THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Cinzia Spencer-Cervato

ABSTRACT

For 30 years the Deep Sea Drilling Project (DSDP) and the Ocean Drilling Program (ODP) have been drilling the ocean floors and retrieving sediment cores. This study presents a relational micropaleontological and stratigraphic database, Neptune, where a selection of the published studies made on these sediments is available. The selected sites and their stratigraphic extent represent a statistically reproducible subset of the whole DSDP and ODP data set as of 1995 (up to Leg 135). Cenozoic sediments from 165 globally distributed holes were dated with age/depth plots using biochronology of four marine plankton groups (diatoms, nannofossils, foraminifera, and radiolarians). Each hole's location is available with paleogeographic coordinates. A taxonomic revision of the 8000+ reported species names was also made. The database is searchable and a variety of routines are available. Data can be exported to produce age range charts, geographic distribution maps, and occurrence charts.

A rigorous evaluation of the database potentials and limitations is presented together with a summary of the published studies that have been carried on with the data. These include stratigraphic studies (diachrony of Neogene plankton, hiati distribution in Cenozoic sediments) and evolution studies (cladogenesis and evolution of one foraminiferal lineage). Unpublished data on macroevolutionary patterns (species longevity and richness, speciation and extinction rates) are presented as example of Neptune's potential for paleobiological research. Finally, some suggestions are presented as to how Neptune can be more fully exploited through the addition of sedimentologic and isotopic data. A variety of critical sedimentologic and paleoceanographic questions could be addressed with this extended database.

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KEY WORDS: Cenozoic, relational database, plankton, evolution, age models

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PLAIN LANGUAGE SUMMARY:

Marine sediments contain the most complete record of the evolution of life on Earth. After the mass extinction event of the Cretaceous/Tertiary boundary, 65 million years' worth of sediments have accumulated on the sea floor. The Deep Sea Drilling Project and its successor, the Ocean Drilling Program, have drilled, retrieved and analysed kilometers of cores, as well as described their paleontological content. The Neptune database was established to compile the most valuable and significant data, and to use them to study the evolution of marine plankton. The global geographic coverage (165 holes), the high number of species described (1400+) from four marine plankton groups, the improved age control on the sediments, and the relatively high sample resolution (a few hundred thousand years) make this relational database the most complete paleontological data set currently available.

The analysis of these data has shown different evolutionary patterns in different plankton groups. On average a plankton species 'survives' 7 to 10 millions of years. Siliceous plankton (diatoms and radiolarians) tend to speciate and become extinct at distinct climatic and oceanographic boundaries independently from their nutritional habits (photosynthetic algae or plankton feeders). On the other hand, calcareous plankton seems to be more independent from these conditions. The results also show that the total number of species preserved in the sediments as fossils (a subset of the total number of species that existed at each given time and location) has gradually increased through time, but has also fluctuated strongly in the last 65 million years perhaps in response to climatic changes. This database has the potential to allow paleontologists to study the complex interactions between marine life and environment at a geological scale.

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1. INTRODUCTION: THE SCOPE OF THE DATABASE AND ORIGINAL PLANNING

Last year (1998) marked the 30th anniversary of the first Deep Sea Drilling Project (DSDP) cruise and the collection of the first cores. The handful of scientists who conceived and initiated this gigantic enterprise in the early 1960s probably did not expect this international project to spur as many controversies and theories on the history of the Earth as it indeed did. At that time, Plate Tectonics, the fundamental theory that unifies most if not all of our geological (and not only geological) knowledge, was still just a controversial hypothesis accepted by only a few scientists. JOIDES (Joint Oceanographic Institute for Deep Earth Science), the program that initiated the DSDP and later the Ocean Drilling Program (ODP), deserves a lot of the credit for the collection and study of the evidence that today practically makes plate tectonics a widely accepted 'truth'.

As a side effect of the wealth of knowledge acquired in these 30 years, scientists have produced an enormous amount of data, so large that I am not aware of any recent estimate after the one done for the first ten years of research (Revelle 1981). Up to recently, all results were first published in reports (also known as 'blue books'). This procedure made most of the raw data available from a centralized and easily accessible printed source. In addition, JOIDES published a CD-ROM containing much of the data produced from the some 1000 holes during the progress of DSDP in electronic format. However, this multitude of data makes sense only to a limited number of scientists that have been involved in their production, and nobody has a concrete overview of what is available. Moreover, the competitiveness of the recent research climate does not encourage the re-evaluation of older data, but leads instead to the production of more new data.

With this background, a group of biostratigraphers at the ETH Zürich initiated the Neptune project in 1990. The group included some veterans from DSDP (Jean-Pierre Beckmann, Katharina von Salis Perch-Nielsen, Hans Thierstein), one participant of the more recent ODP cruises (Dave Lazarus), and some newcomers (Milena Biolzi, Jörg Bollmann, Heinz Hilbrecht, and myself). The project was funded by the Swiss National Science Foundation. The project was, in its initial stages, conceived and led by Dave Lazarus (Lazarus 1994; Lazarus et al. 1995a), while in the later, scientific analysis phase, the effort was carried out by this author (Spencer-Cervato et al. 1993, 1994; Spencer-Cervato and Thierstein 1997; Spencer-Cervato 1998).

The scope of the Neptune project was to evaluate and organize the existing DSDP and ODP data into a relational database that would be accessible to the research community. First, we planned to 'rescue' and compile the micropaleontological information. This information could be used first to establish an updated chronology for selected sites. The micropaleontological data themselves were then to be used for various studies of evolution. The established chronology would also be used to obtain age control on sedimentological and geochemical data. This database would be substantially different from a mere compilation of existing data, as was assembled in the DSDP CD-ROM. The main difference would be in the 'quality control' of the data to be included. Suitable sites would be selected, based on criteria dictated by our experience in biostratigraphy and deep-sea drilling. We decided to limit the number of sites in the database to give preference to an accurate selection and analysis of the data available for each site. We initially planned to include some 100 holes, but this number has been substantially increased in a later phase of the project. The second innovative approach was represented by the search options. The data in the DSDP CD-ROM are not searchable, but are available as a series of gigantic tables with listings of data. As potential end users, we recognized the necessity to create links between the different data sets (e.g., by hole, by age, by geographic location, by fossil group) to optimize the research applications of the database.

In the next chapters, I will provide a description of what is in the Neptune database and how it got there. I will also discuss what we would have liked to do, and why we did not get to it. Some of the published (and in progress) applications of Neptune will be discussed in a separate chapter. I will conclude with some suggestions on possible additions and how Neptune can be used as a tool available to the research community.

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2. THE CONCEPTION OF NEPTUNE AS A STEPPING STONE TOWARD THE MICROPALEONTOLOGIST'S DREAM OF THE IDEAL WORLD

2.1. Stratigraphic and geographic coverage

Marine sediments provide more or less continuous, laterally extensive and correlatable geological archives. The first choice that we made was to limit the database to deep-sea sediments, thereby excluding land sections. Hence the name given by D. Lazarus to the database: Neptune, the Roman god of the sea. The largest amount of data on deep-sea sediments come from ocean drilling, and we began our work by systematically searching through the Initial Reports of the Deep Sea Drilling Project (DSDP) and the Initial and the Proceedings of the Ocean Drilling Program, Scientific Results. Based on a variety of criteria, we rated the holes drilled by DSDP and by ODP up to Leg 135 (the latest leg available in 1995, when I expanded the database to the whole Cenozoic). Ratings were given for each fossil group separately. No rating meant that biostratigraphy was not available, a rating of 'M' (medium) indicated the presence of biostratigraphy limited to a few markers and often the absence of detailed range charts. A rating of 'H' (high) was used for detailed biostratigraphic reports with extensive range charts. Comments on core recovery, preservation, etc. were also added at this point. Given our long-range goal of using the data for a micropaleontological database, we selected to include in Neptune mainly holes that were marked as high or medium priority for more than one biostratigraphic group. Other hole selection criteria included good core recovery, relatively continuous coring, the length of the stratigraphic interval covered, as well as good microfossil preservation. The recognition of magnetostratigraphy, which could be used for further age control, was also determinant in the selection.

To date, Neptune includes stratigraphic data for the whole Cenozoic (last 65 million years). Several reports are available on Cretaceous sediments and some on Jurassic sediments. However, the K/T boundary represents a major stratigraphic boundary that marks a dramatic faunal and floral assemblage turnover. I feel, therefore, justified in the choice of limiting the coverage to the Cenozoic. At the other end of the spectrum, upper Pleistocene and Holocene sediments are not well represented in DSDP and ODP reports and, therefore, in Neptune. This is mainly due to the limited resolution of marine biostratigraphy for recent sediments, the relatively coarse sampling used in most reports and to loss of the upper few meters of sediments in early coring work.

Final additions or changes to the list of holes were done after plotting the geographic location of the selected holes. We aimed to have a broad geographic coverage and at least one complete section for each biogeographic province (Figure 2.1). The coverage of shelf to abyssal sediments was equally considered: the range of water depths



of the sediment/water interface represents a statistically representative subset of all the holes

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drilled by DSDP and ODP as of 1995 (<u>Spencer-Cervato 1998</u>). However, shallow water (shallower than 1000 m) sediments are underrepresented in the DSDP and ODP collection and are, therefore, underrepresented also in Neptune. In total, we did include 165 holes (<u>Table 2.1</u>). More holes would have been desirable and we have possibly excluded holes of considerable importance. This was due to time limitations and the project's goals of creating a 'micropaleontological database'.

2.2. Chronology

The next step consisted in establishing an internally consistent chronology for the selected holes. Because magnetostratigraphic data were available only for some of the holes, biochronology represented the best and often the only way to provide an age model for the holes. Biochronology provides a series of 'calibrated events' which essentially mark the first and last appearance of biostratigraphic markers (taxa). Ideally, these events have been correlated in several locations to an independent stratigraphic method, like magnetostratigraphy or oxygen isotope stratigraphy. These scales have in turn been calibrated to absolute chronology in millions of years through complex procedures. The magnetostratigraphic scale used initially for Neptune was Berggren et al. (1985). We subsequently updated our chronology to Berggren et al. (1995b), which is based on Cande and Kent's magnetostratigraphy (1992, 1995). Berggren et al.'s chronology was chosen because it is the most updated and most comprehensive time scale published to date - it includes biochronological data for several hundred Cenozoic events. Oxygen isotope stratigraphy (in turn calibrated to a magnetostratigraphic scale) was used for only a few of the calibrated events used in Neptune.

Through this two-step approach, numerical ages in million of years (Ma) are given to biostratigraphic events. We assumed that these events are geologically instantaneous and occur simultaneously throughout a given region of the globe (i.e., are globally synchronous and at least regionally widespread, and not dependent on local environment or sediment facies). Berggren et al. (1985) and its recent updates (Berggren et al., 1995a, b) were the source of biochronological events for planktic foraminifera and calcareous nannoplankton. Various regional calibrations were used for siliceous plankton (for radiolarians: Hays and Opdyke 1967; Hays 1970; Theyer et al. 1978; Johnson and Nigrini 1985; Sanfilippo et al. 1985; Goll and Bjørklund 1989; Nigrini 1991; Harwood et al. 1992; Caulet 1991; for diatoms: Barron 1981, 1985a, b; Berggren et al. 1985; Fenner 1984; Koizumi and Tanimura 1985; Gersonde and Burckle 1990; Mikkelsen 1990; and Harwood and Maruyama 1992). Paleogene siliceous plankton biochronology is less well established than the one for the Neogene, so most of the events used were biozonation boundary markers.

Published biochronological events were used to construct the chronology of each

hole. Templates were assembled with all the events that we found in the cited references (Table 2.2). These templates (ASCII files to be used in MS Excel) contained the description of the event, an identification code, and the age interval of the calibration. An excerpt from one of these files is shown in Table 2.3.

2.3. Taxonomy

The articles published in the DSDP and ODP reports are an immense source of evolutionary and biostratigraphic data. Although we were aware of many discrepancies in the subjective nature of taxa and taxonomic names (e.g., <u>Gradstein et al. 1985</u>), we assumed that these factors would be manageable by use of simple synonymy lists in our study. A very extensive taxonomic literature is available for marine plankton, and taxa and nomenclature are quite well defined among the most common microfossil groups. This can be used to reasonably standardize taxonomic usage. Thus if taxon **Ab** is called **Ab** by one author but **Bb** by another, we could standardize the data by creating an equivalence **Bb** = **Ab** in the database. Moreover, the holes that we selected had been extensively studied for biostratigraphy and some of them represented classical micropaleontological studies. We, however, had to assume that taxon names in all the selected holes were uniformly used, in other words, that taxon **Ab** described in Hole 289 was identical to taxon **Ab** described in Hole 747A. More than 8800 taxon names have been used in the selected holes.

2.4. Biostratigraphy

Most micropaleontological studies are limited to one or perhaps two fossil groups. Biostratigraphic studies in DSDP and ODP reports include diatoms, radiolarians, calcareous nannoplankton, planktic and benthic foraminifera, dinoflagellates, silicoflagellates. We decided to consider only planktic organisms and out of the several groups described in the Reports, we selected the four groups that are most abundant in deep-sea sediments, most regularly described in the biostratigraphic literature, and for which extensive event calibration is available: diatoms, radiolarians, calcareous nannoplankton and planktic foraminifera. This selection includes two siliceous (diatoms and radiolarians) and two calcareous (nannoplankton and foraminifera) plankton groups, and at the same time two phytoplankton (diatoms and nannoplankton) and two zooplankton (radiolarians and foraminifera) groups. This approach has several advantages: it would allow us to compare evolutionary trends in multiple groups, but mainly it allowed us to have a better biochronological control on the age models. Planktic foraminifera are probably the most used microfossils for biostratigraphy, and with this approach we were able to compare their resolution and accuracy to the other groups.

The templates were filled in with actual occurrences of the events for each hole. We went through the published range charts or lists of markers and located the events present in the templates. The list of references to the individual reports is given in Table 2.4. Each event was normally recorded as occurring between two samples

within the stratigraphic section. Samples were either recorded as meters below seafloor (mbsf) or as actual sample names, in core-section-interval within section in centimeter format. The sample names were then automatically translated into mbsf by the plotting software. No systematic attempt was made to search the general literature for additional stratigraphic data, although biostratigraphic data for some critical holes (e.g., DSDP 558 and 563) were extracted from charts published outside the DSDP reports. The creation of biostratigraphic files from the templates was initially subdivided among the project participants. In the later phase of the project, I was solely responsible for this task. This eliminated some of the discrepancies in the event identification due to subjective interpretations of range charts in terms of First Occurrence (FO) and Last Occurrence (LO).

The first and last occurrence of a taxon were identified when the taxon was not recorded in two or more samples above or below the first or last recorded occurrence. Because the precision of the actual FO or LO depends on the sample spacing, we recorded each event as the stratigraphic interval between the two samples bracketing the event.

Paleomagnetic stratigraphy was recorded as a set of paleomagnetic polarity interval identifications, as given by the original author. In some cases, it became necessary to revise the original identification scheme to achieve an optimal fit between biostratigraphy and paleomagnetic polarity patterns. However, this was usually apparent only when the events were plotted.

The biostratigraphic files prepared for each group were pasted together and used in the construction of age models. There is a varietyare several of methods available to process stratigraphic event data, including Shaw's plots (Shaw 1964) and Probabilistic Stratigraphy (Hay 1972). However, the most used method of stratigraphic correlation for deep-sea sediments is the age vs. depth plot method. A plot is made of the depth occurrences of previously age-calibrated events in each hole and a line is drawn to correlate depth to age. Although various curve-fitting methods can be used, we have chosen to manually fit a series of straight line segments of varying slopes to the data.

To handle the large volume of data plotting and analysis we used a special-purpose, age-depth plotting program written by Lazarus (1992). The program, written for Macintosh computers, reads the stratigraphic data files and produces an age- vs. versus-depth scatter plot of the data points. The program allows us to draw a line of correlation through the points interactively on the computer screen. Automatic correlation methods were attempted but proven unreliable because they were too easily affected by data outliersdata outliers too easily affected them. The manual construction of the line of correlation allows us to take into consideration recovery gaps and changes in preservation or lithology that may affect the reliability of the age vs. depth plot. The age models are thus subjective and, with a few exceptions, the scatter of data allows for two or more possible interpretations. The use of two or more biostratigraphic groups was intended to minimize the bias introduced by an **a priori** selection of 'good' or 'bad' events.

Age models were initially constructed by several project participants. To eliminate discrepancies in the selection of the line of correlation due to subjective preferences, all Neogene DSDP age models were subsequently revised by Dave Lazarus (Lazarus et al. 1995a) and later by myself (after the addition of Paleogene data and ODP holes, and the update of the chronology). A personal rating of the quality of the age models is given in Table 2.1. Although the results of all these efforts still do not guarantee that the age models are optimally reliable, I hope that they represent a far more consistent and updated data set than available prior to the beginning of the project.

The established chronology provided age control on the 30,000 samples described in the DSDP and ODP reports for the selected holes. Information on the micropaleontological content of these samples is available as range charts. These charts give information on the presence or absence of a taxon, and usually describe its abundance. Properly formatted MS Excel range charts were either extracted from the DSDP CD-ROM by the Neptune database program, typed by us, or provided directly from ODP (Table 2.4). These were then imported into Neptune and represent the bulk of data available. We planned to use this information for various studies (species occurrence patterns, longevity and diversity, identification of temporal distribution of biogeographic provinces) which are described in <u>Chapter 4</u>.

The age/depth plots and the age models (text files) are given in the <u>Appendix A</u>. The stratigraphic data files used to construct the age vs. depth plots, are not published here because of space considerations and the complexity of having such a large number of files and links. They are, however, available from the author.

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Figure 2.1. Location of DSDP and ODP holes in the Neptune database. Dots mark holes with one or more hiati, squares holes with a continuous stratigraphy, within the resolution of the chronology.



Table 2.1.

| | | | | | | Priority | | | R | RCs exist: | | : | |
|-----|------|----------|-----------|-------|---|----------|---|---|----------------|------------|---|---|---|
| Leg | Hole | Latitude | Longitude | Pmag. | D | N | R | F | Loc Quality | D | N | R | F |
| 7 | 62 | 1.9 | 141.9 | N | | | Η | | Р | - | Y | Y | Y |
| 7 | 62A | 1.9 | 141.9 | N | | | Η | | P | - | Y | Y | Y |
| 7 | 63 | 0.8 | 147.9 | N | | | Μ | | Р | - | Y | Y | - |
| 7 | 63A | 0.8 | 147.9 | N | | | Μ | | P | - | Y | Y | - |
| 7 | 63B | 0.8 | 147.9 | N | | | Μ | | P | - | Y | Y | - |
| 7 | 64 | -1.7 | 158.6 | N | | | Μ | | М | - | Y | Y | Y |
| 7 | 64A | 1.5 | 158.4 | N | | | Μ | Η | М | - | Y | Y | Y |
| 13 | 125 | 34.6 | 20.4 | N | | | | Η | Р | - | Y | - | Y |
| 13 | 132 | 40.3 | 11.4 | N | | | | Η | G | - | Y | - | Y |
| 14 | 141 | 19.4 | -24.0 | N | | | | Η | | | | | |

Table 2.2. (Continued next 4 pages.)

| Planktic Foraminifera | | | Berggren et al. 1995a,b | | | | | | |
|--------------------------|--|----------|-------------------------------|------------|-------|--|----------|--------------|------------|
| Group | Event Name | Plotcode | Young Age | Old Age | Group | Event Name | Plotcode | Young Age | Old Age |
| F | TOP Globoquadrina pseudofoliata | Tpft | 0.22 | | F | TOP Globorotalia miozea | Tmza | 15.9 | |
| F | BT G. hirsuta | Bhir | 0.45 | | F | BT Praeorbulina circularis | BPcr | 16 | |
| F | BT G. flexuosa | Bflx | 0.4 | | F | BT Praeorbulina glomerosa | BPgl | 16.1 | |
| F | BT Bolliella calida | BBca | 0.22 | | F | BT Globigerinoides diminutus | Bdim | 16.1 | |
| F | TOP G. flexuosa | Bflx | 0.068 | | F | BT Praeorbulina curva | BPcv | 16.3 | |
| F | TOP Globorotalia tosaensis | Ttos | 0.65 | 1 | F | BT Praeorbulina sicana (= G.es bispher.) | BPsc | 16.4 | |
| F | BT Pulleniatina finalis | BPfn | 1.4 | | F | BT Globorotalia miozea | Bmza | 16.7 | |
| F | TOP Globigerinoides fistulosus | Tfst | 1.6 | | F | TOP Globorotalia zealandica | Tzld | 17.3 | |
| F | TOP Globigerinoides obliquus extremus | Tobe | 1.7 | | F | TOP Globorotalia pseudomiozea | Tpmz | 16.6 | |
| F | BT Globorotalia truncatulinoides | Btrc | 2 | | F | BT Globorotalia birnageae | Bbng | 16.7 | |
| F | TOP Globorotalia exilis | Texl | 2.15 | | F | TOP Catapsydrax stainforthi | TCst | 17.2 | |
| F | reappear.Pulleniatina (local) | RAPu | 2.3 | | F | BT Globorotalia zealandica | Bzld | 17.3 | |
| F | TOP Globorotalia miocenica | Tmio | 2.3 | | F | BT Globorotalia pseudomiozea | Bpmz | 17.3 | |
| F | TOP Neogloboquadrina atlantica | Tatl | 2.41 | | F | TOP Globorotalia semivera | Tsmv | 17.3 | |
| F | TOP G. puncticulata | Tpun | 2.41 | | F | TOP Globorotalia incognita | Tigt | 16.4 | |
| F | TOP Globorotalia pertenuis | Tprt | 2.6 | | F | TOP Catapsydrax dissimilis | TCds | 17.3 | |
| F | TOP Globorotalia multicamerata | Tmtc | 3.09 | | F | BT Globorotalia praescitula | Bpsc | 18.5 | |
| F | TOP Globoquadrina altispira | Talp | 3.09 | | F | TOP Globoquadrina dehiscens f.spinosa | TGqds | 17.9 | |

| F | TOP Sphaeroidinellopsis seminulina | Tsem | 3.12 | F | BT Globigerinoides altiaperturus | Batp | 20.5 |
|---|--|-------|------|--|--|--------|------|
| F | BT Globigerinoides fistulosus | Bfst | 3.33 | F | BT Tenuitella munda | Bmun | 21.4 |
| F | TOP Spheroidinellopsis spp | TSdp | 3.25 | F | TOP Globorotalia kugleri | Tkgl | 21.5 |
| F | BT Spheroidinella dehiscens s.s. | BSdd | 3.25 | 5 F BT Globoquadrina dehiscens f.spinosa | | BGqds | 22.2 |
| F | BT Globorotalia inflata | Bifl | 3.25 | F | TOP Globoquadrina globularis | TGqg | 22.8 |
| F | TOP Globorotalia conomiozea | Tcmz | 3.25 | F | BT Globoquadrina dehiscens | BGqd | 23.2 |
| F | BT Globorotalia tosaensis | Btos | 3.35 | F | BT Globorotalia incognita | Bigt | 21.6 |
| F | BT Globorotalia crassula | Bcrl | 3.3 | F | TOP Globoturborotalita angulisuturalis | Tags | 21.6 |
| F | TOP Pulleniatina (local) | TPlt | 3.45 | F | TOP Globorotalia pseudokugleri | Tpkg | 21.6 |
| F | BT Globorotalia pertenuis | Bprt | 3.45 | F | BT Globigerina euapertura | Beua | 23.8 |
| F | BT Globorotalia miocenica | Bmio | 3.55 | F | B Globorotalia kugleri | bGku | 23.8 |
| F | TOP Globorotalia margaritae | Tmgt | 3.58 | F | T Globorotalia mendacis | tGme | 23.8 |
| F | TOP Pulleniatina primalis | TPpr | 3.65 | F | B Globigerinoides primordius (common) | bGpr-c | 24.3 |
| F | Pulleniatina s->d | sdP | 3.95 | F | B Globigerinoides primordius (rare) | bGpr-r | 26.7 |
| F | TOP Globigerina nepentes | Tnep | 4.2 | F | T Globorotalia opima | tGop | 27.1 |
| F | TOP Pulleniatina spectabilis | TPsp | 4.2 | F | B Globigerina angulisuturalis | bGas | 29.4 |
| F | BT Globorotalia crassaformis s.s. | Bcrs | 4.5 | F | T Globigerina angiporoides | tGap | 30 |
| F | TOP Globigerinoides seiglei | Tsgl | 4.7 | F | T Globigerina ampliapertura | tGam | 30.3 |
| F | BT Globorotalia puncticulata | Bptc | 4.5 | F | T Globorotalia cerroazulensis | tGce | 33.8 |
| F | TOP Globorotalia cibaoensis | Tcbn | 4.4 | F | T Hantkenina | tHan | 33.7 |
| F | BT Spheroidinella dehiscens | BSdd | 5.2 | F | T Porticulasphaera semiinvoluta | tPsi | 35.3 |
| F | BT Globorotalia sphericomiozea | Bsphc | 5.6 | F | T Morozovella spinulosa | tMsp | 38.1 |
| F | BT Globorotalia pliozea | Bpli | 5.6 | F | B Porticulasphaera semiinvoluta | bPsi | 38.4 |

| F | BT Globorotalia tumida | Btum | 5.6 | F | T Subbotina frontosa | tSfr | 39.3 |
|---|--|-------|------|---|------------------------------------|------|-------------|
| F | BT Pulleniatina spectabilis | BPsp | 5.6 | F | T Globigerapsis beckmanni | tGbe | 40.1 |
| F | BT Globigerinoides conglobatus | Bcgb | 5.8 | F | B Globigerapsis beckmanni | bGbe | 40.5 |
| F | BT Globorotalia cibaoensis | Bcbn | 7.8 | F | T Acarinina bullbrooki | tAbu | 40.5 |
| F | TOP Globoquadrina dehiscens | TGqd | 5.8 | F | B Globorotalia pomeroli | bGpo | 42.4 |
| F | TOP Globorotalia lenguaensis | Ting | 6 | F | B Globigerapsis index | bGin | 42.9 |
| F | BT Globorotalia margaritae | Bmgt | 6 | F | B Morozovella lehneri | bMle | 43.5 |
| F | BT Pulleniatina primalis | BPpr | 6.4 | F | T Morozovella aragonensis | tMar | 43.6 |
| F | Neogloboquadrina acostaensis s->d | acsd | 6.2 | F | B Globorotalia possagnoensis | bGpg | 46 |
| F | BT Globorotalia menardii form 5 (dext.) | Bmn5 | 6.4 | F | B Planorotalites palmerae | bPpa | 50.4 |
| F | Neogloboquadrina acostaensis d->s | acds | 6.6 | F | B Morozovella aragonensis | bMar | 52.3 |
| F | Neogloboquadrina atlantica d->s | atds | 6.8 | F | B Morozovella formosa | bMfo | 54 |
| F | BT Globorotalia conomiozea | Bcmz | 7.12 | F | T Morozovella velascoensis | tMve | 54.7 |
| F | TOP Globorotalia menardii form 4 (sin.) | Tmn4 | 7.4 | F | T Planorotalites pseudomenardii | tPps | 55.9 (?) |
| F | BT Globorotalia suterae | Bsut | 7.8 | F | B Morozovella velascoensis | bMve | 60 |
| F | BT Globorotalia juanai | Bjua | 8.1 | F | B Morozovella pusilla | bMpu | 61 (?) |
| F | BT Candeina nitida | Bcnit | 8.1 | F | B Morozovella conicotruncata | bMco | 60.9 |
| F | BT Globigerinoides extremus | BGex | 8.3 | F | B Morozovella angulata | bMan | 61 |
| F | BT Globorotalia plesiotumida | Bplt | 8.3 | F | B Morozovella uncinata | bMun | 61.2 (?) |
| F | BT Neogloborotalia humerosa | Bhum | 8.5 | F | B Planorotalites compressus | bPco | 63 (?) |
| F | TOP Neogloboquadrina nympha | Tnym | 10.1 | F | T Globotruncana | tGtc | 65 |
| F | BT Neogloboquadrina acostaensis | Bac | 10.9 | | | | |

