is no evidence for a variable speciation rate.

| • | | Sampling | 1 shift | 2 shifts | 3 shifts | |
|---|---------|----------|---------|----------|----------|--|
| • | | 10 % | 21.522 | 137.203 | - | |
| | 0 shift | 50 % | 9.729 | 60.878 | 213.075 | |
| | | 72 % | 7.495 | 32.676 | 95.057 | |

Table 7. Matrix of pairwise Bayes Factors (BF), where the BF is the ratio of marginal likelihoods between two models, Mi and Mj. Numerator models are given as rows and denominator models as columns. BF > 3 is considered positive support for Mi, while > 20 is a strong support (Kass and Raftery, 1995).



4. Discussion

4.1. Systematics of *Epimeria*

The monophyly of *Epimeria* was not supported by BI, as two species of Stilipedidae from the subfamily Alexandrellinae (*Bathypanoploea schellenbergi* and *Alexandrella dentata*) are sister to the Antarctic *Epimeria* clade (PP=0.99). The systematics of *Epimeria* should therefore be revised, as the genus may potentially include (at least part of) the Stilipedidae. This latter family was divided into three subfamilies: Astyrinae, Stilipedinae and Alexandrellinae (Holman and Watling, 1983). However, Stilipedidae is not monophyletic, as *Astyra abyssi* (from the subfamily Astyrinae) is not sister to Alexandrellinae neither in the present study nor in a previous phylogeny of Eusiroidea (Verheye et al. 2015). Since the systematics of Stilipedidae remains unclear, we will consider here only the Alexandrellinae as a potential sister clade to the Antarctic *Epimeria*. However, as this relationship is supported by low BV (68) and PP values have been shown to lead to higher type I error rates (Erixon et al., 2003), the support of this node should be verified with additional data before making any taxonomic changes. The main differences between *Epimeria* and *Alexandrella* reside in the mouthpart morphology. Several possible autapomorphies indeed characterize the *Alexandrella*: (1) the mandibular molar is absent, (2) the palp of the maxilla 1 is expanded, (3) the inner and outer plates of the maxilla 2 are expanded and (4) the outer plate of the maxilliped is greatly expanded (Holman and Watling, 1983). Moreover, the non-ommatidian eyes in Alexandrellinae could be indicative of a deep-sea ancestor (Warrant and Locket, 2004). If the Alexandrellinae are indeed nested within *Epimeria*, a shift of trophic niche and/or bathymetric range of their ancestor could have led to these morphological modifications.

4.2. Isolation of Antarctic *Epimeria* or historical dispersals?

As Antarctic *Epimeria* species are monophyletic and the origin of this Antarctic clade is estimated to be older or contemporaneous to the geographical isolation of Antarctica, it is hypothesized that *Epimeria* did not disperse in/out of the shelf throughout its evolutionary history. However, to further test the monophyly of this Antarctic component, it would be interesting to include the single Magellanic (*Metepimeria acanthurus*) and sub-Antarctic (*Epimeria ashleyi*, from the Macquarie Ridge) *Epimeria* species in the phylogenetic tree.

The Antarctic shelf is well isolated from the other ocean's shelves by large distances (> 850 km), deep seas and the most powerful current on earth, the ACC (Clarke et al., 2005; Orsi et al., 2005). Moreover, beyond the APF, surface water temperatures rise by 3-4 °C. Because of their limited tolerance to such temperature change, many Antarctic marine invertebrates would not be able to establish on the other side of the APF (Peck, 2002). Epimeria is presumably cold-adapted, as outside the Antarctic region, it is mainly found at bathyal depths. Temperature is therefore likely to be an important isolating factor for the Antarctic Epimeria clade. However, historical movements of stenothermal taxa in and out of the Antarctic shelf could have been possible during periods of climatic changes (Clarke et al., 1992). A variety of pelagic organisms (or organisms with a pelagic life stage) were reported to cross the APF, supposedly by means of eddies, i.e. water masses transported out of the ACC (Antezana, 1999; Clarke et al., 2005; Li et al., 2002) or by the Antarctic Intermediate water that extends northwards (Antezana, 1999). As such currents do not reach the ocean floor (Clarke et al., 2005), benthic taxa lacking a pelagic larval stage can only disperse through the deep-sea. Historical events of such polar emergences (i.e. colonization of the shelf by deep-sea fauna) and submergences (i.e. colonization of the deep-sea by shelf fauna) did occur in the evolutionary history of some strictly benthic taxa (Held, 2000; Raupach et al., 2007; Strugnell et al., 2008, 2011). Such movement is indeed facilitated by the thermohaline circulation, which connects the Southern Ocean shelf with the deep waters of the other world's oceans through an isothermal water column, and by the unusually deep Antarctic shelf (reaching > 1000 m at places) (Clarke, 2003; Clarke et al., 2009; Rogers, 2000). Polar submergences occurred in the Antarctic Epimeria clade history: whereas the vast majority of Antarctic Epimeria species are found on the shelf and upper slope (< 1200 m), two species (Epimeria larsi and Epimeria sp. 2) nested within shelf clades have bathymetric distributions restricted to slope depths (around 2000 m). However, the Drake Passage has an average depth of 3400 m (Smith and Sandwell, 1997). The bathymetric distributions of Epimeria species worldwide suggest that this might be too deep for most Epimeria to disperse in and/or out of the Antarctic region along the benthos, which would explain the evolutionary isolation of the Antarctic clade. However, two species were sampled in the abyssal plain: Epimeria glaucosa was found at 3710 m around New Zealand (Barnard, 1961) and Epimeria abyssalis at around 5600 m in the Kuril-Kamchatka Trench (Shimomura and Tomikawa, 2016). Additional deep-sea sampling and inclusion of abyssal species in phylogenetic analyses might reveal undetected historical dispersal of *Epimeria* through the deep-sea.

4.3 Origin and historical biogeography of Antarctic Epimeria

The first hypothesis aiming at determining the temporal and geographical origin of Antarctic *Epimeria* was presented by Lörz and Held (2004). Ages of between 34.9 and 15.7 Ma were inferred for the MRCA of *Epimeria*, based on COI distances and rates estimated for cirripeds (3.1 %/Myr) and for alpheid shrimps (2.4 %/Myr), and assuming a strict molecular clock. Applying a relaxed clock model on a combined dataset (COI, 28S and H3), we obtained a mean COI rate estimate of 1.6 %/Myr and an age of 38.04 Ma (95% HPD [48.46 Ma; 28.36 Ma]) for the MRCA of the Antarctic *Epimeria* clade. As the 95 % HPD interval spans an older period than the age estimation of Lörz and Held (2004), the assumption that this clade originated when Antarctica was already isolated from the other fragments of Gondwana is equivocal (Figure 2).

From the Late Cretaceous to the early Cenozoic, Antarctica, South America, Australia and New Zealand were connected in an area of cool temperate shallow seas known as the Weddelian Province (Woodburne and Zinsmeister, 1984; Zinsmeister, 1979, 1984). The South Tasman Saddle, a submarine trough between Tasmania and the South Tasman Rise, existed as a shallow to medium-depth seaway in the late Paleocene to early Eocene. If the Antarctic George V and Oates Coast shelf breaks were closer to the present shoreline at that time, it is possible that a deep seaway between Australia and East Antarctica existed as early as 40 Ma. Otherwise, a deepwater passage only developed after the late Eocene (Lawver et al., 2013, 2014). It has indeed been inferred that the Tasmanian gateway deepened in the period 35.5 to 30.2 Ma (Stickley et al., 2004) and was open to unrestricted deepwater circulation at 32 Ma (Lawver and Gahagan, 2003). An early opening of the Drake Passage to shallow water was suggested to have taken place in the Middle to Late Eocene (50–41 Ma) (Eagles et al., 2006; Livermore et al., 2005; Scher and Martin, 2004). This was followed by a progressive widening and deepening to intermediate depth at 37 Ma, while the deepwater passage is usually dated in the interval 34–30 Ma (Lagabrielle et al., 2009; Latimer and Filippelli, 2002; Lawver and Gahagan, 2003; Livermore et al., 2005; Scher and Martin, 2004) (Figure 2). The opening of the Tasmanian and Drake passages were inferred as important vicariant events, promoting speciation in other Antarctic benthic shelf species (Göbbeler and Klussmann-Kolb, 2010; Lee et al., 2004; Matschiner et al., 2011; Near, 2004).

It cannot be inferred from the current data whether the Antarctic *Epimeria* clade originated in the Weddelian Province or by vicariance when the continent separated from the other Gondwanan fragments. The divergence from its stilipedid sister clade (*Bathypanoploea schellenbergi* and *Alexandrella dentata*) occurred 55.58 Ma (95 % HPD [71.85; 41.02 Ma]), in the Weddelian Province (Figure 2). However, the inclusion of additional non-Antarctic samples — especially from historically connected and geographically closer regions such as South America, Australia and the sub-Antarctic islands — may help to

identify the sister lineage of the Antarctic *Epimeria* clade and shed light on its biogeographic origin. The endemicity of the Antarctic *Epimeria* clade would suggest that it originated *in situ* when the region was already isolated. In the case that this isolation was not yet geographical, it was suggested that a latitudinal gradient in seasonality might have promoted an early Cenozoic divergence of the polar fauna (Crame, 2013).

