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MEASURED SECTIONS OF PALEOGENE ROCKS FROM
THE CALIFORNIA COAST RANGES

By

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MEASURED SECTIONS OF PALEOGENE ROCKS FROM THE CALIFORNIA COAST RANGES

By E. E. Brabb, J. C. Clark and C. K. Throckmorton

INTRODUCTION

The short-term purpose of this report is to provide the geologic and geographic framework for a series of field trips and a symposium on the California Paleogene to be held in Menlo Park, California, October 23-29, 1977, under the auspices of the International Commission on Paleogene Stratigraphy (ICPS). The long-term purpose is to facilitate the correlation of California Paleogene sequences with those in the classic and type areas of western Europe.

The report was conceived in 1968 during a series of field trips held in France, Italy, Belgium, and England as a part of an Eocene Colloquium organized by Professor Ch. Pomerol of the University of Paris, currently Chairman of the ICPS. The great success of these field trips in facilitating exchanges of ideas about the concepts of stratotypes and in correlating sequences from one region to another (see summary by Brabb, 1969) led to subsequent field trips and meetings on the Oligocene in Germany (1969) and the Caribbean region (1973) and the proposal to meet in California during 1977 and in the Soviet Union in 1978.

Ideally, the Paleogene sections selected for study should record continuous deposition from Late Cretaceous to early Miocene time; they should be reasonably accessible and well exposed; they should contain an association of different fossil groups that are useful for intercontinental correlation, such as planktonic foraminifers and nannoplankton; they should contain some of the fossil groups that have proven useful for local correlation, such as mollusks, echinoids, and benthonic foraminifers; and they should be in an area where the structure is reasonably simple.

Accordingly, the senior author in 1974 began a search for sections in California (fig. 1) that would fulfill most, if not all, of the above-mentioned requirements. This search is still continuing, because no perfect sections have been found, but the sections described in this report provide collectively a first approximation of an ideal California Paleogene sequence. Hopefully, additional work in the years ahead will fill gaps in our knowledge about the evolution of the different fossil groups and about the correlation of provincial chronologies with those in other areas of the world.

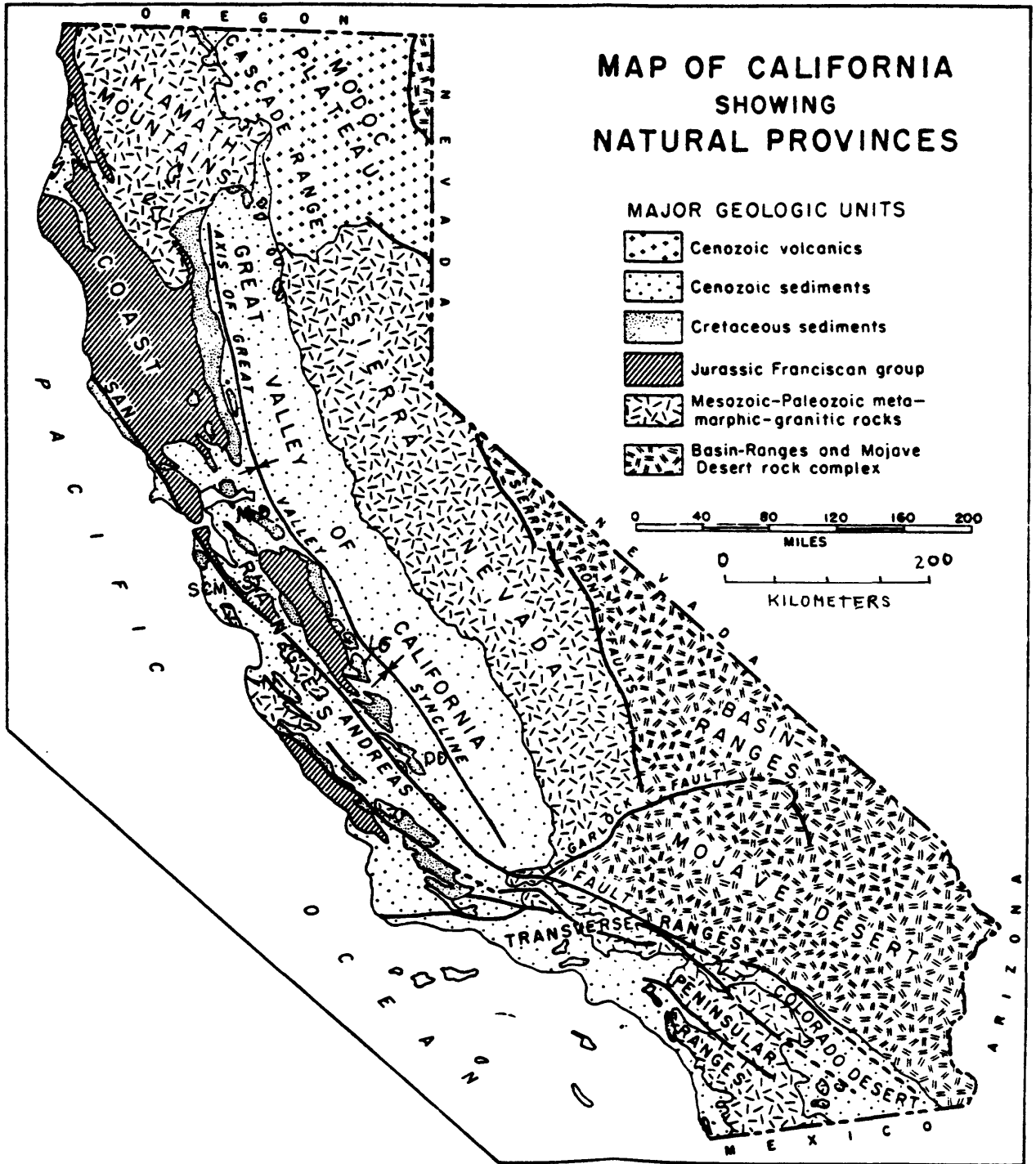


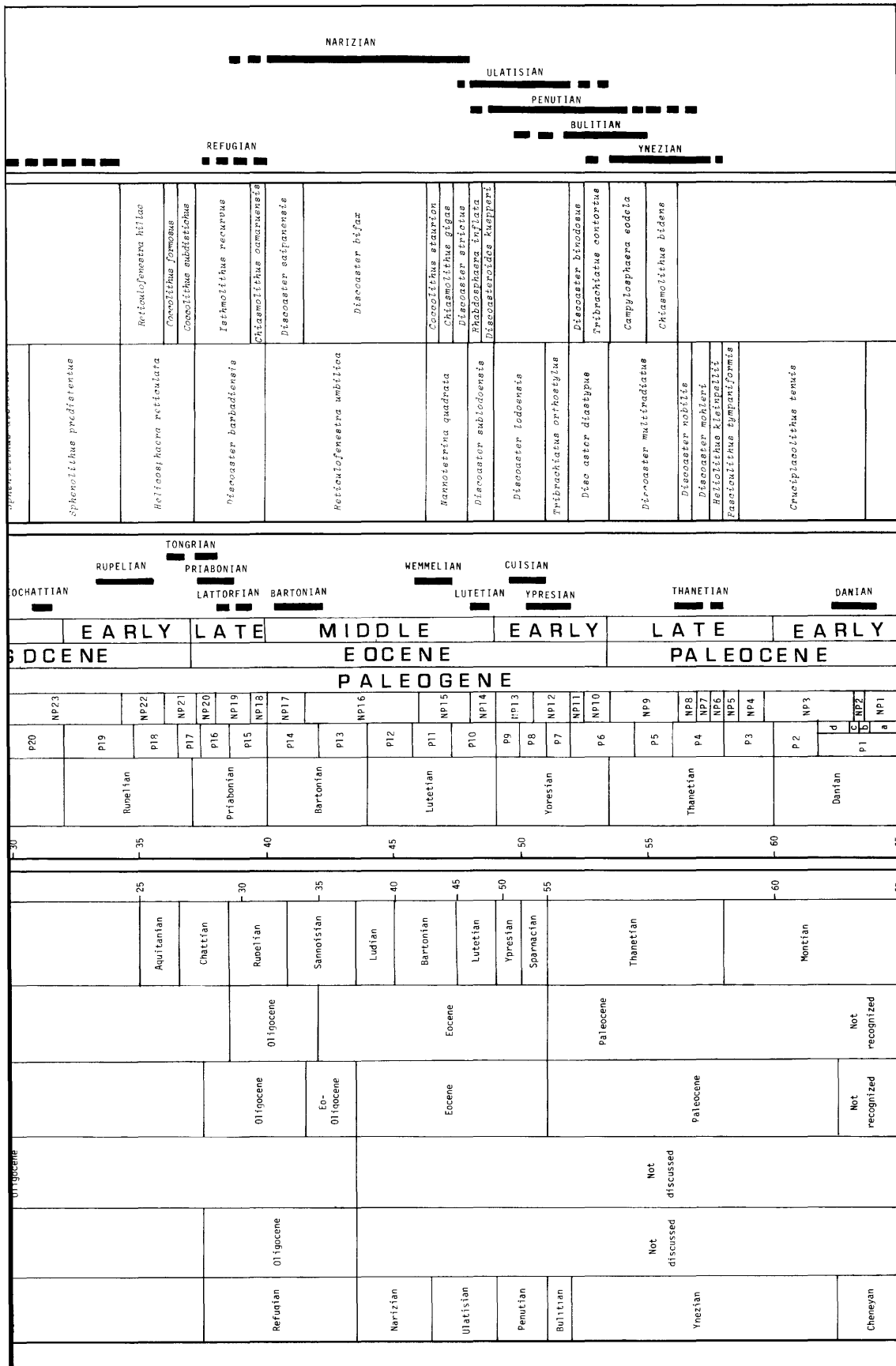
Figure 1.--Natural provinces of California and major geologic units (from Hinds, 1952). MD = Mount Diablo; SJ = San Jose; SCM = Santa Cruz Mountains; LG = Lodo Gulch; DD = Devils Den

Previous Work

The correlation of California rocks of Paleogene age with those in Europe has a long and complex history that can only be highlighted here. Kleinpell (1938, p. 168-181), in his classic work defining the Miocene benthonic foraminiferal stages of California, attempted to correlate the California faunas with those of western European sections and elsewhere. He pointed out that rocks considered traditionally as Lower Miocene in California are probably correlative with those considered Oligocene in Europe (fig. 2). The correlation of the Refugian Stage with the Oligocene is still practiced by some oil company micropaleontologists; see, for example, the report by the International Oil Scouts Association (1973, p. 32). Schenck and Childs (1942) correlated the Saucian Stage and the Vaqueros Formation with the Middle and Upper Oligocene of Europe and elsewhere, based on the stratigraphic occurrence of Lepidocyclina. Their correlation was discussed extensively by a number of paleontologists in the same report, some favoring a Miocene age based on similarities of the Vaqueros molluscan fauna with the Burdigalian faunas of Europe, and on the apparent "Miocene" age of vertebrates from beds below the Vaqueros, whereas others argued that some of the Vaqueros mollusks were similar to those of "Oligocene" age in Trinidad, and that the percent of living mollusks in the Vaqueros (1-2%) is most similar to faunas of "Oligocene" age. These differences in opinion proved to be irreconcilable in the attempt by the U.S. Committee on Stratigraphy to provide a standard correlation chart for marine Cenozoic formations of western North America (Weaver, and others, 1944). Two standards had to be provided, one based largely on species of mollusks, echinoids and corals, and the other on benthonic foraminifers. The present correlation of these separate standards with the European series is shown on figure 2. These correlations were modified somewhat by Turner (1970) and were used by Berggren (1972) and Hardenbol and Berggren (in press) to correlate with planktonic foraminifer and nannoplankton zones shown on figure 2.

The validity of some California Tertiary stages based largely on benthonic foraminifers has been questioned in recent years. Pierce (1972) and Barron (1976) believe that the Delmontian Stage of Kleinpell (1938) is coeval with the lower and middle part of the Mohnian Stage. Steineck and Gibson (1971), Gibson and Steineck (1972), Schmidt (1975), Bandy (1972) and Bukry and others (1977) believe that probably all of the California Eocene and Paleocene stages of Mallory (1959) are time-transgressive when compared to nannoplankton and planktonic foraminifer zonations. Hornaday and Philips (1972), on the other hand, challenge most of these opinions.

California benthonic foraminifer stages *	Kleinpell (1938, fig. 14)		West Coast Cenozoic standard of Weaver and others (1944)		Evernden and others (1964) and Turner (1970)		Time scale of Berggren (1972) and Hardenbol and Berggren (in press).				Range in age of California benthonic foraminifer stages compared to planktonic zones. Data from Bukry, Brabb and Vedder (1977), Poore (1976), Schmidt (1975), Barron (1976), Bands (1972), Lips and Kalisky (1972), Brabb, Bukry and Pierce (1971), and Poore and Brabb (1977).		
	Traditional California usage	Probable European correlation	Correlation based on benthonic foraminifers	Correlation based on mollusks, echinoids, and corals	European stages	Time in M.Y. (not to scale)	Age	Planktonic foraminifer zones	Nannoplankton zones	Epoch	Position of stage in type area	ZONE	SUBZONE
Hallian					Villafranchian	0	Calabrian	N23	N21	PLEIST.		<i>Emiliania huxleyi</i>	
Wheelerian							N22	N20				<i>Gephyrocapsa oerantica</i>	
Venturian	Not discussed	Not discussed	Pliocene	Pliocene	Astian-Plaisancian	-5	Astian	N21	N19	PLIOCENE		<i>Chenailithus doronicoides</i>	Too many subzones (13) to show at scale of figure.
Repettoian	Pliocene						N20	N18				<i>Dicocostes brovanti</i>	
Delmontian	Upper Miocene	Miocene	Miocene	Miocene	Pontian	-10	Messinian and Tortonian	M17	NN11			<i>Reticulolymnaea pseudomonticola</i>	<i>Ammonilithus tricomiculatus</i>
Mohnian	Middle Miocene	Miocene	Miocene	Miocene	Sarmatian	-13.7	Serrevallian	M15	NN9			<i>Dicocostes quinquetumus</i>	<i>Ammonilithus eximus</i>
Luisian	Middle Miocene	Miocene	Miocene	Miocene	Vindobonian	-15.3	Langhian	M14	NN8			<i>Dicocostes neocretus</i>	<i>Dicocostes bellus</i>
Relizian	Lower Miocene	Miocene	Miocene	Miocene	Burdigalian	-20	Burdigalian	M13	NN7			<i>Dicocostes hamatus</i>	<i>Coccolithus calidus</i>
Saucesian	Lower Miocene	Miocene	Miocene	Miocene	Burdigalian	-22.5	Langhian	M12	NN6			<i>Dicocostes estis</i>	<i>Dicocostes rugleri</i>
												<i>Dicocostes bellus</i>	
												<i>Coccolithus micropaginus</i>	
												<i>Sphenolithus heteromorphus</i>	
												<i>Helicobaculites ampliaperta</i>	
												<i>Sphenolithus belemnos</i>	
												<i>Dicocostes druggii</i>	
												<i>Dicocostes deflandrei</i>	
												<i>Cyclonolites abietatus</i>	
												<i>Dicocostes bisectus</i>	
												<i>Cyclonolites floridanus</i>	



* As defined by Scherck and Kiepel (1936); Kiepel (1938); Goukooff (1945); Natland (1953); and Maitory (1959).
 Correlated to planktonic foraminifer and nannoplankton zones by Berggren (1972).