Table 2.3: Example of format used in the biostratigraphic templates. 'Young age' and 'Old age' allow entering two different values and obtain an age range for the event. The last two columns to the right are left blank and will be filled with depth information from each hole. B represents the first appearance datum (FAD), T the last appearance datum (LAD).

| Hole | 2.00 | dd/mm/yy | Author of file | Comments | | ts |
|-------|----------------------|----------|----------------|------------|--------------|-----------------|
| Group | Event name | Code | Young age | Old age | Top Depth | Bottom Depth |
| N | B D. tamalis | B Dta | 3.8 | | | |
| N | B D. asymmetricus | B Das | 4.2 | | | |
| N | T A. primus | T Apr | 4.8 | | | |
| N | B C. rugosus | B Cru | 5 | 5.23 | | |

Table 2.4. References to biostratigraphic files present in Neptune. The 'Group' column indicates the plankton group (D: diatoms; F: foraminifera; N: nannofossils; R: radiolarians) stratigraphy or magnetostratigraphic data (M) presented in the reference.

Vol. Author(s) Group 7 Brönnimann, P.; Resig, J., 1971 F 7 Martini, E.; Worsley, T., 1971 N 7 Riedel. W.R.; Sanfilippo, A., 1971 R 13 Gartner, S., Jr., 1973 N 13 Bukry, D., 1973 N 13 Ryan, W. B. F.; Hsü, K. J., 1973 F 14 Beckmann, J. P., 1972 F 14 Roth, P. H.; Thierstein, H. R., 1972 N 16 Kaneps, A. G., 1973 F 16 Dinkelman, M.G., 1973 R 16 Bukry, D.; Foster, J.H., 1973 D 18 Ingle, J. C., Jr., 1973 F 18 Kling, S. A., 1973 R 18 Schrader, H. J., 1973 D 18 Wise, S. W., 1973 N 19 Koizumi, I., 1973 D 19 Ling, H. Y., 1973 R 19 Worsley, T. R., 1973 N 19 Echols, R. J., 1973 F 19 Akiba, F., 1986 (vol. 87) D 22 Johnson, D. A., 1974 R 22 Gartner, S., Jr., 1974 N 22 McGowran, B., 1974 F 22 Berggren, W.A.; et al., 1974 F 22 Schrader, H.-J., 1974 (vol. 24) D 24 Sanfilippo, A.; Riedel, W. R., 1974 R 24 Roth, P. H., 1974 N 24 Schrader, H. J., 1974 D 24 Vincent, E.; et al., 1974 F 24 Heiman, M.E.; et al., 1974 F 26 Boltovskoy, E., 1974 F 26 Thierstein, H. R., 1974 N 28 McCollum, D. W., 1975 D 28 Chen, P. H., 1975 R 28 Burns, D. A., 1975 N 28 Kaneps, A. G., 1975 F 29 Edwards, A. R.; Perch-Nielsen, K., 1975 N

29 Petrushevskaya, M. G., 1975 R 29 Jenkins, D. G., 1975 F 29 Schrader, H. J., 1976 (vol. 35) D 30 Holdsworth, B. K., 1975 R 30 Shafik, S., 1975 N 31 Koizumi, I., 1975 D 31 Ellis, C. H., 1975 N 31 Ling, H.Y., 1975 R 33 Johnson, D. A., 1976 R 33 Martini, E., 1976 N 33 Takayanagi, Y.; Oda, M., 1976 F 39 Boersma, A., 1977 F 39 Perch-Nielsen, K., 1977 N 39 Sanfilippo, A.; Nigrini, C., 1995 R* 39 Fenner, J., 1978 (supplement) D 40 Jenkins, D. G., 1978 F 40 Proto Decima, F.; et al., 1978 N 40 Toumarkine, M., 1978 F 40 Pisias, N.G.; Moore, T.C. Jr., 1978 R 41 Bukry, D., 1978 N 41 Krasheninnikov, V. A.; Pflaumann, U., 1978 F 41 Krasheninnikov, V. A.; Pflaumann, U., 1978 F 41 Johnson, D. A., 1978 R 41 Schrader, 1978 D 43 Okada, H.; Thierstein, H.R., 1979 N 47 Blechschmidt, G., 1979 N 48 Murray, J.W., 1979 F 48 Müller, C., 1979 N 49 Poore, R. Z., 1979 F 49 Ling, H. Y., 1979 R 49 Steinmetz, J. C., 1979 N 49 Martini, E., 1979 N 49 Schrader, H. J., 1979 D 55 Takayama, T., 1980 N 55 Ling, H. Y., 1980 R 55 Koizumi, I., 1980 D 56/7 Thompson, P. R., 1980 F 56/7 Reynolds, R. A., 1980 R 56/7 Shaffer, B. L., 1980 N 56/7 Keller, G., 1980 F 56/7 Harper, H. E., Jr., 1980 D 56/7 Sakai, T., 1980 R 56/7 Barron, J. A., 1980 D 58 Okada, H., 1980 N 58 Sloan, J., 1980 R

59 Martini, E., 1981 N 59 Theyer, F.; Lineberger, P., 1981 R 59 Heiman, M.E., 1981 F 60 Ellis, C. H., 1982 N 60 Kling, S. A., 1982 R 61 Premoli Silva, I.; Violanti, D., 1981 F 61 Thierstein, H.R.; Manivit, H., 1981 N 61 Sanfilippo, A.; et al., 1981 R 63 Barron, J. A., 1981 D 63 Wolfart, R., 1981 R 63 Poore, R. Z., 1981 F 66 Stradner, H.; Allram, F., 1982 N 66 McMillen, K.J., 1982 R 67 Thompson, P. R., 1982 F 67 Muzylöv, N., 1982 N 67 Westberg, M. J.; Riedel, W. R., 1982 R 67 Harper, H. E., Jr.; et al., 1982 D 67 Jousé, A.P.; et al., 1982 D 68 Kent, D. V.; Spariosu, D. J., 1982 M 68 Riedel, W. R.; Westberg, M. J., 1982 R 68 Keigwin, L. D., Jr., 1982 F 68 Sancetta, C., 1982 D 68 Kent, D. V.; Spariosu, D. J., 1982 M 71 Krasheninnikov, V. A.; Basov, I. A., 1983 F 71 Salloway, J. C., 1983 M 71 Wise, S. W., 1983 N 71 Gombos, A.M.; Ciesielski, P.F., 1983 D 71 Weaver, F. M., 1983 R 71 Ciesielski, P. F., 1983 D 72 Berggren, W.A.; et al., 1983 M 72 Pujol, C., 1983 F 72 Berggren, W.A.; et al., 1983 F 72 Gombos, A.M., Jr., 1983 D 72 Pujol, C.; Duprat, J., 1983 F 73 Percival, S. F., Jr., 1984 N 73 Poore, R. Z., 1984 F 73 Smith, C.C.; Poore, R.Z., 1984 F 73 Tauxe, L.; et al., 1984 M 73 Gombos, A.M., Jr., 1984 D 74 Boersma, A., 1984 F 74 Jiang, M.-J.; Gartner, S., 1984 N 77 Lang, T.H.; Watkins, D.K., 1984 N 80 Snyder, S. W.; Waters, V. J., 1985 F 80 Müller, C., 1985 N 80 Townsend, H. A., 1985 M

80 Pujol, C.; Duprat, J., 1985 F 80 Pujos, A., 1985 N 80 Labracherie, M., 1985 R 81 Krumsiek, K.; Roberts, D. G., 1984 M 81 Backman, J., 1984 N 81 Huddlestun, P. F., 1984 F 81 Baldauf, J.G., 1984 D 81 Westberg-Smith, M.J.; Riedel, W.R., 1984 R 82 Parker, M.E.; et al., 1985 N 82 Bukry, D., 1985 N 82 Miller, K.G.; et al., 1985 N,F* 82 Miller, K.G.; et al., 1994 N,F,M* 85 Weinreich, N.; Theyer, F., 1985 M 85 Saito, T., 1985 F 85 Labracherie, M., 1985 R 85 Baldauf, J. G., 1985 D 85 Barron, J. A., 1985 D 85 Gartner, S.; Chow, J., 1985 N 85 Nigrini, C. A., 1985 R 85 Pujos, A., 1985 N 86 Koizumi, I.; Tanimura, Y., 1985 D 86 Heath, G. R.; et al., 1985 M 86 Bleil, U., 1985 M 86 Monechi, S., 1985 N 86 Morley, J. J., 1985 R 90 Martini, E., 1986 N 89/90 Lohman, W. H., 1986 N 90 Ciesielski, P. F., 1986 D 89/90 Jenkins, D. G.; Srinivasan, M. S., 1986 F 90 Barton, C. E.; Bloemendal, J., 1986 M 89/90 Caulet, J. P., 1986 R 92 Romine, K., 1986 F 92 Knüttel, S., 1986 N 93 Muza, J. P.; et al., 1987 N 93 Lang, T.H.; Wise, S.W., Jr., 1987 N 93 Applegate, J.L.; Wise, S.W., Jr., 1987 N 93 Ma'alouleh, K.; Moullade, M., 1987 F 93 Saint-Marc, P., 1987 F 93 Nishimura, A., 1987 R 93 Gombos, A.M., Jr., 1987 D 93 Canninga, G.; et al., 1987 M 94 Jenkins, D.G., 1987 F 94 Weaver, P. P. E., 1987 F 94 Clement, B. M.; Robinson, F., 1987 M 94 Baldauf, J. G., 1987 D

94 Takayama, T.; Sato, T., 1987 N 94 Westberg-Smith, M.J.; et al., 1987 R 95 Miller, K.G.; Hart, M.B., 1987 F* 95 Palmer, A.A., 1987 R 95 Valentine, P.C., 1987 N 95 Abbott, W.H., 1987 D 105 Knüttel, S.; et al., 1989 N 105 Firth, J.V., 1989 N 105 Aksu, A.E.; Kaminski, M.A., 1989 F 105 Baldauf, J.G.; Monjanel, A.-L., 1989 D 105 Lazarus, D.; Pallant, A., 1989 R 105 Clement, B.M.; et al., 1989 M 108 Manivit, H., 1989 N 108 Weaver, P.P.E.; Raymo, M.E., 1989 F 108 Tauxe, L.; et al., 1989 M 113 Spiess, V., 1990 M 113 Pospichal, J.J.; Wise, S.W., 1990 N 113 Wei, W.; Wise, S.W., 1990 N 113 Abelmann, A., 1990 R 113 Lazarus, D.B., 1990 R 113 Gersonde, R.; Burckle, L.H., 1990 D 113 Stott, L.D.; Kennett, J.P., 1990 F 114 Fenner, J., 1991 D 114 Crux, J.A., 1991 N 114 Madile, M.; Monechi, S., 1991 N 114 Nocchi, M.; et al., 1991 F 114 Hailwood, E.A.; Clement, B.M., 1991 M 114 Hailwood, E.A.; Clement, B.M., 1991 M 115 Okada, H., 1990 N 115 Rio, D.; et al., 1990 N 115 Premoli Silva, I.; Spezzaferri, S., 1990 F 115 Johnson, D.A., 1990 R 115 Fenner, J.; Mikkelsen, N., 1990 D 115 Vincent, E.; Toumarkine, M., 1990 F 115 Schneider, D.A.; Kent, D.V., 1990 M 117 Spaulding, S., 1991 N 117 Sato. T.; et al., 1991 N 117 Nigrini, C., 1991 R 117 Spaulding, S.A.; et al., 1991 F* 117 Hayashida, A.; Bloemendal, J., 1991 M 119 Huber, B.T., 1991 F 119 Wei, W.; Thierstein, H.R., 1991 N 119 Wei, W.; Pospichal, J.J., 1991 N 119 Caulet, J.P., 1991 R 119 Lazarus, D.B., 1992 (vol. 120) R

119 Baldauf, J.G.; Barron, J.A., 1991 D 119 Sakai, H.; Keating, B.H., 1991 M 119 Keating, B.H.; Sakai, H., 1991 M 120 Heider, F.; et al., 1992 M 120 Inokuchi, H.; Heider, F., 1992 M 120 Aubry, M.-P., 1992 N 120 Wei, W.; Wise, S.W., 1992 N 120 Bergaren, W.A., 1992 F 120 Berggren, W.A., 1992 F 120 Harwood, D.M.; Maruyama, T., 1992 D 120 Takemura, A., 1992 R 120 Abelmann, A., 1992 R 120 Lazarus, D.B., 1992 R 122 Siesser, W.G.; Bralower, T.J., 1992 N 122 Galbrun, B., 1992 M 122 Zachariasse, W.J., 1992 F* 122 Tang, C., 1992 M 125 Ciampo, G., 1992 N 125 Xu, Y.; Wise, S.W., Jr., 1992 N 125 Milner, G.J., 1992 F 127 Rahman, A., 1992 N 127 Brunner, C.A., 1992 F 127 Alexandrovich, J.M., 1992 R* 127 Koizumi, I., 1992 D 127 Hamano, Y.; et al., 1992 M 130 Leckie, R.M.; et al., 1993 F* 130 Chaisson, W.P; Leckie, R.M., 1993 F 130 Takayama, T., 1993 N 132 Premoli Silva, I.; et al., 1993 N,F 132 Sager, W.W.; et al., 1993 M 133 Gartner, S.; et al., 1993 N 133 Kroon, D., 1993 F* 134 Zhao, X.; et al., 1994 M 134 Staerker, T.S., 1994 N 134 Perembo, R.C.B., 1994 F 134 Weinheimer, A.L.; et al., 1994 R 135 Nishi, H.; Chaproniere, G.C.H., 1994 F* 135 Quinterno, P.J., 1994 N*

* used only in biostratigraphy files, not available in Neptune

3. THE REALISATION OF NEPTUNE - THE REAL WORLD IS WORSE THAN WE THOUGHT

3.1. Stratigraphic and geographic coverage

The geographic distribution of the 165 holes included in Neptune is uneven. In some areas there is a very detailed coverage, like for example in some parts of the Antarctic Ocean (Fig. 2.1). On the other hand, no holes from the central north and southeastern Pacific Ocean are present in Neptune. (Holes from ODP Leg 145 now provide a transect across the north Pacific.) The mid- and high latitudes in the southern hemisphere and the tropical regions of the Atlantic Ocean are also not well represented. This is due in part to the uneven coverage of DSDP and ODP cruises and in part to the selection made for Neptune, which preferentially included holes with good biostratigraphy.

Each year, ODP organizes five to six drilling cruises which result in as many published **Scientific Results**. Although not all cruises retrieve micropaleontologically significant material, many of them provide a detailed biostratigraphy and data relevant to the scopes of Neptune. The present geographic coverage of holes in Neptune has been last updated in 1995 (Leg 135). Since then more than twenty-five volumes of **Scientific Results** have been published. From the beginning of the project, we were faced with the need to maintain a balance between keeping up with the new data produced by ODP and the need to analyze the data already in Neptune for biostratigraphic or micropaleontological studies. At present, I have decided to keep Neptune at its current, acceptable but not optimal, size in order to complete some of the studies that we had planned. If it will be decided to update Neptune in the future, it will be necessary to:

- select suitable holes from Leg 136 onward, and for the selected holes:
- compile biostratigraphic files and construct age models;
- import the core depth file and the age model file for each hole;
- download from ODP the available range charts;
- format the range chart files to make them compatible with Neptune;
- import the range chart file;
- update the species name list with the new names eventually present in the range charts.

Another limitation of the database is given by the often incomplete often-incomplete recovery of sediments (Fig. 3.1). Before the advent of hydraulic piston coring, few continuously recovered sections were available. Core recovery has drastically improved in the more recent ODP holes but



sediment loss at core breaks is still common even in continuously cored sections (<u>Farrell and Janecek</u> 1991).

In addition, there is an uneven distribution in the temporal coverage of the sections. Whilst Plio-Pleistocene sections are very well represented



Figure 3.1.

in Neptune (as they are in ODP holes overall), the detail of stratigraphic coverage decreases for older time periods, as naturally expected from the drilling procedure (Fig. 3.2, Spencer-Cervato 1998). This might be interpreted as a need to recover more Miocene and older sections, but this pattern actually reflects the number of **studied** sections and not simply the recovered sections. Therefore, I believe that the



problem does not lie only in the 'quantity' of older sections drilled, but also in the 'quality' of their stratigraphy. The reliability of the stratigraphy provided for a section depends strongly on the availability of good calibrations, and these are currently available mainly for Neogene sediments. Figure 3.2 also shows that the number of

Num

well-studied sections does not decrease gradually and regularly with age, but shows peaks (around 2 Ma) and plateaus (e.g., between 20 and 32 Ma). This likely reflects the relative, unequal attention given to the Cenozoic stratigraphy through the history of DSDP and ODP.

3.2. Chronology

For the database, we have chosen to use a comprehensive biochronology based on deep-sea sections, therefore not considering land sections, which represent the type localities where stratigraphic series were first described. This may represent a limitation in the achieved biochronological calibration. The precision of the ages determined with the age models depends on various factors, some subjective and nonquantifiable, and some, like sample spacing, accuracy of biostratigraphic calibration, or core recovery, that can be quantified. A conservative estimate of the age model precision of 0.36 m.y. was determined for Neogene sediments (<u>Spencer-Cervato et al. 1994</u>). For Paleogene sediments it is about 0.66 m.y. (twice the average sample spacing).

Another important factor is the quality of the age model. The Neogene DSDP age/depth plots that we have published so far (<u>Spencer-Cervato et al. 1993</u>; <u>Lazarus et al. 1995a</u>) are a good example of the range of reliability of the line of correlation. The subjective ranking given in <u>Table 2.2</u> varies from very poor or poor (wide scatter of events, straight line of correlation drawn across the middle of the cloud), to moderate (some scatter of a limited number of events, various possible lines of correlation), to good or excellent (40% of the holes: very good agreement of the event ages, abundant events to constrain the line of correlation, good agreement between

magnetostratigraphy and biostratigraphy). Several factors can cause the scatter of events observed in most age/depth plots: reworking, downhole contamination, incorrect entry in the biostratigraphy file, typos in the range charts, diachrony of the calibrated event, , and sample spacing. Whilst most of these causes can be easily double-checked, diachrony is guite difficult to assess. The assumption of 'globally synchronous events' which is at the base of biochronology is validly established only for some selected, well documented events (e.g., Hays and Shackleton 1976; Thierstein et al. 1977; Backman and Shackleton 1983; Wei 1993; Spencer-Cervato et al. 1994). It is likely that more complete data collection and documentation would lead to the identification of more globally synchronous events. But in most cases, a calibration is valid only for the more or less restricted biogeographic province where it is done, and only a few events are truly globally synchronous, within the precision of the method adopted for calibration. The need for localized calibrations has long been known for siliceous plankton stratigraphy, but it is not widely accepted by biostratigraphers using calcareous plankton. To minimize this factor, we intentionally used multiple regional calibrations for diatoms and radiolarians. Even with this approach, the scatter is sometimes too large to provide a reliable line of correlation. For nannofossils and foraminifera only one general (low latitude) calibration is available (Berggren et al. 1985, 1995a, b). The advantage of this calibration is that it is based on several sites, while most of the regional calibrations are based only on one hole. An estimate of the diachrony/synchrony of Neogene events was done with a subset of the holes currently present in Neptune (Spencer-Cervato et al. 1994). This study indicated that calcareous nannofossils provide the most reliable biostratigraphic events, as they are mostly cosmopolitan and, if diachronous, the age margin is relatively small.

Very few sections are actually continuous, and long stratigraphic gaps are common (<u>Spencer-Cervato</u> <u>1998</u>) (Fig. 3.3). Two-thirds of the selected holes contain at least one hiatus, and on average they each contain three hiati of various lengths (<u>Fig.</u> 2.1). The presence of these hiati results in



artificially older or younger ages for the samples adjacent to the gap. This does not allow one to automatically ('blindly') search the database for e.g.,such information as species ages, but requires that every output is be checked and compared with the age models.

The final and probably most necessary improvement of the chronology of Neptune is given by the life-timelifetime of the biochronology selected for the age model calibration. We initially based the age models on <u>Berggren et al. (1985)</u>. An updated magnetostratigraphy was published later (Cande and Kent 1992) but it did not provide the combination of biochronology and magnetostratigraphy available from <u>Berggren et al. (1985)</u> work. We thus decided to continue using <u>Berggren et al. (1985)</u> throughout the first phase of the project (DSDP Neogene sediments). However, ten

years after the first biochronology compilation, a new updated biochronology was published (Berggren et al. 1995b) and the chronology of Neptune became suddenly outdated. The iterations to update Neptune's chronology were greatly helped by additional programming of Neptune by Dave Lazarus and an auxiliary computer program (not part of Neptune and written **ad hoc** by Bernhard Brabec) which created a correlation function between the old and the new master biochronology. This function was applied to all age model files and new revised age models were created. . Then, all biostratigraphy files were updated using a 'find - replace' routine with lookup tables (i.e.: if code in column 3 is equal to xYwz, replace age in column 4 with corresponding value in lookup table). While we could directly use the new calibrations for calcareous plankton as lookup tables, it was necessary to recalibrate to the new time scale all regional templates used for siliceous plankton events. Finally, before the new age models could be imported into Neptune, all the age/depth plots were redone by myself and eventually adjusted to fit the new event ages.

3.3. Taxonomy

Among the other reasons mentioned above, if a bio-event recorded in a specific hole plots far outside the area where the line of correlation can be drawn, it could be due to its taxonomic identification. Many authors have put together the hundreds of range charts that were used for Neptune and not all agree in the detailed taxonomic identification of all the 8800+ taxa included in Neptune. Indeed, taxonomic identification is subjective. The time pressure under which biostratigraphers are during a leg is also an important limiting factor in the number of species described in a range chart, which is often limited to biostratigraphic markers. The extent to which this taxonomic problem has affected the data in Neptune can be judged by experts in particular cases but cannot be easily quantified.