In any way, the initial diversification of the Antarctic Epimeria lineage would have occurred in a cooling environment. Following the early Eocene Climatic Optimum (EECO), high latitude surface water began to cool across the early to middle Eocene boundary (49–48 Ma) (Lawver and Gahagan 2003). A second cooling phase occurred in the late Middle Eocene 44–41 Ma. From the beginning of this cooling trend, the ice sheets progressively grew to culminate at the Eocene-Oligocene boundary (Miller et al., 2008). The transition to an icehouse climate in the earliest Oligocene, known as the Eocene-Oligocene Climate Transition (EOCT), was marked by an abrupt cooling at 33.55 Ma (Miller et al., 1991) and the first continent-wide glaciations (Hambrey et al., 1991; Sorlien et al., 2007; Wilson et al., 2013). Most of the species complexes identified in Verheye et al. (2016) diversified after the MMCT, (~14 Ma), a period marking the beginning of repeated ice sheet advances over the shelf and retreats inshore (Shevenell et al., 2004). Globally, these results suggest that the Antarctic Epimeria clade's diversification would be related to cold waters (Figure 2). Many modern Antarctic lineages likely arose in the time period spanning this transition from Eocene cool-temperate climate to Oligocene polar climate, as the fossil record indicates a fundamental shift in the structure of benthic communities. The extinction of many durophagous (skeleton-crushing) predator taxa — decapods, teleost and cartilaginous fishes — decreased the predation pressure, which allowed the establishment of dense populations of erect sessile suspension feeders (Aronson et al., 2007). These organisms constitute a food source for many amphipods, which are predatory food specialists (e.g. Coleman, 1989, 1991; Klages and Gutt, 1990). Moreover, sessile suspension-feeders form dense assemblages which provide a three-dimensionally structured habitat for an errant fauna of mostly slow-moving invertebrates, such as amphipods (Arntz et al., 2005; Aronson and Blake, 2001; Clarke et al., 2004; Clarke and Crame, 2010). In summary, the adaptation of Epimeria to a cold environment, coupled with the extinction of cold-intolerant taxa and resulting abundant available niche-space likely led to their successful diversification on the Antarctic shelf.

The divergence between a clade entirely composed of Melanesian species and a clade that includes European, African and South-American species occurred 35.81 Ma 95% HPD [47.54 Ma; 25.24 Ma] (Figure 2). The Indo-Pacific islands were likely not colonized by *Epimeria* species from the Weddelian Province or

Australia, but rather from south-eastern continental Asia. Indeed, at the time of origin of this Melanesian clade, the South East Asian gateway that connects the Indian and Pacific oceans was widely opened, likely forming an efficient barrier to dispersion between Australia and the Indo-Pacific region (Hall et al., 2011).

4.4 Diversification of *Epimeria* on the Antarctic shelf

No early diversification burst was detected in the phylogenetic pattern of the Antarctic *Epimeria* clade, nor any shift in diversification rates associated with glacial cycles. The Antarctic *Epimeria* clade has survived through multiple mass extinction events associated with the cooling of the continent and the glacial cycles. The fossil record indeed indicates many biotic turnovers induced by climate change (Beu, 2009; Crame et al., 2014; Hara, 2001). Furthermore, it was shown that following the MMCT, the ice sheet extended to the outer shelf during glacial maxima, at least in some places, thereby erasing shelf habitats and their associated biota (e.g. Chow and Bart, 2003; Hambrey and McKelvey, 2000; Passchier et al., 2003).

Despite the evidence for repeated extinctions of the Antarctic shelf biota, the extinction rate estimates given by the different methods used herein are unrealistically close to zero. Extinction rate estimates from phylogenies of extant taxa were often shown to be unreliable (Nee, 2006; Purvis, 2008; Rabosky, 2010; Rabosky and Lovette, 2008; Ricklefs, 2007). In the present case, the phylogeny may simply lack sufficient information to accurately estimate extinction rates or infer diversity dynamics (Liow et al., 2010; Morlon et al., 2011; Quental and Marshall, 2009; Rabosky, 2010).

Mass extinctions are believed to promote adaptive radiations, manifested by a subsequent sharp increase in the rate of diversification (Benton and Emerson, 2007; Condamine et al., 2013; McInnes et al., 2011). These "rebounds" would be followed by a density-dependent decrease in diversification rates. However, if repeated mass extinctions occurred, the signals of older pulses of diversification tend to be eroded by subsequent extinctions (Phillimore and Price, 2008; Rabosky and Lovette, 2008; Ricklefs and Jønsson, 2014; Weir, 2006). Because of the presumed high turnover rate in the evolutionary history of the Antarctic *Epimeria* clade, even a large early diversification burst could be undetectable (McInnes et al., 2011; Quental and Marshall, 2009; Rabosky and Lovette, 2008). However, without a fossil record, we cannot determine whether this turnover rate was sufficiently high compared to the initial speciation rate to erode an early burst signature (Quental and Marshall, 2009).

If the history of environmental disturbances is responsible for the lack of an early burst signature, then we can expect similar patterns in other benthic shelf taxa. The paucity of branching events between the Oligocene and MMCT in the notothenioid fishes' phylogeny was interpreted as unobserved extinction, possibly eroding the signature of an explosive early radiation. However, the latter study also provided evidence of evolutionary radiations for several younger subclades, coinciding with the MMCT climate change event (Near et al., 2012).

The relatively low number of tips in our phylogeny might decrease the statistical power of the tests to detect lineage-specific shifts. It has been shown that a reduced taxon sampling compromises the detection of discrete rate shifts and slowdowns in diversification (Laurent et al., 2015; Phillimore and Price, 2008, 2009). This could also explain the inability of the model-selection analyses to discriminate between different equally fitting models. When the true diversity of the clade is better known, diversification analyses repeated on a thoroughly sampled phylogeny might reveal unobserved diversification regime changes.

Finally, in order to detect an early burst of diversification, the clade must also be close to its equilibrium diversity (McInnes et al., 2011; Quental and Marshall, 2010). Yet, the Antarctic *Epimeria* clade may still be in the earliest stage of a logistic growth following the Last Glacial Maximum. The exponential growth of diversity that would be observed in such case (Liow et al., 2010; Quental and Marshall, 2010) is consistent with our results of either constant or increasing rates of speciation. Assuming that the dynamics of the Antarctic *Epimeria* clade is governed by diversity-dependent diversification, there is, however, no way of determining in which stage of its diversification trajectory this clade actually stands (Quental and Marshall, 2010).

5. Conclusions

By including species from all the world's oceans and by applying a relaxed clock model to date the phylogeny, this study provides the first spatiotemporal framework for the evolution of Antarctic *Epimeria*. Although the precise timing of origin of the Antarctic clade is still debatable, our data suggest that this lineage originated *in situ* — in the Late Gondwanan Weddellian Province or by vicariance when the plates broke apart — and does not result from a colonization event after the geographical isolation of the continent. The initial diversification of the clade occurred in a cooling environment, which suggests that the adaptation of this genus to cold waters, along with the extinction of cold-intolerant taxa and resulting availability of habitats, would have led to their successful radiation on the Antarctic shelf. The Antarctic *Epimeria* clade appears to have evolved in complete isolation from the other world's oceans,

as our data do not provide any evidence of dispersal in and/or out of the shelf since the isolation of Antarctica. Sampling of additional non-Antarctic *Epimeria*, especially from historically connected regions (South America, Australia and New Zealand), could help identify the sister lineage of the Antarctic clade and clarify its biogeographical origin.