Figure 2.--Correlation chart showing relation between California benthonic foraminifer stages and planktonic zones, and different correlations with European series

Present Investigation

The time is opportune, therefore, for a systematic new look at the validity and correlation of California Tertiary biostratigraphic units. This report and the paleontologic reports that will follow concentrate on the Paleogene. Ideally, the type sections for each of the Paleogene benthonic foraminiferal stages should have been included, but the task proved to be immense and several unanticipated problems were encountered. For instance, there does not seem to be a surface section for the Narizian Stage of Mallory (1959) in its type locality. Instead, several sections representative of deep marine deposition with benthonic foraminifers, planktonic foraminifers, and nannoplankton were measured, and their faunas and floras distributed to paleontologists for study. Their reports will constitute the conference part of the ICPS field meeting in October 1977.

Acknowledgements

Richard Z. Poore has been especially helpful in assembling and reviewing the correlation chart (fig. 2) in determining the abundance of the various microfossil assemblages in several sections, in directing the processing of many of these assemblages, and in offering advice about the most appropriate faunas to collect on the October field trips. Kristin McDougall kindly provided data on the age of some of the benthonic foraminifers. The authors also wish to thank William S. Brabb, James R. LeCompte, Steven W. Moore, Robert L. McCollom, Jr., and Bette R. Hamachi for helping to measure the sections. We are also indebted to Robert M. Kleinpell and A. D. Warren for examining critical Saucesian and Relizian faunas from near ~~At~~ Nuevo Point. A. D. Warren also provided information on the structure and fossils in the Devils Den area.

REGIONAL SETTING

The Paleogene sections described in this report are in the California Coast Ranges (fig. 1). Most are from areas west of the San Andreas fault in the Santa Cruz Mountains of west-central California--a few are from areas east of the San Andreas fault from Mount Diablo to Devils Den. The Paleogene rocks west of the San Andreas fault rest either directly on crystalline metasedimentary and granitic rocks of the Salinian block or on sedimentary strata of Late Cretaceous age. East of the San Andreas fault, the Paleogene rocks commonly rest on Jurassic and Cretaceous strata of the Great Valley sequence, which in turn are faulted against a complex tectonic assemblage of

Mesozoic sedimentary metamorphic, igneous, and volcanic rocks that is commonly referred to by petroleum geologists as "Franciscan basement." The relation of "Franciscan basement" rocks to the granitic basement rocks along which they are juxtaposed by the San Andreas fault has long been one of the most intriguing and complex problems of California geology. A series of lectures edited by Nilsen (1977) gives a synopsis of the problems and some working hypotheses for the tectonic evolution of California during Mesozoic and Cenozoic time. The most significant concept for the purpose of this report is that the San Andreas and other strike-slip faults have displaced the Paleogene sequences hundreds of kilometers with respect to each other. The Santa Cruz Mountains, for example, are thought by Nilsen and Clarke (1975) to have been opposite the Devils Den area, 250 km to the south, during the early Tertiary.

The paleogeography of central California during the Paleogene is not well established. Most published analyses are not based on a firm biostratigraphic framework, so that rocks of middle Eocene age in one area may be compared with rocks of late Eocene age in another. There is also a problem with displacement of the Paleogene sequences along secondary lateral faults. Although extensive lateral displacement along the San Andreas has been considered since the classic report by Hill and Dibblee (1953), many new lateral faults have been discovered that make paleogeographic reconstructions more complex. Nevertheless, much progress has been made, especially by a few geologists who have attempted to apply modern sedimentary concepts to understanding Paleogene depositional patterns. Nilsen (1977) and Dibblee (1977) provided extensive bibliographies listing many of these studies, and they also provided the synthesis that is summarized below.

A generalized paleogeographic map of early Tertiary California based on about 300 km of late Cenozoic right-lateral slip is shown on figure 3. The area between Devils Den and Mount Diablo was characterized by uplands in the Sierra Nevada, a linear basin in the area of the Great Valley in which a westward-thickening sequence of marine strata was deposited, islands on the continental borderland that contributed sediments in local areas, and a trench and subduction zone in what is now the Pacific Ocean. The La Honda basin, where all the field trips will take place, was on the continental borderland, mostly at lower bathyal and abyssal depths with open connections to the ocean.

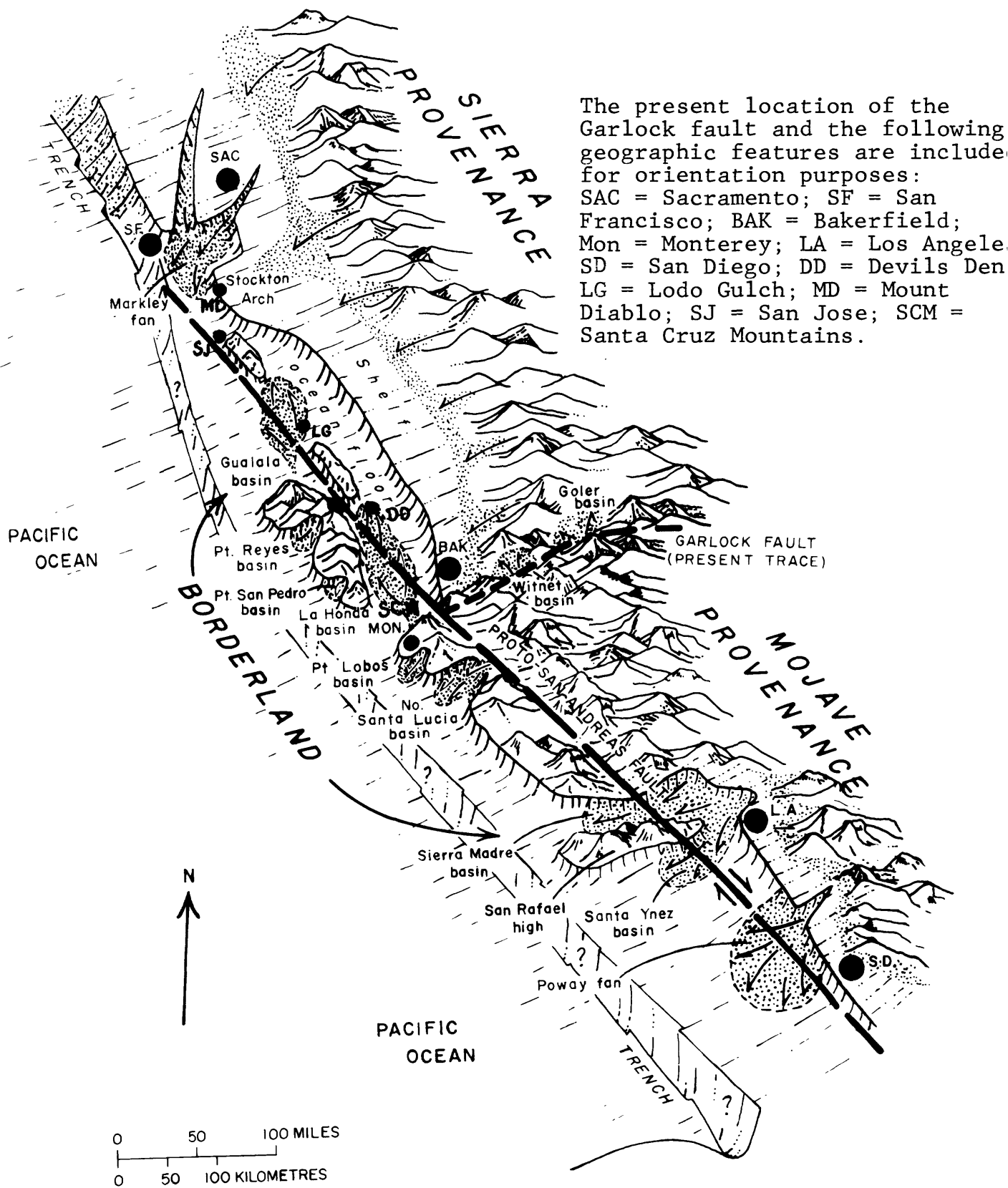


Figure 3.--Generalized paleogeographic map of early Tertiary California, based on reversal of 190 miles (305 km) of late Cenozoic right-lateral slip along the modern San Andreas fault (from Nilsen and Clarke, 1975).

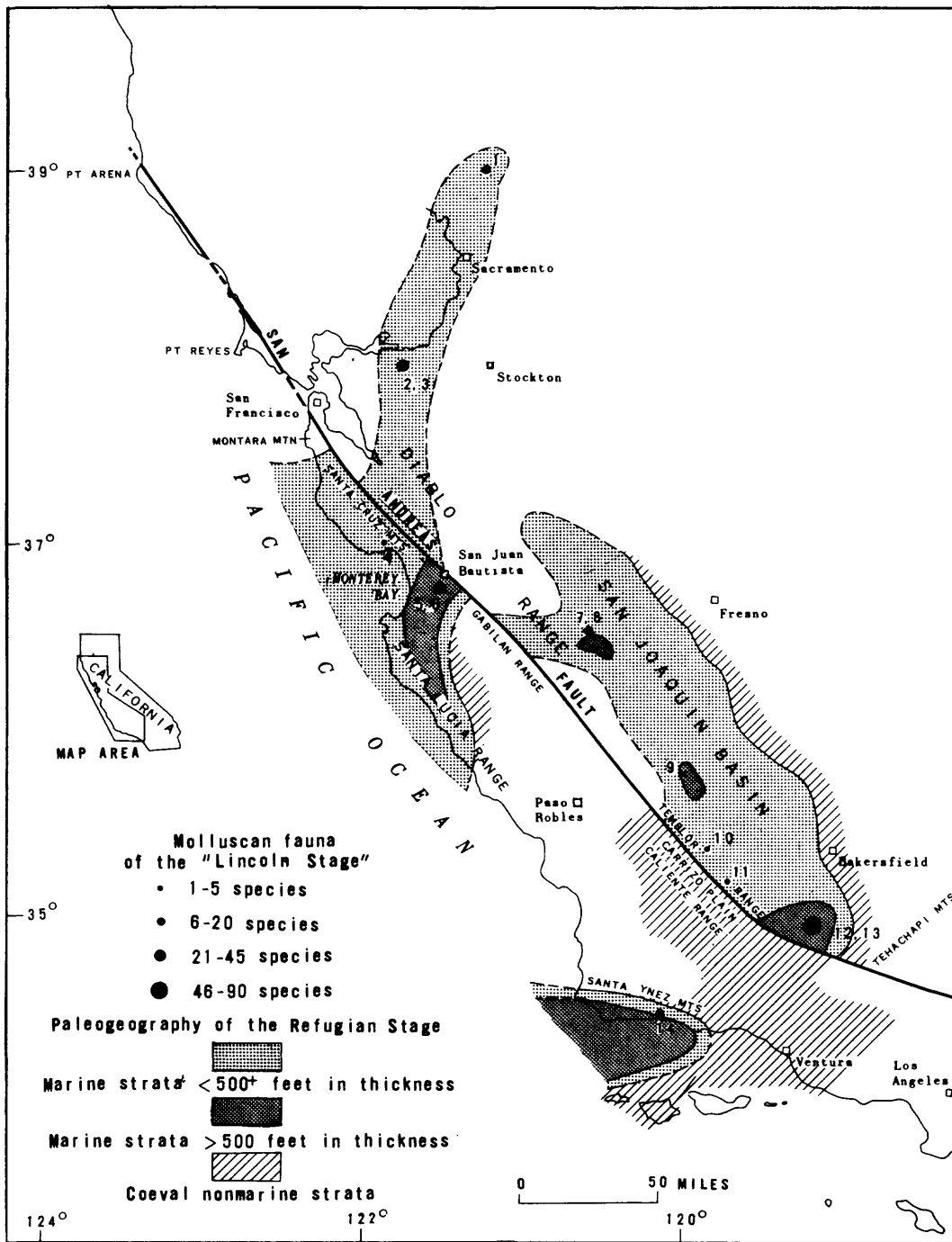


Figure 4.--Index map of California showing generalized Refugian Stage paleogeography and principal occurrences of molluscan assemblages. From Addicott (1968, fig. 2).

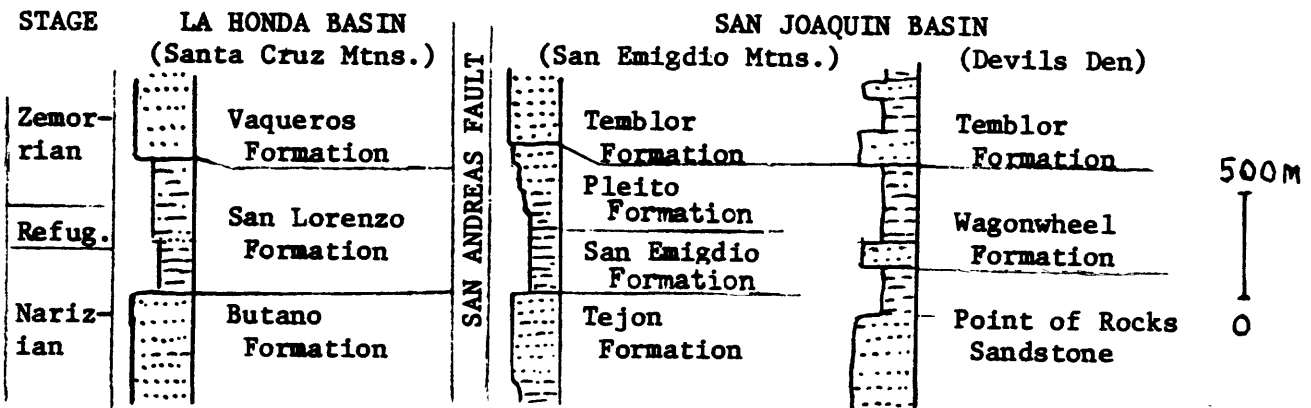
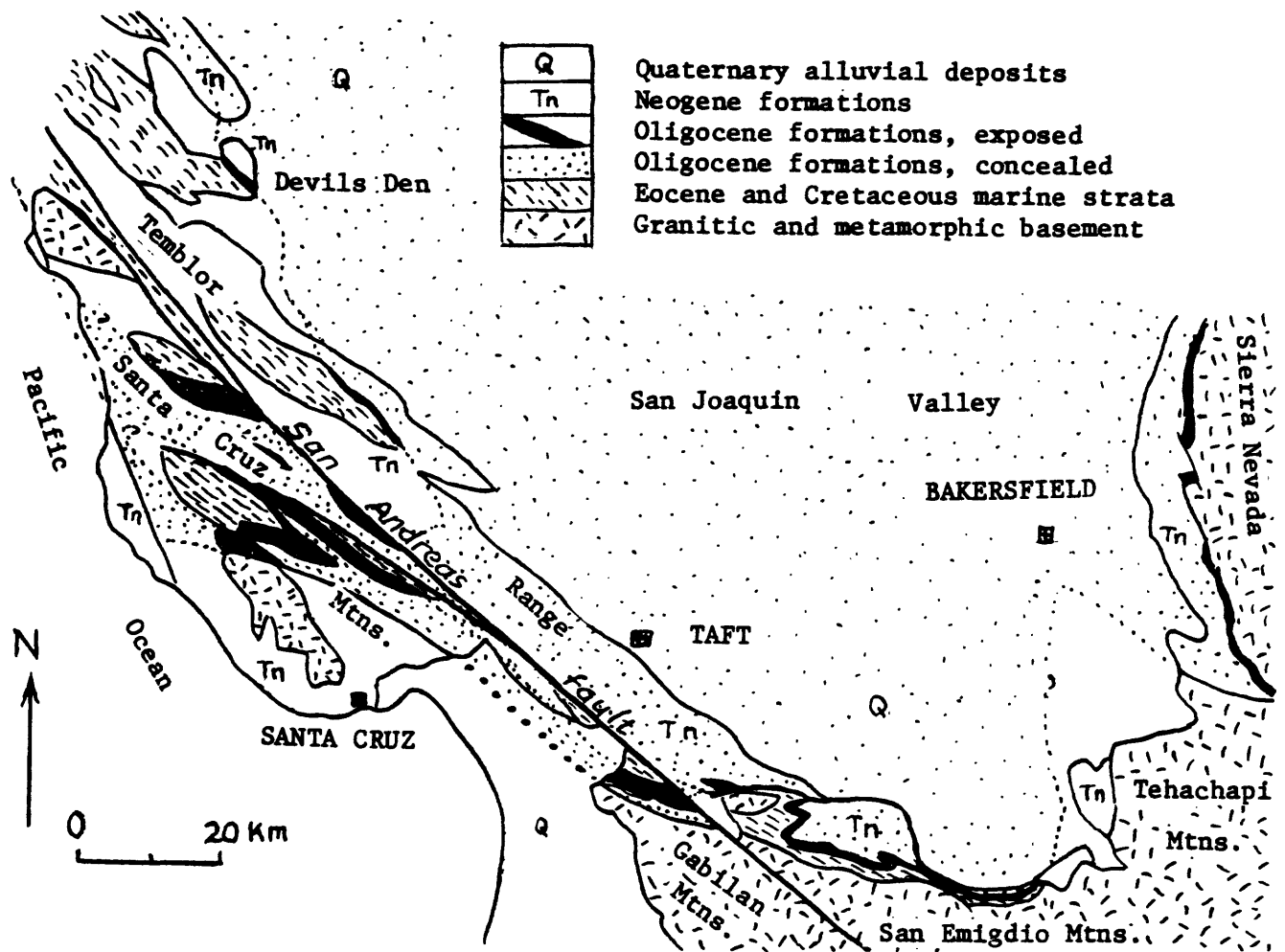


Figure 5.--Palinspastic map showing Oligocene (Refugian-early Zemorrian Stages) and adjacent formations of the La Honda sedimentary basin of Santa Cruz Mountains and of the San Joaquin basin, in inferred original relative positions, prior to about 280 km of post-Oligocene right-slip displacement on the San Andreas fault. Also shown are the nomenclature and stratigraphic positions of the Oligocene units of each area. From Dibblee (1977, fig. 8-6).