Starting from the biostratigraphy filesbiostratigraphic records assembled for the chronology, we assumed that the taxon associated with one event and described in the range chart was the one we were looking for. Further, we needed to consider the occurrence of synonyms. It sometimes happens that the name used by one author for a taxon corresponds either to a different taxon according to another author, or that a different name is used by a second author for this specific taxon (synonymy). For example, the foraminifer species **Globorotalia truncatulinoides** has been also called **Truncorotalia truncatulinoides**. To account for this, we have used the literature, personal experience and extensive consultation with taxonomic experts to identify valid taxon names. Three thousand of the 8810 names listed in Neptune (Fig. 3.4) are

considered valid (i.e., are legal names in the framework of the ICZN and ICBN, and are known to be real to at least one of the experts). Synonyms to these valid names were then identified (with the corresponding valid name). They constitute 31% of the total number of names. In several cases we could not unequivocally identify a specific name and marked it as 'unknown' (15% of all names). Only 43 names (0.5%) were considered invalid. This information is available in the 'Species Names' table of Neptune. The synonymizsation is subjective (the initials of the person who identified each species name is also given in the 'Species



Names' table) and the names list does not at all pretend to be a thorough or complete taxonomic revision of marine plankton. It merely represents a working table that gives us a first approximation of plankton taxonomy. A 'real' taxonomic database would need complete taxonomic descriptions (with history) for each taxon and a series of images to illustrate them. Cathy Nigrini, Jean-Pierre Caulet, and Dave Lazarus are currently working on a detailed taxonomic database for radiolarians, but it is well beyond the scopes of Neptune to even attempt anything like this for all groups. The taxonomic list also needs continuous update: every time a new hole is added to Neptune, the biostratigraphic range charts carry with them new names, sometimes several ones. These need to be added to the 'Species Names' list and identified as valid or not.

3.4. Biostratigraphy

There is an uneven distribution in the number of reports by plankton group in Neptune. Over 60% of the 225 articles from which data for Neptune have been extracted (Table 2.4) are on calcareous plankton, almost equally distributed between nannofossils and foraminifera. Radiolarians follow with about 21% and diatoms trail

with only 16%. At the same time, biostratigraphic work on siliceous plankton is underrepresented in Paleogene sections, and most often limited to the Oligocene and younger sections (Fig. 3.5). This unevenness represents a bias for evolution studies where we would like to compare calcareous and siliceous plankton occurrences. Whether this distribution represents the average abundance of fossil plankton in deep-sea deposits or is instead the reflection of staffing decisions by DSDP and ODP is yet to be determined.



One of the limitations of Neptune as a comprehensive micropaleontological database is given by our decision to include only four plankton groups. The DSDP and ODP

Reports include many articles on benthic foraminifera, silicoflagellates, dinoflagellates as well as palynology. At the moment, there are no plans to include their occurrence data in the database, which in itself would not be a huge task.





Figure 3.1. Total core recovery in percentage in the holes included in Neptune.



Figure 3.2: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 3.2. Age distribution of the sections included in Neptune.



Figure 3.3: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 3.3. Age distribution of hiati in the sections included in Neptune.





Figure 3.4. Number of species names in each plankton group included in Neptune.

Figure 3.5: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window





http://www-odp.tamu.edu/paleo/1999_2/neptune/fig3_5.htm (1 of 2) [10/26/1999 10:02:35 AM]

Figure 3.5: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

| o pper ongoeste | TO MALOUZOAANA | OpperLosene | THREADERING | TOWATTOPAILA | TOTACAMENTA |
|-----------------|----------------|-------------|-------------|--------------|-------------|
| | | | | | |

4. RESEARCH PROJECTS CARRIED OUT WITH NEPTUNE, WHAT THEY TOLD US, AND RECOMMENDATIONS FOR THE FUTURE

4.1. The database structure and search capabilities: a tool to find out what we do or don't know

The Neptune database currently provides rapid retrieval of age information on 165 DSDP and ODP holes; taxonomically corrected species lists and other taxonomic information for calcareous nannofossils, planktic foraminifera, diatoms and radiolarians for the entire Cenozoic; paleogeographic location of the 165 holes (paleolatitude and paleolongitude); extensive distributional data for these fossil groups (e.g., biogeographic occurrence information, computerized microfossil range charts) (Fig. 4.1). The design and implementation of the database software have been described in Lazarus (1994) and in an unpublished guide (Lazarus, personal



Figure 4.1.

commun., 1996). These will be used for this description, with the updates of the data tables based on the present status of the database (after the most recent upgrade).

4.1.1. Overview of database structure. Neptune is designed as a relational database. Macintosh computers and 4th Dimension[™] database software (4D) are used to run the database (<u>Lazarus 1994</u>). The database is implemented as several relational tables that contain (as of February 1998) close to 500,000 records.

Import procedures for range chart data as well as search procedures are available. The search procedures can locate all reported occurrences of any taxon or combination of taxa, automatically identifying occurrences recorded under synonymous names. Searches can also be used to locate other relevant information, such as general hole information, sample age, species occurrences, etc. Commercial mapping software (e.g., Atlas[™]) is used to plot locations of species occurrences, using a Neptune-generated plotting data file with latitude and longitude. A 'composite age range chart' program can also be used with an appropriately formatted file generated by a Neptune search (Lazarus 1994).

Neptune was created as a relational database where the data are separated into simple tables, with relational links between the tables. The structure of the database is shown in Figure 4.2. Five data tables hold the primary data: stratigraphic occurrence data for taxa ('Bug Data' table);



taxonomic data on species' names ('Taxonomy'); biogeographic data on species' occurrences ('Taxa by Hole'); geologic age information ('Age Models'); and paleogeographic information ('Geographic Info'). Paleo-water depths are available for selected samples and have been published



(Spencer-Cervato 1998). As this information is not available for all samples, it is not currently included in Neptune.

The database maintains a strict separation between the primary observational data (occurrences of named taxa at specific depths in holes) and the interpreted meaning of the data (i.e., the species to which the name belongs - 'Taxonomy' - or the age of the section at a specific depth - 'Age Models'). These tables can be in fact modified repeatedly, but the observations remain constant (Lazarus 1994).

The stratigraphic occurrence data form the core of the database (over 380,000 records of 'Bug Data'). A typical range chart is decomposed into a minimum of one species occurrence in one sample. Further data separation is achieved by putting all information about samples and species into separate tables ('Sample Data' and 'Taxonomy'). Samples and species are represented in the 'Bug Data' table only by internal codes, linked to the more detailed records in other tables.

4.1.1.1. Species names ('Taxonomy' table). All names in the database are identified by a separate entry in this table. Each occurrence, including misspellings and questionable names (e.g., A. deflandrei?), is given as a separate entry. Each is identified by up to three words (genus, species, subspecies or gualifier). The gualifier is generally used to identify questionable entries, marked with the letter 'Q'. Each entry is uniquely identified by a 'Taxon code', a combination of nine characters originally given by DSDP. This code is central to the functioning of the database, as it provides links to the other tables. The first five characters are letters, all upper case. The first letter identifies the fossil group (D for Diatoms, N for Nannofossils, etc.). The next four letters are characteristic of the genus. The last four characters are numbers and give the species number in the genus. DSDP started with 0010 and incremented by units of 10 for each new species name. ODP does not use codes to identify species names. Thus, we created new codes for new species names that occur in ODP range charts. To avoid any overlap, we have used the same 5 letters to identify the group and the genus, but started with 5010 to number the new species (DSDP never had numbers higher than 2500).

Every entry has other information attached to it. The Status (or validity) field is a single uppercase letter which states that the name is V-alid, a S-ynonym to another name (with corresponding taxon code entered in the 'synonymous to' field), I-nvalid, Q-uestionable, or U-nknown. A G is used to indicate a genus-level name. Every name has also an author code (initials of person responsible for the entries in the Status field) and a date (mm/dd/yy). Comments of any length are also entered in the 'Comments' field.

Additional species' records are available in the 'Species by Hole' table which comes directly from the DSDP data set. The current table, reformatted from the original data set to save space, contains simply a Taxon Code and a Hole field.

4.1.1.2. Age Models and Hole summary data ('Age Models' and 'Hole Data'

tables). The 'Age Models' table contains the age model developed for each hole with range chart available in the database. The age model was constructed by a broken line composed of straight segments, which can be horizontal in the case of hiati. The extremes of the segments are identified by age and depth and entered in the corresponding field, next to the 'Hole' field. To keep track of which age model is being used in the database, each age model's time of creation date stamp is entered automatically in the 'Hole Data' table ('Age model version'). This latter table contains a variety of information, including latitude and longitude, water depth, ocean basin, hole length and recovery, etc. Holes that have an age model (and therefore range chart data) have an entry in the 'Age Model version' field and a ranking for each fossil group (originally used to select holes).

4.1.1.3. Sample Info ('Sample Data' table). Most of the fields in this table are created directly from computer files or by Neptune. Each sample described in each range chart is identified by a unique digital code and is specific for one fossil group. This means that, if in one sample (identified as depth in a hole (mbsf), but also in three separate fields as core-section-depth interval format - grouped in Figure 4.2 under 'Sample Name') both diatoms and radiolarians ('Taxonomic Group' field) were described, this sample would be described twice in Neptune, each time with a different 'Sample Code'. The age of the sample is derived from the 'Age Model' table through a relational link. If available, information on the preservation and abundance of the specific fossil group in that sample is also given.

4.1.1.4. Paleogeographic data ('Geographic Info' table). With the addition of Paleogene range charts, I considered it necessary to locate species occurrences in their appropriate paleogeographic position. For this purpose, I used a PC-based program kindly provided by Alan Smith (Cambridge University) which uses finite rotations. The program is based on published reconstruction data (Euler rotations and their ages) used to move a given site relative to Africa and then reposition that site in paleomagnetic coordinates (Smith, personal commun., 1997). The input file contained present latitude and longitude: paleolatitudes and paleolongitudes were determined for each hole at 5 m.y. intervals. This approximation was necessary to simplify the entry of these data into Neptune, but I believe that it does not significantly affect the already approximated estimate of paleolatitude and paleolongitude made by the finite rotations program. These paleogeographic data, with hole and age, were imported into a separate table ('Geographic Info') and the Sample Code used to link it to other tables.

4.1.2. Importing data into Neptune (range charts and age models). Data can be imported into Neptune by the 'administrator' (this function is not available in the 'user'
mode or with the runtime version of 4D). Most of the DSDP range charts were imported directly from the DSDP CD-ROM data, but ODP data need to be imported as individual spreadsheet format files. Each procedure creates automatically new sample records for each sample in the range chart data ('Sample Data') and new 'Bug Data' records for each non blank cell of occurrence data in the range chart. No ages are assigned in this procedure and all sample ages are set at zero. Only when the age model is imported, a corresponding age is recorded in the age field of the 'Sample Data'.

Each range chart file needs to hold data for one hole and one fossil group only. Each sample must be entered in one row in a 'leg-hole-core-section-first depth-second depth' format (e.g., 113-689-B-2H-1-115-116). These data are automatically entered in the corresponding fields in the 'Sample Data' table. The depth in mbsf is derived from the 'Core Data' table, where the core depth files for each hole are imported as soon as a hole is selected. Species names must be entered as Species Codes (9 characters, e.g., DACTI0020). Every species code present in the spreadsheet must be already available in the 'Taxonomy' table. The 'import from spreadsheet' procedure in Neptune automatically checks each DSDP Code in the spreadsheet and if it encounters a code that is not present in the 'Taxonomy' table, the procedure aborts.

Age information is present in Neptune in two forms. The 'Age Model' table actually holds all the line of correlations (age models) for each hole. Age for samples are calculated from the line of correlations and stored as calculated fields in the 'Sample Data' table. This calculation is done only once, when the age model is read into the database, and is automatic. Only one age model can be imported at a time. To update an existing age model, it is sufficient to read in the new file and the old ages will be automatically replaced by ages based on the new line of correlation.

4.1.3. Report capabilities and external graphics. Data can be extracted from the database in a variety of ways. The results can be then saved as export files, that can eventually be used with other programs. Procedures that search for taxa, in either the 'Bug Data' table of stratigraphic occurrence information, or in the 'Taxa by Hole' table of biogeographic information, create lists of Taxon Codes (from 'Taxonomy') to search for. These lists include the taxonomic name/s requested by the user, but are supplemented by lists of synonyms to these names. Users can edit these lists to fine-tune searches.

In addition to export formats for statistics and spreadsheet packages (usually in ASCII format), the database exports data in formats specific for two types of graphic data display. Data on the location of specific DSDP/ODP sites can be plotted in a map form using Atlas[™] (WTC Scientific). The most recent version of this program for Macintosh computers does not run reliably on PowerPCs and the use of the PC IBM-compatible version (which



rigure 4.5.

can use the same cross-platform file) is

recommended. A custom application creates

graphic displays of occurrence data for taxa, plotted by age and hole ('Age Range Charts', <u>Lazarus 1994</u>) (Fig. 4.3).



4.1.4. Searching the database. The simplest way to search Neptune is by using the built-in 4D 'Search Editor' (under the 'Select' menu). Any of the tables previously described (Species Names, Hole Info etc.) can be selected from the list in the small window that automatically appears when Neptune is started. This shows a window with all the records in that table. The 'Search Editor' function displays a dialogue window which shows the fields available in the table. Only fields in bold can be selected and additional search criteria (equal to, contains, less than etc.) added. The

results of the search are displayed in a few seconds (Fig. 4.4).

More complex procedures, such as a 'Bug Data' search, allow to locate range chart data about one or more taxa. These predetermined procedures can be selected with the 'Execute Procedure' function under the 'Special' menu (Fig. 4.5). The 'Bug Data Search' procedure first shows a search editor window for 'Species Names' and waits for a taxon entry. This can be formulated as 'Species - is equal to - **name'** or done directly with DSDP codes (Fig. 4.6). This procedure locates all taxa matching the entered criteria, as well as other taxa identified in



the database as synonyms for any of these. The user can then select one or all of the identified taxon names and click the 'done' button at the bottom of the window (Fig. 4.7). The procedure then searches the 'Bug Data' table to locate all records for this list of taxa. This search is done using indices, and only takes a few seconds (Fig. 4.8).



The procedure informs the user via a dialogue box how many records have been found, and then presents the search editor window a second time. At this point the user can enter any other

criteria, such as only samples with ages greater than 0 (i.e., holes with age models), or from holes from a specific geographic location (Fig. 4.9). The 'search in selection' box (lower left corner) is automatically marked allowing to search only among the already identified occurrences (and not the whole database!). The user should then click 'ok' to proceed. The procedure will refine the selection according to these secondary criteria, and present the user with a list (Fig. 4.10). This list can also be edited to refine the selection. Lastly, the user clicks 'done' to exit the procedure. The selected records can then be printed, exported to disc, or summarized in a report.

A search procedure is also available to the automatic search for the ages of all samples recording several taxa given in a list (and their synonyms). This 'Batch Search'

('BugDataSearchBat') allows for the automatic operation of the series of procedures described above (Species name selection, identification of synonyms, bug data search, restriction to holes with age models, sorting of samples by age). It produces one separate output file for each name, as well as a cumulative file. This procedure was used to obtain species longevity data (described below).



Figure 4.8.

Alternatively, samples can be sorted by latitude and longitude to obtain ranges of geographic distribution of taxa through time (e.g., to identify cosmopolitan or endemic taxa).

4.2. Paleontological research based on Neptune: plankton evolution



In this and in the next section (4.3), I am presenting a summary of published paleontological and stratigraphic research conducted with Neptune, as well as some unpublished data on macroevolution. Neptune's potential for paleontological research has been, so far, only marginally exploited. In spite



Figure 4.9.

of the limitations outlined in <u>Chapter 3</u>, the database provides the opportunity for large-scale macroevolutionary studies that could go well beyond presently available studies (e.g., <u>Jablonski</u> 1993; Kammer et al. 1997). The age control and

time resolution, combined with the taxonomic information at species level on four distinct plankton groups, make Neptune a high quality data set.

Currently, the two studies we published on evolution were focused on the evolution of one foraminifer species and were based on Plio-Pleistocene sediments, which are represented in a large number of holes in Neptune (Lazarus et al. 1995b;

Spencer-Cervato and Thierstein 1997). The goal of these studies was to document

patterns of evolution of a new species (the planktic foraminifer, **Globorotalia truncatulinoides)** from its ancestors and to identify speciation and migration in distinct biogeographic provinces by using the tests' morphometry. In addition, we attempted to determine the environmental conditions (water depth, thermal structure of water column) at the time of speciation or immigration with stable isotope geochemistry. Whether changes in these environmental conditions were a determinant factor in the speciation or migration, even after these detailed studies, remains still speculative.

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Figure 4.10.

For these studies, Neptune was used in the selection of sites by identifying the occurrences by hole and FADs/LADs (first appearance datum/last appearance datum) of the species and its ancestors. An age range chart was produced from the search for all **G. truncatulinoides** and related species occurrences (Fig. 4.3). This was used



to identify the oldest first occurrences and to have an overview of the age distribution, which shows a distinct diachrony (Fig. 4.11). This search lead to the selection of suitable DSDP and ODP sites from which samples were requested. The samples were then analyzed morphometrically and isotopically to determine patterns of evolution (in this case, cladogenesis or phylogenetic branching) and species migration (Lazarus et al. 1995b; Spencer-Cervato and Thierstein 1997).

The earliest first occurrences are found in several sites in the southwest Pacific. Gradual cladogenesis was documented in this region during the late Pliocene in sympatric or parapatric populations (<u>Lazarus et al. 1995b</u>). Based on qualitative observations, similar but younger, gradual transitions had been reported from other areas of the world's oceans. Therefore, the hypothesis arose that this gradual evolutionary branching might have occurred in response to changing environments at different times in different ocean areas. To evaluate this hypothesis, we studied the morphological transitions of the three taxa, using image analytical techniques, in several deep-sea sections from various areas, identified with Neptune (Spencer-Cervato and Thierstein 1997). The morphometric analyses showed that **G**. **truncatulinoides** evolved between 2.8-2.3 Ma sympatrically in large populations from its ancestor **G**. **crassaformis** in the southwest Pacific. Differentiated morphotypes of **G**. **truncatulinoides** subsequently immigrated into the Indian and Atlantic Oceans between 2.3 and 1.9 Ma. Our morphometric data show these younger appearances outside the southwest Pacific to be punctuated, and representing migration events (Spencer-Cervato and Thierstein 1997).

One of the most crucial, yet elusive issues in evolution is the role played by the environment in the appearance of a new species or its extinction. Planktic foraminifera are ideally suited for these studies because of the large populations, widespread occurrence of tests in marine sediments, relatively large size that allows for detailed identification with traditional microscopic techniques, abundance of information on living populations and their habitats, conceivably rapid colonization of biogeographic provinces under suitable conditions, and the possibility to reconstruct these conditions (e.g., water depth, temperature, nutrients) with stable isotope geochemistry.

Globorotalia truncatulinoides is an ideal species for the study of the environmental conditions at the time of speciation. We hypothesized that the global cooling of surface waters, coinciding with the northern hemisphere glaciation, led to the formation of oceanographic barriers that could have retarded the expansion of **G**. **truncatulinoides** up to 2.3 Ma. At this time, a relative warming and subsequent transgression could have spurred the migration from the southwest Pacific into the Indian Ocean, possibly through the Indonesian passage. A direct link between the speciation and surface water changes linked to the northern hemisphere glaciation has not been proven so far and seems unlikely. In fact, stable isotope data in **G**. **truncatulinoides** and its ancestors indicate that the three species' depth habitat preferences remained unchanged through the speciation and migration of **G**. **truncatulinoides** and that all three species were dominantly deep-dwellers, in agreement with their present environmental preferences (Spencer-Cervato and Thierstein 1997).

One of the original goals of the Neptune project was to perform macroevolutionary studies. Macroevolution is a major area of paleontology that developed during the 1970s and 1980s, inspired by the apparent success of the taxic approach to evolutionary patterns (Smith 1994). Macroevolution covers various concepts and processes. These studies differ from the previously described, 'microevolutionary' ones - which concentrated on the heritable variations of a population composed of one species and its immediate ancestors - mainly in the scale. Macroevolution studies

large-scale patterns of diversification and extinction arising from processes active at or above the species level (<u>Smith 1994</u> and citations therein). Some workers have instead defined macroevolution as the extrapolation of microevolutionary processes into geological time (e.g., <u>Levinton 1988</u>). <u>Smith</u> (1994, Chapter 4) presents a comprehensive review of macroevolutionary concepts and theories.

Existing studies mainly consider fossil records of marine invertebrates in high hierarchical groupings (orders, families) with low stratigraphic age resolution (e.g., Jablonski 1993) from punctual, geographically restricted sources. Amongst the various causes of artifacts in macroevolutionary patterns, sampling resolution seems to be an important biasing factor (e.g., Alroy 1998). The chronological control of Neptune and its large amount of paleontological data with taxonomic accuracy at species level, combined with its comparably high sampling resolution (on average, 185 k.y. for the Neogene (Spencer-Cervato et al. 1994) and 330 k.y. for the Paleogene) holds promise for potentially significant contributions to this field of paleontology. Because the quality of the results of macroevolutionary studies is strongly dependent on a sound basis of chronology and taxonomy, we have first exploited the stratigraphic data set (see Section 4.3) and revised the taxonomy of the paleontological records (Section 3.3). We have thus left the study of macroevolutionary patterns in marine plankton to the final phase of the project.

I am presenting here some examples of data searches conducted with Neptune to answer some typical paleobiological questions. They cover the overall longevity and speciation/extinction distribution of Cenozoic marine plankton. I am purposely leaving the discussion and interpretation of these data to an absolute minimum. My goal is in fact to show what type of data can be obtained from the database and the potential of Neptune for paleobiological research.

To optimize the diverse paleontological data set in Neptune, the analyses have been conducted separately for the four plankton groups and the results interpreted in terms of similarities or differences among the groups. The data presented here are based on more than 1400 valid species names (Table 4.1), and include the occurrences of their synonyms. The output of every species' search consisted of their oldest first appearance and their youngest last appearance. Every result was checked to eliminate false entries caused by, for example, occurrences near hiati, typos, occurrences reported in one single sample, etc. Species that were reported only in one hole were not considered to eliminate the bias of single geographic data points. Finally, one table was produced for each group including the species name, the number of times it had been reported in a sample, the location (paleolatitude and paleolongitude) and age of its first appearance, and the location and age of its last occurrence. From these ages, the species' longevity was calculated.

• What is the distribution of the longevity of plankton species? Are there substantial differences or similarities among the four plankton groups?

Figure 4.12 shows the species' longevity

in the second second

distribution of the four groups with a 1 m.y. resolution. Comparing the four groups, three simple observations can be made: (1) all groups show an asymmetric, unimodal distribution, with a mode around 7 m.y. (diatoms and radiolarians), 14 m.y. (foraminifera) and 19 m.y. (nannofossils), and a tail towards higher longevity values; (2) the median for all distributions is around 10 m.y., except for diatoms, where it is around 7 m.y. - comparing these values with the mode, the peak of the



distribution of diatoms is narrower and has a higher symmetry than the other groups; (3) a few phytoplankton species (diatoms and nannofossils) are very longevous (more than 40 m.y.), whilst zooplankton species (foraminifera and radiolarians) live all less than 43 m.y. These observations point to similarities between phyto- or zooplankton in one case, and between siliceous or calcareous plankton in another. However, they also show that diatoms are quite distinct from the other groups.

• What is the average species' longevity? Are there substantial differences between extinct and extant species' longevities?

Table 4.2 shows the average longevity (and standard deviations) of both extant and extinct plankton species. It is noticeable that the longevity of extinct species is consistently shorter than the one of extant species. This could be due to the different sizes of the populations considered (less than 30% of the species are extant), which might also explain the larger standard deviations of extant species' longevity. Alternatively, this could be the effect of differential preservation. Or it could be caused by the artificial boundary set at the beginning of the Cenozoic - the data might include species originated in the Mesozoic, giving them a shorter-than-real longevity. However, only a very small number of species (e.g., the extant nannofossils **Braarudosphera bigelowii** and **Scapholithus fossilis**, the extinct nannofossil **Placozygus sigmoides**; <u>Perch-Nielsen 1985</u>) are reported also from the Mesozoic. Statistically, they should not significantly affect the data set.

For demographic reasons, one would expect a gradually decreasing longevity instead of the asymmetrical peaks shown in Fig. 4.12. Is this lower-than-expected number of short-lived species an artifact of the analysis or a real signal? On the other hand, the differences seen in the longevity data could be real, suggesting for example that species which evolved in the Neogene (the majority of the extant species) are more likely to live longer. One can only speculate on the cause of this, such as larger surface water temperature gradients linked to growth of ice caps in polar regions? However, the Neogene climate mode, characterized by abrupt shifts from glacials to interglacials and vice versa, would seem to provide stressful environmental conditions that intuitively should increase species turnover, i.e., shorter longevities. This question requires further analyses (e.g., longevity plots at selected critical times) before a viable hypothesis can be formulated. To help answer the various open hypotheses on species' longevities, an important factor that should be considered is the geographic distribution of species throughout their duration. This parameter allows to identify endemic versus cosmopolitan species and is an important factor in ecological studies. By comparing this parameter with species longevity, one would test if a species restricted to a narrow geographic region is more likely to survive longer than a globally widespread species, or vice versa.