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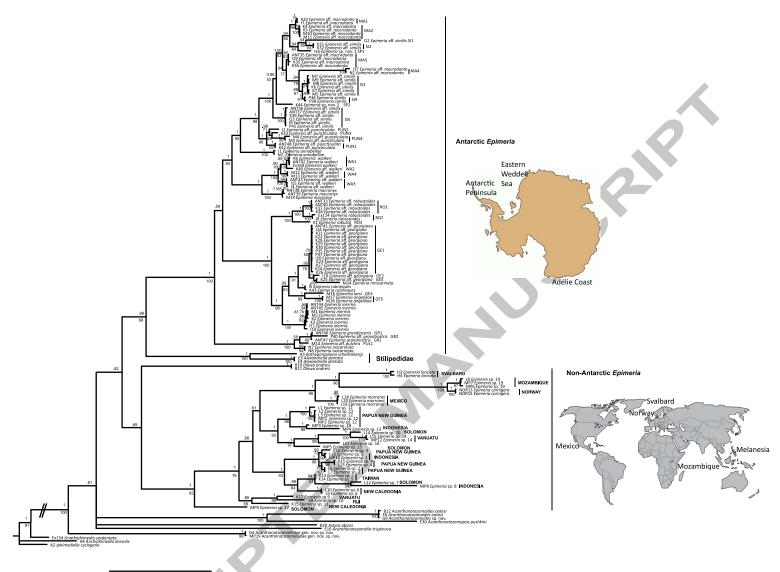
Captions

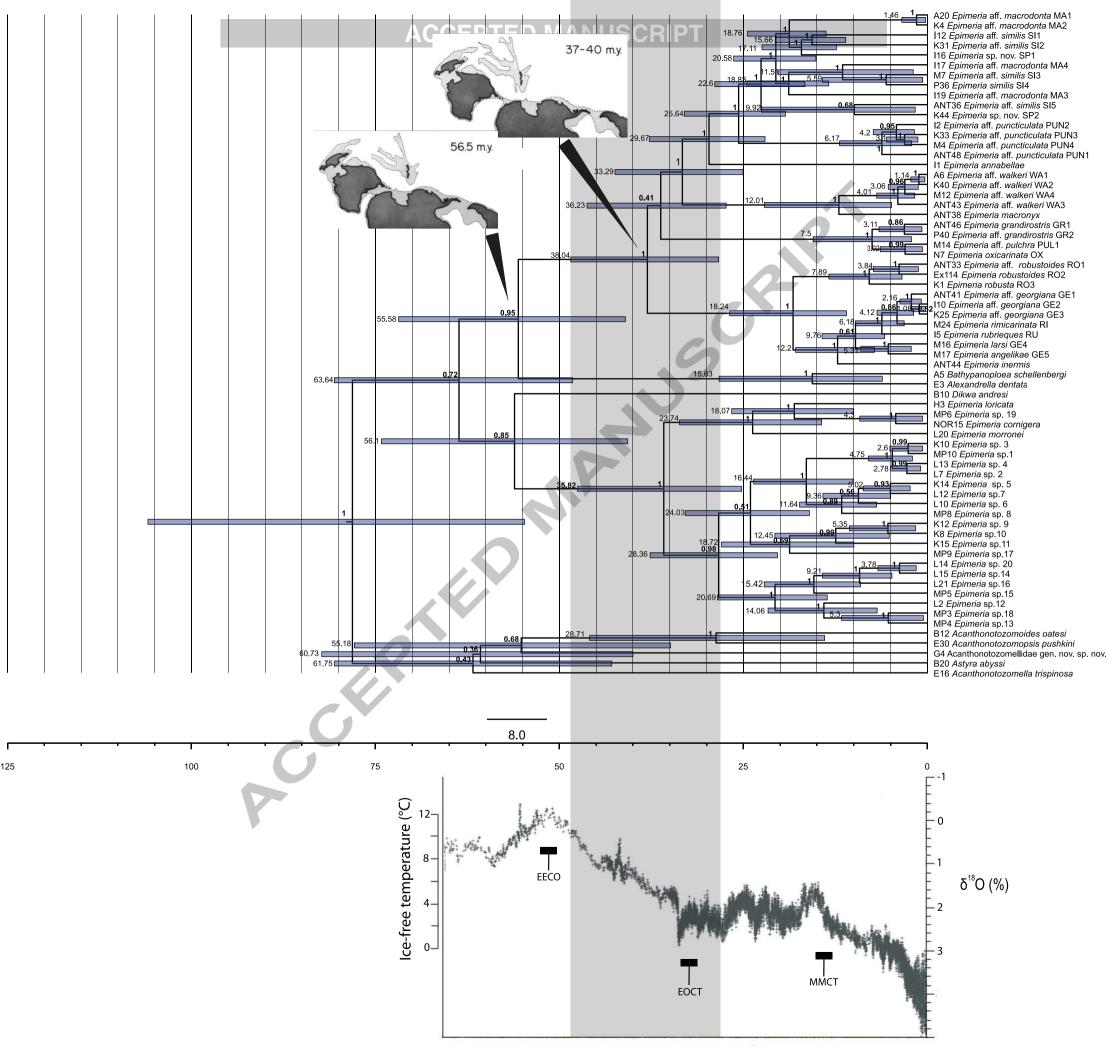
Figure 1. Bayesian phylogenetic tree of the concatenated dataset (COI, 28S, H3). Posterior probabilities are indicated above the nodes and Bootstrap values (> 50) from the Maximum Likelihood analysis below the nodes. MOTUs names (putative species) from Verheye et al. (2016) are indicated for Antarctic *Epimeria* specimens.

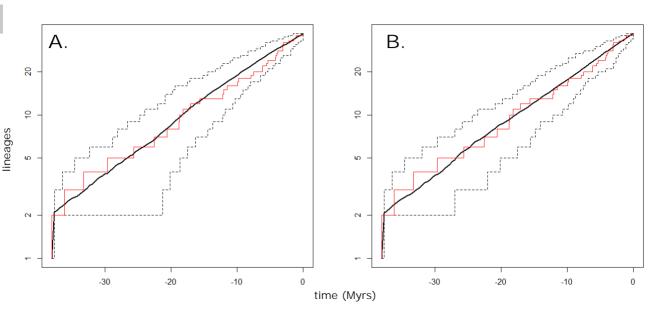
Figure 2. Maximum clade credibility chronophylogenetic tree from the BEAST analysis of the concatenated dataset. Blue bars on the tree indicate 95 % confidence intervals for estimated node ages. Mean node ages are indicated in front of the bars. Posterior probabilities are indicated above the nodes. The grey shaded area is the 95 % confidence interval of the initial diversification of the Antarctic *Epimeria* clade. The paleogeographic figures (from Zinsmeister, 1982) on the tree show the Late Gondwana's breakup. Shaded area on these figures represent inferred areas of shallow marine conditions. The graph below (modified from Zachos et al. 2008) shows the paleotemperatures over the past 65 million years, inferred from the benthic foraminiferal δ^{18} O curve, which is based on records from Deep Sea Drilling Project and Ocean Drilling Program sites. EECO = Early Eocene Climate Optimum; EOCT = Eocene-Oligocene Climate Transition; MMCT = Middle Miocene Climate Transition.

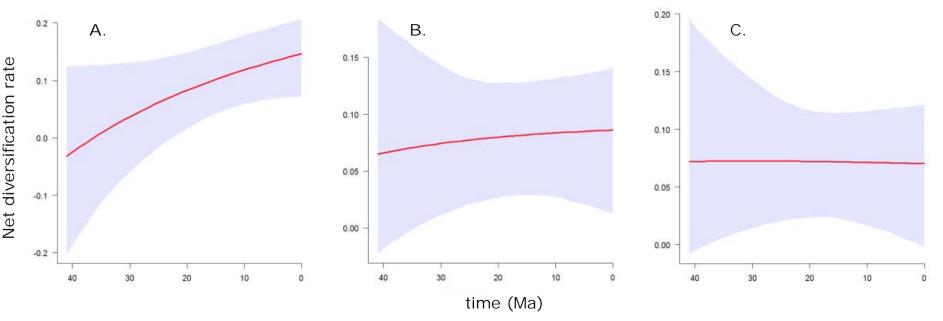
Figure 3. Lineage diversification through time plot of the Antarctic *Epimeria* clade, generated with the dated tree from the BEAST analysis (in red). Dotted lines represent the 95% confidence interval of lineage diversification simulated 100 times under PB (A) and BD (B) models, with λ = 0.09 (PB) and with λ = 0.09 and μ = 0.02 (BD), and assuming a sampling fraction of 0.72.