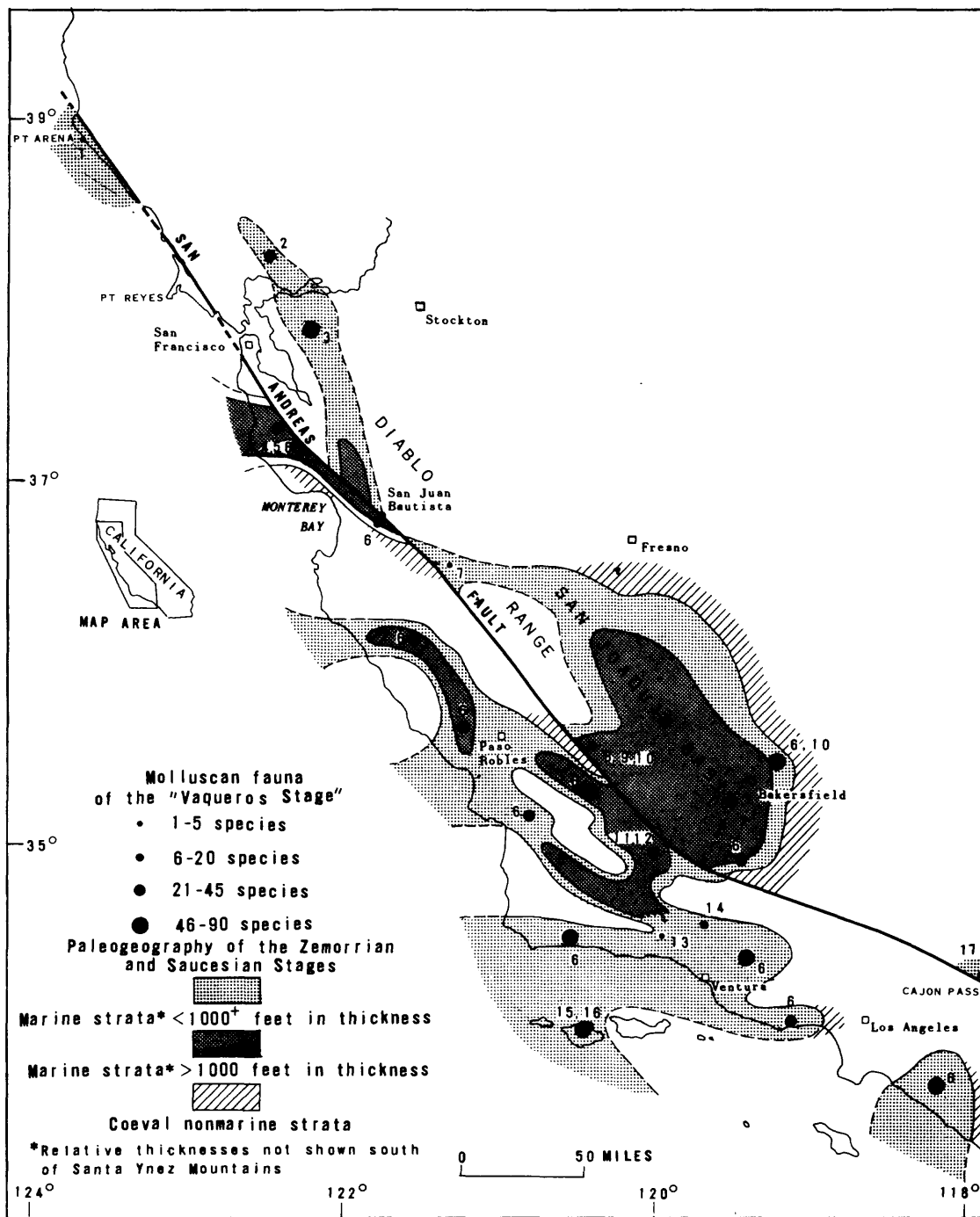


Figure 6.--Index map of California showing generalized Zemorrian and Saucesian Stages paleogeography and principal occurrences of molluscan assemblages. From Addicott (1968, fig. 3).

The Oligocene Epoch was characterized by major changes in California tectonics and paleogeography. Significant dip-slip displacement along several major faults produced widespread emergence of the Salinian block and the deposition in down-dropped blocks of hundreds of meters of nonmarine clastic sediments. Volcanism was extensive. Marine conditions continued in the La Honda basin with the deposition of mud that now forms the San Lorenzo Formation and Lambert Shale, which was interrupted by the local extrusion of as much as 700m of volcanic rocks (Mindego Basalt) and the northerly influx of arkosic sand and conglomerate that now form the Vaqueros Sandstone and Zayante Sandstone. Some paleogeographic reconstructions during the Oligocene are shown on figures 4, 5, and 6.

DESCRIPTION OF THE SECTIONS

Most of the sections described in this report are shown on the geologic map, figure 7. A composite geologic column for the sections in the Santa Cruz Mountains between the San Andreas and San Gregorio faults is shown on figure 8. Additional geologic and geographic maps and other information about the sections are provided in the individual description of each section. The field trip route is not shown on any map because it will likely change according to weather, tide height, and other access problems.

Exposures along the creek and river sections vary considerably from year to year depending primarily on rainfall. Most sections were best exposed after the 1955 flood waters had moved enormous quantities of gravel, sand and silt into the sea. Since then, construction and logging activities upstream and less runoff from the winter rain storms have resulted in the deposition of considerable sediment and the covering of bedrock exposures.

The availability of fresh rock exposures in the river and creek bed beneath the water may determine the relative chances of obtaining calcareous foraminifers. The tests of the calcareous foraminifers seem to be weakened or entirely dissolved by circulating ground water along the river and creek banks, whereas rocks collected beneath the summer water level have yielded well-preserved assemblages.

SYSTEM	SERIES	SEDIMENTARY SEQUENCE	FORAMINIFERAL STAGE	FORMATION	LITHOLOGY	THICKNESS Meters (feet)	DESCRIPTION	
TERTIARY	PLIOCENE	Pliocene to Pliocene		Purisima Formation		1725 (5650)	Very thick bedded yellowish-gray tuffaceous and diatomaceous siltstone, thick bedded and thickly crossbedded bluish-gray semifriable andesitic sandstone; and thin bedded medium gray siliceous mudstone	
				Upper Miocene	Santa Cruz Mudstone of Clark (1966)		0-2700 (0-8850)	Medium- to thick-bedded and faintly laminated, pale-yellowish-brown siliceous mudstone with scattered spheroidal dolomite concretions, locally grading to sandy siltstone
					Santa Margarita Sandstone		0-130 (0-430)	Very thick bedded and thickly crossbedded, yellowish-gray to white friable arkosic sandstone
	MIOCENE	Middle Miocene	Luisian	Monterey Formation		810 (2700)	Medium- to thick-bedded and laminated olive-gray subsiliceous organic mudstone and sandy siltstone with few thick dolomite interbeds	
				Lompico Sandstone of Clark (1966)		0-240 (0-800)	Thick-bedded to massive yellowish-gray arkosic sandstone	
				Lambert Shale		460 (1500)	Thin- to medium-bedded and faintly laminated olive-gray to dusky-yellowish-brown organic mudstone with phosphatic laminae and lenses in lower part	
	OLIGOCENE	Lower Miocene	Zemorian	Mindego Basalt		0-1200 (0-4000)	Predominantly olivine-bearing basaltic pillow lava and flow breccia	
				Vaqueros Sandstone		0-910 (0-3000)	Thick-bedded to massive yellowish-gray arkosic sandstone with thin interbeds of medium-gray siltstone	
				Zayante Sandstone		0-550 (0-1800)	Thick to very thick bedded, yellowish-orange arkosic sandstone with thin interbeds of green and red siltstone and lenses and thick interbeds of pebble and cobble conglomerate	
	EOCENE	Lower Eocene	Nanzian	San Lorenzo Formation		400-810 (1300-2700)	Upper part is nodular light-gray mudstone, locally grading to fine-grained arkosic sandstone, lower part is very thin bedded olive-gray clay shale	
				Butano Sandstone		2400+ (8000+)	Medium-bedded to massive yellowish-gray arkosic sandstone with thin interbeds of olive-gray siltstone and thick interbeds of sandy pebble conglomerate in lower part	
	PALEOGENE	Paleocene	Ynezian	Locatelli Formation of Cummings and others (1962)		75-270 (250-900)	Nodular olive-gray to pale-yellowish-brown micaceous siltstone; massive arkosic sandstone locally at base	
				Salinian basement rocks			Predominantly hornblende-biotite quartz diorite, ranging from gabbro to granite. Intrusive into schists, quartzites, marbles, and calc-silicate rocks	

Figure 8.--Composite stratigraphic column for the area between the San Andreas and San Gregorio faults

The distribution of the foraminifers is also uneven. They seem to be concentrated in laminae or thin beds, and sparse or absent in similar rocks nearby. They are generally less abundant in sandstones and siltstones compared to claystones. Several hours of careful searching with a hand lens may be required to find the same foraminifer-bearing bed shown on the various stratigraphic columns. To help geologists find the exact bed, metal tags were put on the rock at many localities. Unfortunately, the tags do not seem to last long in populated areas.

The sections are reasonably accessible, unless otherwise noted. Permission to cross private property can usually be obtained from residents in the area. There are generally no problems taking small samples for microfossil analyses except in State parks and nature preserves, such as in the Aho Nuevo area, where no samples can be collected without a permit.

GLAUCONITE

The upper and lower boundaries of the Refugian Stage are marked by glauconitic sandstones. In Little Boulder Creek, the basal 30 cm (1 foot) of the Refugian Stage (the basal 30 cm of the Rices Sandstone Member of the San Lorenzo Formation) is 90 percent glauconite. The glauconite was bright green in outcrop before it was covered by material caved in from the stream bank. The other 10 percent of the rock consists of well-rounded quartz and feldspar grains and abraded fish teeth.

About 2.3 km southeast of Little Boulder Creek, in an unnamed tributary of Pescadero Creek, the glauconite-rich basal bed of the Refugian Stage is about 15 cm (6 inches) thick. Some of the glauconite is in ovoid tubes, presumably worm burrows, that extend several centimeters into the underlying Twobar Shale Member.

At other localities the glauconite is not as abundant in the basal bed, but it generally comprises at least 5 percent of the basal sandstone. The glauconitic sandstone marks not only the base of the Refugian Stage but also the contact between the laminated shales characteristic of the Twobar Shale Member and the more massive and spheroided weathering mudstones and siltstones characteristic of Rices Mudstone Member. It also marks a change from open-sea faunas represented predominantly by the Buliminids, Globigerinids, robust arenaceous foraminifers, and a few thin-shelled mud pectens to assemblages in the Rices Mudstone Member dominated by Lagenids Anomalinids, and more abundant and more robust mollusks like Bruclarkia, Priscofusus, and Modiolus. As mentioned by Brabb (1964, p. 677), these changes suggest significant tectonic adjustments requiring considerable time. The glauconite, therefore, may indicate a stratigraphic break or disconformity.

The glauconitic sandstone at the base of the Zemorrian Stage contains phosphatic pellets that are 1-2 cm long at most localities. Some of the phosphatic pellets are coated with manganese and are black, whereas others are pale brown and appear to be entirely phosphate.

There are at least two glauconitic sandstone beds within the Vaqueros Sandstone. Burchfiel (1964) and McCollom (1959) used these beds to correlate sections of the Vaqueros from one valley to another. The significance of these glauconites with respect to foraminiferal and nannoplankton zonations has not yet been established.

No glauconite-rich beds have been found in the Butano Sandstone or the Locatelli Formation. Glauconite-rich sandstone and conglomerate beds are common in Neogene sequences in the Santa Cruz Mountains, but few of these have been described or dated by paleontologic methods. The Paleogene glauconites seem to have been too deeply buried for meaningful K-Ar dating, but one of the Neogene glauconites (the basal Purisima of late Miocene or early Pliocene age) has been dated by this method.

LITTLE BOULDER CREEK SECTION

The Little Boulder Creek section (fig. 9) on the north side of Butano Ridge was designated the type section for the upper 518 m (1700 feet) of the Butano Sandstone by Cummings and others (1962, p. 184). The section was also described briefly and designated as a reference section for the Twobar Shale Member of the San Lorenzo Formation by Brabb (1964).

After this section was measured in 1958, landslide debris and sediments deposited in the creek have reduced the amount of section exposed. Trees felled by snow have made access along the creek more difficult. The access road is not suitable for a large bus, and part of the road may be impassable when it rains. Permission to visit the section must be obtained from the Santa Cruz Lumber Company.

The stratigraphic column (fig. 10) provides information on the lithology and texture of the Butano Sandstone and San Lorenzo Formation. Table 1 provides information on the microfaunas of these formations. The planktonic foraminifers have not been examined by specialists, and no attempt has yet been made to find nannoplankton.

Nilsen and Simoni (1973) studied the sedimentary structures in the Butano Sandstone along Butano Ridge and elsewhere. They believe that the Butano Sandstone is a deep-sea fan deposit derived from a southern provenance area adjacent to Monterey Bay.

Table 1.--Check list of Foraminifera from the Little Boulder Creek section

Note: Foraminifera washed and identified by oil company micropaleontologists. Many identifications are provisional.