• Are there periods in the Cenozoic with a high concentration of species' appearances or extinctions? Are there geographically defined speciation centers or survival refugia?

Appearance and extinction rates were calculated for the four groups to eliminate the bias of the sample size (Wei and Kennett 1983). The rates are calculated as the ratio between the number of extinctions or appearances and the total diversity (number of species) in each 1 m.y. time slice. The Cenozoic appearance rates are shown in Fig. 4.13. Appearances are widespread throughout the Cenozoic and no specific time interval is characterized by anomalously high appearance rates, with the exception of the Paleocene. The graphs show that diatom, radiolarian and foraminifer species appeared all during the Cenozoic, with 100% peaks in the Paleocene, while only 50% of the nannofossil species present in the first million year of the Cenozoic appeared then the remaining 50% existed already in the Mesozoic (see above for some examples). The apparent late appearance of radiolarians in the early Cenozoic is probably an artifact of the data set: no radiolarian reports are available for the Paleocene (Fig. 3.5).

Average appearance rates are less than 10% and only rarely reach 30%, and are characterized by



short fluctuations with a somewhat random frequency. In some instances (e.g., at 61 Ma, 35 Ma, 10 Ma), peaks of appearances in one group correspond to peaks in other groups, but no consistent pattern is apparent.

Extinction rate values are much lower than appearance rates and show a more random distribution (Fig. 4.14). Diatoms show a distinct peak in extinctions at the Paleocene/Eocene boundary which is not clearly reproduced in the other groups. On the other hand, radiolarians, nannofossils and foraminifera show a minor peak around the Oligocene/Miocene boundary (25-22 Ma), while all groups (with the exception of radiolarians) show exceptionally high extinctions in the past 3 m.y.

It is interesting to notice that these trends do not



correspond to peaks in appearances (Fig. 4.13), but there seems to be a time lapse of a couple of million years between the peaks in extinctions and appearances as the two curves are mostly out of phase. The increase in extinctions in the last 5-7 Ma might be related to the onset of highly variable environmental conditions, which apparently did not cause a corresponding increase in the rate of species' appearances.

Speciation centers and survival refugia are discrete geographic regions with high concentrations of appearances or extinctions. These are often associated with particularly favorable or stressed environmental conditions and may be limited by biogeographic or oceanographic boundaries (e.g., Jablonski 1993). One simple way to identify these regions is by plotting the latitude of the location of the earliest first appearance or the latest last appearance (Fig. 4.15). The latitudinal distribution

of FADs (and of LADs, not shown

here, but with an identical pattern to FADs) is clearly different in siliceous and calcareous plankton. Appearances and extinctions of diatom and radiolarian species are concentrated in three belts, around the equator and at mid- to high northern and southern latitudes respectively. These belts are bound by well established nutrient boundaries, like e.g., the polar front. This pattern also reflects the present distribution of siliceous



plankton in marine sediments (e.g., <u>Leinen et al. 1986</u>), suggesting that the environmental preferences of these organisms did not change through time. A different scenario is presented by calcareous plankton groups, whose appearances (and extinctions) are more uniformly distributed throughout the latitudinal range.

• How did plankton diversity change through time? How do the patterns for the four plankton groups compare?

Even with the limitations summarized in <u>Chapter 3</u>, I attempted to estimate the distribution of plankton species' diversity during the Cenozoic. This 'partial' diversity, limited mainly to occurrences of biostratigraphic markers and biased by the low number of extensive range charts published for DSDP and ODP holes, is still a very comprehensive estimate, even though not a 'real' diversity. I present here some preliminary results based on the data included in Neptune.

The total species richness for the four plankton groups was calculated at one million-year intervals. To eliminate the bias of the uneven distribution of the number of sections in Neptune (progressively more sections in younger times, Fig. 3.2), I have normalized the diversity by dividing it by the total number of sections in each time interval. The



results are shown in Figure 4.16. The normalized diversity patterns shown by siliceous plankton are quite similar and clearly distinct from what is shown by calcareous plankton. The two distinct patterns shown by the siliceous and calcareous plankton groups are exactly out of phase, with diatoms and radiolarians showing maximum diversity in the Oligocene to Recent, when nannofossils and foraminifera show their minimum values. Both diatoms and radiolarians show a gradual increase in diversity peaking around the Eocene/Oligocene boundary, followed by a relatively stable plateau during the Neogene. Diatoms also show a peak of diversity in the late Paleocene, when radiolarians are not reported. Diversity of nannofossils and foraminifera, instead, peaked during the early to middle Eocene, decreased in the late Eocene, and has remained more of less constant since the Oligocene.

While it is possible that much of the general variability is due to taphonomy, several further speculations could be made on these patterns. However, potential biases would have to be examined first. For example, what is the lithology of the sections in Neptune through time? Are siliceous sediments more common in the Neogene, thereby explaining the higher siliceous plankton diversity? And how do the absolute normalized values compare? The highest values are recorded in nannofossils while the lowest ones are given for diatoms. The number of valid species names in the two groups is almost identical (Table 4.1), but nannofossil names (valid and non valid) are overall slightly more abundant than diatom names in Neptune (Fig. 3.4). However,



foraminifera and radiolarian names are the most abundant ones of the four groups, while their normalized diversity is intermediate between diatoms' and nannofossils' diversity. Is there a consistent bias in the published range charts, with more reports available on siliceous plankton than on calcareous plankton in the Neogene? The number of reports on Paleogene sections (Fig. 3.5) shows a relatively lower number of reports on siliceous groups than on calcareous groups. The high correlation shown between diatoms' and radiolarians' species richness (and to a lesser extent foraminifera) and the total number of sections available for each time interval (Fig. 4.17)

suggests that absolute values of species diversity are strongly biased by the size of the data set (i.e., more species are described when more sections, and therefore reports, are available). On the other hand, the species richness of nannofossils shows a complex polynomial correlation with the number of sections but a completely random linear correlation ($R_2 = 0.07$). This may be due to the fact that the correlation is made with the total number of sections and not with the number of sections that contain nannofossil stratigraphy. For nannofossils and foraminifera the latter may be the significant parameter which would perhaps show a higher correlation with the species richness, similarly to the one shown for siliceous plankton.

Finally, how strong is the bias caused by the dominant presence of stratigraphic markers in the reports? Siliceous plankton biostratigraphy is better developed for Neogene sediments than it is for the Paleogene, while it is more uniform for calcareous plankton groups. One approach to this question would be to separate the species included in the distribution into stratigraphic markers, other common taxa and rare taxa, and see if the diversity patterns remain the same or change substantially.

These are only some of the factors that one must consider before a feasible interpretation of these trends can be formulated, and some of these require the addition of data to Neptune which are not currently available (e.g., distribution of siliceous versus calcareous sediments). But however preliminary and partial, these results are still quite encouraging and represent a more detailed data set than what is available from the paleobiological literature.

4.2.1. Availability of relational databases for the paleontological community: the ODP database JANUS versus Neptune. At present, the ODP database, JANUS, which is currently available onboard the JOIDES Resolution and through the WWW, does not represent a viable substitute for Neptune. I must point out, however, that JANUS is very new and that the import of data has just begun. My experience with JANUS is limited to a superficial browsing through ODP's database WWW site, which provided me with the following information. Site data (water depth, coordinates, length)

drilled, length recovered, etc.), physical properties (e.g., GRAPE, magnetic susceptibility), and chemical results (e.g., carbonate content) represent the bulk of the database and are available for most ODP Legs. Age model and paleontological information are part of the database structure, but (as of March 1998) are given only for a handful of sites. There is one general grouping ('Paleontology') which is divided into four searchable tables: Age Model, Paleontological Investigation, Range Table, and Species Information. Age Model information is currently (as of March 1998) available only for one hole in Leg 105 and consists of two points, the top of the drilled section and the bottom. ODP is probably planning to progressively add more detailed age model data for all ODP sites.

I used a simple, predetermined query to search the database for paleontological information and I only obtained very preliminary information, such as the name/s of the paleontologist/s who did the shipboard study, the depth in meters below seafloor of the samples analyzed, their relative stratigraphic position (e.g., middle Eocene), and the abundance and preservation of the microfossil group. As this information was available only from Leg 171 onward, legs for which no reports are published as yet, I do not know if it is planned to add more detailed paleontological information (e.g., the range charts that are available in Neptune) from the **Scientific Results**, once they become available.

Very basic taxonomic information is also available. For example, the search for **'Globorotalia truncatulinoides'** resulted in the name of the author who named the species (d'Orbigny), when the species was first described (1839), and the stratigraphic interval it is found in (Neogene).

I finally attempted to develop a customized 'Power Query' to search JANUS but I did not succeed. No instructions were given on how to select the various items present in the relational tables and the query routine was neither user-friendly nor intuitive.

While JANUS is undoubtedly a very valuable resource for site information and shipboard results (mainly physical properties), the preliminary search of the paleontological content of JANUS suggests that Neptune is still clearly a more valuable source of paleontological information. Although I do not see how JANUS and Neptune could be easily integrated, the two databases certainly complement each other. As shown in the studies outlined above (Lazarus et al. 1995b; Spencer-Cervato and Thierstein 1997), Neptune can be extremely helpful to biostratigraphers during ODP cruises, for example for the identification of the taxa previously recorded in a specific region during a certain time interval, thereby restricting the field of species identification to likely occurrences.

4.3. Stratigraphic research with Neptune: diachrony and hiati distribution. The field where Neptune's potential has been already quite thoroughly exploited is stratigraphy. The chronology of Neptune's holes has been revised several times and even if the quality of age models is quite varied (<u>Table 2.1</u>), it still represents the most complete and reliable data set available for stratigraphic studies. Two major groups of

information have been derived from this data set, the first directly applicable to biostratigraphy, the second of a stratigraphic and paleoceanographic significance.

The goal for the first group of studies (Spencer-Cervato et al. 1993; 1994) was to

determine the reliability of biostratigraphic markers, in terms of their regional versus global significance and of their synchrony or diachrony. As mentioned in <u>Chapter 3</u>, siliceous biostratigraphy is based on several regional calibrations of events, whilst calcareous biostratigraphy relies on a single, mainly low-latitude calibration (Berggren et al. 1995a, b). The use of the latter approach (one calibration for all holes, irrelative of their biogeographic location) implies a global synchrony of biostratigraphic events, that has actually been demonstrated only in very few cases. The first study (Spencer-Cervato et al. 1002) was simed to aplibrate acvers! Neogone

al. 1993) was aimed to calibrate several Neogene radiolarian events in the north Pacific and to study the degree of diachrony within this biogeographic region (Fig. 4.18). The projected ages of radiolarian first and last occurrences derived from the line of correlation of age/depth plots from the North Pacific have been computed from twelve North Pacific sites, and 28 radiolarian events have thereby been newly cross-calibrated to North Pacific diatom and other stratigraphy. Several of the North Pacific radiolarian events are older than in previously published equatorial Pacific calibrations (Johnson and Nigrini 1985) (Fig. 4.18), and some may be diachronous within the North Pacific. We hypothesized that these patterns may be due to complex latitudinal patterns of clinal variation in morphotypes within lineages, or to migration events from the North Pacific towards the Equator.

The second, more comprehensive study (Spencer-Cervato et al. 1994) evaluated the synchrony and diachrony of 124 commonly used Neogene biostratigraphic events in 35 globally distributed DSDP and ODP holes. Global mean age estimates based on combined biostratigraphy and magnetostratigraphy were calculated for each event. The ages' standard deviations were used as an estimate of synchrony/diachrony. Average standard deviations for event ages by fossil group are: calcareous nannofossil first appearance



datums (FADs): 0.57 m.y. (21 events), calcareous nannofossil last appearance datums (LADs): 0.60 m.y. (25 events), diatom FADs: 0.57 m.y. (7 events), diatom LADs: 0.85 m.y. (14 events), planktic foraminifera FADs: 0.88 m.y. (22 events),



foraminifera LADs: 0.68 m.y. (16 events), radiolarian FADs: 0.30 m.y. (9 events), radiolarian LADs: 0.31 m.y. (10 events). 53 of the 124 events can be considered synchronous, within the resolution of the method (± two average sample spacings, i.e., 360 k.y.). The remaining diachronous events were analyzed for true patterns of diachrony and other biases. Generally, diachrony is more frequent among cosmopolitan than among endemic taxa (Fig. 4.19). Also, the precision of age calibrations decreases with increasing age. Some diachrony patterns may be due to investigator bias (see examples shown in <u>Spencer-Cervato et al. 1994</u>), but in general they appear to be, at least in part, real phenomena. Thus, they could provide opportunities for exploration of paleobiological processes (see for example the study on **G. truncatulinoides** described above, <u>Spencer-Cervato and Thierstein 1997</u>).

A similar study of diachrony was not attempted for Paleogene events and is not recommended either. The age control on the chronology of Paleogene sediments is poorer than what is available for Neogene sediments. Moreover, fewer sections were analyzed for magnetostratigraphy, which provides the independent control on the age models selected for the Neogene study described above. I expect that the patterns of diachrony that could be obtained for Paleogene events would be largely biased by the data set and, therefore, would not provide a scientifically sound basis for further studies.

The chronology of the 165 holes in Neptune was the subject of the third stratigraphic study originated from Neptune. It was mentioned in <u>Chapter 3</u> that continuous stratigraphic sequences were very rare and that most age models were characterized by hiati. Hiati are commonly recognized in shelf sediments, but regional deep-sea hiati have also been extensively studied (e.g., <u>Keller and Barron 1983</u>). The reason for the interest in the timing and geographic distribution of hiati lies in the processes that cause them. A hiatus is a stratigraphic gap caused by erosion, dissolution, corrosion, nondeposition, rate of sediment supply versus dissolution (corrosion) of sediments (controlled by fluctuations in the calcite compensation depth - CCD), or shallow to deep water sediment fractionation (<u>Berger 1970</u>). Several studies have interpreted the occurrence of deep-water hiati in terms of changes in deep water circulation and corrosiveness (e.g., <u>Keller and Barron 1987</u>). Other studies have interpreted the mwithin a framework of sea-level fluctuations (e.g., <u>Vail et al. 1977</u>; <u>Haq et al.</u> 1987).

For this study (Spencer-Cervato 1998) I have identified 'hiatus events' during the Cenozoic, based on the occurrence of individual hiati both in shelf and deep-sea sediments. The goal of the study was to test if there is a causal link between sea-level fluctuations (and climate change) and global occurrences of hiati, which are linked to oceanic circulation through a variety of complex processes. I initially attempted to reproduce the 'global eustatic sea-level curve' of <u>Haq et al. (1987)</u> with a curve of hiati distribution. This sea-level curve was constructed by the Exxon Exploration Group and

is based on proprietary seismic data collected mainly on the eastern Atlantic passive continental margin. This curve has been a source of controversy since its publication, mainly because scientists had failed to reproduce it and because it was difficult to find physical mechanisms that could cause rapid sea-level fluctuations of more than 250 m, such as the ones implied in the curve. Drilling off New Jersey during ODP Leg 150X has recovered stratigraphic sequences which contain gaps that can be correlated to the ones used to construct the sea-level curve. These results (e.g., Miller et al., 1996) seem to have sedated the debates on the reliability of the sea-level curve, but the dispute on the magnitude of the fluctuations is still unresolved.

Compared to previous compilations of hiatus distribution in the DSDP stratigraphic record (e.g., Moore et al., 1978), the curve that I have obtained (Fig. 4.20c) has a better resolution (0.5 m.y.), contains more recent holes with better recovery, and is based on a more reliable and updated biochronology. Other studies (e.g., Keller and Barron 1983; Ramsay et al. 1994) instead focused on specific regions, e.g., the Indian or the Atlantic



Ocean, while my study (Spencer-Cervato 1998) is of global extent.

To help in the interpretation of the record of hiati, I have estimated the paleo-water depth at which the hiati occurred and constructed three individual curves for shallow (0-2000 m), intermediate (2000-3000 m) and deep (> 3000 m) water (Fig. 4.21). The curves show that the Paleogene is characterized by few, several million-years long hiati, while the Neogene is punctuated by short, frequent hiati events (Fig. 4.20), occurring nearly synchronously in shallow and deep water sediments. The most significant Cenozoic hiatus event spans most of the Paleocene. Epoch boundaries are characterized by peaks in deepwater hiati possibly caused by an increased circulation of corrosive bottom water and sediment dissolution. The Plio-Pleistocene is

characterized by a gradual decrease in the frequency of hiati. This could be caused by several factors, including the better recovery of younger sediments and therefore a lower chance of recording artificial hiati. Alternatively, this can indicate that sediment erosion and corrosion is time dependent and thus that there has been insufficient time to create hiati in the youngest sections. However, this smooth drop can also be an artifact of the time interval chosen for this analysis, which masks the high-frequency cycles of Quaternary alacio-eustatic sea level change possibly characterized by short (<0.5 m.y.) hiati, not recorded in this study.



Although some speculations were advanced on the causes of these hiatus events, their regional significance and possible causes will be the topic of future, more detailed studies. Among these, of particular interest would be the geographic distribution of hiati within ocean basins (e.g., latitudinal distribution of hiati versus latitudinal distribution of DSDP and ODP holes in the database and western versus eastern margins to identify the temporal evolution of oceanic gyre circulation) and their comparison to detailed isotopic records of deepwater circulation (e.g., Wright and Miller 1993). The depth distribution of hiati in mid-ocean and aseismic ridge sites versus continental shelf and slope sites must be also analyzed separately. These areas should be affected differently by sea level changes.

In summary, Neptune's data have been used for five published studies of plankton evolution and stratigraphy. While the stratigraphic studies provide a quite complete overview of the potential of Neptune, the study of plankton evolution has so far been limited to biostratigraphic applications. The analysis of plankton longevity and diversity has been shown here as raw data. This is the field where Neptune's data still have much to offer to the paleontological community. In the following section, I will explore the possibility of expanding Neptune beyond the paleontological field and will suggest possible future avenues of research based on this database.

4.4. Potential additional data for Neptune for sedimentological and paleoceanographic research

In my opinion, the potential of Neptune for future research extends considerably beyond analyses of micropaleontological data. Neptune's chronology and relatively large number of holes are its greatest assets and they should be properly exploited. A significant step forward would be represented by the addition of sedimentological data, which would open up a whole new range of research possibilities. The expansion of Neptune would benefit the research community by providing interdisciplinary links and correlations that are at present rarely possible to scientists working on ODP material. Time pressure and poor funding force ODP-participating scientists to limit their post-cruise research to very limited, mainly isolated goals (James D. Wright, personal commun., 1998). Once their duty as sedimentologists or micropaleontologists is fulfilled and their report is submitted, scientists move quickly to the next 'hot' research topic, and the potential for correlations between data sets and large-scale research studies is left largely untouched. In this scenario, the opportunities provided by Neptune's chronology and paleontological data would be greatly enhanced by other data that would allow to make large-scale, interdisciplinary (e.g., modeling) studies, or at least would provide an easily accessible source of a large amount of quality data from which to start such studies. Only very few of these studies based on deep-sea sediments are available at present (e.g., Delaney and Boyle 1988).

Among the data that should be included in Neptune, and that are consistently available at least for the more recent ODP holes, are: lithology (percentage carbonate, percentage silica); organic carbon content; physical properties (e.g., bulk density,

grain density, porosity); and grain-size distribution. The field of paleoceanography would be the primary beneficiary of the combination of the existing paleontological and chronological data in Neptune with sedimentological information and physical properties data. I will mention here only a couple of the several, current research questions that are debated in the paleoceanographic community and that could be addressed with these additional data.

During the middle Miocene, important changes occurred in the climate of the Earth, an important step toward the establishment of cold polar climates and the modern climate mode characterized by glacial and interglacial cycles. These changes are documented in the oxygen isotope records (e.g., Miller et al. 1987; Fig. 4.20e), and indicate the onset of a progressive global cooling. It is not yet known what causes the abrupt shifts in climate mode that the Earth has experienced in the last million years, even though some recent evidence (Zachos et al. 1997) suggest that these shifts might have characterized the earth's climate already since the Oligocene. It is, therefore, obvious why so many studies have focused on middle Miocene sediments and have led to the formulation of various hypotheses. Several hypotheses have linked climate changes to large-scale deepwater circulation (e.g., Shackleton et al. 1983), but the causal relationship between middle Miocene changes in deepwater circulation and the establishment of a permanent ice sheet in eastern Antarctica, is still uncertain (e.g., Kennett and Barker 1990).

Keller and Barron (1983) proposed that a "silica switch" occurred between the Atlantic and Pacific Ocean in the middle Miocene, around 15 Ma and contemporary to the δ 18O increase. Based on the relative abundance of siliceous sediments in nine tropical Atlantic sites, they suggested that prior to 15 Ma, Atlantic sediments were silica-rich, but that after that time, silica sedimentation switched to the Pacific Ocean. Predominantly carbonatic sediments have apparently been accumulating in the Atlantic since then. This switch would have been caused by the initiation of the Northern Component Water (NCW) circulation in the north Atlantic. However, Wright et al. (1992) have raised some concerns on the selection of data on which Keller and Barron (1983) based their hypothesis, and argue that NCW's production began earlier (around 19 Ma) and had actually shut down during the 15 Ma δ 18O event. Wright et al. (1992) propose that the middle Miocene δ 18O increase does not correlate with deepwater circulation changes and does not represent the transition from an ice-free to an ice-house world, but is part of two or three glacial/interglacial cycles.

How could Neptune help solve this controversy? The cause of disagreement in the interpretations is the data on which <u>Keller and Barron (1983)</u> based their hypothesis. The possibility of modeling the results of a larger number of chronologically well constrained holes in the Atlantic and Pacific Oceans would give the 'middle Miocene controversy' a strong, potentially unbiased basis of data. The data that would be needed are the concentrations of carbonate and opal. These data have not been incorporated in Neptune yet, because data from different holes are potentially

incompatible, due to the different methods used in their collection.

Since the beginning of the DSDP project, carbonate concentration has been a routine analysis performed on the sediments. Biogenic opal data are also available for DSDP holes, but only in the more recent ODP holes has it been estimated analytically. It would be conceptually simple to add a new table to Neptune, which would include the sample number, percentage carbonate and percentage opal. This table would be linked to the Sample Data table, which would provide an age estimate for each sample. In reality, this task is far from being trivial. The main reason for this is due to the analytical methodology used to determine the concentration of carbonate and opal. During DSDP, several different methods were used, giving results that are not comparable to each other. Routinely, percentage carbonate and opal have been estimated from smear slides, a method that has a maximum accuracy of ±10% (Hsü, Montadert et al. 1978). These data are highly subjective and not useful for rigorous quantitative studies. Keller and Barron (1983) used these counts for their "silica switch" hypothesis, on sites where the biosiliceous component was actually minimal and easily affected by dissolution of diatoms, the primary components of siliceous productivity (Wright et al. 1992). Quantitative analytical measurements of carbonate and opal would be required to map the distribution of these carbonatic and siliceous sediments in the entire ocean basins and to test the "silica switch" hypothesis.

The shipboard-based 'carbonate bomb' method (Müller and Gastner 1971) has been used relatively early in the DSDP on a few selected samples to provide a control on the smear slide estimates. The accuracy of this method, the most common analytical method used during DSDP cruises, is between 1 and 5%, lower for carbonate-rich sediments. In some instances, other shorebase methods (e.g., the LECO method; Hsü, Montadert et al. 1978) were compared to the results of the 'carbonate bomb' method and systematic differences observed. In more recent ODP Legs, a shipboard Coulometer is used to determine percent carbonate (and percent organic carbon)(e.g., Leg 121), with a precision of approx. 1%. I am not aware of a study that compares this method to the previous methods. The silica content has been quantitatively analyzed only for some ODP holes, using X-Rayx-ray fluorescence and the normative equation of Leinen (1977) (e.g., Littke et al. 1991) or, more recently, a single-step alkaline extraction method (Mortlock and Froelich 1989). The precision of the latter method is ±4%. The carbonate bomb, LECO and Coulometer estimates for carbonate and the X-ray fluorescence and alkaline extraction estimates of opal content from holes where sampling frequency is sufficient (e.g., one sample every 0.5 m.y.), could be selectively incorporated into Neptune. Some of these data are available for ODP holes from the JANUS database, but they would probably need to be added manually for the DSDP holes.

These same lithological data, combined with the paleo-water depth estimates of <u>Spencer-Cervato (1998)</u> (not yet included in Neptune but available at the NOAA-WDCA for Paleoclimatology Data Contr. Series #97-030: http://ftp.ngdc.noaa.gov/paleo/paleocean/by_contributor/spencer-cervato1998/) could

be used to reconstruct the depth fluctuations of the Calcite Compensation Depth (CCD) during the Cenozoic, for which the curve by <u>van Andel (1975)</u> is still being used. The very smooth fluctuations of the CCD curve of <u>van Andel (1975)</u> do not agree with the abrupt changes that have been recently shown to characterize the earth's climate and ocean systems in the Cenozoic. This is probably caused by the low resolution of stratigraphic studies in the 1970s. The better resolution (around 0.5 m.y. or better) of the Neptune data would allow us to refine the curve and to make it more compatible with, for example, the isotopic data currently being produced.