Figure 4. Net diversification rate through time inferred with BAMM for the Antarctic Epimeria clade, assuming sampling fractions of A. 0.1 B. 0.5 and C. 0.72. The shaded blue area represents the 95 % credible interval on the rate values.









| Specimens | Consister | | | | | Genba | number | |
|-------------|------------------------------|---|--------------|---------------|-------------------|----------|----------|---------|
| id code | Species | Locality | Latitude | Longitude | Voucher ID | COI | 285 | Н3 |
| EPIMERIIDAE | | | | | | | | |
| A20 | Epimeria aff. macrodonta MA1 | Peninsula, Larsen B | 65° 57.51′ S | 60° 28.15' W | RBINS INV.132655 | KU870817 | KU759589 | KY82581 |
| I7 | Epimeria aff. macrodonta MA1 | Peninsula, Larsen B | 65° 57.51' S | 60° 28.15' W | RBINS INV.132975 | KU870851 | KU759628 | KY82586 |
| K4 | Epimeria aff. macrodonta MA2 | Adélie Coast | 66° 38.42′ S | 139° 49.72′ E | MNHN-IU-2009-2570 | KU870872 | KU759652 | KY82589 |
| K5 | Epimeria aff. macrodonta MA2 | Adélie Coast | 66° 36.37' S | 140°05.07' E | MNHN-IU-2009-2563 | KU870876 | KU759657 | KY82589 |
| M10 | Epimeria aff. macrodonta MA2 | Adélie Coast | 66° 41.12' S | 139° 56.69' E | MNHN-IU-2014-4299 | KU870878 | KU759661 | KY82592 |
| M11 | Epimeria aff. macrodonta MA2 | Adélie Coast | 66° 40.12' S | 139° 55.93' E | MNHN-IU-2014-4296 | N/A | KU759662 | KY82592 |
| I19 | Epimeria aff. macrodonta MA3 | Peninsula, Dundee Island | 63° 37.29' S | 56° 09.11' W | RBINS INV.132974 | KU870844 | KU759621 | KY82585 |
| К36 | Epimeria aff. macrodonta MA3 | Peninsula, North of Joinville Island | 62° 33.79' S | 56° 27.81′ W | RBINS INV.122929B | KU87086 | KU759648 | KY82588 |
| K35 | Epimeria aff. macrodonta MA3 | Peninsula, North of Joinville Island | 62° 33.79' S | 56° 27.81′ W | RBINS INV.122929A | N/A | KU759647 | KY82588 |
| ANT35 | Epimeria aff. macrodonta MA3 | Peninsula, Erebus and Terror Gulf | 63° 58.78′ S | 56° 46.24′ W | RBINS INV.122940 | N/A | KU759594 | KY82582 |
| I17 | Epimeria aff. macrodonta MA4 | Bransfield Strait East | 62° 43.73′ S | 57° 29.04' W | RBINS INV.132660 | N/A | KU759619 | KY82585 |
| N1 | Epimeria aff. macrodonta MA4 | Bransfield Strait East | 62° 43.73′ S | 57° 29.04' W | RBINS INV.132973 | N/A | KU759677 | KY82592 |
| l12 | Epimeria aff. similis SI1 | Eastern Weddell Sea | 70° 50.48' S | 10° 35.28' W | RBINS INV.132664 | N/A | KU759614 | KY82585 |
| K31 | Epimeria aff. similis SI2 | Bransfield Strait Central | 62° 53.45′ S | 58° 13.06′ W | RBINS INV.122931A | KU870865 | KU759644 | KY82588 |
| K32 | Epimeria aff. similis SI2 | Bransfield Strait Central | 62° 53.45′ S | 58° 13.06′ W | RBINS INV.122935 | KU870866 | KU759645 | KY82588 |
| К6 | Epimeria aff. similis SI3 | Adélie Coast | 66° 38.00' S | 140° 42.00' E | MNHN-IU-2009-2532 | N/A | KU759658 | KY82589 |
| M7 | Epimeria aff. similis SI3 | Adélie Coast | 66° 10.57' S | 143° 20.75' E | MNHN-IU-2014-4342 | N/A | KU759674 | KY82593 |
| M8 | Epimeria aff. similis SI3 | Adélie Coast | 66° 45.14′ S | 145° 20.07' E | MNHN-IU-2014-4333 | N/A | KU759675 | KY82593 |
| M9 | Epimeria aff. similis SI3 | Adélie Coast | 66° 45.14′ S | 145° 20.07' E | MNHN-IU-2014-4322 | N/A | KU759676 | KY82593 |
| M5 | Epimeria aff. similis SI3 | Adélie Coast | 65° 59.83' S | 143° 38.99' E | MNHN-IU-2014-4340 | N/A | KU759672 | KY82593 |

| K7 | Epimeria aff. similis SI3 | Adélie Coast | 66° 38.00' S | 140° 42.00' E | MNHN-IU-2009-2539 | N/A | KU759659 | KY825900 |
|-------|---------------------------------------|---|--------------|---------------|-------------------|---------------------------------------|----------|----------|
| P36 | Epimeria similis SI4 | Bransfield Strait East | 62° 43.73′ S | 57° 29.04' W | RBINS INV.122956A | N/A | KU759680 | KY825935 |
| P38 | Epimeria similis SI4 | Bransfield Strait Central | 62° 53.45′ S | 58° 13.06′ W | RINBS INV.122922B | N/A | KU759682 | KY825953 |
| ANT37 | Epimeria aff. similis SI5 | Bransfield Strait East | 62° 47.80′ S | 57° 05.35′ W | RBINS INV.122942 | KU870823 | KU759596 | KY825822 |
| 113 | Epimeria aff. similis SI5 | Bransfield Strait East | 62° 43.73′ S | 57° 29.04' W | RBINS INV.132976 | KU870839 | KU759615 | KY825852 |
| 19 | Epimeria aff. similis SI5 | Eastern Weddell Sea | 70° 47.34' S | 10° 40.39' W | RBINS INV.132665 | KU870853 | KU759630 | KY825865 |
| P41 | Epimeria aff. similis \$15 | Eastern Weddell Sea | 70° 23.94' S | 08° 19.14' W | RBINS INV.132977 | KU870895 | KU759684 | KY825955 |
| K39 | Epimeria aff. similis SI5 | Bransfield Strait Central | 62° 53.45′ S | 58° 13.06′ W | RBINS INV.122922A | KU870871 | KU759651 | KY825892 |
| ANT36 | Epimeria aff. similis SI5 | Bransfield Strait East | 62° 43.50′ S | 57° 27.92' W | RBINS INV.132666 | KU870822 | KU759595 | KY825821 |
| l16 | Epimeria sp. nov. 1 SP1 | Eastern Weddell Sea | 70° 38.66′ S | 10° 28.16' W | RBINS INV.132667 | KU870842 | KU759618 | KY825855 |
| K44 | Epimeria sp. nov. 2 SP2 | Eastern Weddell Sea | 70° 05.13′ S | 03° 23.50' W | RBINS INV.132663 | KU870875 | KU759656 | KY825897 |
| ANT48 | Epimeria aff. puncticulata PUN1 | Bransfield Strait | 62° 45.05′ S | 57° 26.68′ W | RBINS INV.122947 | N/A | KU759607 | KY825832 |
| K42 | Epimeria aff. puncticulata PUN1 | Peninsula, Dundee Island | 63°36.84' S | 56° 10.28' W | RBINS INV.122934 | N/A | KU759655 | KY825896 |
| 12 | Epimeria aff. puncticulata PUN2 | Peninsula, Larsen B | 65° 57.51' S | 60° 28.15' W | RBINS INV.132651 | KU870845 | KU759622 | KY825859 |
| К33 | Epimeria aff. puncticulata PUN3 | Adélie Coast | 66° 39.3' S | 139° 55.8' E | MNHN-IU-2009-2578 | KU870867 | KU759646 | KY825887 |
| M4 | Epimeria aff. puncticulata PUN4 | Adélie Coast | 65° 48.48' S | 143° 03.76' E | MNHN-IU-2014-4288 | KU870888 | KU759671 | KY825933 |
| M3 | Epimeria aff. puncticulata PUN4 | Adélie Coast | 65° 48.48′ S | 143° 03.76' E | MNHN-IU-2014-4288 | N/A | KU759670 | KY825932 |
| A6 | Epimeria walkeri WA1 | King George Island | 62° 18.21′ S | 58° 39.90' W | RBINS INV.132667 | KU870819 | KU759591 | KY825818 |
| Ex169 | Epimeria walkeri WA1 | Eastern Weddell Sea | 70° 56.40' S | 10° 32.60' W | Specimen missing | KU870836 | KU759610 | KY825842 |
| ANT42 | Epimeria walkeri WA1 | Bransfield Strait Central | 62° 53.64′ S | 58° 12.52' W | RBINS INV.122944 | KU870828 | KU759601 | KY825826 |
| K40 | Epimeria aff. walkeri WA2 | Bransfield Strait East | 62° 44.73′ S | 57° 26.79' W | RBINS INV.122932 | KU870873 | KU759653 | KY825894 |
| ANT43 | Epimeria aff. walkeri WA3 | Drake Passage West | 62° 17.36' S | 61° 12.06' W | RBINS INV.122949 | KU870829 | KU759602 | KY825827 |
| l15 | Epimeria aff. walkeri WA3 | Peninsula, South of Joinville Island | 63° 50.92′ S | 55° 37.66' W | RBINS INV.132656 | KU870841 | KU759617 | KY825854 |
| 14 | Epimeria aff. walkeri WA3 | Elephant Island | 61° 20.76′ S | 55° 12.14' W | RBINS INV.132959 | KU870848 | KU759625 | KY825861 |
| | · · · · · · · · · · · · · · · · · · · | | | · | | · · · · · · · · · · · · · · · · · · · | | · |