V = Very abundant (>500) A = Abundant (101-500) C = Common (10-100) R = Rare (1-9) X = Abundance not estimated ? = Questionable identification	Approximate stratigraphic distance (in feet) from glauconite at Twobar-Rices contact.	Formation																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																																														
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2238	EB 2239	EB 2240	EB 2241	EB 2242	EB 2243	EB 2244	EB 2245	EB 2246	EB 2247	EB 2248	EB 2249	EB 2250	EB 2251	EB 2252	EB 2253	EB 2254	EB 2255	EB 2256	EB 2257	EB 2258	EB 2259	EB 2260	EB 2261	EB 2262	EB 2263	EB 2264	EB 2265	EB 2266	EB 2267	EB 2268	EB 2269	EB 2270	EB 2271	EB 2272	EB 2273	EB 2274	EB 2275	EB 2276	EB 2277	EB 2278	EB 2279	EB 2280	EB 2281	EB 2282	EB 2283	EB 2284	EB 2285	EB 2286	EB 2287	EB 2288	EB 2289	EB 2290	EB 2291	EB 2292	EB 2293	EB 2294	EB 2295	EB 2296	EB 2297	EB 2298	EB 2299	EB 2300	EB 2301	EB 2302	EB 2303	EB 2304	EB 2305	EB 2306	EB 2307	EB 2308	EB 2309	EB 2310	EB 2311	EB 2312	EB 2313	EB 2314	EB 2315	EB 2316	EB 2317	EB 2318	EB 2319	EB 2320	EB 2321	EB 2322	EB 2323	EB 2324	EB 2325	EB 2326	EB 2327	EB 2328	EB 2329	EB 2330	EB 2331	EB 2332	EB 2333	EB 2334	EB 2335	EB 2336	EB 2337	EB 2338	EB 2339	EB 2340	EB 2341	EB 2342	EB 2343	EB 2344	EB 2345	EB 2346	EB 2347	EB 2348	EB 2349	EB 2350	EB 2351	EB 2352	EB 2353	EB 2354	EB 2355	EB 2356	EB 2357	EB 2358	EB 2359	EB 2360	EB 2361	EB 2362	EB 2363	EB 2364	EB 2365	EB 2366	EB 2367	EB 2368	EB 2369	EB 2370	EB 2371	EB 2372	EB 2373	EB 2374	EB 2375	EB 2376	EB 2377	EB 2378	EB 2379	EB 2380	EB 2381	EB 2382	EB 2383	EB 2384	EB 2385	EB 2386	EB 2387	EB 2388	EB 2389	EB 2390	EB 2391	EB 2392	EB 2393	EB 2394	EB 2395	EB 2396	EB 2397	EB 2398	EB 2399	EB 2400	EB 2401	EB 2402	EB 2403	EB 2404	EB 2405	EB 2406	EB 2407	EB 2408	EB 2409	EB 2410	EB 2411	EB 2412	EB 2413	EB 2414	EB 2415	EB 2416	EB 2417	EB 2418	EB 2419	EB 2420	EB 2421	EB 2422	EB 2423	EB 2424	EB 2425	EB 2426	EB 2427	EB 2428	EB 2429	EB 2430	EB 2431	EB 2432	EB 2433	EB 2434	EB 2435	EB 2436	EB 2437	EB 2438	EB 2439	EB 2440	EB 2441	EB 2442	EB 2443	EB 2444	EB 2445	EB 2446	EB 2447	EB 2448	EB 2449	EB 2450	EB 2451	EB 2452	EB 2453	EB 2454	EB 2455	EB 2456	EB 2457	EB 2458	EB 2459	EB 2460	EB 2461	EB 2462	EB 2463	EB 2464	EB 2465	EB 2466	EB 2467	EB 2468	EB 2469	EB 2470	EB 2471	EB 2472	EB 2473	EB 2474	EB 2475	EB 2476	EB 2477	EB 2478	EB 2479	EB 2480	EB 2481	EB 2482	EB 2483	EB 2484	EB 2485	EB 2486	EB 2487	EB 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2863	EB 2864	EB 2865	EB 2866	EB 2867	EB 2868	EB 2869	EB 2870	EB 2871	EB 2872	EB 2873	EB 2874	EB 2875	EB 2876	EB 2877	EB 2878	EB 2879	EB 2880	EB 2881	EB 2882	EB 2883	EB 2884	EB 2885	EB 2886	EB 2887	EB 2888	EB 2889	EB 2890	EB 2891	EB 2892	EB 2893	EB 2894	EB 2895	EB 2896	EB 2897	EB 2898	EB 2899	EB 2900	EB 2901	EB 2902	EB 2903	EB 2904	EB 2905	EB 2906	EB 2907	EB 2908	EB 2909	EB 2910	EB 2911	EB 2912	EB 2913	EB 2914	EB 2915	EB 2916	EB 2917	EB 2918	EB 2919	EB 2920	EB 2921	EB 2922	EB 2923	EB 2924	EB 2925	EB 2926	EB 2927	EB 2928	EB 2929	EB 2930	EB 2931	EB 2932	EB 2933	EB 2934	EB 2935	EB 2936	EB 2937	EB 2938	EB 2939	EB 2940	EB 2941	EB 2942	EB 2943	EB 2944	EB 2945	EB 2946	EB 2947	EB 2948	EB 2949	EB 2950	EB 2951	EB 2952	EB 2953	EB 2954	EB 2955	EB 2956	EB 2957	EB 2958	EB 2959	EB 2960	EB 2961	EB 2962	EB 2963	EB 2964	EB 2965	EB 2966	EB 2967	EB 2968	EB 2969	EB 2970	EB 2971	EB 2972	EB 2973	EB 2974	EB 2975	EB 2976	EB 2977	EB 2978	EB 2979

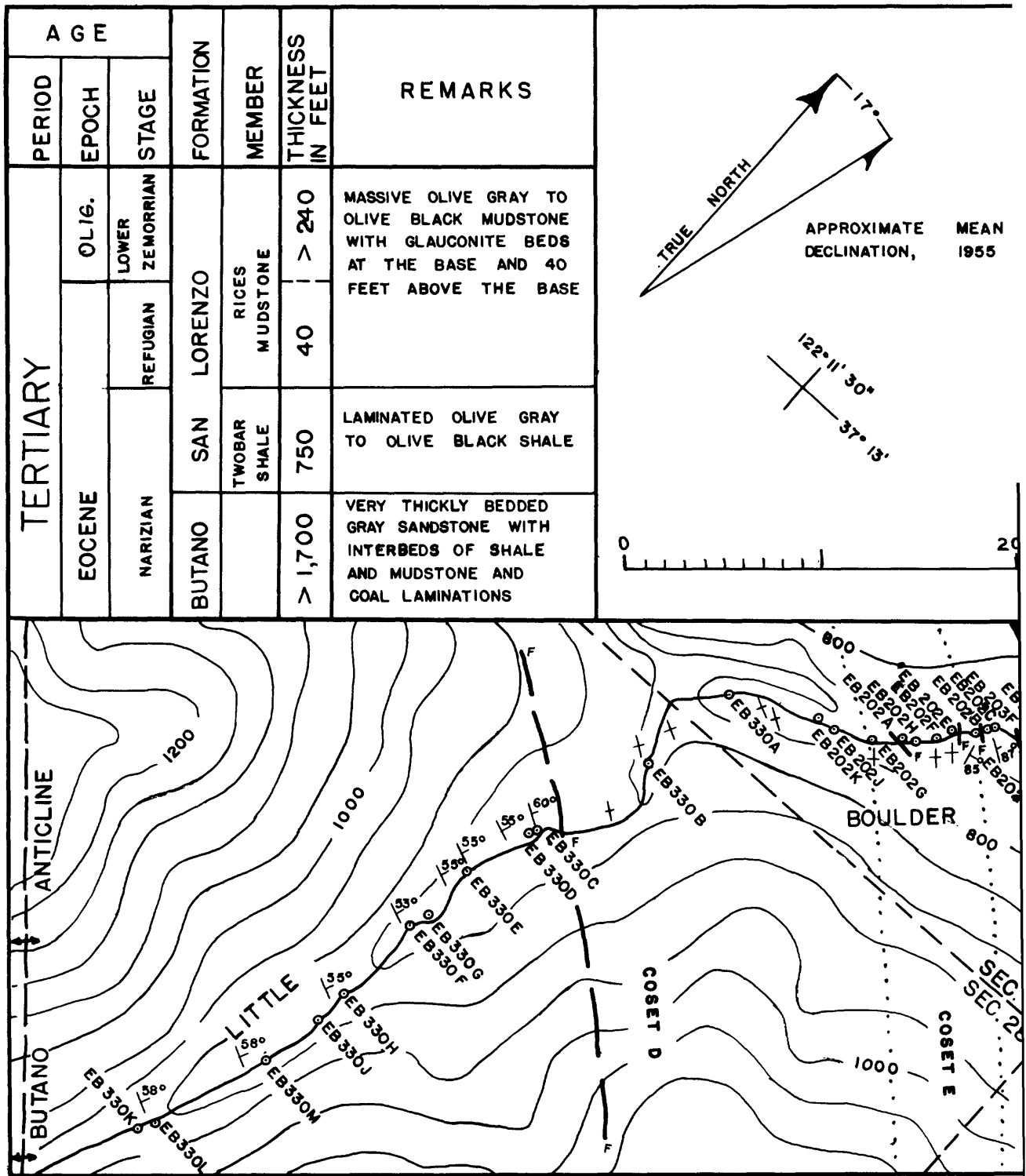


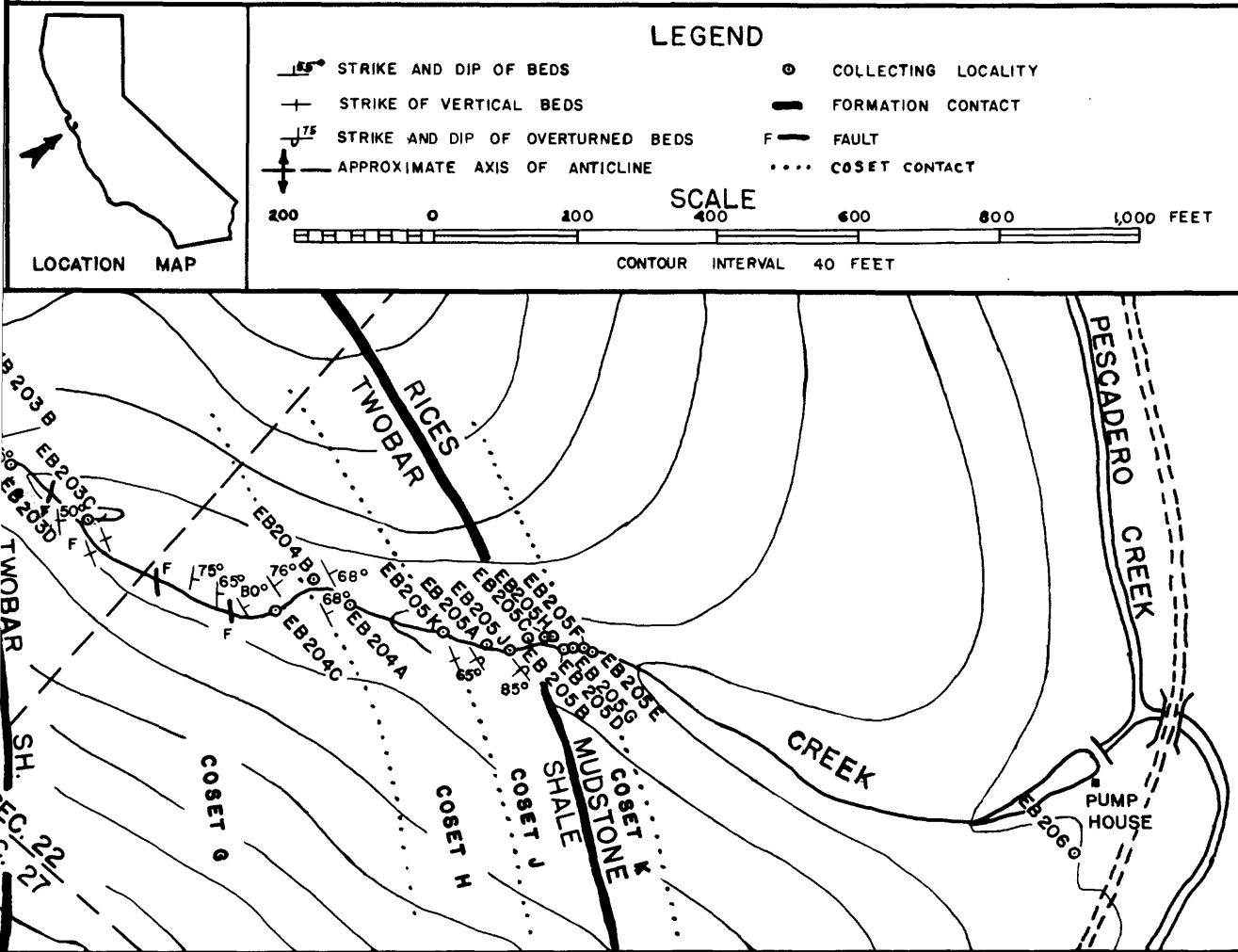
Figure 9.--Collecting localities along Little Boulder Creek

COLLECTING LOCALITIES ALONG LITTLE BOULDER CREEK IN
 A PART OF THE TYPE LOCALITY OF THE BUTANO SANDSTONE
 SECS. 21, 22, 27, AND 28, T. 8 S., R. 3 W. M. D. B. M.

FROM REPORT BY EARL E. BRABB, 1960

"GEOLOGY OF THE BIG BASIN AREA, SANTA CRUZ MOUNTAINS, CALIFORNIA"

BASE MAP ENLARGED FROM U. S. GEOLOGICAL SURVEY BIG BASIN 7.5 QUADRANGLE, 1955
 EDITION. COLLECTING LOCALITIES AND DETAILS OF CREEK SURVEYED BY TAPE AND
 BRUNTON COMPASS



STRATIGRAPHIC COLUMN

OF THE SAN LORENZO FORMATION AND TYPE BUTANO SANDSTONE
EXPOSED IN LITTLE BOULDER CREEK, CALIFORNIA

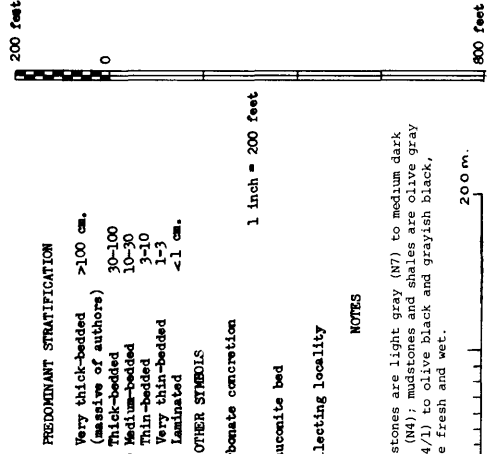
Secs. 21, 22, and 28, T. 8 S., R. 2 W., Mount Diablo Base and Meridian

From report by Earl E. Brabb, 1960

"Geology of the Big Basin area, Santa Cruz Mountains, California"

LEGEND

- PREDOMINANT PARTICLE SIZE**
- Fine- to medium-grained sandstone (Median diameter 0.50-0.12 mm.)
 - Very fine-grained sandstone (Median diameter 0.12-0.06 mm.)
 - Interbedded sandstone and mudstone or shale
 - Mudstone (Median diameter 0.06 mm.)
 - Shale (Median diameter 0.06 mm.)
 - Fault - part of section missing
 - Covered
- PREDOMINANT STRATIFICATION**
- V = Very thick-bedded (>100 cm. (massive of authors))
 - T = Thick-bedded (30-100)
 - M = Medium-bedded (10-30)
 - H = Thin-bedded (3-10)
 - E = Very thin-bedded (1-3)
 - L = Laminated (<1 cm.)
- OTHER SYMBOLS**
- Carbonate concretions
 - Glauconite bed
 - Collecting locality
- 1 inch = 200 feet



NOTES

(1) Sandstones are light gray (N7) to medium dark gray (N4); mudstones and shales are olive gray (5Y 4/1) to olive black and grayish black, where fresh and wet.