I briefly mentioned earlier the importance of deepwater circulation reconstruction for the understanding of the climate/ocean systems. Sediment accumulation rates and dissolution profiles can be reconstructed from physical property data and grain-size distributions, all data that are routinely produced onboard ship and that could be quite easily added to Neptune. Recent studies (Zachos et al. 1997) use a record of percent coarse fraction to demonstrate that glacial/interglacial cycles existed as far back as the late Oligocene. This parameter, combined with a high resolution δ 180 record from ODP Leg 154 (Ceara Rise - south Atlantic), shows a 40-k.y. periodicity, indicating a high-latitude orbital control on ice volume and temperature. This isotopic record suggests that there is an orbital control on deepwater circulation, which had not yet been shown so early probably because of the low resolution of previous studies and the paucity of deep-sea sections with high sedimentation rates and a long stratigraphic record, like the ones recovered by Leg 154 on the Ceara Rise.

The trend for future paleoceanographic studies is toward high (tens of thousands of years) resolution studies. Is Neptune ready for these studies? The answer is: not yet. The chronology of Neptune, its biggest strength and the most updated record available, is based on biostratigraphy and magnetostratigraphy, which provide an accuracy on the order of hundreds of thousands of years, at best. Oxygen isotope stratigraphy is currently the only means to obtain a better age resolution than this for Neogene sediments. Results from ODP Leg 154 provide the longest and most complete isotopic record for the whole Neogene which extends into the Oligocene (Weedon et al. 1997). A calibration of the sporadic isotopic records available for DSDP holes and the more common records from ODP holes to this recent isotopic calibration would allow us to refine the (mainly) late Neogene chronological resolution for some of the sites in Neptune. My biggest concern about this calibration is that correlations with standard isotopic records are still done by 'wiggle matching'. Because the absolute isotopic values vary depending on the foraminifer species or sediment fraction used to obtain the record, the shape of the curve (which remains substantially the same) is used for the correlation. I am not aware of any comprehensive study that has carefully pinpointed and tabulated some of the ages of the 200+ (the total number is actually unknown: some 140 are recognized only in the Plio-Pleistocene) isotopic stages to magnetostratigraphy beyond the late Miocene (Hodell et al. 1994). This would provide fixed reference points for stratigraphic interpretations. 'Eye-balled' graphic correlation is, in my opinion, too inaccurate for the type of studies that it is used for (unless one can actually count back all stages

starting from the Holocene) and greatly reduces the potential resolution of isotope stratigraphy.





Figure 4.1. Overview of database project (modified after Lazarus, 1994). Combining existing data (boxes with thin outlines) with new information generated by project members (rounded boxes with italic text), we have produced new information (bold outline boxes).



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Figure 4.2. Simplified sketch of the database structure (modified after Lazarus 1994). Each box represents one relational table, with field names listed in the box. Relational links between tables are shown by thin lines. Arrows point from a table with many records with the same value for the relating field, to a table with unique values for each record in this field. The primary key for each table (the combination of one or more fields which makes each record in the table unique) is underlined. The approximate number of records (as of February, 1998) is shown below each box. The records in the 'Taxa by Hole' table were extracted from the DSDP CD-ROM. The other tables were created by Neptune.



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Figure 4.3. Age Range Chart produced from Neptune for the occurrences of Globorotalia truncatulinoides and Globorotalia tosaensis. The plot shown here represents only a subset of the data obtained from Neptune for this search. The symbols are explained in the caption at the bottom of the figure. The charts are produced in colour, with the colour of the bullets corresponding to the species name on the right. Small black dots represent samples that were examined but where none of the selected species was found. This figure is an unmodified reproduction of the format produced by the Age Range Chart program (Lazarus, 1994). The large number of data in the figure and its format make difficult to understand it when it is reduced to a one-page size and does not clearly show the differences between the different species' occurrences.

Composite Age-Range Chart



http://www-odp.tamu.edu/paleo/1999_2/neptune/fig4_3.htm (1 of 2) [10/26/1999 10:02:40 AM]

Figure 4.3: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

KEY: | HIATUS • Absent • Trace-Rare • Few-Common • Abundant-Dominant

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Figure 4.4. Screen snapshot of a simple query in Neptune. The table in the lower right corner is the list of tables in Neptune that are directly searchable with a 'Search Editor' window (under the 'Select' menu). Selecting one of the tables shows the records included in it and their number (in the heading). The larger table in the background is the 'Sample Info' table, which contains 30800 records.

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| 414 | | 173 | 23 | 2 | 57 | 207.07 | 12.73 | R | С | м | | | |
| 208 | | 158 | 1 | 1 | 34 | 0.34 | 0.24 | R | | P | | | |
| 209 | | 158 | 1 | 1 | 132 | 1.32 | 0.28 | R | | P | | | |
| 210 | | 158 | 1 | 2 | 135 | 2.85 | 0.33 | R | | P | | | |
| 211 | | 158 | 1 | 3 | 70 | 3.70 | 0.37 | R | | P | | | |
| 212 | | 158 | 1 | 9 | 0 | 2.90 | 0.34 | R | | P | | | |
| 213 | | 158 | 2 | 1 | 86 | 9.86 | 0.59 | R | | P | | | - |
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Figure 4.4: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE



Close Window

Figure 4.5. Screen snapshot of the 'Execute Procedure' window (under the 'Special' menu). This allows to search a combination of tables using one of the predefined procedures. The following figures show the progression of a 'Bug Data Search' procedure for the occurrence of the planktic foraminifer **Globorotalia truncatulinoides** in the Neptune holes. A 'Bug Data Search Batch' (described in <u>Section 4.1.4</u>) is one of the other procedures shown on the list.

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Figure 4.5: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE



Figure 4.6: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.6. Once the type of procedure is chosen, a 'Search Editor' window is shown. In this example, a search is made for the species' name - 'Species' 'is equal to' '**truncatulinoides**'.

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Figure 4.6: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE



Close Window

Figure 4.7. The selected species' name is found in the 'Species Names' table together with linked synonyms. After one has highlighted the names that should be used, one can proceed with the Search. Note the format of the data in the 'Species Names' table, described in <u>Section 3.3</u> and characteristic of Neptune.

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Figure 4.8: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.8. The result of the search is shown in a window that gives the number of occurrences in all the holes present in Neptune (779 in this example). The next step allows to refine the selection with additional criteria.

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Figure 4.8: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE



Figure 4.9: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.9. In this example I have chosen to refine the search to obtain only data from holes for which we have age models. To do this, I have selected 'Age' in the 'Sample Info' table (left) and specified that it has to be greater than 0 (all samples from holes without age models have ages set to 0). The 'Search in selection' box, which allows to search only in the selection found through the previous queries, is automatically marked.

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Figure 4.10: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.10. The results of this procedure are shown in a table containing 485 records (of the 779 found earlier in the search). This table and related data can be sorted, printed out, exported, or saved in various formats.

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| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 6.05 | 1.41 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 7.70 | 1.51 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 9.20 | 1.61 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 10.70 | 1.70 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 12.20 | 1.80 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 13.70 | 1.89 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 141 | 13.00 | 1.85 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 233 | 17.20 | 0.74 | A | | | | | | | | |
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| GLOBOROTALIA | | 265 | 8.55 | 0.09 | R | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 265 | 26.95 | 0.18 | F | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 265 | 27.00 | 0.18 | P | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 281 | 7.64 | 0.89 | P | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 281 | 10.10 | 1.19 | Р | | | | | | | | |
| GLOBOROTALIA | TRUNCATULINOIDES | 281 | 11.50 | 1.36 | Р | | | | | | | | |

http://www-odp.tamu.edu/paleo/1999_2/neptune/fig4_10.htm (1 of 2) [10/26/1999 10:02:43 AM]

Figure 4.10: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE



Figure 4.11. Age versus latitude plot of the first occurrences of **G. truncatulinoides** in selected Neptune sites (from <u>Spencer-Cervato and Thierstein 1997</u>).



Figure 4.12. Distribution of diatom, foraminifera, nannofossil and radiolarian longevity by species (in m.y.), including both extant and extinct species. Note that, while most phytoplankton species are less longevous than zooplankton species, extreme longevity values are exclusively found in phytoplankton species.



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Figure 4.13. Appearance rates (calculated as the ratio between number of FADs versus number of species) of the four groups during the Cenozoic Note that the rates are biased by species richness in low diversity intervals (see Fig. 4.9).



http://www-odp.tamu.edu/paleo/1999_2/neptune/fig4_13.htm (1 of 2) [10/26/1999 10:02:45 AM]





Figure 4.14. Extinction rates (calculated as the ratio between number of LADs versus number of species) of the four groups during the Cenozoic.

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Figure 4.15: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.15. Geographic distribution of species FADs for the four groups. Southern latitudes are given as negative numbers. The latitude given for each FAD is a calculated paleolatitude at the time of the appearance (see 4.1.1.4 for details on the method).



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| Latitude | | | | | | | | | | | | Latitu | de | | | | | | |

Figure 4.16. Species richness (total number of species - diamond symbols and black line) and diversity normalised on the number of sections in Neptune (ratios between diversity and number of sections - crosses and grey line) for the four plankton groups during the Cenozoic. Species richness is calculated at 1 m.y. intervals. Note that the scale of the normalised diversity is different in the four graphs.



http://www-odp.tamu.edu/paleo/1999_2/neptune/fig4_16.htm (1 of 2) [10/26/1999 10:02:47 AM]



Figure 4.17: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window

Figure 4.17. Species richness versus the total number of sections in Neptune. The R2 refers to the simple linear correlation shown in the graphs.



http://www-odp.tamu.edu/paleo/1999_2/neptune/fig4_17.htm (1 of 2) [10/26/1999 10:02:47 AM]



Figure 4.18. Age differences between equatorial Pacific calibration (defined as 0) and North Pacific age ranges. The horizontal bar represents the suggested calibration for the single events. The events are divided into three groups, from left to right: the first, largest group includes events that are older in the North Pacific than they are in the Equatorial Pacific, the second group of events is synchronous in the equatorial and North Pacific, the third group includes events that are younger in the North Pacific. The number of sites included in the age range is written below the bar (from <u>Spencer-Cervato et al. 1993</u>).



Figure 4.19. Latitudinal range versus standard deviation (in m.y.) for four plankton groups (from Spencer-Cervato et al. 1994). Note the higher standard deviations (interpreted as higher diachrony) for events with broader latitudinal ranges (i.e., cosmopolitan events). Also, note the relatively low standard deviation values for calcareous nannofossils, which are largely cosmopolitan. This suggest that they are particularly well suited for biostratigraphic correlations.



Latitudinal range

Latitudinal range

Figure 4.20: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

Close Window



Figure 4.20. (a) Number of Cenozoic DSDP and ODP sections per 0.5 m.y. in the Neptune database. (b) Number of hiati recorded in the Neptune database during the last 65 m.y. (c) Curve of the frequency of hiati. The vertical line represents the average frequency (30%). Shaded intervals mark periods characterised by a higher than average frequency of hiati (hiatus events). (d) "Eustatic" sea level curve of <u>Hag et al. (1987</u>). (e) Benthic foraminifera oxygen isotope curve (<u>Miller et al. 1987</u>). All curves are calibrated to <u>Berggren et al.'s (1995b</u>) biochronology (from <u>Spencer-Cervato 1998</u>).

Figure 4.21. The lines with symbols represent the frequency of hiati recorded at various paleodepth intervals: (a) shallower than 2000 m, (b) between 2000 and 3000 m, and (c) deeper than 3000 m. The continuous lines represent the percentage of sections in the Neptune database in that interval of paleo-water depth (from <u>Spencer-Cervato 1998</u>).





Table 4.1. Names of 1418 species included in the macroevolution analyses.

Diatoms

ABAS WITTII ACHTINOPTYCHUS BIPUNCTATUS ACHTINOPTYCHUS SENARIUS ACHTINOPTYCHUS SPLENDENS ACTINOCYCLUS CURVATULUS ACTINOCYCLUS ELLIPTICUS ACTINOCYCLUS ELONGATUS ACTINOCYCLUS INGENS ACTINOCYCLUS LANCEOLATUS ACTINOCYCLUS MORONENSIS ACTINOCYCLUS OCHOTENSIS **ACTINOCYCLUS OCTONARIUS** ACTINOCYCLUS OCULATUS ACTINOCYCLUS TSUGARUENSIS **ACTINOCYCLUS ACTINOCHILUS** ACTINOCYCLUS FASCICULATUS ACTINOCYCLUS FRYXELLAE ACTINOCYCLUS KARSTENII ANNELLUS CALIFORNICUS ASTEROLAMPRA ACUTILOBA ASTEROLAMPRA AFFINIS ASTEROLAMPRA GREVILLEI ASTEROLAMPRA INSIGNIS ASTEROLAMPRA MARYLANDICA

ASTEROLAMPRA PUNCTIFERA ASTEROLAMPRA VULGARIS ASTEROLAMPRA SCHMIDTII **ASTEROMPHALUS ARACHNE** ASTEROMPHALUS DARWINII ASTEROMPHALUS FLABELLATUS **ASTEROMPHALUS HEPTACTIS** ASTEROMPHALUS HOOKERI ASTEROMPHALUS IMBRICATUS ASTEROMPHALUS OLIGOCENICUS ASTEROMPHALUS PARVULUS ASTEROMPHALUS PETERSONII ASTEROMPHALUS ROBUSTUS ASTEROMPHALUS SYMMETRICUS ASTEROMPHALUS KENNETTII **AZPEITIA NODULIFER** AZPEITIA ENDOI **AZPEITIA TABULARIS AZPEITIA GOMBOSI** AZPEITIA OLIGOCENICA **AZPEITIA PRAENODULIFER** AZPEITIA VETUSTISSIMA **BACTERIOSIRA FRAGILIS** BACTERIASTRUM HYALINUM **BAXTERIA BRUNII BOGOROVIA VENIAMINI BOGOROVIA PALEACEA BOGOROVIA PRAEPALEACEA**

CESTODISCUS CONVEXUS **CESTODISCUS MUKHINAE** CESTODISCUS PEPLUM CESTODISCUS PULCHELLUS CESTODISCUS RETICULATUS CESTODISCUS ROBUSTUS CESTODISCUS STOKESIASUS CESTODISCUS TROCHUS CESTODISCUS ANTARCTICUS CESTODISCUS DEMERGITUS CESTODISCUS PARMULA CHAETOCEROS FURCELLATUS CHAETOCEROS ASYMMETRICUS CHAETOCEROS BULBOSUM CLADOGRAMMA DUBIUM COCCONEIS CALIFORNICA COCCONEIS COSTATA COCCONEIS PSEUDOMARGINATA COCCONEIS SCUTELLUM CORETHRON CRIOPHILUM COSCINODISCUS AFRICANUS COSCINODISCUS APICULATUS COSCINODISCUS ASTEROMPHALUS COSCINODISCUS BLYSMOS COSCINODISCUS CENTRALIS COSCINODISCUS CRENULATUS COSCINODISCUS DEFORMANS

COSCINODISCUS ELEGANS COSCINODISCUS EXCAVATUS COSCINODISCUS FLEXUOSUS **COSCINODISCUS GIGAS** COSCINODISCUS KOLBEI COSCINODISCUS LEWISIANUS COSCINODISCUS MARGINATUS COSCINODISCUS NODULIFER COSCINODISCUS OCCULUS-IRIDIS COSCINODISCUS PLICATUS COSCINODISCUS PRAEYABEI COSCINODISCUS PUSTULATUS COSCINODISCUS RADIATUS COSCINODISCUS RHOMBICUS COSCINODISCUS STELLARIS COSCINODISCUS SUPERBUS COSCINODISCUS TEMPEREI COSCINODISCUS TUBERCULATUS COSCINODISCUS YABEI COSCINODISCUS CRUXII **COSCINODISCUS HAJOSIAE** COSCINODISCUS VULNIFICUS **COSMIODISCUS INSIGNIS** CRASPEDODISCUS COSCINODISCUS CRASPEDODISCUS UMBONATUS **CRUCIDENTICULA NICOBARICA** CRUCIDENTICULA PUNCTATA CRUCIDENTICULA KANAYAE

CUSSIA LANCETTULA CYCLOTELLA STRIATA CYMATOGONIA AMBLYOCERAS CYMATOSIRA COMPACTA CYMATOSIRA DEBYI CYMATOTHECA WEISSFLOGII DACTYLIOSOLEN ANTARCTICUS **DELPHINEIS SURIRELLA** DELPHINEIS ISCHABOENSIS DENTICULOPSIS DIMORPHA DENTICULOPSIS HUSTEDTII DENTICULOPSIS HYALINA DENTICULOPSIS LAUTA DENTICULOPSIS MACCOLLUMII DENTICULOPSIS MIOCENICA DENTICULOPSIS NORWEGICA DENTICULOPSIS PRAEDIMORPHA DENTICULOPSIS PUNCTATA DENTICULOPSIS SEMINAE DENTICULOPSIS KATAYAMAE DENTICULOPSIS MERIDIONALIS DIMEROGRAMMA FOSSILE **DIPLONEIS BOMBUS** DIPLONEIS COFFAEIFORMIS **DIPLONEIS SMITHII DIPLONEIS WEISSFLOGII** ETHMODISCUS REX

EUCAMPIA ANTARCTICA GONIOTHECIUM DECORATUM GONIOTHECIUM ODONTELLA **GRUNOWIELLA GEMMATA GRUNOWIELLA PALAEOCAENICA HEMIDISCUS CUNEIFORMIS** HEMIDISCUS KARSTENII **HEMIDISCUS OVALIS** HEMIDISCUS SIMPLICISSIMUS **HEMIDISCUS TRIANGULARUS HEMIAULUS ALTAR** HEMIAULUS BARBADENSIS **HEMIAULUS CARACTERISTICUS HEMIAULUS EXIGUUS HEMIAULUS INCISUS HEMIAULUS POLYMORPHUS HEMIAULUS PUNGENS HEMIAULUS SUBACUTUS HEMIAULUS TAURUS HEMIAULUS FRAGILIS HEMIAULUS GRACILIS** HEMIAULUS KRISTOFFERSENII **HEMIAULUS NOCCHIAE HEMIAULUS PERIPTERUS** HEMIAULUS POLYCYSTINORUM **HEMIAULUS RECTUS HEMIAULUS ROSSICUS** HYALODISCUS OBSOLETUS

HYALODISCUS AMBIGUUS

IKEBEA TENUIS

KATATHIRAIA ASPERA

KISSELEVIELLA CARINA

KOZLOVIELLA MINOR

LISITZINIA ORNATA

LITHODESMIUM CORNIGERUM

LITHODESMIUM MINUSCULUM

LITHODESMIUM REYNOLDSII

LITHODESMIUM UNDULATUM

MACRORA STELLA

MEDIARIA SPLENDIDA

MELOSIRA ALBICANS

MELOSIRA SOL

NAVICULA LYRA

NAVICULA UDENTSEVII

NAVICULA WISEI

NEOBRUNIA MIRABILIS

NEODENTICULA KAMTSCHATICA

NEODENTICULA KOIZUMII

NITZSCHIA AEQUATORIALIS

NITZSCHIA ANGULATA

NITZSCHIA BICAPITATA

NITZSCHIA BRAARUDII

NITZSCHIA CHALLENGERI

NITZSCHIA CLAVICEPS

NITZSCHIA CURTA

NITZSCHIA CYLINDRICA NITZSCHIA GRANULATA NITZSCHIA GROSSEPUNCTATA NITZSCHIA GRUNOWII NITZSCHIA INFLATULA NITZSCHIA INTERFRIGIDARIA NITZSCHIA INTERRUPTESTRIATA NITZSCHIA JANUARIA NITZSCHIA JOUSEAE NITZSCHIA KERGUELENSIS NITZSCHIA KOLACZECKII NITZSCHIA MARINA NITZSCHIA MIOCENICA NITZSCHIA PANDURIFORMIS NITZSCHIA PORTERI NITZSCHIA PRAEINTERFRIGIDARIA NITZSCHIA PRAEREINHOLDII NITZSCHIA PSEUDOKERGUELENSIS NITZSCHIA PUNCTATA NITZSCHIA PUSILLA NITZSCHIA REINHOLDII NITZSCHIA RITSCHERII NITZSCHIA ROLANDII NITZSCHIA SEPARANDA NITZSCHIA SERIATA NITZSCHIA SICULA NITZSCHIA SUIKOENSIS

NITZSCHIA UMAOIENSIS

NITZSCHIA WEAVERI

NITZSCHIA CYLINDRUS

NITZSCHIA DENTICULOIDES

NITZSCHIA DIETRICHII

NITZSCHIA DONAHUENSIS

NITZSCHIA EFFERANS

NITZSCHIA EVENESCENS

NITZSCHIA EXTINCTA

NITZSCHIA FOSSILIS

NITZSCHIA MALEINTERPRETARIA

ODONTELLA AURITA

ODONTELLA TUOMEYI

PARALIA SULCATA

PARALIA ARCHITECTURALIS

PARALIA CLAVIGERA

PLAGIOGRAMMA STAUROPHORUM

PLANKTIONELLA SOL

POROSIRA GLACIALIS

PSEUDOEUNOTIA DOLIOLUS

PSEUDOPODOSIRA ELEGANS

PSEUDOPODOSIRA SIMPLEX

PSEUDOPYXILLA AMERICANA

PSEUDOPYXILLA RUSSICA

PSEUDOTRICERATIUM CHENEVIERI

PSEUDOTRICERATIUM RADIOSORETICULATUM

PTEROTHECA ACULEIFERA

PTEROTHECA CLAVATA

PTEROTHECA EVERMANNI PTEROTHECA KITTONIANA PYXILLA RETICULATA RHABDONEMA ARCUATUM RHABDONEMA JAPONICUM RHAPHIDODISCUS MARYLANDICUS RHAPHONEIS AMPHICEROS RHAPHONEIS ANGUSTATA RHAPHONEIS MARGARITALIMBATA **RHIZOSOLENIA ALATA** RHIZOSOLENIA BERGONII RHIZOSOLENIA CRETACEA RHIZOSOLENIA CURVIROSTRIS **RHIZOSOLENIA HEBETATA** RHIZOSOLENIA INTERPOSITA RHIZOSOLENIA MATUYAMAI **RHIZOSOLENIA MIOCENICA** RHIZOSOLENIA PRAEBERGONII RHIZOSOLENIA STYLIFORMIS RHIZOSOLENIA ANTARCTICA RHIZOSOLENIA OLIGOCENICA RHIZOSOLENIA SETIGERA **RIEDELIA CLAVIGER ROCELLA GELIDA ROCELLA VIGILANS ROCELLA PRAENITIDA ROPERIA PRAETESSELATA ROPERIA TESSELATA**

ROSSIELLA PALEACEA ROSSIELLA PRAEPALEACEA ROSSIELLA TATSUNOKUCHIENSIS ROSSIELLA SYMMETRICA ROUXIA CALIFORNICA ROUXIA GRANDA ROUXIA HETEROPOLARA **ROUXIA ISOPOLICA ROUXIA NAVICULOIDES ROUXIA OBESA** ROUXIA PERAGALLI ROUXIA YABEI SCEPTRONEIS GRUNOWII SCEPTRONEIS HUMUNCIA SCEPTRONEIS PESPLANUS SCEPTRONEIS TENUE SCEPTRONEIS LINGULATUS SIMONSENIELLA BARBOI SIMONSENIELLA PRAEBARBOI SKELETONEMA BARBADENSE SPHYNCTOLETHUS HEMIAULOIDES STEPHANOPYXIS DIMORPHA STEPHANOPYXIS GRUNOWII STEPHANOPYXIS HORRIDUS STEPHANOPYXIS MARGINATA STEPHANOPYXIS SPINOSISSIMA STEPHANOPYXIS TURRIS

STEPHANOPYXIS ORNATA STEPHANOGONIA HANZAWAE **STELLARIMA MICROTRIAS** STELLARIMA PRIMALABIATA SYNEDRA INDICA SYNEDRA JOUSEANA SYNEDRA MIOCENICA THALASSIOSIRA ANTARCTICA THALASSIOSIRA ANTIQUA THALASSIOSIRA BURCKLIANA THALASSIOSIRA CONVEXA THALASSIOSIRA DECIPIENS THALASSIOSIRA DELICATULA THALASSIOSIRA ECCENTRICA THALASSIOSIRA FRAGA THALASSIOSIRA GRACILIS THALASSIOSIRA GRAVIDA THALASSIOSIRA HYALINA THALASSIOSIRA JACKSONII THALASSIOSIRA KRYOPHILA THALASSIOSIRA LACUSTRIS THALASSIOSIRA LEPTOPUS THALASSIOSIRA LINEATA THALASSIOSIRA NIDULUS THALASSIOSIRA NODULOLINEATA THALASSIOSIRA NORDENSKIOELDII THALASSIOSIRA OESTRUPII THALASSIOSIRA OPPOSITA