| M12 | Epimeria aff. walkeri WA4 | Adélie Coast | 66° 45.14′ S | 145° 20.07' E | MNHN-IU-2014-4331 | KU870879 | KU759663 | KY825923 |
|-------|-----------------------------|---|--------------|---------------|-------------------|----------|----------|----------|
| M13 | Epimeria aff. walkeri WA4 | Adélie Coast | 66° 23.99' S | 140° 32.35′ E | MNHN-IU-2014-4336 | KU870880 | KU759664 | KY825924 |
| ANT38 | Epimeria macronyx MX1 | Drake Passage West | 62° 22.65′ S | 61° 17.63' W | RBINS INV.122943 | KU870824 | KU759597 | КҮ907661 |
| ANT39 | Epimeria macronyx MX1 | Drake Passage West | 62° 22.65′ S | 61° 17.63' W | RBINS INV.122943 | KU870825 | KU759598 | KY825823 |
| M19 | Epimeria macronyx MX2 | Adélie Coast | 66° 44.86′ S | 145° 26.66′ E | MNHN-IU-2014-4276 | KU870885 | KU759668 | KY825930 |
| ANT41 | Epimeria aff. georgiana GE1 | Bransfield Strait East | 62° 44.73′ S | 57° 26.79' W | RBINS INV.122867 | KU870827 | KU759600 | KY825825 |
| 114 | Epimeria aff. georgiana GE1 | Drake Passage West | 62° 17.36' S | 61° 12.06' W | RBINS INV.132970 | KU870840 | KU759616 | KY825853 |
| 120 | Epimeria aff. georgiana GE1 | Peninsula, Dundee Island | 63° 51.53′ S | 55° 40.74' W | RBINS INV.132971 | KU870846 | KU759623 | KY825860 |
| K21 | Epimeria aff. georgiana GE1 | Bransfield Strait West | 63° 00.53′ S | 58° 35.67' W | RBINS INV.122926 | KU870855 | KU759633 | KY825874 |
| K22 | Epimeria aff. georgiana GE1 | Bransfield Strait West | 63° 00.53′ S | 58° 35.67' W | RBINS INV.122924 | KU870856 | KU759634 | KY825875 |
| K23 | Epimeria aff. georgiana GE1 | Peninsula, East of Joinville Island | 63° 10.57′ S | 54° 06.66' W | RBINS INV.122930A | KU870857 | KU759635 | КҮ825876 |
| K24 | Epimeria aff. georgiana GE1 | Peninsula, East of Joinville Island | 63° 10.57' S | 54° 06.66' W | RBINS INV. 122933 | KU870858 | KU759636 | КҮ825877 |
| K26 | Epimeria aff. georgiana GE1 | Peninsula, South of Joinville Island | 61° 56.05′ S | 60° 05.56′ W | RBINS INV.122921A | KU870860 | KU759638 | КҮ825879 |
| K27 | Epimeria aff. georgiana GE1 | Peninsula, North of Joinville Island | 62° 33.79′ S | 56° 27.81′ W | RBINS INV.122920 | KU870861 | KU759639 | KY825880 |
| K28 | Epimeria aff. georgiana GE1 | Peninsula, North of Joinville Island | 62° 33.79′ S | 56° 27.81′ W | RBINS INV.122920 | KU870862 | KU759640 | KY825881 |
| K29 | Epimeria aff. georgiana GE1 | Peninsula, Dundee Island | 63° 37.28′ S | 56° 09.11′ W | RBINS INV.122923 | KU870863 | KU759641 | KY825882 |
| K30 | Epimeria aff. georgiana GE1 | Peninsula, Dundee Island | 63° 37.28′ S | 56° 09.11′ W | RBINS INV.122925 | KU870864 | KU759643 | KY825884 |
| P35 | Epimeria aff. georgiana GE1 | Peninsula, South of Joinville Island | 63° 51.34′ S | 55° 41.11′ W | RBINS INV.122921B | KU870893 | KU759679 | КҮ825927 |
| P37 | Epimeria aff. georgiana GE1 | Peninsula, South of Joinville Island | 63° 51.34′ S | 55° 41.11′ W | RBINS INV.122921C | KU870894 | KU759681 | КҮ825952 |
| 110 | Epimeria aff. georgiana GE2 | South Orkney Islands | 62° 53.45' S | 58° 13.06′ W | RBINS INV.132658 | KU870838 | KU759612 | KY825849 |
| K25 | Epimeria aff. georgiana GE3 | Drake Passage East | 61° 56.05′ S | 60° 05.56' W | RBINS INV.122936 | KU870859 | KU759637 | KY825878 |
| M16 | Epimeria aff. georgiana GE4 | Adélie Coast | 65° 43.12' S | 143° 03.61' E | MNHN-IU-2014-4344 | KU870882 | KU759665 | KY825926 |
| | | | | | | | | |

| M17 | Epimeria angelikae GE5 | Adélie Coast | 65° 28.85' S | 139° 24.18' E | MNHN-IU-2014-4278 | KU870883 | KU759666 | KY825928 |
|-------|---------------------------------|---------------------------|--------------|---------------|----------------------------|----------|----------|----------|
| M18 | Epimeria angelikae GE5 | Adélie Coast | 65° 52.74′ S | 144° 10.92' E | MNHN-IU-2014-4281- IO15 | KU870884 | KU759667 | KY825929 |
| M24 | Epimeria rimicarinata RI | Prydz Bay | 66° 55.75' S | 74° 04.19' E | MNHN-IU-2014-4265 | KU870887 | N/A | N/A |
| 15 | Epimeria rubrieques RU | Eastern Weddell Sea | 70° 47.34' S | 10° 40.39' W | RBINS, INV.132668 | KU870849 | KU759626 | KY825862 |
| K41 | Epimeria rubrieques RU | Eastern Weddell Sea | 70° 23.94' S | 08° 19.14' W | RBINS INV.132643 | KU870874 | KU759654 | KY825895 |
| ANT44 | Epimeria inermis IN1 | Bransfield Strait West | 62° 55.83′ S | 58° 41.09' W | RBINS INV.122948 | KU870830 | KU759603 | KY825828 |
| ANT45 | Epimeria inermis IN1 | Bransfield Strait West | 63° 00.53′ S | 58° 35.67' W | RBINS INV.122945 | KU870831 | KU759604 | KY825829 |
| I18 | Epimeria inermis IN1 | Bransfield Strait Central | 62° 57.22' S | 58° 14.60′ W | RBINS INV.132953 | KU870843 | KU759620 | KY825857 |
| I11 | Epimeria inermis IN2 | Eastern Weddell Sea | 70° 48.93′ S | 10° 32.69' W | RBINS INV.132655 | N/A | KU759613 | KY825850 |
| K2 | Epimeria inermis IN2 | Adélie Coast | 66° 39.30' S | 140° 01.60' E | MNHN-IU-2009-2531 | N/A | KU759632 | KY825873 |
| К3 | Epimeria inermis IN2 | Adélie Coast | 66° 40.50' S | 139° 55.20' E | MNHN-IU-2009-2569 | N/A | KU759642 | KY825883 |
| M1 | Epimeria inermis IN2 | Adélie Coast | 66° 53.36' S | 142° 38.90' E | MNHN-IU-2014-4272 | KU870877 | KU759660 | KY825919 |
| M2 | Epimeria inermis IN2 | Adélie Coast | 66° 23.99' S | 140° 32.35' E | MNHN-IU-2014-4338 | KU870886 | KU759669 | KY825931 |
| ANT33 | Epimeria aff. robustoides RO1 | Bransfield Strait East | 62° 47.80′ S | 57° 05.35′ W | RBINS INV.122937A | KU870820 | KU759592 | KY825819 |
| ANT40 | Epimeria aff. robustoides RO1 | Bransfield Strait West | 62° 55.83′ S | 58° 41.09' W | RBINS INV.122939 | KU870826 | KU759599 | KY825824 |
| K37 | Epimeria aff. robustoides RO1 | Bransfield Strait Central | 62° 55.83′ S | 58° 41.09' W | RBINS INV.122927 | KU870869 | KU759649 | KY825890 |
| K38 | Epimeria aff. robustoides RO1 | Bransfield Strait Central | 62° 55.83′ S | 58° 41.09' W | RBINS INV.122928 | KU870870 | KU759650 | KY825891 |
| Ex114 | Epimeria robustoides RO2 | Eastern Weddell Sea | 70° 47.34′ S | 10° 40.39' W | RBINS-INV.122894 | KU870834 | KU759608 | KY825840 |
| 18 | Epimeria robustoides RO2 | Eastern Weddell Sea | 70° 23.94′ S | 08° 19.14' W | RBINS INV.132969 | KU870852 | KU759629 | KY825864 |
| K1 | Epimeria robusta RO3 | Adélie Coast | 66° 38.40′ S | 140° 01.80' E | MNHN-IU-2009-2571 | KU870854 | KU759631 | KY825866 |
| ANT46 | Epimeria grandirostris GR1 | Peninsula, Dundee Island | 63° 51.34' S | 55° 41.11′ W | RBINS INV.122946 | KU870832 | KU759605 | KY825830 |
| ANT47 | Epimeria grandirostris GR1 | Bransfield Strait East | 62° 45.05' S | 57° 26.68' W | RBINS INV.122950 | KU870833 | KU759606 | KY825831 |
| P40 | Epimeria aff. grandirostris GR2 | Adélie Coast | 66° 20.33′ S | 143° 41.13' E | MNHN-IU-2014-4327 | N/A | KU759683 | KY825954 |
| M14 | Epimeria aff. pulchra PUL1 | Adélie Coast | 65° 59.78' S | 143° 02.95' E | MNHN-IU-2014-4284 | KU870881 | N/A | KY825925 |
| | | | | | | | - | |