PERCENT OF SET CEMENTED BY CARBONATE	0	50	100
MAX. PERM., MDCYS.			
MAX. POROSITY, %			
TRASK INDEX (S ₀)			
MEDIAN DIAMETER OF SAMPLES, IN MM.			
ONE PERCENTILE (C) IN MM.			
STRATIFICATION			
THICKNESS OF COSET IN FEET			?
FORMATION THICKNESS			MORE 240
SET			
COSET			L
MEMBER	RICES MUDSTONE		
FORMATION	SAN LORENZO		
STAGE	ZEMORRIAN		
EPOCH	OLIGOCENE		

GRAPHIC COLUMN	COLLECTING LOCALITIES	REMARKS
PERCENT OF SANDSTONE IN SET		
		Top of section at pumphouse near intersection of Little Boulder Creek and Pascadero Creek
		EB 206 <i>Ungerina yellowyi</i> <i>Plectrogoniataria vaughani</i> <i>Acila</i> sp., and others

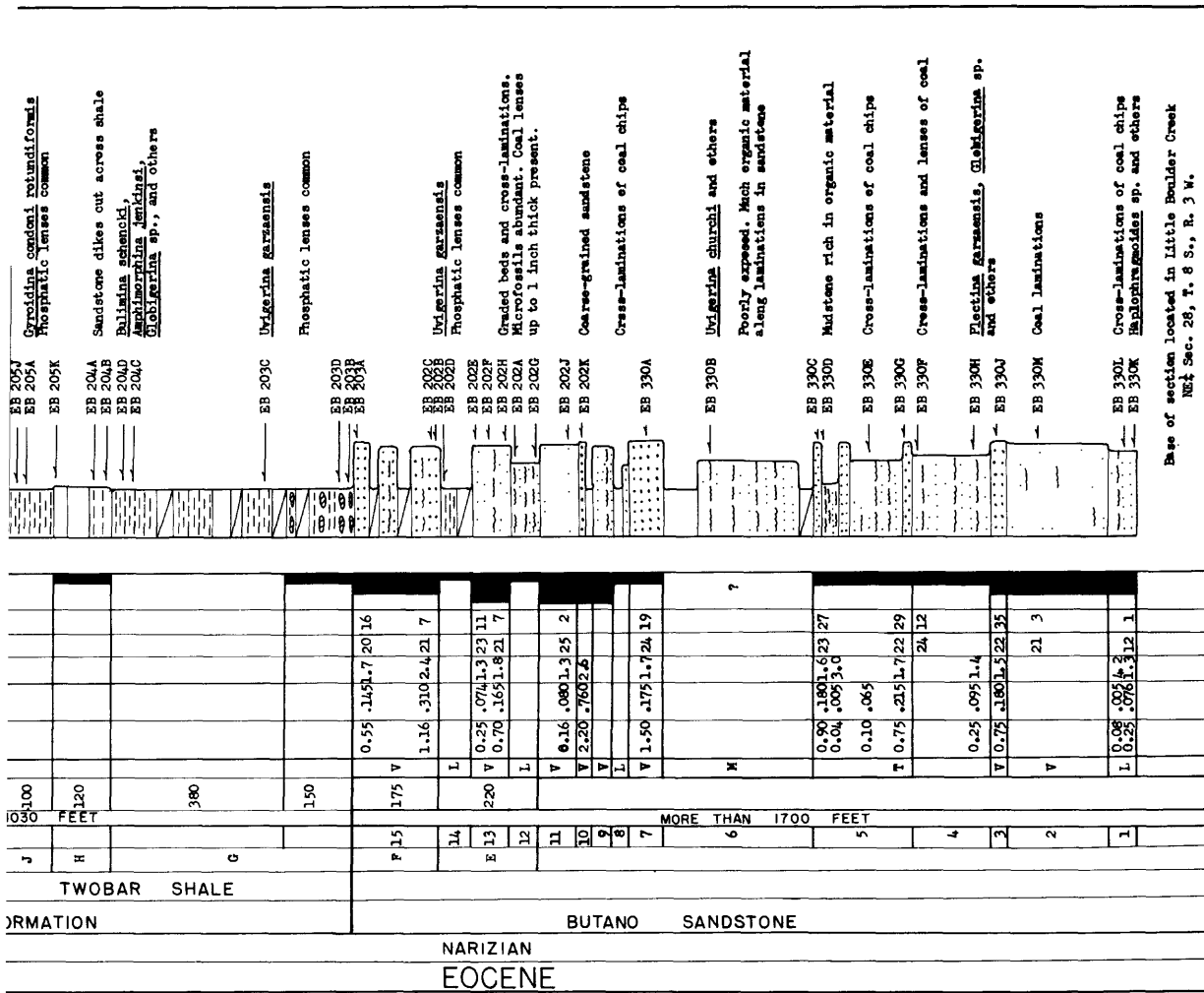


Figure 10.--Stratigraphic column, Little Boulder Creek

SAN LORENZO RIVER SECTION

The San Lorenzo River area (fig. 11) includes the lower part of the San Lorenzo Formation in its type section (Arnold, 1906). The upper part of the formation is folded and has few reliable attitudes, so that it is difficult to construct a complete stratigraphic column. The column by Sullivan (1962, fig. 4) includes about 170 m (550 feet) of repeated section, according to Brabb (1964, p. 675).

Benthonic foraminifers and a few planktonic foraminifers from the San Lorenzo River section were described by Sullivan (1962). Poore and Brabb (1977) described the planktonic foraminifers in greater detail.

A stratigraphic column (fig. 12) shows the position of faunas collected by various geologists. Sample B2251, a critical sample in Sullivan's faunal sequence (1962, p. 290), could not be located precisely enough to show on the map.

In the San Lorenzo River and other measured sections, the following symbols are used to indicate abundance of foraminifers and nannoplankton:

A	Abundant
C	Common
F	Few
S	Sparse
X	None

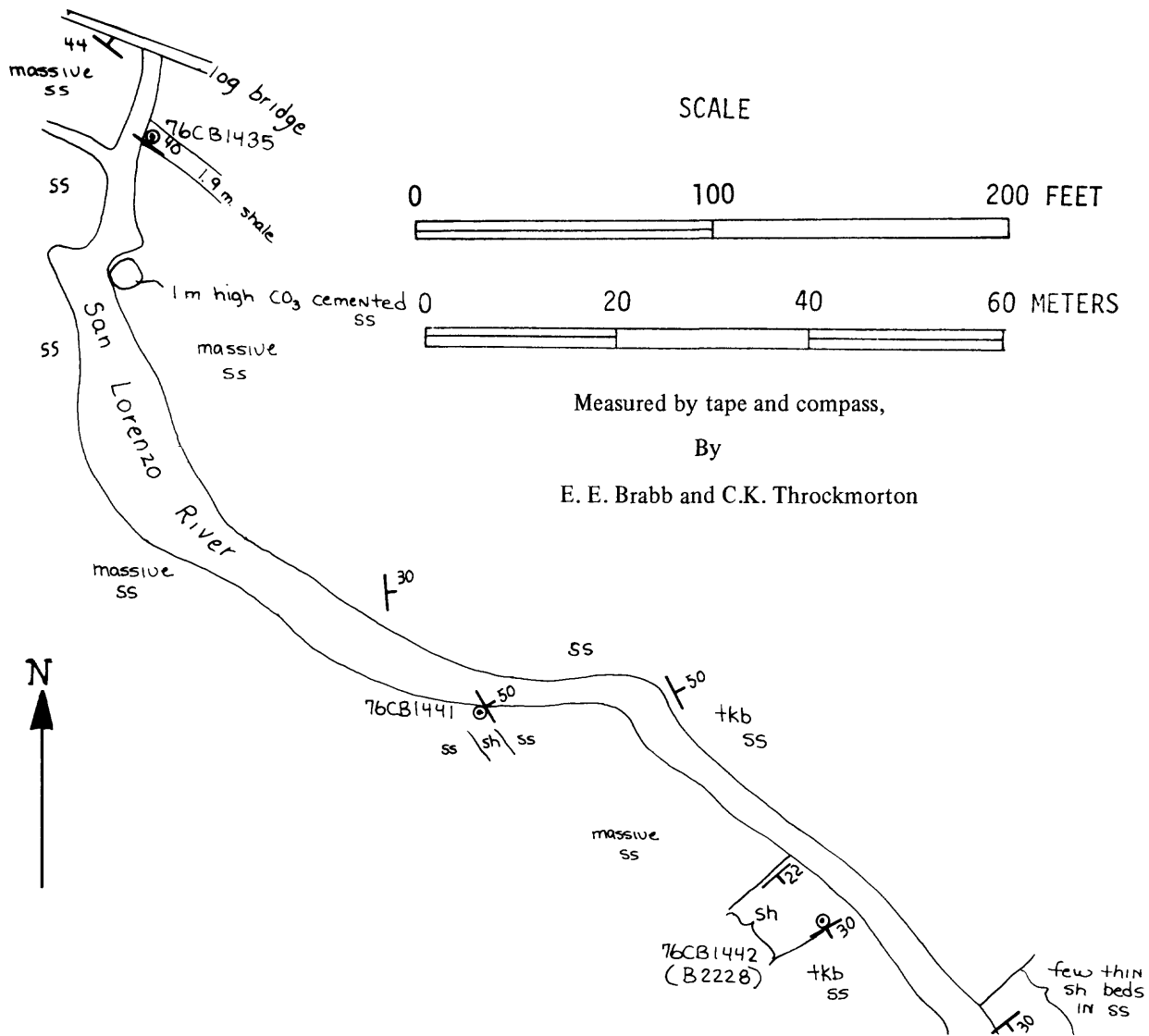


Figure 11.--Geologic sketch map of the San Lorenzo River near Riverside Grove showing formation contacts and localities where fossils and rocks were collected

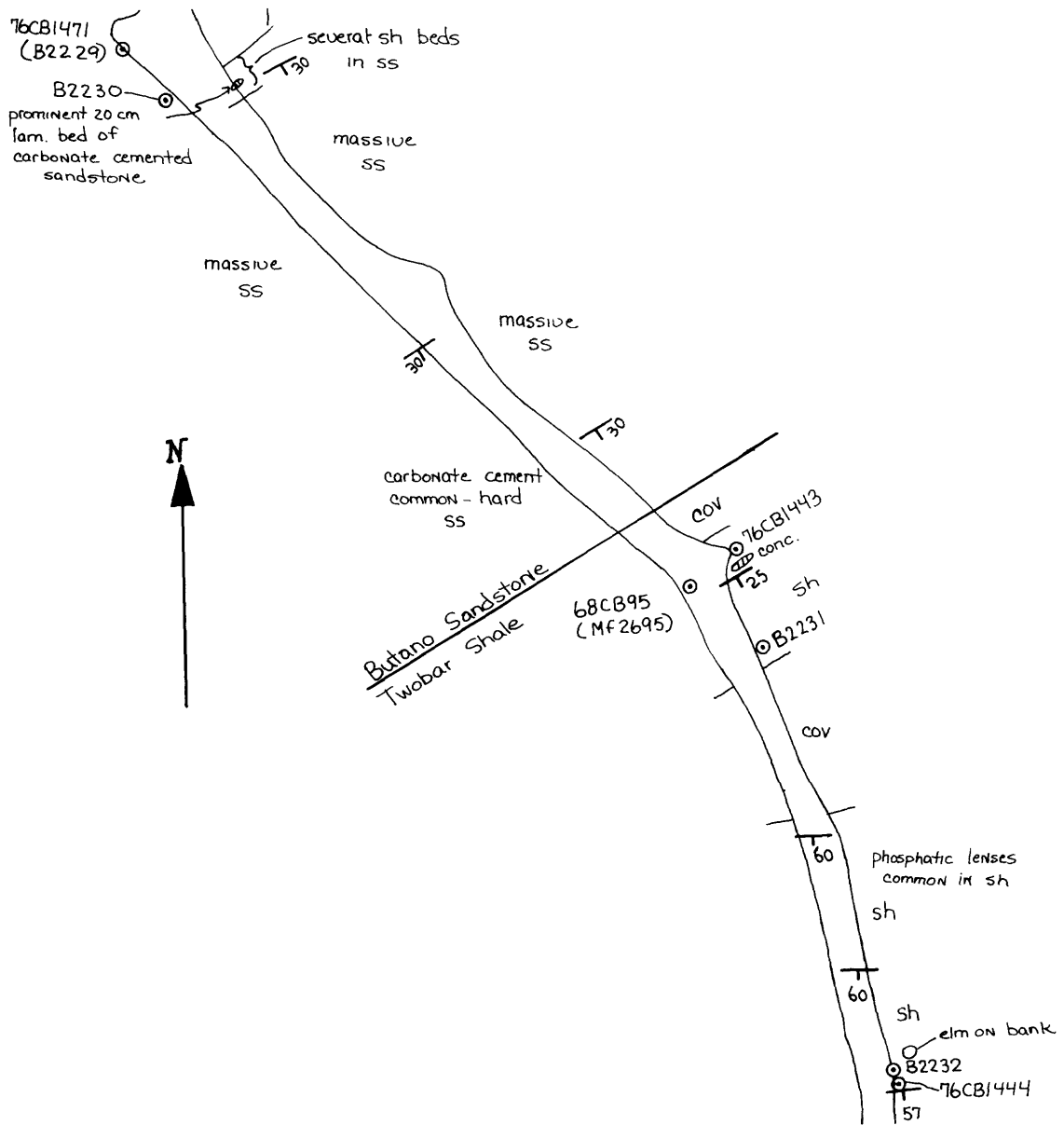
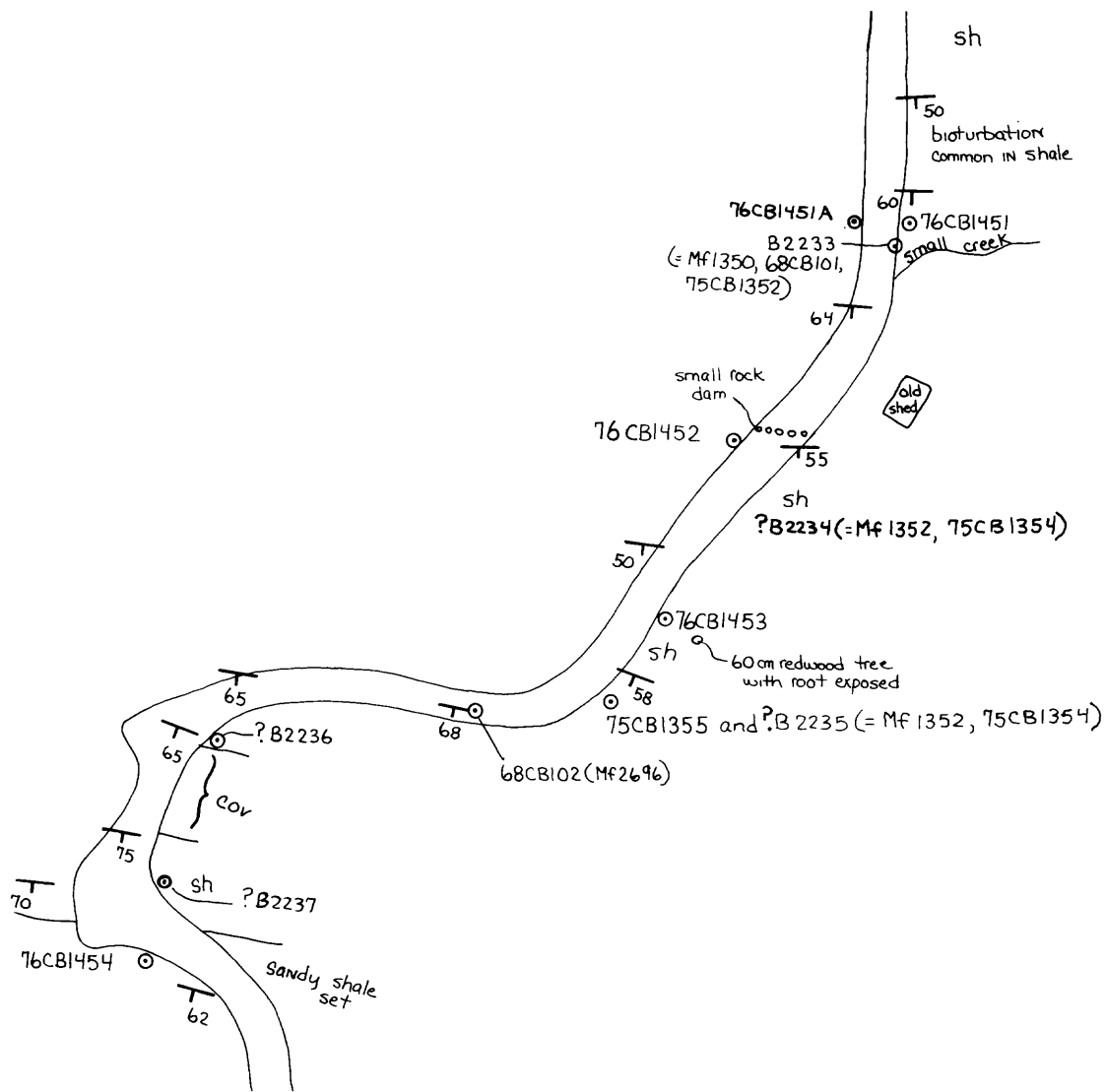