THALASSIOSIRA PACIFICA THALASSIOSIRA PLICATA THALASSIOSIRA PRAECONVEXA THALASSIOSIRA PUNCTATA THALASSIOSIRA SINGULARIS THALASSIOSIRA SPINOSA THALASSIOSIRA SPUMELLAROIDES THALASSIOSIRA SYMBOLOPHORA THALASSIOSIRA SYMMETRICA THALASSIOSIRA TEMPEREI THALASSIOSIRA TAPPANES THALASSIOSIRA TRIFULTA THALASSIOSIRA TUMIDA THALASSIOSIRA USATSCHEVII THALASSIOSIRA YABEI THALASSIOSIRA ZABELINAE THALASSIOSIRA BUKRYI THALASSIOSIRA ELLIPTIPORA THALASSIOSIRA GERSONDEI THALASSIOSIRA LENTIGINOSA THALASSIOSIRA MANIFESTA THALASSIOSIRA MARUJAMICA THALASSIOSIRA MIOCENICA THALASSIOSIRA NATIVA THALASSIOSIRA OLIVERANA THALASSIOSIRA STRIATA THALASSIOSIRA WEBBI

THALASSIONEMA BACILLARIS THALASSIONEMA HIROSAKIENSIS THALASSIONEMA NITZSCHIOIDES THALASSIONEMA ROBUSTA THALASSIONEMA SCHRADERI THALASSIOTHRIX FRAUENFELDII THALASSIOTHRIX LONGISSIMA THALASSIOTHRIX MIOCENICA TRACHYNEIS ASPERA **TRICERATIUM ACUTANGULUM** TRICERATIUM BARBADENSE TRICERATIUM CINNAMOMEUM TRICERATIUM CONDECORUM **TRICERATIUM GRONINGENSIS** TRICERATIUM SCHULZII TRICERATIUM CELLULOSUM **TRICERATIUM GOMBOSII TRICERATIUM MIRABILE** TRICERATIUM UNGUICULATUM TRINACRIA EXCAVATA TRINACRIA PILEOLUS TRINACRIA REGINA TRINACRIA SIMULACRUM TRINACRIA SUBCAPITATA TRINACRIA CONIFERA TRINACRIA DECIUSII TRINACRIA SENTA TRINACRIA SIMULACROIDES
TROCHOSIRA CONCAVA

TROCHOSIRA SPINOSA

TROCHOSIRA GRACILLIMA

TROCHOSIRA MARGINATA

TROCHOSIRA RADIATA

XANTHIOPYXIS OBLONGA

XANTHIOPYXIS ACROLOPHA

Foraminifera

ACARININA ASPENSIS

ACARININA BROEDERMANNI

ACARININA BULLBROOKI

ACARININA CONVEXA

ACARININA DECEPTA

ACARININA ESNAENSIS

ACARININA INTERMEDIA

ACARININA MATHEWSAE

ACARININA MCKANNAI

ACARININA PENTACAMERATA

ACARININA PRIMITIVA

ACARININA ROTUNDIMARGINATA

ACARININA RUGOSOACULEATA

ACARININA SOLDADOENSIS

ACARININA SPINULOINFLATA

ACARININA SPIRALIS

ACARININA TRIPLEX

ACARININA WILCOXENSIS

ACARININA APPRESSOCAMERATA

ACARININA CUNEICAMERATA ACARININA ECHINATA **ACARININA PRAEPENTACAMERATA** ACARININA PRAETOPILENSIS **BEELLA DIGITATA CANDEINA NITIDA** CASSIGERINELLOITA AMEKIENSIS **CASSIGERINELLA CHIPOLENSIS** CATAPSYDRAX DISSIMILIS CATAPSYDRAX PARVULUS CATAPSYDRAX STAINFORTHI CHILOGUEMBELINA CUBENSIS CHILOGUEMBELINA MIDWAYENSIS CHILOGUEMBELINA WILCOXENSIS CHILOGUEMBELINA CRINITA CRIBROHANTKENINA INFLATA DENTOGLOBIGERINA GALAVISI EOGLOBIGERINA EOBULLOIDES EOGLOBIGERINA TAURICA EOGLOBIGERINA FRINGA **GLOBOROTALIA AEMILIANA** GLOBOROTALIA ACROSTOMA **GLOBOROTALIA ANFRACTA** GLOBOROTALIA ARCHAEOMENARDII **GLOBOROTALIA BIRNAGAE GLOBOROTALIA BONONIENSIS GLOBOROTALIA CIBAOENSIS GLOBOROTALIA COLLACTEA**

GLOBOROTALIA CONICA GLOBOROTALIA CONOIDEA GLOBOROTALIA CONOMIOZEA **GLOBOROTALIA CONTINUOSA GLOBOROTALIA CRASSAFORMIS GLOBOROTALIA CRASSATA GLOBOROTALIA EXILIS GLOBOROTALIA FOHSI GLOBOROTALIA GEMMA GLOBOROTALIA HIRSUTA GLOBOROTALIA INCONSTANS GLOBOROTALIA INCREBESCENS GLOBOROTALIA INFLATA GLOBOROTALIA INSOLITA GLOBOROTALIA JUANAI GLOBOROTALIA KUGLERI GLOBOROTALIA ICHINOSEKIENSIS** GLOBOROTALIA LENGUAENSIS **GLOBOROTALIA MARGARITAE** GLOBOROTALIA MARGINODENTATA **GLOBOROTALIA MAYERI GLOBOROTALIA MENARDII GLOBOROTALIA MENDACIS GLOBOROTALIA MEROTUMIDA GLOBOROTALIA MINIMA GLOBOROTALIA MINUTISSIMA GLOBOROTALIA MIOCENICA**

GLOBOROTALIA MIOTUMIDA GLOBOROTALIA MIOZEA GLOBOROTALIA MULTICAMERATA GLOBOROTALIA NANA GLOBOROTALIA OBESA GLOBOROTALIA OPIMA GLOBOROTALIA NYMPHA GLOBOROTALIA PANDA GLOBOROTALIA PERIPHEROACUTA GLOBOROTALIA PERIPHERORONDA GLOBOROTALIA PERMICRA **GLOBOROTALIA PERTENUIS GLOBOROTALIA PLESIOTUMIDA GLOBOROTALIA PLIOZEA GLOBOROTALIA POSTCRETACEA** GLOBOROTALIA PRAEMENARDII **GLOBOROTALIA PSEUDOMENARDII GLOBOROTALIA PSEUDOMIOCENICA GLOBOROTALIA PSEUDOPUMILIO GLOBOROTALIA PUNCTICULATA GLOBOROTALIA QUADRATA GLOBOROTALIA QUETRA GLOBOROTALIA REISSI GLOBOROTALIA SCITULA GLOBOROTALIA SEMIVERA GLOBOROTALIA SIAKENSIS GLOBOROTALIA SUTERAE GLOBOROTALIA TADJIKITANENSIS**

GLOBOROTALIA THEYERI GLOBOROTALIA TOSAENSIS GLOBOROTALIA TRINIDADENSIS GLOBOROTALIA TRUNCATULINOIDES GLOBOROTALIA TUMIDA GLOBOROTALIA UNCINATA GLOBOROTALIA UNGULATA GLOBOROTALIA WHITEI GLOBOROTALIA ZEALANDICA GLOBOROTALIA SHERICOMIOZEA GLOBOCONUSA DAUBJERGENSIS GLOBIGERINOIDES ALTIAPERTURUS **GLOBIGERINOIDES BISPHERICUS** GLOBIGERINOIDES BOLLII **GLOBIGERINOIDES CONGLOBATUS GLOBIGERINOIDES DIMINUTUS GLOBIGERINOIDES FISTULOSUS** GLOBIGERINOIDES HIGGINSI **GLOBIGERINOIDES MITRA GLOBIGERINOIDES PRIMORDIUS GLOBIGERINOIDES RUBER GLOBIGERINOIDES SACCULIFER GLOBIGERINOIDES SICANUS** GLOBIGERINOIDES SUBQUADRATUS **GLOBIGERINOIDES TYRRHENICUS GLOBIGERINOIDES OBLIQUUS GLOBIGERINATELLA INSUETA**

GLOBIGERINA AMPLIAPERTURA GLOBIGERINA ANGULIOFFICINALIS GLOBIGERINA ANGULISUTURALIS GLOBIGERINA APERTURA GLOBIGERINA ATLANTICA GLOBIGERINA BINAIENSIS GLOBIGERINA BRAZIERI **GLOBIGERINA BREVIS GLOBIGERINA BULBOSA GLOBIGERINA BULLOIDES GLOBIGERINA CARIACOENSIS GLOBIGERINA CIPEROENSIS GLOBIGERINA CORPULENTA GLOBIGERINA CRYPTOMPHA GLOBIGERINA DRURYI GLOBIGERINA EAMESI GLOBIGERINA EUAPERTA GLOBIGERINA FALCONENSIS GLOBIGERINA FOLIATA GLOBIGERINA GLOBULARIS GLOBIGERINA GORTANII GLOBIGERINA HAGNI GLOBIGERINA INAEQUISPIRA GLOBIGERINA JUVENILIS GLOBIGERINA LABIACRASSATA GLOBIGERINA LINAPERTA GLOBIGERINA MEGASTOMA GLOBIGERINA MUNDA**

GLOBIGERINA NEPENTHES GLOBIGERINA OFFICINALIS GLOBIGERINA OUACHITAENSIS GLOBIGERINA PATAGONICA GLOBIGERINA PRAEBULLOIDES GLOBIGERINA PRAEDIGITATA GLOBIGERINA PRAETURRITILINA GLOBIGERINA PRASAEPIS GLOBIGERINA PSEUDOAMPLIAPERTURA **GLOBIGERINA PSEUDOCIPEROENSIS GLOBIGERINA PSEUDOBULLOIDES GLOBIGERINA PSEUDOVENEZUELANA GLOBIGERINA RUBESCENS GLOBIGERINA SELLII GLOBIGERINA SENNI GLOBIGERINA TAPURIENSIS GLOBIGERINA TENELLA GLOBIGERINA TRILOCULINOIDES GLOBIGERINA TRIPARTITA GLOBIGERINA TRIVIALIS GLOBIGERINA UTILISINDEX GLOBIGERINA VARIANTA GLOBIGERINA WINKLERI GLOBIGERINA WOODI GLOBIGERINA YEGUAENSIS GLOBIGERINELLOIDES SUBCARINATUS GLOBIGERINELLOIDES MULTISPINATUS**

GLOBIGERINELLA CALIDA GLOBIGERINELLA SIPHONIFERA GLOBOQUADRINA ALTISPIRA GLOBOQUADRINA BAROEMOENENSIS **GLOBOQUADRINA CONGLOMERATA GLOBOQUADRINA DEHISCENS** GLOBOQUADRINA LARMEUI **GLOBOQUADRINA PSEUDOFOLIATA** GLOBOQUADRINA VENEZUELANA **GLOBIGERINATHEKA BARRI GLOBIGERINATHEKA INDEX GLOBIGERINATHEKA MEXICANA GLOBIGERINATHEKA SEMIINVOLUTA GLOBIGERINATHEKA SUBCONGLOBATA** GLOBOROTALOIDES SUTERI GLOBOROTALOIDES TESTARUGOSUS **GLOBOROTALOIDES TREMA GLOBOROTALOIDES VARIABILIS GLOBOROTALOIDES HEXAGONA** GLOBOROTALOIDES OREGONENSIS **GLOBIGERINITA GLUTINATA GLOBIGERINITA HOWEI GLOBIGERINITA PERA GLOBIGERINITA UVULA** HANTKENINA ALABAMENSIS HANTKENINA PRIMITIVA HASTIGERINA PELAGICA **IGORINA PUSILLA**

MOROZOVELLA ACUTA MOROZOVELLA ACUTISPIRA MOROZOVELLA AEQUA MOROZOVELLA ANGULATA MOROZOVELLA ARAGONENSIS MOROZOVELLA CONICOTRUNCATA MOROZOVELLA FORMOSA MOROZOVELLA LEHNERI MOROZOVELLA LENSIFORMIS MOROZOVELLA OCCLUSA MOROZOVELLA SPINULOSA MOROZOVELLA SUBBOTINAE MOROZOVELLA VELASCOENSIS MOROZOVELLA CRATER MOROZOVELLA EDITA MOROZOVELLA NICOLI MOROZOVELLA PRAECURSORIA MURICOGLOBIGERINA AQUIENSIS MURICOGLOBIGERINA CHASCANONA NEOGLOBOQUADRINA ACOSTAENSIS NEOGLOBOQUADRINA ASANOI NEOGLOBOQUADRINA DUTERTREI NEOGLOBOQUADRINA HUMEROSA NEOGLOBOQUADRINA PACHYDERMA NEOGLOBOQUADRINA PSEUDOPIMA **ORBULINA SUTURALIS ORBULINA UNIVERSA**

PLANOROTALITES AUSTRALIFORMIS PLANOROTALITES CHAPMANI PRAEORBULINA GLOMEROSA PRAEORBULINA TRANSITORIA PROTENTELLA CLAVATICAMERATA PROTENTELLA NAVAZUELENSIS PSEUDOHASTIGERINA BARBADOENSIS PSEUDOHASTIGERINA MICRA PSEUDOHASTIGERINA NAGUEWICHIENSIS **PSEUDOHASTIGERINA WILCOXENSIS** PSEUDOHASTIGERINA DANVILLENSIS PULLENIATINA OBLIQUILOCULATA PULLENIATINA PRAECURSOR PULLENIATINA PRIMALIS PULLENIATINA SPECTABILIS SPHAEROIDINELLA DEHISCENS SPHAEROIDINELLOPSIS DISJUNCTA SPHAEROIDINELLOPSIS SEMINULINA SPHAEROIDINELLOPSIS MULTILOBA STREPTOCHILUS GLOBIGERUM SUBBOTINA ANGIPOROIDES SUBBOTINA EOCAENA SUBBOTINA EOCAENICA SUBBOTINA TRIANGULARIS SUBBOTINA TRILOCULARIS SUBBOTINA TURGIDA SUBBOTINA BAKERI SUBBOTINA CROCIAPERTURA

SUBBOTINA EUAPERTURA SUBBOTINA PSEUDOEOCAENA **TENUITELLINATA IOTA TENUITELLA ANGUSTIUMBILICATA** TRUNCOROTALOIDES PSEUDOTOPILENSIS TRUNCOROTALOIDES ROHRI TRUNCOROTALOIDES TOPILENSIS **TURBOROTALIA CERROAZULENSIS** TURBOROTALIA COMPRESSA **TURBOROTALITA HUMILIS** TURBOROTALITA QUINQUELOBA ZEAGLOBIGERINA INCISA ZEAGLOBIGERINA MICROSTOMA EOGLOBIGERINA EOBULLOIDES Nannofossils AMAUROLITHUS AMPLIFICUS AMAUROLITHUS BIZZARUS AMAUROLITHUS DELICATUS AMAUROLITHUS PRIMUS AMAUROLITHUS TRICORNICULATUS AMAUROLITHUS SIGMUNDII ANGULOLITHINA ARCA **BIANTHOLITHUS SPARSUS BICOLUMNUS OVATUS BIRKELUNDIA STAURION** BRAARUDOSPHAERA BIGELOWII BRAARUDOSPHAERA DISCULA

BRAMLETTEIUS SERRACULOIDES CALCIDISCUS LEPTOPORUS CALCIDISCUS MACINTYREI CALCIDISCUS PREMACINTYREI CALCIDISCUS PROTOANNULUS CAMPYLOSPHAERA DELA CAMPYLOSPHAERA EODELA CATINASTER CALYCULUS CATINASTER COALITUS **CERATOLITHUS ACUTUS** CERATOLITHUS ARMATUS CERATOLITHUS CRISTATUS CERATOLITHUS RUGOSUS CERATOLITHUS TELESMUS CHIASMOLITHUS ALTUS CHIASMOLITHUS BIDENS CHIASMOLITHUS CALIFORNICUS CHIASMOLITHUS CONSUETUS CHIASMOLITHUS DANICUS CHIASMOLITHUS EOGRANDIS CHIASMOLITHUS EXPANSUS CHIASMOLITHUS FREQUENS CHIASMOLITHUS GIGAS CHIASMOLITHUS GRANDIS CHIASMOLITHUS OAMARUENSIS CHIASMOLITHUS SOLITUS CHIASMOLITHUS TITUS CHIPHRAGMALITHUS ACANTHODES CHIPHRAGMALITHUS AUSTRIACUS CHIPHRAGMALITHUS CALATHUS CLAUSICOCCUS CRIBELLUM CLAUSICOCCUS FENESTRATUS COCCOLITHUS EOPELAGICUS COCCOLITHUS MIOPELAGICUS COCCOLITHUS PELAGICUS COCCOLITHUS TENUISTRIATUS COCCOLITHUS CRASSIPONS COCCOLITHUS FUSCUS CORANNULUS GERMANICUS CORONOCYCLAS PRIONION CORONOCYCLUS NITESCENS CORONOCYCLUS SERRATUS CRASSIDISCUS BACKMANII CRASPEDOLITHUS DECLIVUS CRENALITHUS PRODUCTELLUS CRIBROCENTRUM RETICULATUM CRUCIPLACOLITHUS EDWARDSII CRUCIPLACOLITHUS PRIMUS CRUCIPLACOLITHUS SUBROTUNDUS CRUCIPLACOLITHUS TENUIS CRUCIPLACOLITHUS NOTUS CRUCIPLACOLITHUS CRIBELLUM **CRUCIPLACOLITHUS CRUCIFORMIS** CRUCIPLACOLITHUS LATIPONS CRUCIPLACOLITHUS TARQUINIUS

CRUCIPLACOLITHUS VANHECKAE CYCLICARGOLITHUS ABISECTUS CYCLICARGOLITHUS FLORIDANUS CYCLICARGOLITHUS PSEUDOGAMMATON CYCLICARGOLITHUS LUMINIS CYCLOCOCCOLITHINA KINGII DICTYOCOCCITES ANTARCTICUS DICTYOCOCCITES BISECTUS DICTYOCOCCITES CALLIDUS DICTYOCOCCITES HESSLANDII DICTYOCOCCITES ONUSTUS DICTYOCOCCITES PRODUCTUS DICTYOCOCCITES SCRIPPSAE DICTYOCOCCITES DAVIESII DISCOASTER ADAMANTEUS DISCOASTER ASYMMETRICUS **DISCOASTER AULAKOS** DISCOASTER BARBADIENSIS **DISCOASTER BELLUS** DISCOASTER BERGGRENII **DISCOASTER BIFAX** DISCOASTER BINODOSUS DISCOASTER BLACKSTOCKAE DISCOASTER BOLLII DISCOASTER BRAARUDII DISCOASTER BRAMLETTEI DISCOASTER BROUWERI **DISCOASTER CALCARIS**

DISCOASTER CRUCIFORMIS DISCOASTER DECORUS DISCOASTER DEFLANDREI **DISCOASTER DELICATUS** DISCOASTER DIASTYPUS **DISCOASTER DISTINCTUS** DISCOASTER DIVARICATUS **DISCOASTER DRUGGII** DISCOASTER ELEGANS **DISCOASTER EXILIS** DISCOASTER EXTENSUS DISCOASTER FALCATUS DISCOASTER FORMOSUS DISCOASTER GEMMIFER DISCOASTER GERMANICUS **DISCOASTER HAMATUS** DISCOASTER INTERCALARIS DISCOASTER KUEPPERI DISCOASTER KUGLERI DISCOASTER LENTICULARIS **DISCOASTER LIDZII DISCOASTER LODOENSIS** DISCOASTER LOEBLICHII DISCOASTER MEDIOSUS **DISCOASTER MIRUS** DISCOASTER MISCONCEPTUS

DISCOASTER MOHLERI

DISCOASTER MOOREI DISCOASTER MULTIRADIATUS DISCOASTER NEOHAMATUS DISCOASTER NEORECTUS DISCOASTER NEPHADOS **DISCOASTER NOBILIS DISCOASTER NODIFER** DISCOASTER NONARADIATUS DISCOASTER PANSUS **DISCOASTER PENTARADIATUS** DISCOASTER PREPENTARADIATUS DISCOASTER PSEUDOVARIABILIS DISCOASTER QUADRAMUS **DISCOASTER QUINQUERAMUS** DISCOASTER ROBUSTUS **DISCOASTER SAIPANENSIS DISCOASTER SALISBURGENSIS** DISCOASTER SANMIGUELENSIS **DISCOASTER SAUNDERSI** DISCOASTER SEPTEMRADIATUS **DISCOASTER SIGNUS** DISCOASTER STELLULUS **DISCOASTER STRICTUS** DISCOASTER SUBLODOENSIS DISCOASTER SUBSURCULUS DISCOASTER SURCULUS DISCOASTER TAMALIS **DISCOASTER TANII**

DISCOASTER TRIDENUS DISCOASTER TRINIDADENSIS **DISCOASTER TRIRADIATUS** DISCOASTER TRISTELLIFER **DISCOASTER VARIABILIS** DISCOASTER WEMMELENSIS DISCOASTER CHALLENGERI **DISCOASTEROIDES MEGASTYPUS DISCOSPHAERA TUBIFERA ELLIPSOLITHUS DISTICHUS** ELLIPSOLITHUS LAJOLLAENSIS ELLIPSOLITHUS MACELLUS EMILIANIA ANNULA EMILIANIA HUXLEYI **ERICSONIA CAVA** ERICSONIA FORMOSA **ERICSONIA OBRUTA ERICSONIA OVALIS** ERICSONIA ROBUSTA **ERICSONIA SUBDISTICHA ERICSONIA SUBPERTUSA FASCICULITHUS BILLII FASCICULITHUS INVOLUTUS FASCICULITHUS MITREUS FASCICULITHUS PILEATUS FASCICULITHUS SCHAUBII FASCICULITHUS TYMPANIFORMIS**

FASCICULITHUS ULII GEMINILITHELLA ROTULA GEPHYROCAPSA APERTA GEPHYROCAPSA CARIBBEANICA **GEPHYROCAPSA ERICSONIA** GEPHYROCAPSA OCEANICA GEPHYROCAPSA PROTOHUXLEYI GEPHYROCAPSA SINUOSA HAYASTER PERPLEXUS **HAYELLA SITULIFORMIS** HELICOSPHAERA AMPLIAPERTA HELICOSPHAERA BRAMLETTEI HELICOSPHAERA CARTERI HELICOSPHAERA COMPACTA HELICOSPHAERA DINESENII **HELICOSPHAERA EUPHRATIS** HELICOSPHAERA GRANULATA HELICOSPHAERA HEEZENI HELICOSPHAERA INTERMEDIA HELICOSPHAERA INVERSA HELICOSPHAERA KAMPTNERI HELICOSPHAERA LOPHOTA HELICOSPHAERA NEOGRANULATA HELICOSPHAERA OBLIQUA HELICOSPHAERA RECTA HELICOSPHAERA RETICULATA HELICOSPHAERA SELLII **HELICOSPHAERA SEMINULUM**

HELICOSPHAERA OMANICA HELICOSPHAERA PERCH-NIELSENIAE **HELICOSPHAERA WILCOXONII HELIOLITHUS CANTABRIAE HELIOLITHUS CONICUS** HELIOLITHUS KLEINPELLII HELIOLITHUS RIEDELII HOLODISCOLITHUS MACROPORUS HORNIBROOKINA AUSTRALIS HORNIBROOKINA TEURIENSIS **ILSELITHINA FUSA** ISTHMOLITHUS RECURVUS LANTERNITHUS MINUTUS LITHOSTROMATION PERDURUM LOPHODOLITHUS ACUTUS LOPHODOLITHUS NASCENS LOPHODOLITHUS MOCHLOPORUS MARKALIUS INVERSUS MICRANTHOLITHUS FLOS MICRANTHOLITHUS ALTUS MINYLITHA CONVALLIS NANNOTETRINA ALATA NANNOTETRINA CRISTATA NANNOTETRINA FULGENS NANNOTETRINA QUADRATA NEOCHIASTOZYGUS CEARAE NEOCHIASTOZYGUS CONCINNUS

NEOCHIASTOZYGUS CHIASTUS NEOCHIASTOZYGUS DISTENTUS NEOCHIASTOZYGUS JUNCTUS NEOCHIASTOZYGUS MODESTUS NEOCHIASTOZYGUS PERFECTUS NEOCOCCOLITHES DUBIUS **NEOCOCCOLITHES PROTENUS OOLITHOTUS FRAGILIS ORTHORHABDUS SERRATUS** ORTHOZYGUS AUREUS **PEDINOCYCLUS LARVALIS** PERITRACHELINA JOIDESA PLACOZYGUS SIGMOIDES PONTOSPHAERA ANISOTREMA PONTOSPHAERA DISCOPORA PONTOSPHAERA DISTINCTA PONTOSPHAERA JAPONICA PONTOSPHAERA JONESII PONTOSPHAERA MULTIPORA PONTOSPHAERA OVATA PONTOSPHAERA PECTINATA PONTOSPHAERA PACIFICA PONTOSPHAERA PLANA PONTOSPHAERA RIMOSA PONTOSPHAERA SCUTELLUM PONTOSPHAERA SEGMENTA PONTOSPHAERA FORMOSA PONTOSPHAERA SCISSURA