| N7 | Epimeria oxicarinata OX | Elephant Island | 61°20.27'S | 55° 30.92' W | RBINS INV.122468 | KU870891 | N/A | KY825949 |
|------|-------------------------|---|--------------|----------------|--------------------|----------|----------|----------|
| N8 | Epimeria oxicarinata OX | Elephant Island | 61° 20.33'S | 55° 31.53' W | RBINS INV.122483 | KU870892 | N/A | N/A |
| I1 | Epimeria annabellae AN | Eastern Weddell Sea | 70° 48.93' S | 10° 32.69' W | RBINS INV.132652 | KU870837 | KU759611 | KY825848 |
| N2 | Epimeria annabellae AN | Eastern Weddell Sea | 70° 30.99' S | 08° 48.08' W | RBINS INV.122476 | KU870890 | KU759678 | KY825948 |
| L16 | Epimeria sp. 1 | Indonesia | 05° 14.00' S | 133° 00.00′ E | MNHN-IU-2009-2493 | KY907637 | N/A | KY825908 |
| MP10 | Epimeria sp. 1 | Indonesia | 05°15.00 S | 133°01.00 E | MNHN-IU-2009-2491 | KY907648 | KY907501 | KY825939 |
| L6 | Epimeria sp. 2 | Papua New Guinea, Bismarck Sea, Dogreto Bay | 03° 18.00' S | 143° 00.00' E | MNHN-IU-2013-1583 | KY907643 | KY907496 | KY825915 |
| L7 | Epimeria sp. 2 | Papua New Guinea, Brokenwater Bay | 03° 52.64′ S | 144° 40.60 ′ E | MNHN-IU-2013-11797 | KY907645 | KY907497 | KY825916 |
| K10 | Epimeria sp. 3 | Papua New Guinea | 09° 06.00' S | 152° 19.00' E | MNHN-IU-2009-2487 | N/A | КҮ907483 | KY825867 |
| K11 | Epimeria sp. 4 | Papua New Guinea | 04° 16.00' S | 152° 18.00' E | MNHN-IU-2009-2486 | KY907627 | N/A | KY825868 |
| L13 | Epimeria sp. 4 | Papua New Guinea, Solomon Sea, South East of Tuam Island | 06° 04.25′ S | 148° 10.42′ E | MNHN-IU-2013-18091 | KY907635 | КҮ907488 | KY825905 |
| K13 | Epimeria sp. 5 | Taiwan | 24° 08.70′ N | 122° 09.90′ E | MNHN-IU-2009-2520 | КҮ907629 | N/A | KY825870 |
| K14 | Epimeria sp. 5 | Taiwan | 24° 08.70′ N | 122° 09.90′ E | MNHN-IU-2009-252 | КҮ907630 | KY907484 | KY825871 |
| L9 | Epimeria sp. 6 | New Caledonia | 22° 23.40′ S | 167° 21.60′ E | MNHN-IU-2013-1666 | КҮ907644 | КҮ907499 | KY825918 |
| L10 | Epimeria sp. 6 | New Caledonia | 22° 21.80′ S | 167° 20.80′ E | MNHN-IU-2011-6489 | KY907634 | КҮ907486 | KY825903 |
| L12 | Epimeria sp. 7 | Solomon Islands | 07° 42.00' S | 157° 43.00' E | MNHN-IU-2009-2504 | N/A | КҮ907487 | KY825904 |
| MP8 | Epimeria sp. 8 | Indonesia | 05° 46.00′ S | 132° 10.00′ E | MNHN-IU-2009-2495 | KY907657 | N/A | KY825946 |
| K12 | Epimeria sp. 9 | Vanuatu | 16° 02.14′ S | 166° 38.39' E | MNHN-IU-2009-2514 | KY907628 | N/A | KY825869 |
| K8 | Epimeria sp. 10 | Fiji Islands | 17° 16.00′ S | 179° 35.00' W | MNHN-IU-2009-2517 | KY907632 | N/A | KY825901 |
| K15 | Epimeria sp. 11 | New Caledonia | 23° 54.00′ S | 167° 42.00' E | MNHN-IU-2007-2501 | KY907631 | N/A | KY825872 |
| L1 | Epimeria sp. 12 | Papua New Guinea, Bismarck sea, West Kairiru Island | 03° 21.00' S | 143° 26.00' E | MNHN-IU-2013-18106 | КҮ907633 | KY907485 | KY825902 |
| L2 | Epimeria sp. 12 | Papua New Guinea, Woodlark Islands | 09° 09.00' S | 152° 18.00' E | MNHN-IU-2014-3478 | KY907640 | KY907493 | KY825911 |
| L3 | Epimeria sp. 12 | Papua New Guinea, Woodlark Islands | 09° 07.00' S | 152° 14.00' E | MNHN-IU-2014-4270 | KY907642 | N/A | KY825914 |
| MP1 | Epimeria sp. 12 | Papua New Guinea | 09° 09.00' S | 152° 18.00' E | MNHN-IU-2009-2481 | KY907647 | KY907500 | N/A |
| - | | | | | | | | |

| MP2 | Epimeria sp. 12 | Papua New Guinea | 09° 09.00' S | 152° 18.00' E | MNHN-IU-2009-2482 | KY907652 | KY907503 | KY825941 |
|------------|--------------------------------------|------------------------------|--------------|---------------|-------------------|----------|----------|----------|
| MP4 | Epimeria sp. 13 | Indonesia | 05° 15.00′ S | 132° 59.00′ E | MNHN-IU-2009-2496 | KY907653 | N/A | KY825943 |
| L14 | Epimeria sp. 20 | Solomon Islands | 08° 24.40′ S | 159° 22.55′ E | MNHN-IU-2009-2506 | КҮ907636 | KY907489 | KY825906 |
| L15 | Epimeria sp. 14 | Solomon Islands | 08° 24.40′ S | 159° 22.55′ E | MNHN-IU-2009-3479 | KY907649 | KY907490 | KY825907 |
| MP12 | Epimeria sp. 14 | Solomon Islands | 08° 24.40′ S | 159° 22.55′ E | MNHN-IU-2009-2511 | КҮ907650 | N/A | KY825940 |
| MP5 | Epimeria sp. 15 | Solomon Islands | 08° 19.60′ S | 160° 01.95′ E | MNHN-IU-2009-2513 | KY907654 | KY907505 | KY825956 |
| L21 | Epimeria sp. 16 | Solomon Islands | 09° 55.00′ S | 161° 33.00′ E | MNHN-IU-2014-4315 | N/A | KY907495 | KY825913 |
| MP9 | Epimeria sp. 17 | Solomon Islands | 08° 19.60′ S | 159° 22.55′ E | MNHN-IU-2009-2507 | KY907658 | N/A | KY825947 |
| MP3 | Epimeria sp. 18 | Papua New Guinea | 05° 40.00' S | 154° 29.00' E | MNHN-IU-2011-2345 | N/A | KY907504 | KY825942 |
| Н3 | Epimeria loricata | Svalbard, Erik Eriksenstrait | N/A | N/A | RBINS INV. 132637 | KY907625 | KT808709 | KY825846 |
| H4 | Epimeria loricata | Svalbard, Hinlopen | N/A | N/A | RBINS INV. 138044 | KY907626 | KY907482 | KY825847 |
| L18 | Epimeria morronei | Mexico, Pacific | 23° 09.91′ N | 115° 51.00′ W | None | KY907638 | KY907491 | KY825909 |
| L19 | Epimeria morronei | Mexico, Pacific | 29° 20.90′ N | 115°51.00′ W | None | КҮ907639 | KY907492 | KY825910 |
| L20 | Epimeria morronei | Mexico, Pacific | 30° 48.40′ N | 116° 47.80′ W | None | KY907641 | KY907494 | KY825912 |
| L8 | Epimeria sp. 19 | Mozambique | 24° 22.29′ S | 35° 41.86′ E | MNHN-IU-2009-22 | KY907646 | КҮ907498 | KY825917 |
| MP6 | Epimeria sp. 19 | Mozambique, Vizconde de Eza | 24° 02.06′ S | 35° 40.66′ E | MNHN-IU-2009-2527 | KY907655 | N/A | N/A |
| MP7 | Epimeria sp. 19 | Mozambique, Visconde de Eza | 24° 05.28′ S | 35° 41.73′ E | MNHN-IU-2009-2525 | KY907656 | КҮ907506 | KY825945 |
| NOR13 | Epimeria cornigera | Norway | 68° 11.27' N | 14° 59.87' E | RBINS INV. 138043 | KY907659 | КҮ907507 | KY825950 |
| NOR15 | Epimeria cornigera | Norway | N/A | N/A | RBINS INV. 132630 | KY907660 | KT808708 | KY825951 |
| ACANTHONOT | OZOMELLIDAE | | | | | | | |
| B12 | Acanthonotozomoides oatesi | Sub-Antarctic, Shag Rocks | 53° 24.81' S | 42° 40.03′ W | RBINS INV. 132669 | KY907616 | KT808686 | KY825834 |
| F6 | Acanthonotozomoides oatesi | Sub-Antarctic, Shag Rocks | 53° 23.94' S | 42° 40.10' W | RBINS INV. 138046 | KY907622 | KY907480 | KY825843 |
| G9 | Acanthonotozomoides oatesi | Sub-Antarctic, Shag Rocks | 53° 24.53' S | 42° 40.70' W | RBINS INV. 138045 | KY907624 | KY907481 | KY825845 |
| MP19 | Acanthonotozomellidae n. gen. n. sp. | New Caledonia | 23° 43.00' S | 168° 02.00' E | MNHN-IU-2009-2499 | KY907651 | КҮ907502 | N/A |
| G4 | Acanthonotozomellidae n. gen. n. sp. | New Caledonia | 23° 22.76′ S | 167° 51.60' E | MNHN-IU-2009-2497 | КҮ907623 | KT808685 | KY825844 |
| E16 | Acanthonotozomella trispinosa | Antarctica, Larsen B | 65° 57.51' S | 60° 28.15' W | RBINS INV. 132673 | KY907618 | KT808684 | KY825836 |
| DIKWIDAE | | | | | | | | |
| B10 | Dikwa andresi | Sub-Antarctic, Burdwood Bank | 54° 34.04' S | 56° 10.64' W | RBINS INV. 132666 | KY907614 | KY907478 | KY825833 |
| | | | | | | | | |