Figure 11.--



continued

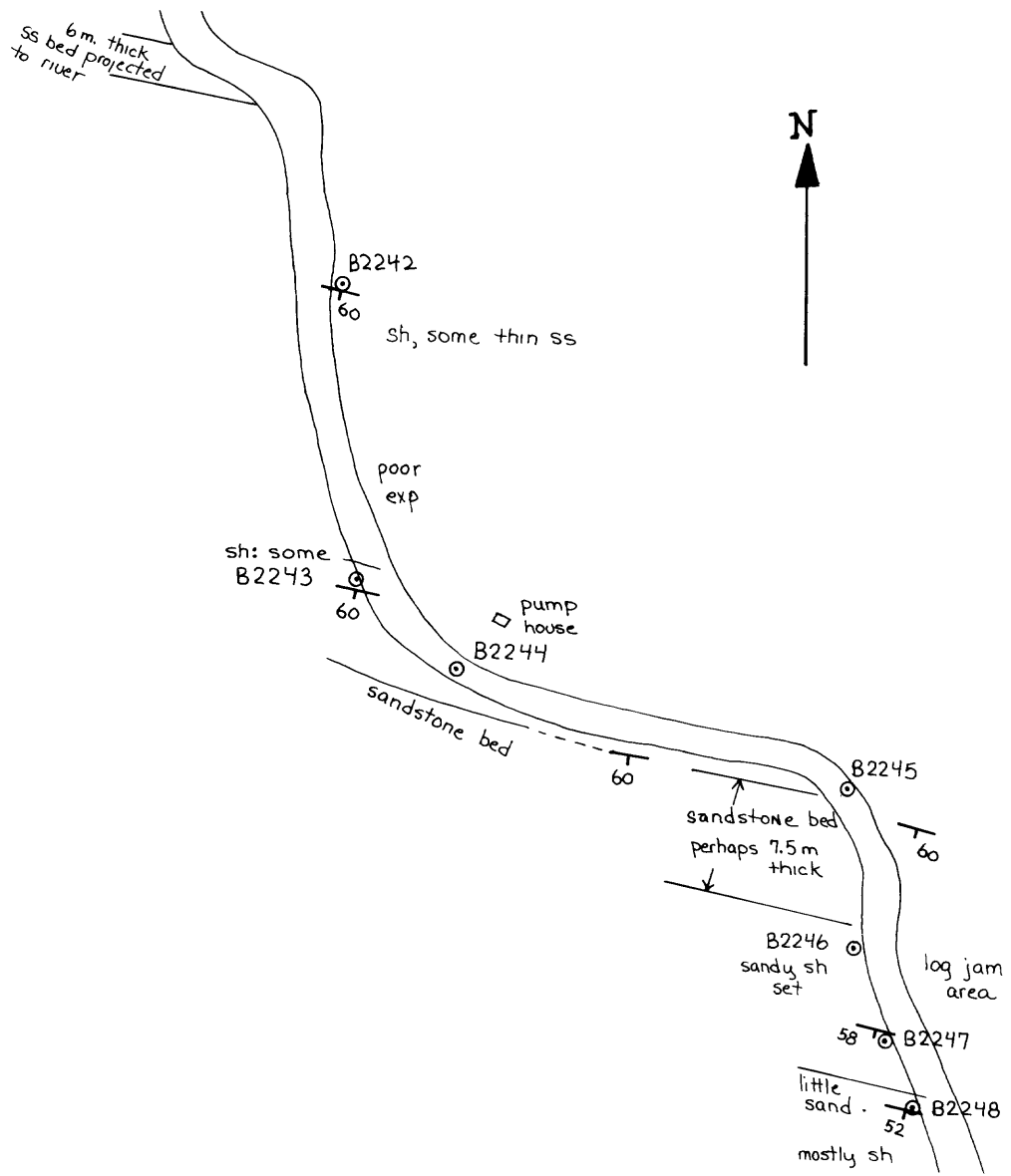
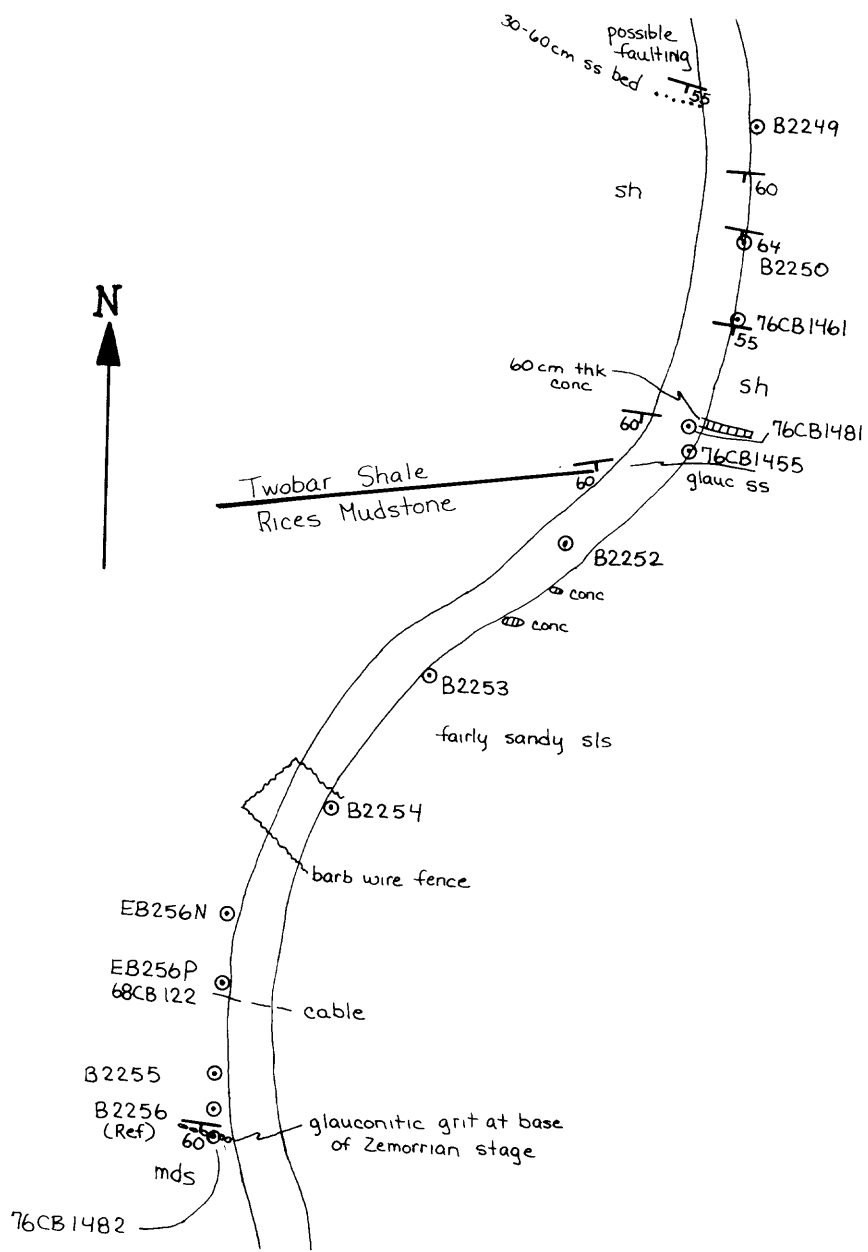


Figure 11.-



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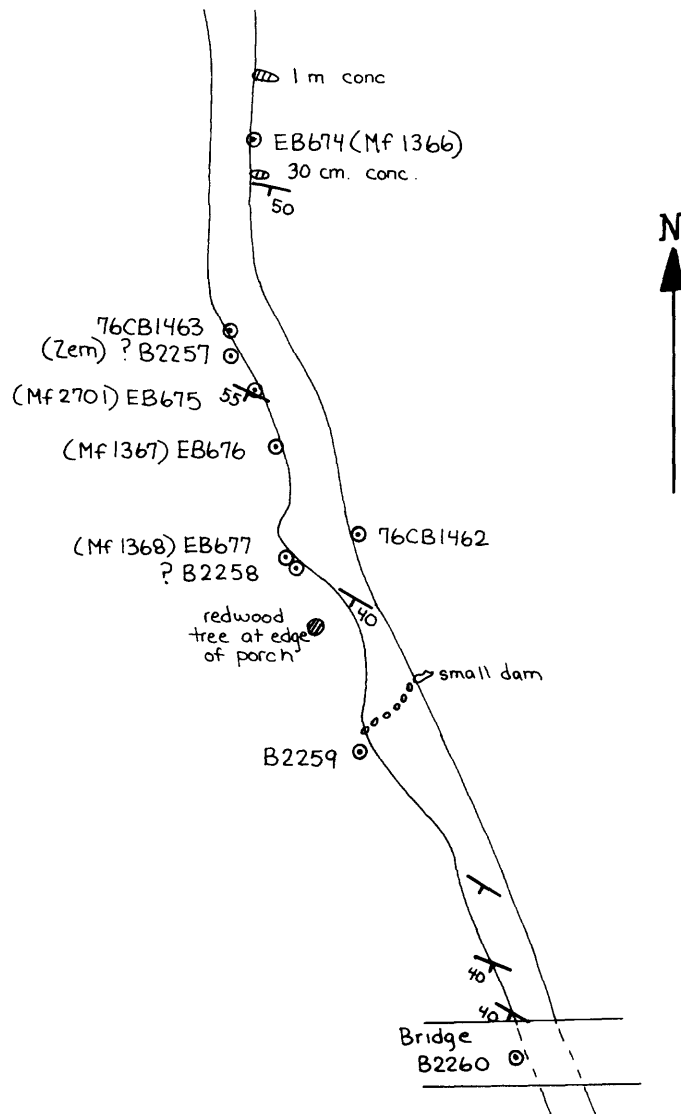


Figure 11.--continued

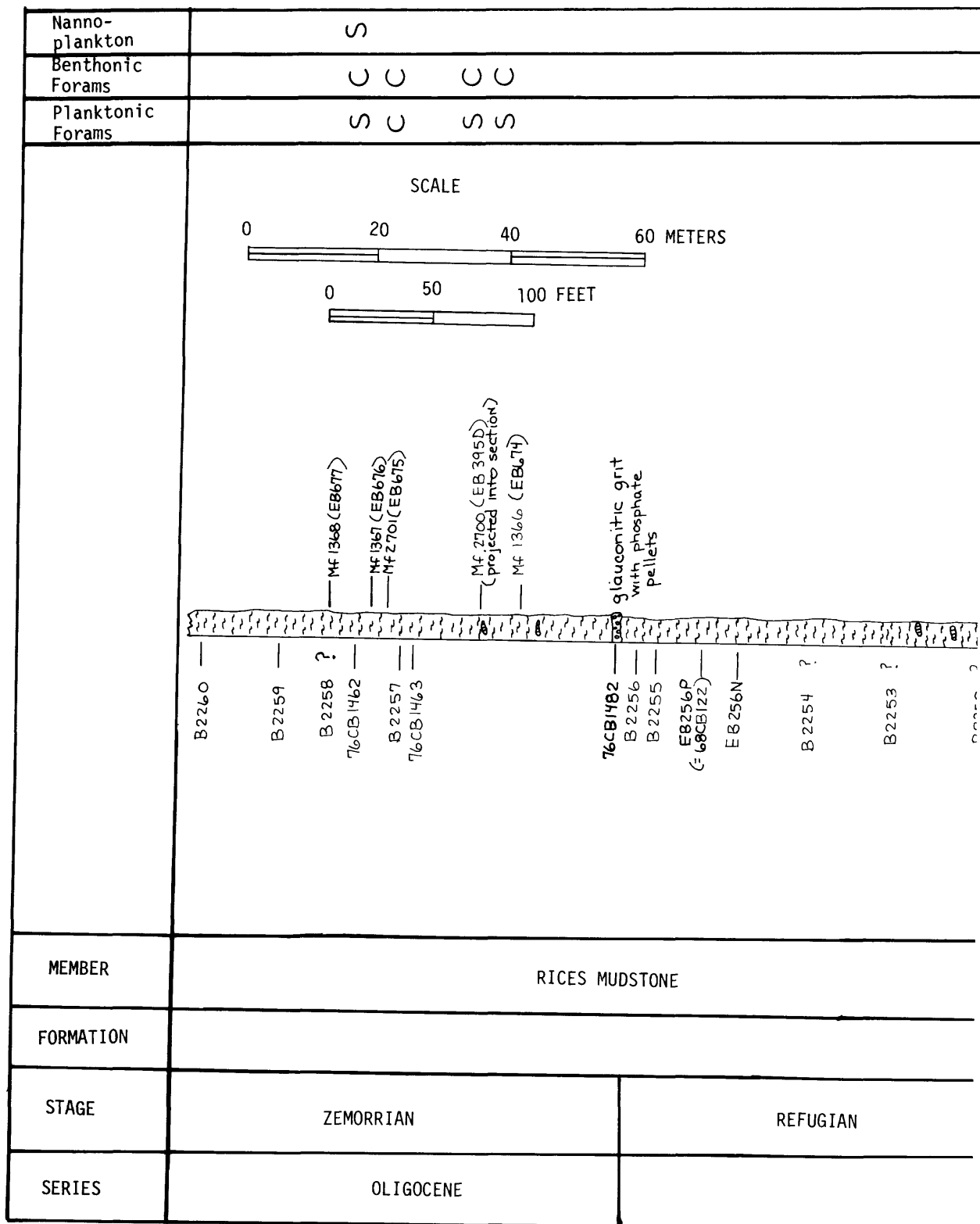


Figure 12.--Stratigraphic column for San Lorenzo River section near Riverside Grove