PRINSIUS BISULCUS PRINSIUS DIMORPHOSUS PRINSIUS MARTINII **PSEUDOEMILIANIA LACUNOSA PSEUDOTRIQUETRORHABDULUS INVERSUS PYROCYCLUS INVERSUS** PYROCYCLUS ORANGENSIS **RETICULOFENESTRA DICTYODA** RETICULOFENESTRA HAMPDENENSIS **RETICULOFENESTRA HAQII RETICULOFENESTRA HILLAE RETICULOFENESTRA INSIGNATA** RETICULOFENESTRA LOCKERI **RETICULOFENESTRA MINUTA RETICULOFENESTRA MINUTULUS** RETICULOFENESTRA OAMARUENSIS **RETICULOFENESTRA PSEUDOUMBILICA** RETICULOFENESTRA SAMODUROVII **RETICULOFENESTRA UMBILICA RETICULOFENESTRA AMPLA RETICULOFENESTRA ASANOI RETICULOFENESTRA GELIDA RETICULOFENESTRA LONGISTYLIS** RHABDOLITHUS STYLIFER RHABDOLITHUS TENUIS RHADBOSPHAERA CLAVIGERA RHADBOSPHAERA INFLATA

RHADBOSPHAERA PROCERA RHADBOSPHAERA VITREA RHOMBOASTER CUSPIS SCAPHOLITHUS FOSSILIS SCAPHOLITHUS RHOMBIFORMIS SCYPHOSPHAERA AMPHORA SCYPHOSPHAERA APSTEINII SCYPHOSPHAERA CAMPANULA SCYPHOSPHAERA CANTHARELLA SCYPHOSPHAERA CONICA SCYPHOSPHAERA CYLINDRICA SCYPHOSPHAERA EXPANSA SCYPHOSPHAERA GLADSTONENSIS SCYPHOSPHAERA GLOBULATA SCYPHOSPHAERA INTERMEDIA SCYPHOSPHAERA MAGNA SCYPHOSPHAERA PROCERA SCYPHOSPHAERA PULCHERRIMA SCYPHOSPHAERA RECTA SCYPHOSPHAERA RECURVATA SEMIHOLOLITHUS KERABYI SPHENOLITHUS ABIES SPHENOLITHUS ANARRHOPUS SPHENOLITHUS BELEMNOS SPHENOLITHUS CIPEROENSIS SPHENOLITHUS CONICUS SPHENOLITHUS DELPHIX SPHENOLITHUS DISSIMILIS

SPHENOLITHUS DISTENTUS SPHENOLITHUS EDITUS SPHENOLITHUS FURCATOLITHOIDES SPHENOLITHUS HETEROMORPHUS SPHENOLITHUS INTERCALARIS SPHENOLITHUS MORIFORMIS SPHENOLITHUS NEOABIES SPHENOLITHUS OBTUSUS SPHENOLITHUS ORPHANKNOLLI SPHENOLITHUS PREDISTENTUS SPHENOLITHUS PRIMUS SPHENOLITHUS PSEUDORADIANS SPHENOLITHUS RADIANS SPHENOLITHUS SPINIGER SPHENOLITHUS CAPRICORNUTUS STRIATOCOCCOLITHIS PACIFICANUS SYRACOSPHAERA HISTRICA SYRACOSPHAERA PULCHRA THORACOSPHAERA ALBATROSSIANA THORACOSPHAERA DEFLANDREI THORACOSPHAERA HEIMII THORACOSPHAERA OPERCULATA THORACOSPHAERA SAXEA **TOWEIUS CALLOSUS TOWEIUS CRATICULUS TOWEIUS EMINENS** TOWEIUS MAGNICRASSUS

TOWEIUS GAMMATION **TOWEIUS TOVAE TOWEIUS OCCULTATUS TOWEIUS PERTUSUS TOWEIUS CRASSUS** TRANSVERSOPONTIS FIMBRIATUS TRANSVERSOPONTIS OBLIQUIPONS TRANSVERSOPONTIS PULCHER TRIBRACHIATUS CONTORTUS **TRIBRACHIATUS NUNNII** TRIBRACHIATUS ORTHOSTYLUS TRIQUETRORHABDULUS CARINATUS TRIQUETRORHABDULUS MILOWII TRIQUETRORHABDULUS RUGOSUS UMBELLOSPHAERA IRREGULARIS UMBILICOSPHAERA JAFARII **UMBILICOSPHAERA MIRABILIS** UMBILICOSPHAERA SIBOGAE **ZYGODISCUS ADAMAS** ZYGODISCUS PLACTOPONS **ZYGODISCUS SPIRALIS** ZYGODISCUS HERLYNII ZYGRHABLITHUS BIJUGATUS

Radiolarians ACROCUBUS OCTOPYLUS ACROBOTRYS TRITUBUS ACROSPHAERA MURRAYANA ACROSPHAERA SPINOSA ACROSPHAERA AUSTRALIS ACROSPHAERA LABRATA ACROSPHAERA MERCURIUS **ACTINOMMA BEROES** ACTINOMMA GOLOWNINI **ACTINOMMA HOLTEDAHLI ACTINOMMA MEDIANUM** ACTINOMMA MEDUSA ACTINOMMA DELICATULUM ACTINOMMA KERGUELENSIS ACTINOMMA LEPTODERMUM ACTINOMMA POPOFSKII ACTINOMMA MAGNIFENESTRA AMPHYMENIUM CHALLENGERAE AMPHYMENIUM SPLENDIARMATUM AMPHICRASPEDUM PROLIXUM AMPHISTYLUS ANGELINUS AMPHIRHOPALUM YPSILON AMPHIRHOPALUM VIRCHOWII AMPHISPYRIS ROGGENTHENI ANDROSPYRIS ANTHROPISCUS ANOMALACANTHA DENTATA ANTARCTISSA CYLINDRICA ANTARCTISSA DENTICULATA ANTARCTISSA LONGA ANTARCTISSA ROBUSTA

ANTARCTISSA STRELKOVI ANTARCTISSA DEFLANDREI ANTHOCYRTELLA CALLOPISMA ANTHOCYRTIDIUM EURYCLATHRUM ANTHOCYRTIDIUM JENGHISI ANTHOCYRTIDIUM PLIOCENICA ANTHOCYRTIDIUM ANGULARE ANTHOCYRTIDIUM EHRENBERGII ANTHOCYRTIDIUM OPHIRENSE ANTHOCYRTIDIUM MICHELINAE ANTHOCYRTIDIUM NOSICAAE ARTOPHORMIS BARBADENSIS **ARTOPHORMIS GRACILIS** ARTOSTROBUS ANNULATUS AXOPRUNUM ANGELINUM **AXOPRUNUM PIERINAE** BATHROPYRAMIS WOODRINGI **BEKOMIFORMA MYNX** BOTRYOPYLE DICTYOCEPHALUS BOTRYOPYLE DIONISI **BOTRYOPERA TRILOBA** BOTRYOSTROBUS AQUILONARIS **BOTRYOSTROBUS AURITUS-AUSTRALIS** BOTRYOSTROBUS MIRALESTENSIS BOTRYOSTROBUS TUMIDULUS BOTRYOSTROBUS BRAMLETTEI BOTRYOSTROBUS KERGUELENSIS BOTRYOSTROBUS REDNOSUS

BUCCINOSPHAERA INVAGINATA BURYELLA CLINATA CALOCYCLETTA ACANTHOCEPHALA CALOCYCLETTA CAEPA CALOCYCLETTA COSTATA CALOCYCLETTA ROBUSTA CALOCYCLETTA SERRATA **CALOCYCLETTA VIRGINIS** CALOCYCLAS ASPERUM CALOCYCLAS DISPARIDENS CALOCYCLAS HISPIDA CALOCYCLAS TURRIS CALOCYCLAS BANDYCA CALOCYCLOMA AMPULLA CARPOCANOPSIS BRAMLETTEI CARPOCANOPSIS CINGULATA CARPOCANOPSIS FAVOSA CARPOCANOPSIS CRISTATA **CENOSPHAERA CRISTATA** CENOSPHAERA OCEANICA **CENTROBOTRYS GRAVIDA** CENTROBOTRYS PETRUSHEVSKAYAE CENTROBOTRYS THERMOPHILA CERATOCYRTIS AMPLUS CERATOCYRTIS MASHAE CERATOCYRTIS HISTRICOSA CERATOCYRTIS STIGI

CERATOSPYRIS PENTAGONA **CIRCODISCUS ELLIPTICUS** CLATHROCANIUM SPHAEROCEPHALUM CLATHROCYCLAS BICORNIS CLATHROCYCLAS UNIVERSA COLLOSPHAERA ORTHOCONUS COLLOSPHAERA TUBEROSA CORYTHOSPYRIS FISCELLA CORYTHOMELISSA HORRIDA CORNUTELLA PROFUNDA CRYPTOCARPIUM AZYX CRYPTOCARPIUM ORNATUM CYCLADOPHORA GOLLI CYCLADOPHORA DAVISIANA CYCLADOPHORA HUMERUS CYCLADOPHORA SPONGOTHORAX CYCLADOPHORA PLIOCENICA CYCLADOPHORA BICORNIS CYCLADOPHORA ANTIQUA CYCLADOPHORA CABRILLOENSIS CYCLADOPHORA CONICA CYMAETRON SINOLAMPAS CYRTOCAPSELLA CORNUTA CYRTOCAPSELLA ELDHOLMI CYRTOCAPSELLA ELONGATA CYRTOCAPSELLA JAPONICA CYRTOCAPSELLA TETRAPERA CYRTOCAPSELLA ROBUSTA

CYRTOCAPSELLA LONGITHORAX CYRTOLAGENA LAGUNCULA DENDROSPYRIS BURSA DENDROSPYRIS DAMAECORNIS DENDROSPYRIS MEGALOCEPHALIS DENDROSPYRIS STABILIS DENDROSPYRIS RHODOSPYROIDES **DESMOSPYRIS SPONGIOSA** DIARTUS HUGHESI DIARTUS PETTERSSONI DICTYOPRORA AMPHORA DICTYOPRORA ARMADILLO DICTYOPRORA MONGOLFIERI DICTYOPRORA OVATA DICTYOPRORA PIRUM DICTYOPRORA URCEOLUS DICTYOPRORA PHYSOTHORAX **DICTYOPHIMUS CALLOSUS** DICTYOPHIMUS INFABRICATUS DICTYOPHIMUS ARCHIPILIUM DICTYOPHIMUS CRATICULA DICTYOPHIMUS CRISIAE **DICTYOPHIMUS HIRUNDO** DICTYOPHIMUS POCILLUM DICTYOPHIMUS SPLENDENS DICTYOCORYNE ONTONGENSIS DIDYMOCYRTIS ANTEPENULTIMA

DIDYMOCYRTIS AVITA DIDYMOCYRTIS DIDYMUS DIDYMOCYRTIS LATICONUS DIDYMOCYRTIS MAMMIFERA DIDYMOCYRTIS PENULTIMA DIDYMOCYRTIS PRISMATICA DIDYMOCYRTIS TETRATHALAMUS **DIDYMOCYRTIS TUBARIA** DIDYMOCYRTIS VIOLINA **DIDYMOCYRTIS BASSANII** DORCADOSPYRIS ALATA DORCADOSPYRIS ARGISCA DORCADOSPYRIS ATEUCHUS DORCADOSPYRIS CIRCULUS DORCADOSPYRIS CONFLUENS DORCADOSPYRIS DENTATA DORCADOSPYRIS FORCIPATA DORCADOSPYRIS PAPILIO DORCADOSPYRIS PENTAGONA DORCADOSPYRIS PLATYACANTHA DORCADOSPYRIS QUADRIPES DORCADOSPYRIS RIEDILI DORCADOSPYRIS SIMPLEX DORCADOSPYRIS SPINOSA DORCADOSPYRIS TRICEROS DRUPPATRACTUS HASTATUS EUCECRYPHALUS CRASPEDOTA **EUCHITONIA FURCATA**

EUCYRTIDIUM ACUMINATUM EUCYRTIDIUM ANOMALUM EUCYRTIDIUM CALVERTENSE EUCYRTIDIUM CIENKOWSKII EUCYRTIDIUM DIAPHANES EUCYRTIDIUM HEXAGONATUM **EUCYRTIDIUM INFLATUM EUCYRTIDIUM MATUYAMAI** EUCYRTIDIUM PSEUDOINFLATUM **EUCYRTIDIUM BICONICUM EUCYRTIDIUM ANTIQUUM** EUCYRTIDIUM CHENI EUCYRTIDIUM INFUNDIBULUM **EUCYRTIDIUM MARIAE** EUCYRTIDIUM TEUSCHERI EUSYRINGIUM FISTULIGERUM EUSYRINGIUM LAGENA GONDWANARIA DEFLANDREI **GONDWANARIA DOGELI GONDWANARIA HISTER GONDWANARIA JAPONICA** HAECKELIELLA INCONSTANS HALIOMMETTA MIOCENICA HELODISCUS ASTERISCUS HELOTHOLUS PRAEVEMA **HISTIASTRUM MARTINIANUM** LAMPROMITRA CORONATA

LAMPROCYCLAS AEGLES LAMPROCYCLAS HANNAI LAMPROCYCLAS JUNONIS LAMPROCYCLAS MARGATENSIS LAMPROCYCLAS MARITALIS LAMPROCYRTIS DANIELLAE LAMPROCYRTIS HETEROPOROS LAMPROCYRTIS NEOHETEROPOROS LAMPROCYRTIS NIGRINIAE LARCOSPIRA QUADRANGULA LARCOPYLE BUTSCHLII LIPMANELLA DICTYOCERAS LIRIOSPYRIS ELEVATA LIRIOSPYRIS GENICULOSA LIRIOSPYRIS MUTUARIA LIRIOSPYRIS OVALIS LIRIOSPYRIS PARKERAE LIRIOSPYRIS STAUROPORA LITHELIUS MINOR LITHELIUS NAUTILOIDES LITHATRACTUS TIMMSI LITHOCIRCUS TOXARIA LITHOMITRELLA MINUTA LITHOCYCLIA ANGUSTA LITHOCYCLIA ARISTOTELIS LITHOCYCLIA CRUX LITHOCYCLIA OCELLUS LITHOCARPIUM FRAGILIS

LITHOMELISSA CHALLENGERAE LITHOMELISSA EHRENBERGII LITHOMELISSA ROBUSTA LITHOMELISSA SPHAEROCEPHALIS LITHOMELISSA TRICORNIS LITHOMELISSA CHENI LITHOMELISSA DUPLIPHYSA LITHOPERA BACCA LITHOPERA BAUERI LITHOPERA NEOTERA LITHOPERA RENZAE LITHOPERA THORNBURGI LITHOCHYTRIS VESPERTILIO LOPHOCYRTIS BIAURITA LOPHOCYRTIS JACCHIA LOPHOCYRTIS BRACHYTORAX LOPHOCYRTIS LEPTETRUM LOPHOCYRTIS LONGIVENTER LOPHOCYRTIS MILOWI LOPHOCYRTIS NEATUM LOPHOCYRTIS PEGETRUM LOPHOCYRTIS TANYTHORAX LYCHNOCANOMA AMPHITRITE LYCHNOCANOMA BELLUM LYCHNOCANOMA ELONGATA LYCHNOCANOMA GRANDE LYCHNOCANOMA TRIFOLIUM

LYCHNOCANOMA CONICA MITROCALPIS ARANEAFERA **OTOSPHAERA AURICULATA** PENTAPYLONIUM IMPLICATUM PERIPHAENA DECORA PERIPHAENA HELIASTERISCUS PERIPHAENA TRIPYRAMIS PERIPHAENA CIRCUMTEXTA PHORMOCYRTIS EMBOLUM PHORMOCYRTIS STRIATA PHORMOSTICHOARTUS CORBULA PHORMOSTICHOARTUS DOLIOLUM PHORMOSTICHOARTUS MULTISERIATUS PHORMOSTICHOARTUS FISTULA PHORMOSTICHOARTUS FURCASPICULATA PHORMOSTICHOARTUS MARYLANDICUS PHORMOSTICHOARTUS PLATYCEPHALA PHORMOSTICHOARTUS PITOMORPHUS PHORMOSTICHOARTUS CRUSTULA PHORTICIUM CLEVEI PHORTICIUM PYLONIUM PODOCYRTIS AMPLA PODOCYRTIS CHALARA PODOCYRTIS DIAMESA PODOCYRTIS DORUS PODOCYRTIS FASCIOLATA PODOCYRTIS GOETHEANA **PODOCYRTIS MITRA**

PODOCYRTIS PAPALIS PODOCYRTIS SINUOSA PODOCYRTIS TRACHODES **PRUNOPYLE ANTARCTICA** PRUNOPYLE FRAKESI PRUNOPYLE HAYESI **PRUNOPYLE TETRAPILA PRUNOPYLE TYTAN** PRUNOPYLE MONIKAE PRUNOPYLE POLYACANTHA PRUNOPYLE TRYPOPYRENA PERICHLAMYDIUM PRAETEXTUM PSEUDOCUBUS VEMA PSEUDODICTYOPHIMUS GRACILIPES PSEUDODICTYOPHIMUS GALEATUS PTEROCORYS CAMPANULA PTEROCORYS CLAUSUS PTEROCORYS HERTWIGII PTEROCORYS ZANCLEUS PTEROCANIUM CHARYBDEUM PTEROCANIUM KOROTNEVI PTEROCANIUM PRAETEXTUM PTEROCANIUM PRISMATIUM PTEROCANIUM AUDAX **RHABDOLITHIS PIPA** RHIZOSPHAERA ANTARCTICA RHOPALOCANIUM ORNATUM

SACCOSPYRIS ANTARCTICA SACCOSPYRIS CONITHORAX SACCOSPYRIS PREANTARCTICA SATURNALIS CIRCULARIS SETHOCHYTRIS BABYLONIS SETHOCHYTRIS TRICONISCUS SIPHOCAMPE ACEPHALA SIPHOCAMPE ARACHNEA SIPHOCAMPE LINEATA SIPHOCAMPE NODOSARIA SIPHOCAMPE IMBRICATA SIPHOCAMPE PACHYDERMA SIPHONOSPHAERA VESUVIUS SIPHONOSPHAERA MAGNISPHAERA SIPHOSTICHARTUS CORONA SOLENOSPHAERA OMNITUBUS SPHAEROPYLE LANGII SPHAEROPYLE ROBUSTA SPIROCYRTIS GYROSCALARIS SPIROCYRTIS SCALARIS SPIROCYRTIS SUBSCALARIS SPIROCYRTIS SUBTILIS SPONGODISCUS AMBUS SPONGODISCUS CRATICULATUS SPONGODISCUS OSCULOSUS SPONGODISCUS KLINGI SPONGOMELISSA DILLI SPONGASTER BERMINGHAMI
SPONGASTER PENTAS SPONGASTER TETRAS SPONGOCORE PUELLA SPONGOTROCHUS GLACIALIS SPONGOTROCHUS VENUSTUM SPONGURUS PYLOMATICUS STAUROXIPHOS COMMUNIS STICHOCORYS ARMATA STICHOCORYS DELMONTENSIS STICHOCORYS PEREGRINA STICHOCORYS JOHNSONI STICHOCORYS WOLFII STYLATRACTUS CORONATUS STYLATRACTUS NEPTUNUS STYLATRACTUS SANTANNAE STYLATRACTUS UNIVERSUS STYLODICTYA TARGAEFORMIS STYLODICTYA VALIDISPINA STYLODICTYA HASTATA STYLODICTYA OCELLATA STYLOSPHAERA ANGELINA STYLOSPHAERA HISPIDA STYLOSPHAERA RADIOSA STYLACONTARIUM ACQUILONIUM STYLACONTARIUM BISPICULUM THEOCOTYLISSA FICUS THEOCYRTIS ANNOSA

THEOCYRTIS DIABLOENSIS THEOCYRTIS TUBEROSA THEOCYRTIS ROBUSTA THEOCORYS REDONDENSIS THEOCORYS SPONGOCONUS THEOCOTYLE CONICA THEOCOTYLE CRYPTOCEPHALA THEOCORYTHIUM TRACHELIUM THEOCORYTHIUM VETULUM **THOLOSPYRIS RHOMBUS** THYRSOCYRTIS BROMIA THYRSOCYRTIS CLAUSA THYRSOCYRTIS RHIZODON THYRSOCYRTIS TENSA THYRSOCYRTIS TETRACANTHA THYRSOCYRTIS TRICANTHA TRICERASPYRIS ANTARCTICA TRICERASPYRIS CORONATA TRICOLOCAPSA PAPILLOSA **TRIPILIDIUM CLAVIPES** TRISTYLOSPYRIS TRICEROS **VELICUCULLUS ALTUS** ZYGOCIRCUS BUTSCHLII

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Table 4.2: Average species longevities for 1418 Cenozoic marine plankton species contained in Neptune. Longevities and standard deviations are in million years.

| Plankton Group | Number of species | Extant species (in %) | Mean Longevity of Extant Species | Standard Deviation | Mean Longevity of Extinct Species | Standard Deviation |
|-------------------|-------------------------|-----------------------------|---|-----------------------|--|-----------------------|
| Diatoms | 389 | 26.5 | 14.4 | 13.7 | 10.2 | 10.3 |
| Foraminifera | 281 | 13.9 | 13.9 | 9.2 | 11.1 | 7.8 |
| Radiolarians | 383 | 29.8 | 13.4 | 9.2 | 12.9 | 7.8 |
| Nannofossils | 365 | 8.5 | 17.8 | 19.8 | 13.1 | 10.6 |
| All groups | 1418 | 20.2 | 14.3 | 12.4 | 11.9 | 9.4 |

5. A MORE REALISTIC ASSESSMENT: NEPTUNE AS A TOOL IN SUPPORT OF DATA IMPROVEMENT

Eight years after the first planning meetings at the ETH, I have finally achieved a balance between the initial expectations/dreams and the realization of what Neptune really is and its limitations. I have described the scopes, accomplishments, and drawbacks of this project in the previous chapters and now it is time to answer the question: what next? What is the future of Neptune? In the previous chapters, I tried to convey that Neptune is not a sufficient data set to base several (biostratigraphic) publications on, but mainly a tool to make collecting new data more focused and more efficient. And this is the legacy of this project to the research community.

The chronology of the Neogene sediments of some 100 DSDP holes was published as an Ocean Drilling Program Technical Note (Lazarus et al. 1995a) and is currently available through the WWW site of the NOAA- National Geographic Data Center (NGDC) (<u>http://www.ngdc.noaa.gov/mgg/geology/ lazarus.html</u>). The age models for these sites were based on <u>Berggren et al. (1985</u>). The updated models (to <u>Berggren</u> <u>et al. 1995a, b</u>) extended to the whole Cenozoic, as well as additional ODP holes, are published here in graphic form and as text files in the <u>Appendix A</u>. At the same time, a link to them will be deposited at the widely used archival site of the NGDC.

At present, the database is accessible through the author, at the ETH Zürich, and at the Natural History Museum in Basel, Switzerland. It is still unclear how the whole database with its search options will be made accessible to the community. Among the options discussed are a CD-ROM (which would, however, require the relatively expensive 4th Dimension® program to run) and a server at the Micropaleontological Reference Center (MRC) in Basel. The optimal solution for this second option would be an interactive WWW site that could be remotely accessible world-wide. Recently, the program NetLink/4DTM has been made available on the market. This program apparently makes databases searchable through the Internet (Lazarus, personal commun., 1998). However, I have not seen nor tried the program yet and do not know how user-friendly it is. However, I suspect that the large size of Neptune would make even the simplest searches very slow and time consuming through the Internet. A more modest, but immediately feasible alternative, would be to have one person at the MRC in charge of the use of Neptune. Requests for searches could be e-mailed to the MRC and the results (in print or as computer files) mailed to the requester. Among the various possibilities, searches for presence/absence and location or number of occurrence of single or multiple taxa would require only a few minutes. This search would also provide information on the taxonomic validity of the taxa and the lists of synonyms. A more extensive search would distinguish between stratigraphically and thematically well and poorly covered intervals. The identification of significant gaps in the biostratigraphic record (e.g., Paleogene biochronology of siliceous microfossils and



revisitation of suitable Paleogene sediments for detailed biostratigraphy) (Fig. 5.1) would be the basis for the logical, objective planning of future research. It could spur clearly aimed detailed micropaleontological studies, instead of random studies that generate a lot of repetition and overlap (e.g., Moore 1972).

Considering that Neptune contains selected, good quality holes, it is still notable how small the number of useful holes (well cored, well analyzed and well documented with modern biochronological methods and modern taxonomy) has remained. An enormous amount of re-analysis of older sections could be quite profitable. Fig. 5.1 indicates that a lot can still be done on sections older than the late Miocene, especially on siliceous plankton groups. New coring needed to fill the existing

coverage gaps could be identified with three dimensional (latitude vs. longitude vs. time) maps of the oceans produced with Neptune. Another way to identify stratigraphic coverage gaps is given by the rate of success in recovering drilled sections (Fig.