| B11 | Dikwa andresi | Sub-Antarctic, Burdwood Bank | 54° 34.04' S | 56° 10.64' W | RBINS INV. 132666 | KY907615 | KT808704 | N/A |
|--------------|------------------------------|------------------------------------|--------------|--------------|-------------------|----------|----------|----------|
| STILIPEDIDAE | | | | | | | | |
| A5 | Bathypanoploea schellenbergi | Antarctica, Eastern Weddell Sea | 70° 23.94' S | 08° 19.14' W | RBINS INV. 132623 | КҮ907613 | KT808699 | KY825817 |
| E3 | Alexandrella dentata | Antarctica, Eastern Weddell Sea | 70° 56.52' S | 10° 34.84' W | RBINS INV. 132653 | кү907619 | КҮ907479 | KY825837 |
| E4 | Alexandrella dentata | Antarctica, Larsen A | 64° 54.75' S | 60° 39.01' W | RBINS INV. 132653 | KY907621 | KT808688 | KY825839 |
| B20 | Astyra abyssi | Norway, Vestfjorden | 68° 11.27' N | 14° 59.87' E | RBINS INV. 132658 | KY907617 | KT808694 | KY825835 |
| VICMUSIIDAE | | | | | | | | |
| E30 | Acanthonotozomopsis pushkini | Sub-Antarctic, Shag Rocks | 53° 24.53' S | 42° 40.70' W | RBINS INV. 132656 | КҮ907620 | KT808687 | KY825838 |
| OUTGROUP | | | | | | | | |
| Ex154 | Gnathiphimedia sexdentata | Eastern Weddell Sea | 70° 50.64′ S | 10° 36.11' W | RBINS INV.132752 | KU870835 | KU759609 | KY825841 |
| A4 | Anchiphimedia dorsalis | Antarctica, King George Island | 62° 18.21' S | 58° 39.90' W | RBINS INV. 132625 | KY907612 | KT808690 | KY825816 |
| A2 | Iphimediella cyclogena | Antarctica, Larsen A | 64° 55.58' S | 60° 33.37' W | RBINS INV. 132632 | KY907611 | KT808722 | KY825814 |
| | | | | | | | | |

Table 1. Sampling details for the sequenced *Epimeria* specimens including sample location, geographical coordinates, voucher number and GenBank accession number. "N/A" (not available) indicates unobtainable data. GenBank accession numbers of sequences obtained in this study were indicated in bold.

Table A1. Mean parameter estimates (with standard deviations) and comparison of the fit of different lineage diversification models to the empirical and semi-empirical datasets (assuming 10 %, 50 % and 72 % sampling). LH is the log-likelihood value of the model; the shift times (Ma) are indicated for the models implying discrete shifts in diversification rates; λ is the speciation rate for the constant-rate models and the initial speciation rate for the rate-variable models; μ is the extinction rate; x is the rate change parameter of the density-dependent exponential (DDX) model; K is the carrying capacity parameter of the density-dependent linear (DDL) model. For each model, the Akaike Information Criterion (AIC) was computed. The best-fitting model with the lowest AIC score is highlighted in dark grey. The 95 % confidence intervals of the AIC values are indicated. dAIC is the difference between the AIC score of the model and the AIC score of the best fitting model. Minimal and maximal values for dAIC were calculated based on the AIC 95% confidence intervals. The best-fitting model with the lowest mean AIC score is highlighted in dark grey. The evaluated models for which the dAIC range contains zero are highlighted in light grey. wAIC are the Akaike weights.

| Dataset | Model | LH | Rate shift times | λ | μ | x | К | AIC | dAIC | wAIC |
|-------------------------|--------------|-----------------|--|---|--|--------------|--------------------|---|-----------------------------|---------------------|
| | РВ | 1010 (17.79) | - | 0.34 (0.02) | - | - | - | -2018 (35.58) [-2022.93 ;-2013.07] | 128.13 (117.25 ; 139.01) | 9 10 ⁻²⁹ |
| | BD | 1075.07 (21.46) | - | 0.83 (0.07) | 0.74 (0.08) | - | | -2146.13 (42.92) [-2152.08 ; -2140.18] | 0.00 | 0.62 |
| | DDL | 1008.87 (19.75) | - | 0.34 (0.02) | - | - | 6085859 (1026968) | -2013.74 (39.5) [-2019.21 ; -2008.27] | 132.39 (120.97 ; 143.81) | 1 10 ⁻²⁹ |
| | DDX | 1073.7 (21.54) | - | 0.02 (0.01) | - | -0.58 (0.05) | - | -2143.4 (43.08) [-2149.37 ; -2137.43] | 2.73 (-9.19 ; 14.65) | 0.16 |
| Semi-emp 10 % | Yule2rate | 1063.62 (20.5) | 0.01 (0.01) | 0.34 (0.02) 0.81 (0.61) | - | | - | -2121.25 (41) [-2126.93 ; -2115.57] | 24.88 (13.95 ; 25.85) | 2 10 ⁻⁶ |
| | BD- 1 shift | 1077.7 (21.39) | 19.45 (4.86) | 0.16 (0.04) 0.8 (0.08) | 0.01 (0.04) 0.68 (0.11) | | - | -2142.26 (43.08) [-2148.23 ; -2136.29] | 3.87 (-8.05; 15.79) | 0.09 |
| | BD- 2 shifts | 1077.92 (21.45) | 19.23 (6.04) 7.63 (5.13) | 0.13 (0.06) 0.7 (0.75) 0.78 (0.09) | 0 (0.01) 0.52 (0.75) 0.61 (0.17) | - | - | -2139.83 (42.89) [-2145.77 ; -2133.89] | 6.3 (-5.59 ; 18.19) | 0.03 |
| | BD- 3 shifts | 1079.28 (21.5) | 21.69 (5.18) 12.79 (7.04) 6.5 (5.59) | 0.12 (0.06) 0.38 (0.02) 0.62 (0.5) 0.78 (0.09) | 0 (0.01) 0.14 0.45 (0.54) 0.6 (0.2) | - | - | -2142.57 (43) [-2148.53 ; -2136.61] | 3.56 (-8.35 ; 15.47) | 0.10 |
| | РВ | 14.58 (3.36) | - | 0.12 (0.01) | - | - | - | -27.16 (6.73) [-28.09 ; -26.23] | 0.97 (-0.99 ; 2.93) | 0.17 |
| Causi aura 50 0/ | BD | 15.88 (3.89) | - | 0.16 (0.02) | 0.08 (0.03) | - | - | -27.76 (7.77) [-28.84 ; -26.68] | 0.37 (-1.74 ; 2.48) | 0.22 |
| Semi-emp 50 % | DDL | 14.57 (3.38) | | 0.12 (0.01) | - | - | 1202874 (229958.4) | -25.14 (6.75) [-26.07 ; -24.20] | 4.88 (1.03 ; 4.96) | 0.06 |
| | DDX | 15.86 (3.83) | / / | 0.06 (0.01) | - | -0.21 (0.07) | - | -27.73 (7.67) [-28.79 ; -26.67] | 2.29 (-1.69 ; 2.49) | 0.22 |