S
C
S

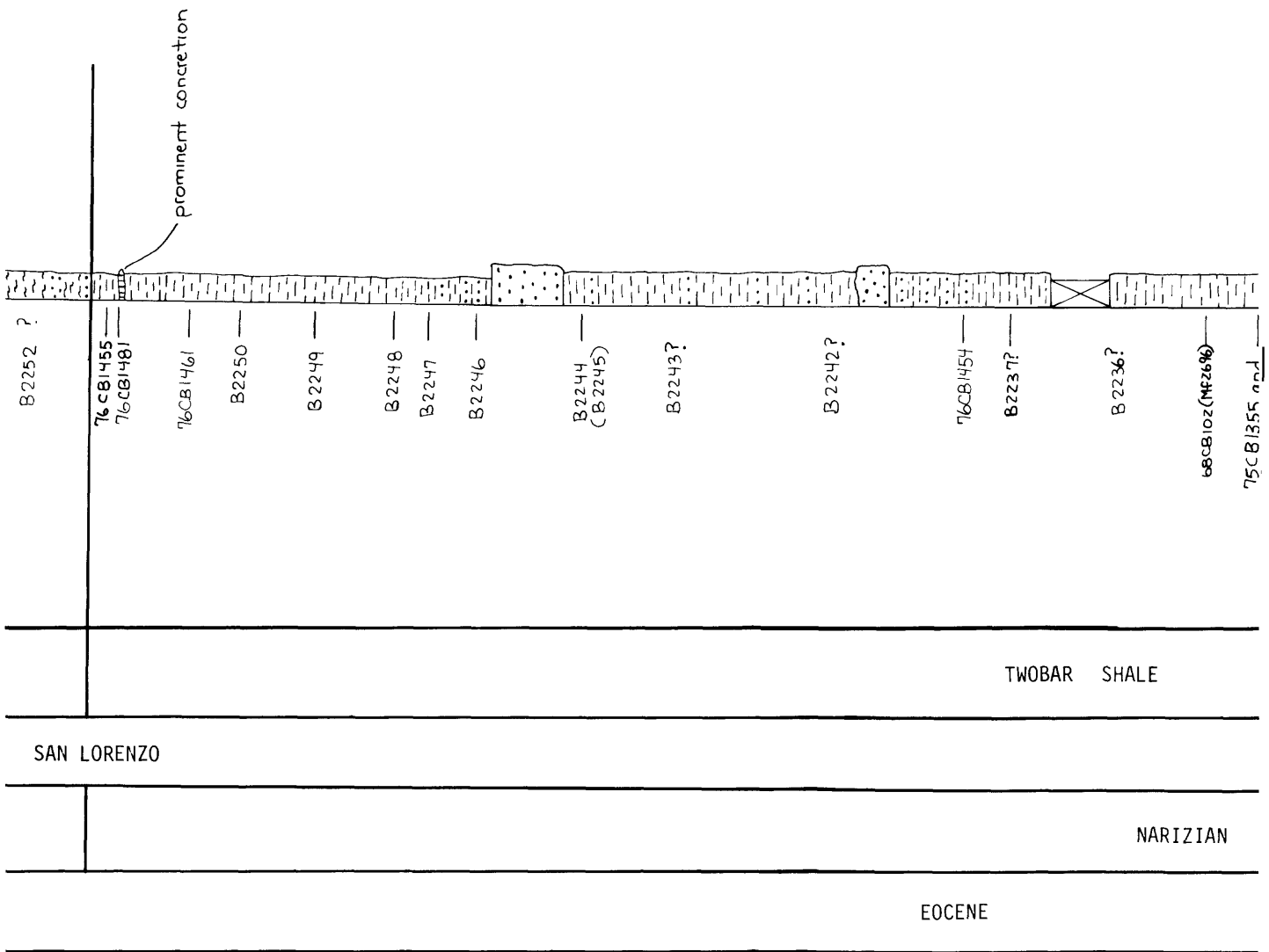
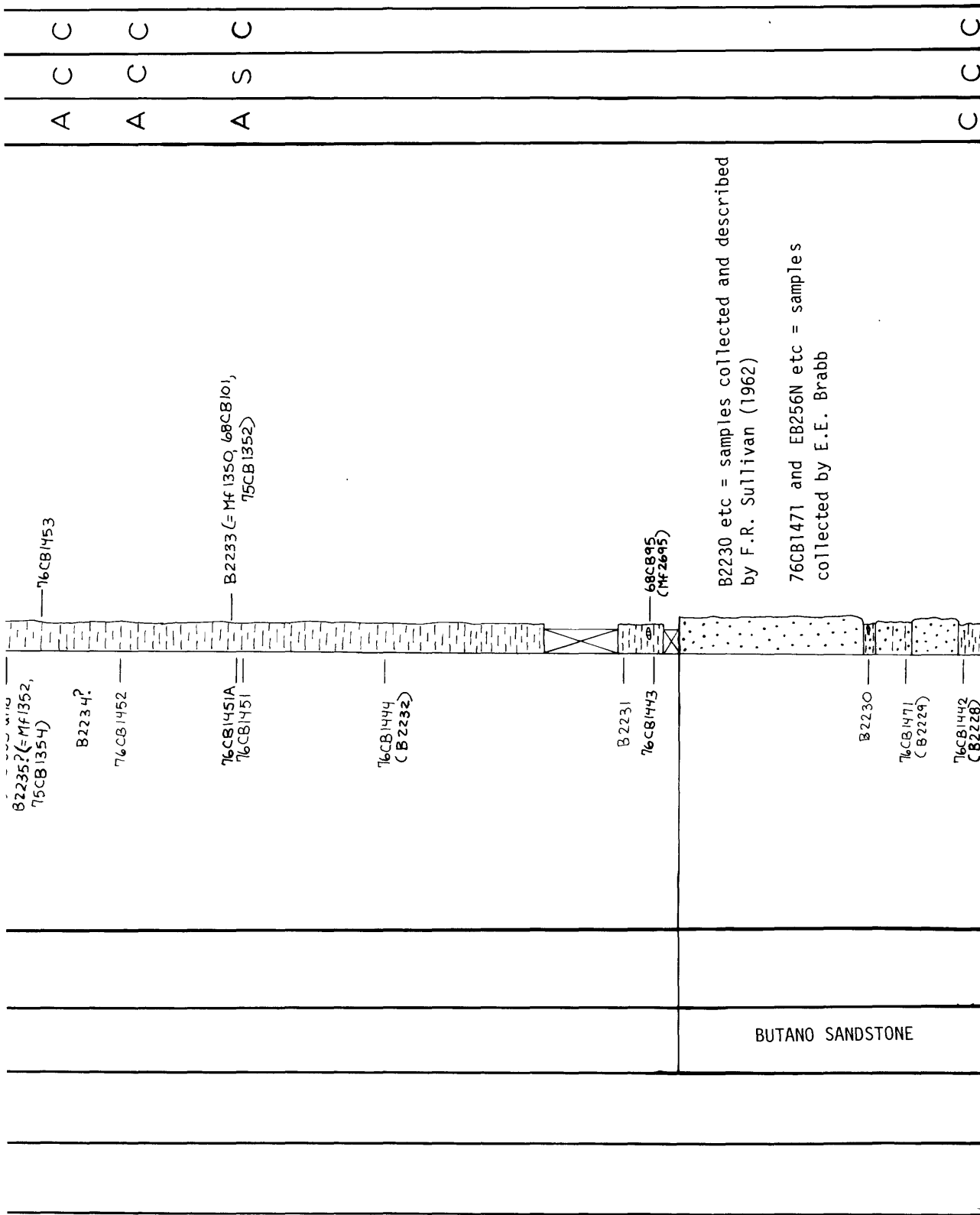


Figure 1



tinued

KINGS CREEK SECTION

The Kings Creek area (fig. 13) was designated by Brabb (1964) as the type section for the Twobar Shale and Rices Mudstone Members of the San Lorenzo Formation. After the section was measured in 1958, sediment from road construction along the upper part of Kings Creek was washed into the creek, covering part of the section that was formerly exposed. Heavy snowfalls in the 1970's felled many trees into the creek, thereby making access difficult in some places. Nevertheless, the section still has excellent exposures of the San Lorenzo, Butano and Vaqueros Formations.

The stratigraphic column for the Kings Creek area is shown on figure 14. Foraminifers from the formations are listed on table 2. Three samples (B4357, B4358, and B4369) collected and described by Fairchild and others (1969) are not shown. The section was recollected by the senior author in 1968 but these samples have not been completely processed. Preliminary examination by R. Z. Poore indicates that planktonic foraminifers and nannoplankton are not as abundant or as well preserved as in the San Lorenzo River section.

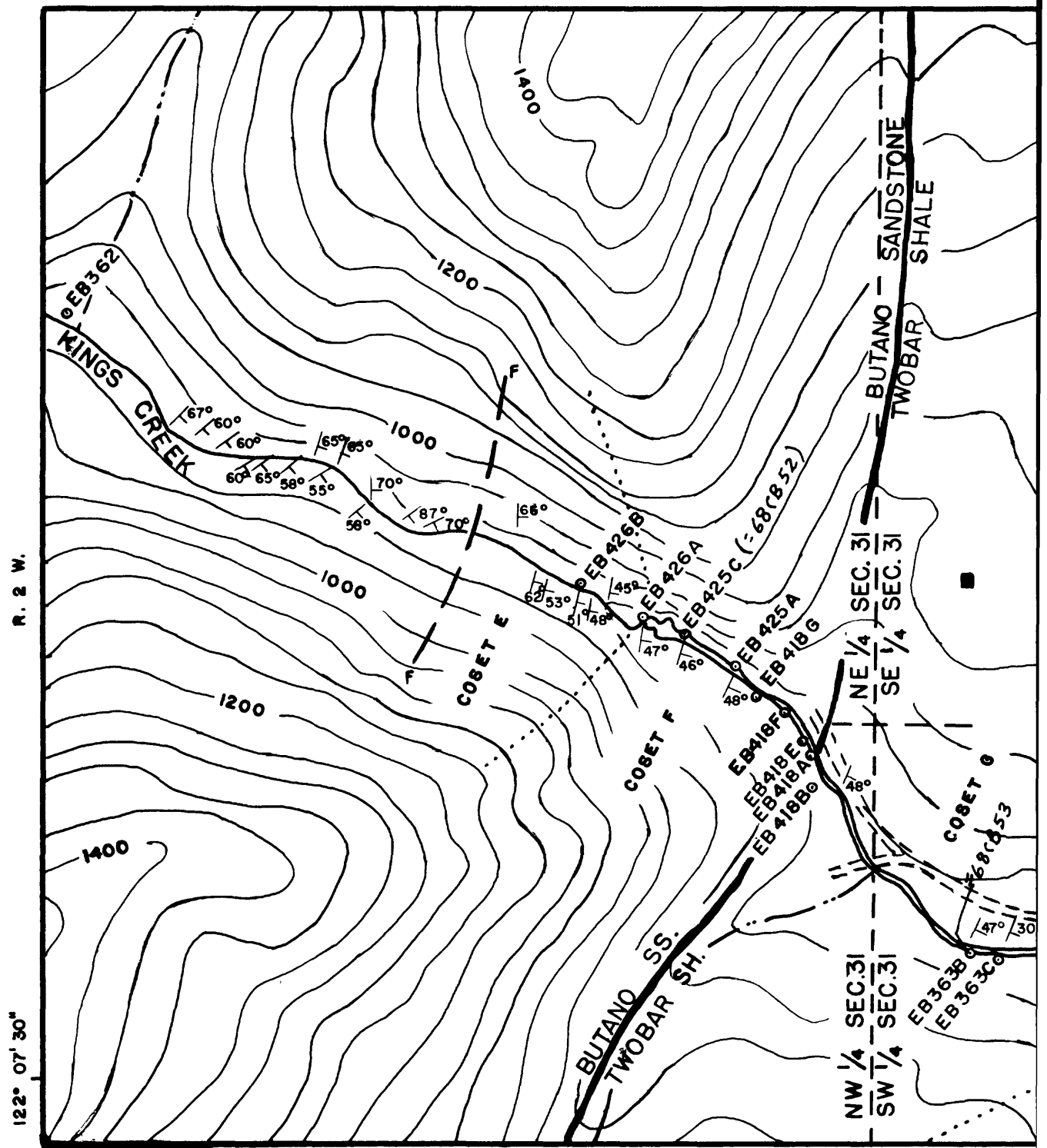
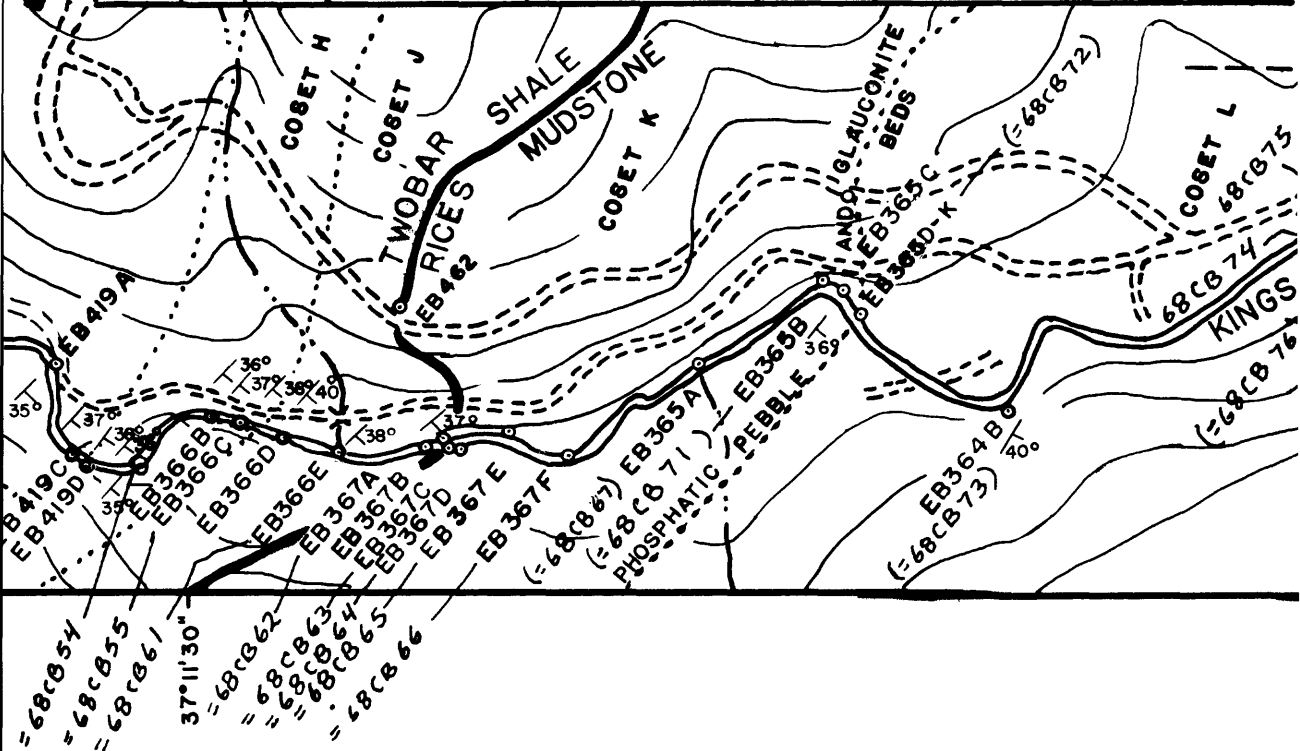


Figure 13.--Type section of the Rices Mudstone and Twoba

37°11'30"