5.2). This curve indicates that the early and middle Eocene, as well as large parts of the Miocene, have



been less well recovered than e.g., the Plio-Pleistocene or the Eocene/Oligocene boundary. This might be one of the causes of the poor Eocene biochronology for certain plankton groups. Recent ODP Legs (e.g., <u>171B</u>) have recovered long Eocene sections: these should be studied in detail to cover this recovery gap.

These are just a couple of examples of the utility of Neptune in designing goal-oriented studies aimed to obtain a complete picture of the oceans' history during the Cenozoic, necessary for a better understanding of the complex processes that control the Earth systems. This approach would, however, require the change in nature of the DSDP and ODP projects from 'leg oriented' to 'overview oriented', which in my opinion is a more effective investment of resources. This step would certainly represent the most valuable contribution of Neptune to the geological community.



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Figure 5.1. Number of reports on Cenozoic biostratigraphy in Neptune by plankton group. Reports that cover the whole or only a part of the series are included. No selection based on the detail or the quality of the reports was made. Note the overall better availability of nannofossil biostratigraphy in comparison with the other plankton groups. The anomalously low number of reports for foraminifera, diatoms and radiolarians may be in part due to the lower number of Pleistocene sections available in Neptune (Fig. 3.2). However, it likely represents also an artifact of the analysis caused by the scarcity of biostratigraphic events of these three groups calibrated for the Pleistocene.



http://www-odp.tamu.edu/paleo/1999_2/neptune/fig5_1.htm (1 of 2) [10/26/1999 10:03:01 AM]

Figure 5.1: THE CENOZOIC DEEP SEA MICROFOSSIL RECORD: EXPLORATIONS OF THE DSDP/ODP SAMPLE SET USING THE NEPTUNE DATABASE

| 10 | | | | | | | | | | | | |
|----|-------------|----------|---------------|----------------|---------------|-----------------|-----------------|--------------|--------------|--------------|-----------|--|
| | Pleistocene | Pliccene | Upper Miccene | Middle Miccene | Lower Miccene | Upper Oligooene | Lower Oligooene | Upper Eccene | Middle Econe | Lower Eccene | Paleocene | |

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Figure 5.2. Ratio between the number of recovered sections and the total number of times recovery was attempted at 1-m.y. resolution. Sections that recovered sediments containing at least 5% of the one m.y. interval are included. The horizontal line represents the average recovery ratio for all the sections analysed (0.89). The data are limited to the sections included in Neptune and for which we have age control. This means that in some instances, deeper sediments were also recovered, but no biostratigraphic study done on them to provide us with an age model. This graph does not reflect the 'true' availability of sediments but only the time interval that was recovered in each section - it includes also intervals that are represented by hiati (see Fig. 3.3), which are e.g., quite frequent in Miocene sediments. Note that Paleocene sediments are overall quite well recovered even though scarce (Fig. 3.2), an improvement from the early study by Moore (1972) which identified this interval as the least well represented in the Cenozoic sediments recovered by the early DSDP legs.



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Appendix A. Age/Depth Plots and Age Model files

One or two plots are given for each hole analysed. Where two plots are available, they cover the Neogene (N in the file name) and the Paleogene (P in the file name) sections, respectively.

All plots were created with the ADP program described in Lazarus (1992), which gives additional details on plotting procedures, conventions, and usage. The x-axis represents the age in million years. There is also a reproduction at the bottom of the plot of the Berggren et al. (1995b) magnetostratigraphic scale. On the left y-axis is the depth in meters below sea floor, on the right a representation of the cores (numbered boxes, with heights proportional to actual recovery) and (short lines on right side of boxes) 1.5 m section breaks. The symbols corresponding to the microfossil groups are on the top of the plot: D=diatoms, F=foraminifera, M=magnetostratigraphy, N=nannofossils, R=radiolarians. Each event is labelled only with a short plotcode to reduce visual clutter in the figure. Translations for the plotcodes are given in Table 2.2. Error bars for the depth level of each event are represented as vertical lines crossing the symbols, but sometimes the error bars are not visible because they are smaller than the symbol itself.

ASCII files of the age models are also given. Only one file is given for each hole. The files consist of two columns. The first row shows the hole name (e.g. 62A) on the left and the date the age model was created on the right (in the format YYMMDD, e.g., 19950725). The second row gives the headings of the data columns, age to the left and depth to the right. The third row gives the number of points in the age model. The coordinates of the age model points follow from the fourth row. The files are exactly in the format generated by Lazarus's (1994) ADP program. This format is required by the ADP program and can, therefore, be used directly to create an age/depth plot with this software. Alternatively, the age and depth coordinates can be used to construct a simple xy line graph. The files can be opened with any word processing or spreadsheet program.

To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number).

DIRECTORY: adps_app

Next Section...

Holes 62A-192

To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: <u>adps_app</u>

62A 19950725



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen1.htm (1 of 5) [10/26/1999 10:03:07 AM]


0.26997723 3.89784

0.50953233 9.27418





158 19952507

AGE DEPTH

6

0.23538477 1.55E-07



Aga (Ma)

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5

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To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: <u>adps_app</u>

Holes 213-289

213 19952507



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen2.htm (1 of 8) [10/26/1999 10:03:09 AM]

| 60.021 | 327.642 |
|---------|---------|
| 61.292 | 331.853 |
| 65.2206 | 331.853 |
| 65.5095 | 333.002 |



233A 19952507

AGE DEPTH

3

0 0

2.7295732 155.484



1.88222182 12.6344



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen2.htm (3 of 8) [10/26/1999 10:03:09 AM]





24.9632 420.833

253 19950725



265 19952507





9.08219e-2 .755





203 13330123

AGE DEPTH

22

.98226 1.51514



To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: adps_app

Holes 317-398D

317B 19952407

| AGE DEPTH | |
|------------|----------|
| 11 | |
| 0.08559784 | 0.589391 |
| 1.82097273 | 10.0196 |
| 3.78595 | 25.3438 |
| 4.940436 | 56.5815 |
| 7.88500493 | 87.2298 |
| 9.74485053 | 113.752 |
| 11.8377352 | 131.434 |
| 12.8819848 | 147.348 |
| 16.7658775 | 173.87 |
| 19.0025073 | 195.678 |
| 23.72128 | 237.397 |



354 19950709

AGE DEPTH

- -7.35239e-3 .378792
- 2.20589 94.3182
- 3.08351 142.424
- 5.59244 142.424
- 6.60294 239.



AGE DEPTH

13

- 7.79154e-3 .159091
- 4.84164 17.9848

23.6765 39.9697





http://www-odp.tamu.edu/paleo/1999_2/neptune/appen3.htm (3 of 7) [10/26/1999 10:03:11 AM]



366A 19950710











14.7848705 104.126





http://www-odp.tamu.edu/paleo/1999_2/neptune/appen3.htm (6 of 7) [10/26/1999 10:03:11 AM]



To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: adps_app

Holes 406-499





http://www-odp.tamu.edu/paleo/1999_2/neptune/appen4.htm (1 of 11) [10/26/1999 10:03:14 AM]







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| AGE | DEPTH | |
|------------|----------|----|
| 11 | | |
| 2.86E-06 | 0.785854 | |
| 0.51356375 | 33.0059 | |
| 1.73792353 | 88.0157 | 14 |
| 2.11830588 | 110.685 | - |
| 2.9214148 | 158.743 | Å |
| 4.77656625 | 191.748 | |
| 5.79136621 | 247.544 | |
| 7.15975577 | 247.544 | |
| 8.34119649 | 310.413 | |
| 12.1370338 | 310.413 | |
| 13.2941932 | 346.562 | |



438A 19952507

AGE DEPTH





440B 19952507





AGE DEPTH

8

23.5137 43.6447

24.5116 76.5697





AGE DEPTH





| AGE | DEPTH |
|-----------|---------|
| 14.342376 | .589391 |
| .534961 | 21.8075 |



470A 19952507









495 19950725





To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: <u>adps_app</u>

Holes 572C-594

572C 19952507



Appendix A: Holes 572C-594

573A 19952507



573B 19952407







| AGE | DEPTH | Age Depth Piol of Hole 578 | 1.0 |
|------------|---------|---|------------------|
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| 0.29917808 | 0.51 | | |
| 0.78361317 | 12.7419 | A CONTRACT OF A | - |
| 1.02575086 | 17.1774 | | |
| 1.60206176 | 22.0521 | | - |
| 1.78543909 | 25.0383 | | |
| 1.97989229 | 27.5651 | | -94 |
| 2.20811108 | 31.853 | | |
| 2.5934644 | 35.8346 | en Les la la companya de la | |
| 3.17876222 | 43.7903 | | n N Nerval |
| 3.324797 | 45.5645 | | |
| 3.9908125 | 56.2097 | | |
| 4.624995 | 66.7688 | | |
| 5.27888414 | 76.0337 | | |
| 5.58053476 | 76.0337 | | |

- 6.115724 93.9516
- 7.459755 124.866
- 9.12527575 159.812

574C 19950725



. –

| 15 | | |
|---------|---------|-----|
| .235068 | 3.445 | |
| .295065 | 4.5023 | |
| 1.34111 | 4.5023 | |
| 1.7856 | 7.81011 | 144 |
| 2.66708 | 12.0368 | - |
| 3.53242 | 18.1048 | |
| 4.24741 | 25.6355 | |
| 5.43856 | 35.4391 | |
| 6.56324 | 45.8805 | |
| 7.81077 | 56.6256 | |
| 8.90698 | 68.9522 | |
| 9.73687 | 78.3333 | |
| 12.91 | 130.2 | |
| 14.1958 | 151.387 | |
| 16.6846 | 207.504 | |
| | | |

45

6.27







| 4.49847 | 45.9677 | |
|--------------|---------|---|
| 4.9905 | 51.3441 | |
| 5.23801379 | 54.0995 | |
| 5.87812124 | 58.5299 | |
| 34.9487527 | 58.5299 | • |
| 36.446 | 64.7652 | - |
| 37.0691 | 66.9394 | |
| 48.5955 | 66.9394 | |
| 52.3592 | 75.5758 | |
| 52.6194 | 76.697 | |
| 52.7843 | 78.6515 | |
| 52.8268 | 79.7121 | |
| 54.8247 | 81.9394 | |
| 55.9299 | 85.6515 | |
| 56.3975 | 87.4545 | |
| 57.6302 | 92.6515 | |
| 58.4506 | 96.0833 | |
| 59.7374 | 98.25 | |
| 62.3911 | 102.409 | |
| 64.644 | 108.561 | |
| 65.578 | 112.5 | |
| 578 19952507 | | |
| AGE | NEDTH | |



AGE П

- -0.0095231 2.8229E-08
- 0.78635833 27.3813
- 0.99368914 32.5268
- 1.55028235 48.147
| 2.0707349 | 61.6935 |
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| 2.8150594 | 78.1011 |
| 3.642725 | 88.6559 |
| 5.83493834 | 109.709 |
| 6.61827383 | 115.957 |
| 7.81642502 | 122.177 |
| 12.3023485 | 135.618 |
| 15.155 | 144.91 |
| | |





AGE DEPTH

6

4.6785675 181.011





586B 19952507



9.8853 239.839







27.0723 457.955



| AGE | DEPTH | | Age Depin Piol of Hole 586 | 11 M 11 |
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| 6.003971 | 124.579 | | | |
| 6.61926783 | 144.355 | | | |

- 7.16871308 158.576 7.4171053 174.885 7.69418242 181.048 8.27060278 186.371 9.02869 204.747 9.24030467 206.585
- 9.56333029 215.773
- 10.9826883 238.441

9



590B 19950803

AGE DEPTH

- .768099 11.5591
- 1.75824 31.6667
- 1.95369 35.8064
- 2.57121 45.8602
- 3.03085 57.3925



7.68561484 294.793

| AGE | DEPTH | | Age Depth Plot of Hole 591 | an. |
|------------|---------|----------------------|--|----------------------|
| 16 | | Þ | and the second second | • |
| 0.0779516 | 2.82259 | | | 2 + 8 |
| 0.29754436 | 3.89784 | | 2 1973 306 10 1 CE 1967 306 306 | 7 6 9 0 |
| 0.7785554 | 16.6667 | D -100 - 6 P | T H | L U L |
| 0.995248 | 23.3871 | н - | a) an | 10 17 14 |
| 1.07318809 | 25.4032 | т 6 8 - 200 | and the second se | 19 20 21 0 |
| 1.77666818 | 44.8925 | | | 50 54 55 26 |
| 1.95368975 | 49.5968 | | | 57 33 59 51 |
| 2.5816528 | 67.0699 | -500 | - TEF | |
| 3.04082 | 82.5269 | L004e:9916 | e a Age (Ma) and the set of the s |) (1 |
| 3.324797 | 89.2473 | | | |
| 3.50827273 | 95.7121 | | | |
| 3.7136375 | 117.917 | | | |
| 4.58829 | 189.893 | | | |
| | | | | |

9.78029667 358.346

14.8491822 449.464



Appendix A: Holes 572C-594

593 19950802





0.45419721 43.951

| 1.76051912 | 104.747 |
|------------|---------|
| 2.25941459 | 104.747 |
| 2.7064498 | 127.565 |
| 3.268774 | 127.565 |
| 3.6024875 | 158.652 |
| 4.79643375 | 172.818 |
| 5.44970386 | 184.763 |
| 6.47276105 | 184.763 |
| 7.59911086 | 209.265 |
| 8.48852682 | 209.265 |
| 8.95168155 | 222.818 |
| 9.31007238 | 227.412 |
| 9.81005227 | 227.412 |
| 9.84377733 | 273.354 |
| 11.8602283 | 395.559 |
| 13.522587 | 497.397 |
| 14.2159278 | 533.844 |
| 15.4730108 | 533.844 |
| 17.4801073 | 633.537 |
| | |

| AGE | DEPTH |
|-----------|-----------|
| 17 | |
| 5.6594e-7 | 4.7047e-8 |
| .339205 | 46.8606 |
| 1.76052 | 104.747 |
| 2.25941 | 104.747 |
| 2.63589 | 144.87 |
| 3.26877 | 144.87 |





To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: <u>adps_app</u>

Holes 603C-699A

603C 19952507



603 19950726



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen6.htm (1 of 14) [10/26/1999 10:03:20 AM]

| 4.05271 | 215.726 |
|---------|---------|
| 4.49092 | 239.919 |
| 4.81995 | 263.4 |
| 5.28041 | 296.478 |
| 9.81513 | 667.534 |
| 12.1318 | 798.775 |



| AGE | DEPTH |
|------------|---------|
| 15 | |
| 0 | 0 |
| 0.3232117 | 5.91397 |
| 0.51143318 | 10.6183 |
| 0.78361433 | 20.0269 |
| 0.9662945 | 23.3871 |
| 1.07776794 | 24.0591 |
| 1.762825 | 52.957 |







| AGE | DEPTH | |
|------------|---------|---|
| 19 | | |
| 0.41637148 | 14.5161 | |
| 0.77760658 | 34.2742 | |
| 1.08509426 | 46.5054 | 1 |
| 1.762825 | 72.8494 | |
| 2.57905353 | 109.504 | |
| 3.0362974 | 127.419 | |
| 3.35394318 | 138.71 | |
| 3.57870455 | 145.661 | |
| 4.19835778 | 176.344 | |
| 4.29730769 | 180.372 | |
| 4.48796 | 191.942 | |
| 4.6131 | 194.22 | |
| 4.80438429 | 202.066 | |
| 4.900989 | 208.333 | |



| 4.976319 | 210.744 |
|------------|---------|
| 5.23312538 | 223.76 |
| 5.37928552 | 238.507 |
| 5.9916185 | 238.507 |
| 6.7584515 | 279.764 |



608 19950802

| -7.1163e-3 | .134415 |
|------------|---------|
| 1.07777 | 20.0269 |
| 1.76831 | 48.9247 |
| 1.94845 | 55.6451 |
| 2.59073 | 84.543 |
| 3.01543 | 92.6075 |
| 3.09687 | 101.344 |
| 3.586 | 110.081 |



AGEDEPTH190000.7836131743.9516

0.99805371 56.8181

| 1.06558629 | 64.2562 | |
|------------|---------|-----|
| 1.76831176 | 118.952 | |
| 1.94845364 | 132.258 | |
| 2.57882893 | 170.968 | |
| 3.12222222 | 219.355 | 0.4 |
| 3.6004625 | 241.129 | - |
| 4.17875 | 256.855 | |
| 4.31889462 | 261.364 | |
| 4.47003231 | 265.083 | |
| 4.62102375 | 272.521 | |
| 4.79940375 | 277.419 | |
| 4.896543 | 281.818 | |
| 5.8926949 | 332.025 | |
| 6.131789 | 337.603 | |
| 6.2635968 | 346.901 | |
| 6.905599 | 374.793 | |





http://www-odp.tamu.edu/paleo/1999_2/neptune/appen6.htm (6 of 14) [10/26/1999 10:03:20 AM]

4.894887 199.902

610 19952407

| AGE | DEPTH |
|------------|---------|
| 14 | |
| 7.1044E-07 | 9.6876E |
| 0.78048067 | 42.7308 |
| 1.77002455 | 91.8467 |
| 1.96329378 | 100.196 |
| 2.5878952 | 127.701 |
| 3.04705 | 151.747 |
| 3.5879625 | 168.959 |
| 4.20923556 | 175.269 |
| 4.4921 | 191.398 |
| 4.77957 | 199.018 |
| 16.7284281 | 653.831 |
| 17.3728691 | 671.12 |
| 18.2655609 | 678.978 |
| 19.6353798 | 721.414 |

612 19950718

AGE DEPTH

10

- 22.423 119.142
- 23.7832 138.285
- 33.4325 138.285
- 35.5579 182.695
- 42.1041 182.695
- 44.6121 320.521

49.0755 320.521







659A 19950803

Appendix A: Holes 603C-699A



689B 19950802



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen6.htm (9 of 14) [10/26/1999 10:03:20 AM]

| 5.31542 | 17.7795 |
|---------|---------|
| 7.09735 | 17.7795 |
| 7.16589 | 18.8208 |
| 7.42861 | 19.0046 |
| 8.26817 | 20.1685 |
| 8.70223 | 22.2511 |
| 9.31906 | 23.7825 |
| 9.5932 | 23.7825 |
| 9.73598 | 24.7014 |
| 9.83824 | 29.8333 |
| 9.88235 | 34.3333 |
| 9.88971 | 37.3333 |
| 11.3971 | 37.3333 |
| 11.9485 | 46. |
| 12.83 | 48.55 |
| 13.01 | 48.8 |
| 13.2 | 49.55 |
| 13.46 | 50.77 |
| 13.69 | 51.8 |
| 14.08 | 52.27 |
| 14.2 | 52.55 |
| 15.1471 | 57.2727 |
| 15.7604 | 58.7862 |
| 17.7237 | 58.7862 |
| 18.2553 | 61.2273 |
| 19.5841 | 67.0455 |
| 24.8388 | 67.0455 |
| 25.625 | 68.7576 |
| 26.0063 | 70.3333 |

| 79.303 | | |
|----------------|--|--|
| 84.8788 | | |
| 89.8485 | | |
| 91.8182 | | |
| 100.273 | | |
| 106.818 | | |
| 116.818 | | |
| 128.182 | | |
| 139.667 | | |
| 145 | | |
| 152.273 | | |
| 153.182 | | |
| 162.727 | | |
| 170.909 | | |
| 183.182 | | |
| 201 | | |
| 201 | | |
| 227 | | |
| 227 | | |
| 239.545 | | |
| 690 B 19950802 | | |
| | | |

| AGE | DEPTH |
|---------|---------|
| 48 | |
| 0 | 0 |
| .172814 | 1.93308 |
| 2.41719 | 1.93308 |
| 2.59424 | 4.96212 |



| 36.3456 | 91.5454 |
|---------|---------|
| 36.6197 | 96.5682 |
| 37.6155 | 99.4545 |
| 38.2742 | 101.553 |
| 40.4107 | 101.553 |
| 42.5389 | 106.273 |
| 43.6712 | 118.636 |
| 47.2416 | 118.636 |
| 49.1711 | 131.7 |
| 51.6419 | 131.7 |
| 52.678 | 139.74 |
| 52.8905 | 154.671 |
| 55.9724 | 185.299 |
| 56.3975 | 195.635 |
| 58.1509 | 216.692 |



Appendix A: Holes 603C-699A

699A 19950719



Next Section...

To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: adps_app

Holes 700B-736A

700B 19950719



703A 19950719

| AGE | DEPTH |
|-----|-------|
| | |

- 11
- 24.1977 46.8606
- 24.7556 53.5988
- 28.2625 53.5988
- 28.7673 66.7687
- 29.4049 78.1818
- 30.1222 86.3636
- 30.9724 95.8652
- 31.1583 97.5
- 32.1945 97.5







707C 19950720

AGE DEPTH



1.4958 15.5455



709C 19950720



710 A 19950802

AGE DEPTH

23

.414918 3.



1.60177 14.697









714A 19950720





721B 19950731



722A 19950731

AGE DEPTH







724B 19950729

AGE DEPTH

10

8.40446e-3 1.83767





736A 19950730

AGE DEPTH



To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number). DIRECTORY: <u>adps_app</u>

Holes 737A-797B

737A 19950730








AGE DEPTH 26 .919121 2.63636 2.57353 15.1818





AGE DEPTH

12

.182774 11.8683





747A 19950801

AGE DEPTH

41

-3.27093e-2 9.09087e-2

2.32165 21.2955





| 25.1765 | 132.925 |
|---------|---------|
| 25.6282 | 138.285 |
| 25.8908 | 138.285 |
| 26.5525 | 141.654 |
| 28.2857 | 154.671 |
| 29.543 | 170.917 |
| 49.9628 | 170.917 |
| 53.3581 | 173.197 |
| 63.4006 | 173.197 |
| 64.8831 | 182.53 |



2.59554 2.78939



AGE DEPTH

8

26.8704 176.187



749B 19950723



751A 19950801

- AGE DEPTH 22
- 2.54424e-2 8.33337e-2

.642265 1.75

| 1.69315 | 1.75 |
|---------|---------|
| 2.01298 | 7.41667 |
| 3.42437 | 7.41667 |
| 4.81153 | 35.6667 |
| 6.20509 | 35.6667 |
| 6.6163 | 41.8333 |
| 9.16355 | 41.8333 |
| 9.32347 | 44.9167 |
| 9.57477 | 44.9167 |
| 10.2201 | 70.2727 |
| 10.4989 | 70.2727 |
| 12.2526 | 103.364 |
| 12.9685 | 103.364 |
| 13.74 | 109.833 |
| 14.0063 | 109.833 |
| 14.9963 | 131.333 |
| 15.959 | 131.333 |
| 17.2941 | 147.841 |
| 18. | 147.841 |
| 18 8655 | 159 848 |



| AGE | DEPTH |
|------------|---------|
| 14 | |
| 3.67681e-2 | .757579 |
| 1.06618 | 13.4848 |
| 2.57353 | 37.7273 |
| 3.22269 | 51.2121 |
| 4.81092 | 84.9394 |





| AGE | DEPTH | -12-1 | Age Depth Plot of Hole 7690 | 1999 |
|---------|---------|--------------|---|--------|
| 10 | | | | - |
| 33.8718 | 168.182 | 20 - | | 4 |
| 37.0231 | 263.864 | | • Winter (* 1997) | |
| 40.458 | 277.727 | 0 + - 320 | Contraction of the second s | 2 |
| 43.7513 | 277.727 | | | |
| 47.8108 | 330.475 | an a | | |
| 52.3592 | 369.525 | | a star | |
| 55.9013 | 420.455 | 90 - | | |
| 61.3157 | 519.045 | - | | 64 4 A |
| 62.4076 | 519.045 | | 10 0 1 0 1 0 1 0 1 0 1 0 1 0 1 0 1 0 1 | |
| 64.8514 | 553.273 | | | |

786A 19950724

AGE DEPTH



794A 19950731



795A 19950731

AGE DEPTH

5









To download full-size diagrams, open the following directory and choose the file (named by hole number). File naming convention is 62A_PICT.GIF (62A = hole number).

DIRECTORY: adps_app

Holes 803B-841B

803D 19950731



806B 19950731

AGEDEPTH81.2524.65543.7594.33385.51471161.8688.63975284.07310.6748365.68213.1987578.867





| 5.24685 | 73.5454 |
|---------|---------|
| 5.89953 | 76.8182 |
| 6.55374 | 81.7045 |
| 7.43224 | 85.0758 |
| 8.06776 | 87.197 |
| 8.25935 | 88.1818 |
| 8.69393 | 88.5606 |
| 9.74055 | 90.9697 |
| 13.1513 | 98.8182 |
| 18.2983 | 98.8182 |
| | |

18.7447 99.1818

815A 19950731



10





841B 19950802



http://www-odp.tamu.edu/paleo/1999_2/neptune/appen8.htm (4 of 5) [10/26/1999 10:03:26 AM]

38.5767 607.81