| | Yule2rate | 17.06 (3.72) | 0.35 (0.21) | 0.12 (0.01) | - | - | - | -28.13 (7.43) | 0.00 | 0.27 |
|---------------|--------------|---------------|---|--|--|--------------|-------------------|------------------------------------|------------------------|--------------------|
| | BD- 1 shift | 17.36 (3.79) | 15.35 (4.61) | 0.12 (0.1) 0.07 (0.03) | 0.01 (0.02) | | | [-29.16 ; -27.10] -24.73 (7.57) | 3.4 | 0.05 |
| | DD 2 shifts | 10 (/2 71) | 17.00 (4.04) | 0.14 (0.02) | 0.02 (0.03) | | | [-25.78 ; -23.68] | (1.32 ; 5.48) | 8 10 ⁻³ |
| | BD- 2 shifts | 18.6 (3.71) | 17.69 (4.64) 10.31 (6.95) | 0.07 (0.02) 0.13 (0.14) 0.13 (0.03) | 0 (0) 0.05 (0.14) 0.02 (0.05) | - | - | -21.19 (7.41) [-22.22 ; -20.16] | 6.94 (4.88 ; 9) | 8 10 |
| | BD- 3 shifts | 19.59 (3.71) | 8.09 (6.63) 14.13 (6.33) 20.43 (5.94) | 0.07 (0.03) 0.16 (0.28) 0.14 (0.19) 0.13 (0.04) | 0 (0) 0.06 (0.26) 0.09 (0.5) 0.03 (0.1) | - | - | -17.19 (7.42) [-18.22 ; -16.16] | 10.94 (8.88; 13) | 1 10 ⁻³ |
| | РВ | -16.9 (1.93) | - | 0.09 (0) | - | - | - | 35.79 (3.85) [35.26 ; 36.32] | 0.00 | 0.33 |
| | BD | -16.71 (2.07) | - | 0.11 (0.01) | 0.02 (0.02) | - | - | 37.43 (4.14) [36.86 ; 38.00] | 1.64 (0.54 ; 2.74) | 0.15 |
| | DDL | -16.87 (1.9) | - | 0.1 (0.01) | - | - | 506016 (454041.2) | 37.74 (3.8) [37.21 ; 38.27] | 1.95 (0.89; 3.01) | 0.13 |
| | DDX | -16.71 (2.06) | - | 0.07 (0.01) | - | -0.09 (0.06) | | 37.42 (4.13) [36.85 ; 37.99] | 1.63 (0.53 ; 2.73) | 0.15 |
| Semi-emp 72 % | Yule2rate | -15.41 (1.94) | 0.9 (0.3) | 0.1 (0) 0.06 (0.04) | - | - | | 36.82 (3.88) [36.28 ; 37.36] | 1.03 (-0.04 ; 2.1) | 0.20 |
| | BD- 1 shift | -15.32 (2.05) | 4.73 (4.67) | 0.07 (0.03) 0.09 (0.03) | 0 (0.01) 0 (0) | | - | 40.64 (4.1) [40.07 ; 41.21] | 4.85 (3.75 ; 5.95) | 0.03 |
| | BD- 2 shifts | -14.02 (2.01) | 19.01 (4.48) 8.65 (7.99) | 0.06 (0.02) 0.09 (0.08) 0.07 (0.03) | 0 (0) 0.01 (0.03) 0 (0.01) | | - | 44.04 (4.01) [43.48 ; 44.59] | 8.25 (7.16 ; 9.33) | 5 10 ⁻³ |
| | BD- 3 shifts | -13.19 (1.93) | 21.39 (5.46) 17.23 (5.96) 7.41 (7.57) | 0.06 (0.03) 0.06 (0.09) 0.11 (0.08) 0.07 (0.03) | 0 (0) 0.01 (0.03) 0.01 (0.07) 0 (0.01) | - | - | 48.39 (3.86) [47.85 ; 48.92] | 12.6 (11.53; 13.66) | 6 10 ⁻⁴ |
| | РВ | -28.47 | - | 0.078 | - | - | - | 58.94 | 0.00 | 0.31 |
| | BD | -28.47 | - | 0.078 | 0.000 | - | - | 60.94 | 2.00 | 0.12 |
| Empirical | DDL | -28.31 | - | 0.093 | - | - | 120.040 | 60.62 | 1.68 | 0.14 |
| Empirical | DDX | -28.47 | - | 0.082 | - | 0.016 | - | 60.94 | 2.00 | 0.12 |
| | Yule2rate | -26.62 | 2.09 | 0.088 0.027 | - | - | - | 59.24 | 0.30 | 0.27 |

| BD- 1 shift | -26.79 | 2.00 | 0.088 0.029 | 0.001 0.000 | - | - | 63.58 | 4.64 | 0.03 |
|--------------|--------|------------------------|----------------|----------------|---|---|-------|-------|-------|
| BD- 2 shifts | -25.51 | 17.00 2.00 | 0.049 | 0.000 | - | - | 67.02 | 8.08 | 0.005 |
| | | 2.00 | 0.102 0.028 | 0.000 | | | | | |
| BD- 3 shifts | -24.77 | 19.00 17.00 2.00 | 0.058 | 0.000 | - | - | 71.54 | 12.60 | 0.001 |
| | | 2.00 | 0.103 0.029 | 0.001 0.000 | | | | | |
| | | | | | | | | | |

Table A2. Parameter estimates and comparison of the fit of the lineage diversification models from Morlon et al. (2011) to the empirical dataset, assuming sampling fractions of 0.72, 0.5 and 0.1. λ is the speciation rate (when constant) or the initial speciation rate (when variable); α is the rate change parameter for the exponential variation of the speciation rate; μ is the extinction rate (when constant) or the initial extinction rate (when variable); β is the rate change parameter for the exponential variation of the extinction rate. For each model, the Akaike Information Criterion (AIC) was computed. dAIC is the difference between the AIC score of the evaluated model and the AIC score of the best-fitting model. wAIC is the Akaike weight. The best-fitting model with the lowest AIC score is highlighted in dark grey. When the wAIC of the best model < 0.5, the second best-fit is highlighted in light grey.

| Dataset | Model | AIC | dAIC | wAIC | λ | α | μ | β |
|----------------|----------|--------|-------|------|-------|--------|-------|-------|
| | BcstDcst | 245.55 | 2.11 | 0.16 | 0.10 | - | 0.02 | - |
| | BvarDcst | 247.26 | 3.82 | 0.07 | 0.11 | -0.015 | 0.00 | - |
| Compling 72 9/ | BcstDvar | 247.08 | 3.64 | 0.07 | 0.10 | | 0.007 | 0.08 |
| Sampling 72 % | BvarDvar | 249.62 | 6.18 | 0.02 | 0.10 | 0.002 | 0.008 | 0.08 |
| | Bcst | 243.44 | 0.00 | 0.45 | 0.093 | - | = | = |
| | Bvar | 244.88 | 1.44 | 0.22 | 0.11 | -0.015 | - | = |
| | BcstDcst | 245.55 | 1.37 | 0.16 | 0.15 | - | 0.07 | - |
| | BvarDcst | 246.57 | 2.39 | 0.10 | 0.14 | -0.03 | 0.00 | - |
| Sampling 50 % | BcstDvar | 246.65 | 2.47 | 0.09 | 0.14 | - | 0.02 | 0.06 |
| Sampling 50 % | BvarDvar | 249.07 | 4.89 | 0.03 | 0.14 | -0.02 | 0.004 | 0.07 |
| | Bcst | 244.34 | 0.16 | 0.30 | 0.11 | = | = | = |
| | Bvar | 244.18 | 0.00 | 0.32 | 0.14 | -0.03 | - | - |
| | BcstDcst | 245.55 | 2.83 | 0.13 | 0.75 | - | 0.66 | - |
| | BvarDcst | 245.11 | 2.39 | 0.16 | 0.37 | -0.07 | 0.00 | = |
| Compling 10 9/ | BcstDvar | 245.93 | 3.21 | 0.11 | 0.55 | - | 0.39 | 0.01 |
| Sampling 10 % | BvarDvar | 247.59 | 4.87 | 0.05 | 0.42 | -0.06 | 0.09 | -0.05 |
| | Bcst | 259.60 | 16.88 | 0.00 | 0.19 | = | = | = |
| | Bvar | 242.72 | 0.00 | 0.54 | 0.37 | -0.07 | - | - |

Highlights

- A molecular phylogeny of *Epimeria* amphipods from every world's oceans is provided.
- Antarctic and non-Antarctic Epimeria form 2 distinct clades.
- The Antarctic clade arose from a Late Gondwanan ancestor.
- The diversification of the Antarctic *Epimeria* clade appears related to cold-waters.