STRATIGRAPHIC UNITS

AGE			FORMATION	MEMBER	THICKNESS IN FEET	REMARKS
PERIOD	EPOCH	STAGE				
TERTIARY	EOCENE	OLIG.	VAQUEROS SANDSTONE		>1170	VERY FINE TO MEDIUM-GRAINED SANDSTONE WITH INTERBEDS OF MUDSTONE AND SHALE
		REFUGIAN	SAN LORENZO FORMATION	RICES MUDSTONE	1030	MASSIVE FOSSILIFEROUS OLIVE GRAY MUDSTONE WITH GLAUCONITE AND PHOSPHATIC PEBBLE BEDS AND CARBONATE CONCRETIONS
	MARIZIAN	TWOBAR SHALE		790	LAMINATED OLIVE BLACK FORAMINIFERAL SHALE AND MINOR VERY FINE-GRAINED SANDSTONE	
			BUTANO SANDSTONE		>600	SANDSTONE WITH INTERBEDS OF SHALE AND MUDSTONE



Shale Members of the San Lorenzo Formation along Kings Creek

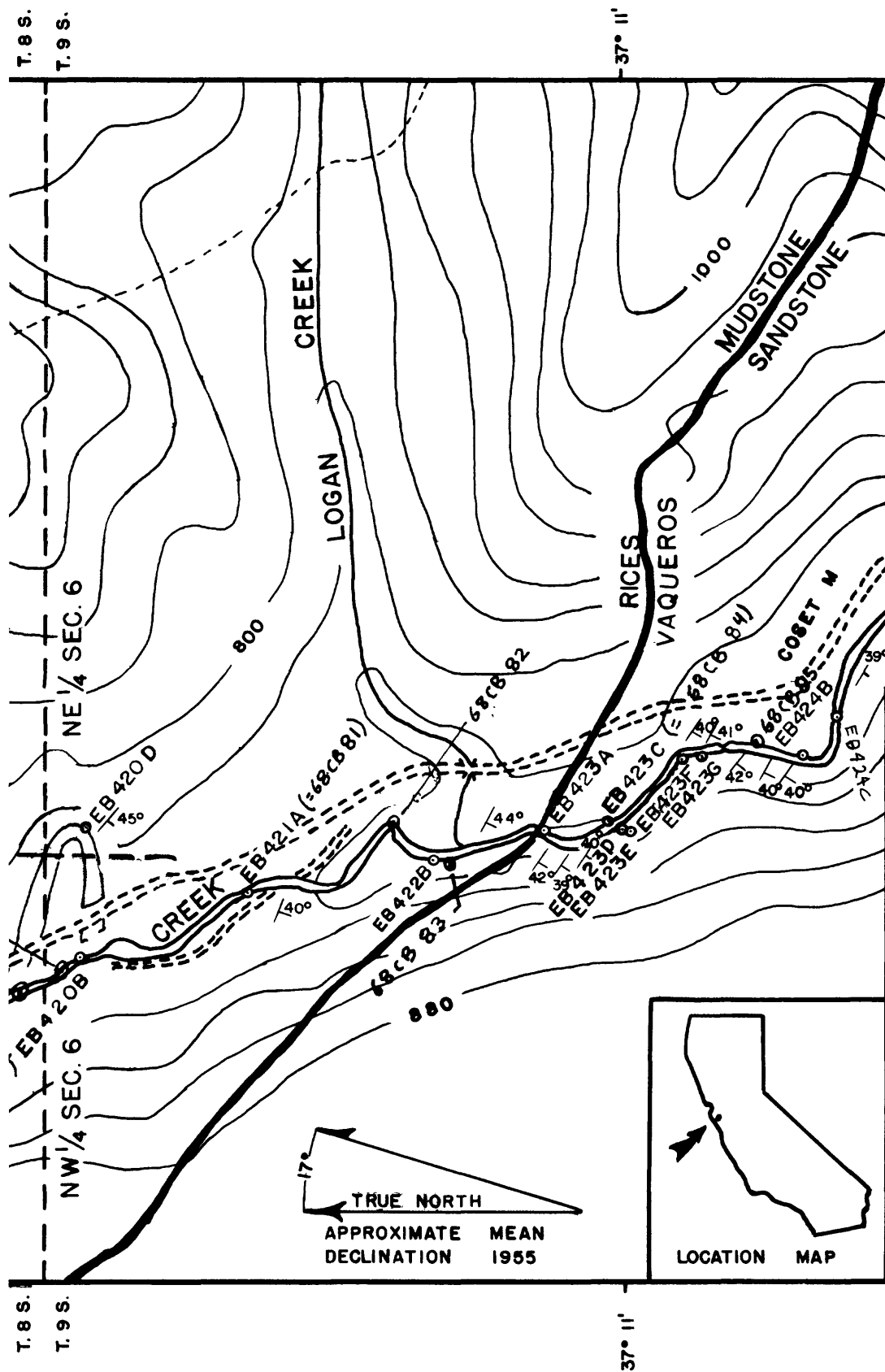
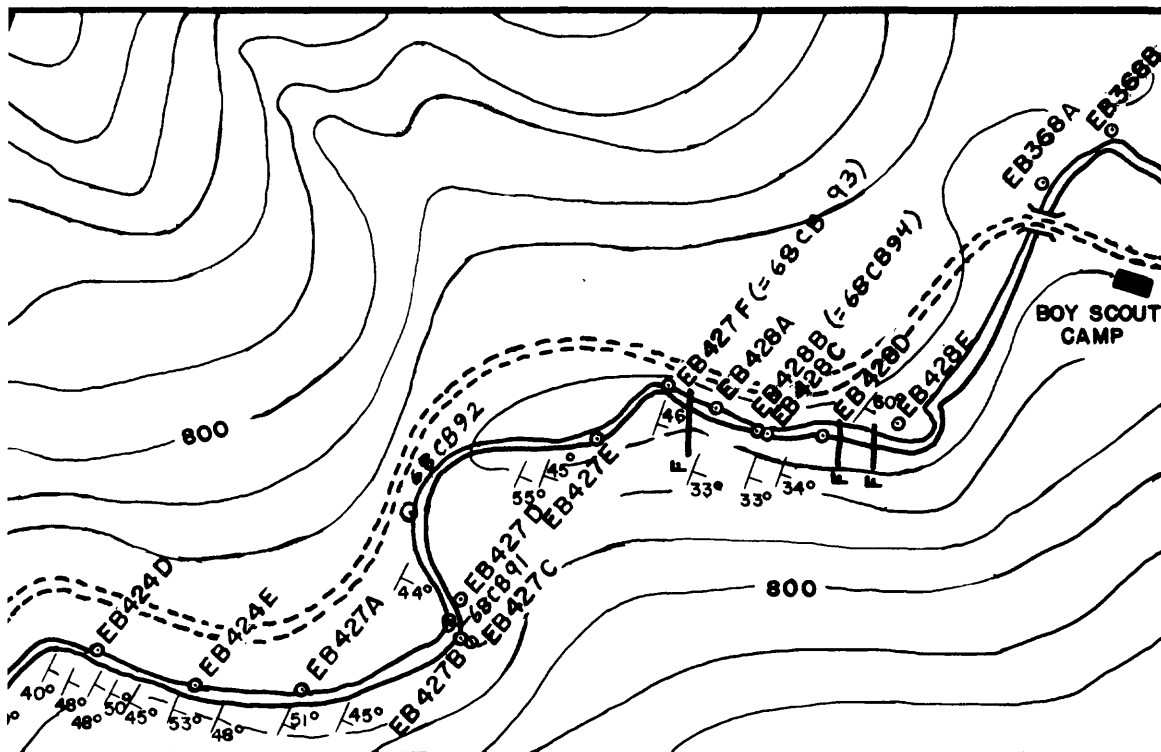


Figure 13.



TYPE SECTION OF THE RICES MUDSTONE
AND TWOBAR SHALE MEMBERS OF THE
SAN LORENZO FORMATION, KINGS CREEK,
CALIFORNIA

SECS. 31 AND 6, T.8 S. AND T.9 S., R.2 W., M. D. B. M.

FROM REPORT BY EARL E. BRABB, 1960

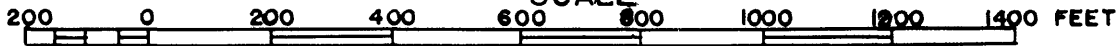
"GEOLOGY OF THE BIG BASIN AREA, SANTA CRUZ MOUNTAINS, CALIFORNIA"

BASE MAP ENLARGED FROM U.S. GEOLOGICAL SURVEY BIG BASIN AND CASTLE ROCK
RIDGE 7.5 QUADRANGLES, 1955 EDITIONS. COLLECTING LOCALITIES AND
DETAILS OF KINGS CREEK SURVEYED WITH BRUNTON COMPASS AND TAPE

LEGEND

- 30° STRIKE AND DIP OF BEDS
- COLLECTING LOCALITIES
- F FAULT
- FORMATION OR MEMBER CONTACT
- ... COSET CONTACT
- CONTOUR INTERVAL 40 FEET

SCALE



R. 2 W.

122°07'30"

ntinued

STRATIGRAPHIC COLUMN

OF THE TYPE TWOBAR SHALE AND RICES MUDSTONE MEMBERS OF THE
SAN LORENZO FORMATION, AND ADJACENT UNITS, KINGS CREEK,
CALIFORNIA

Sec. 31, T.8 S., R.2 W. and Sec. 6, T.9 S., R.2 W., Mount Diablo Base and Meridian

From report by Earl E. Brabb, 1960

"Geology of the Big Basin area, Santa Cruz Mountains, California"

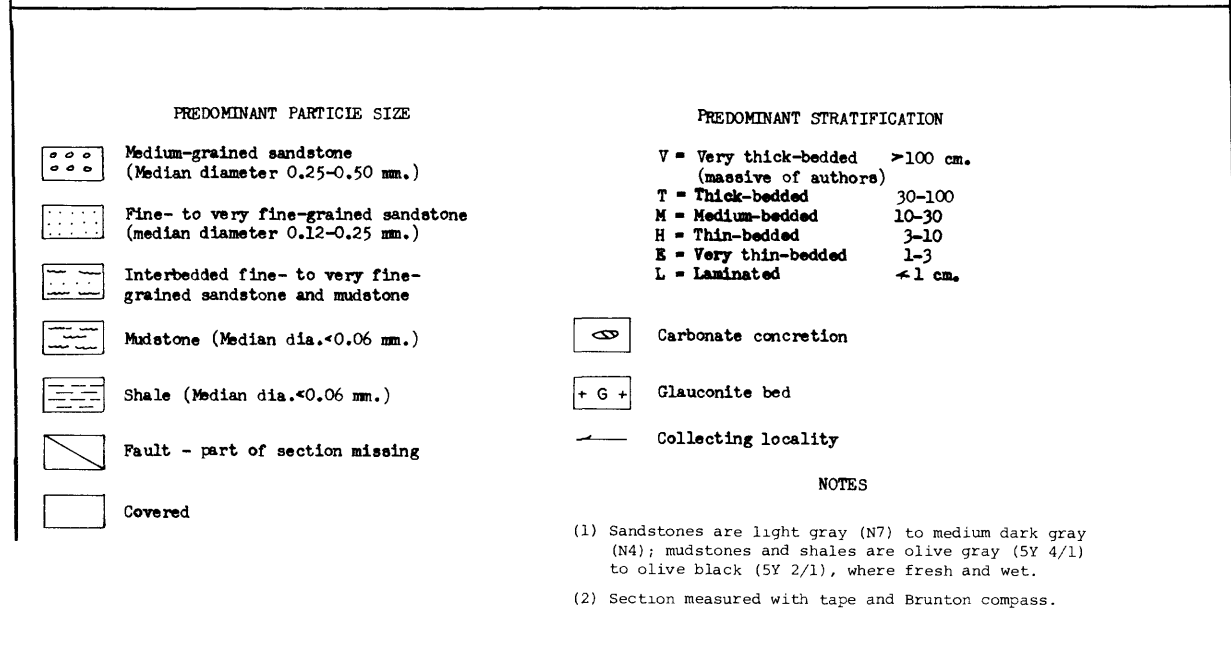


Figure 14.--Stratigraphic column, Kings Creek area

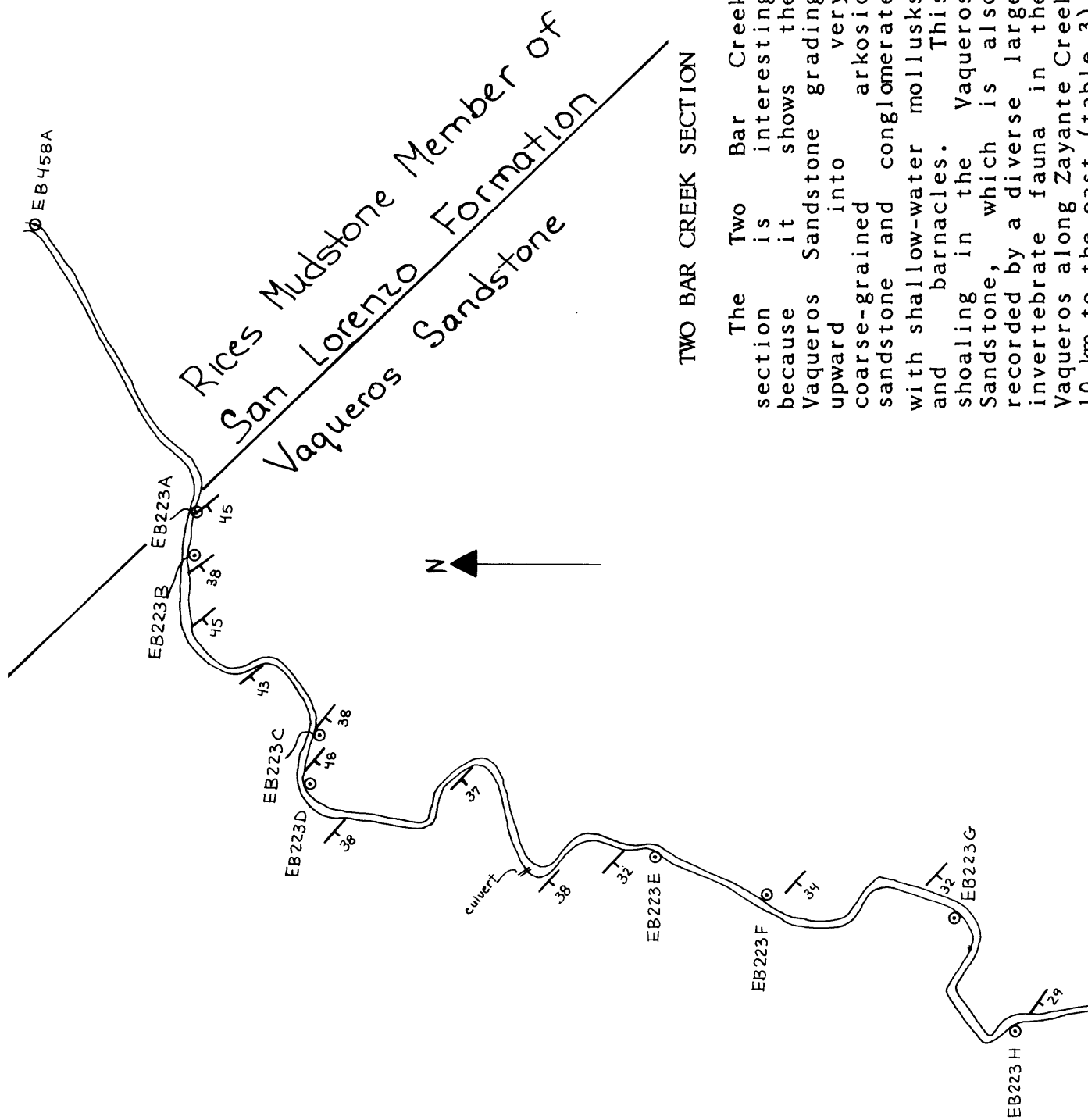
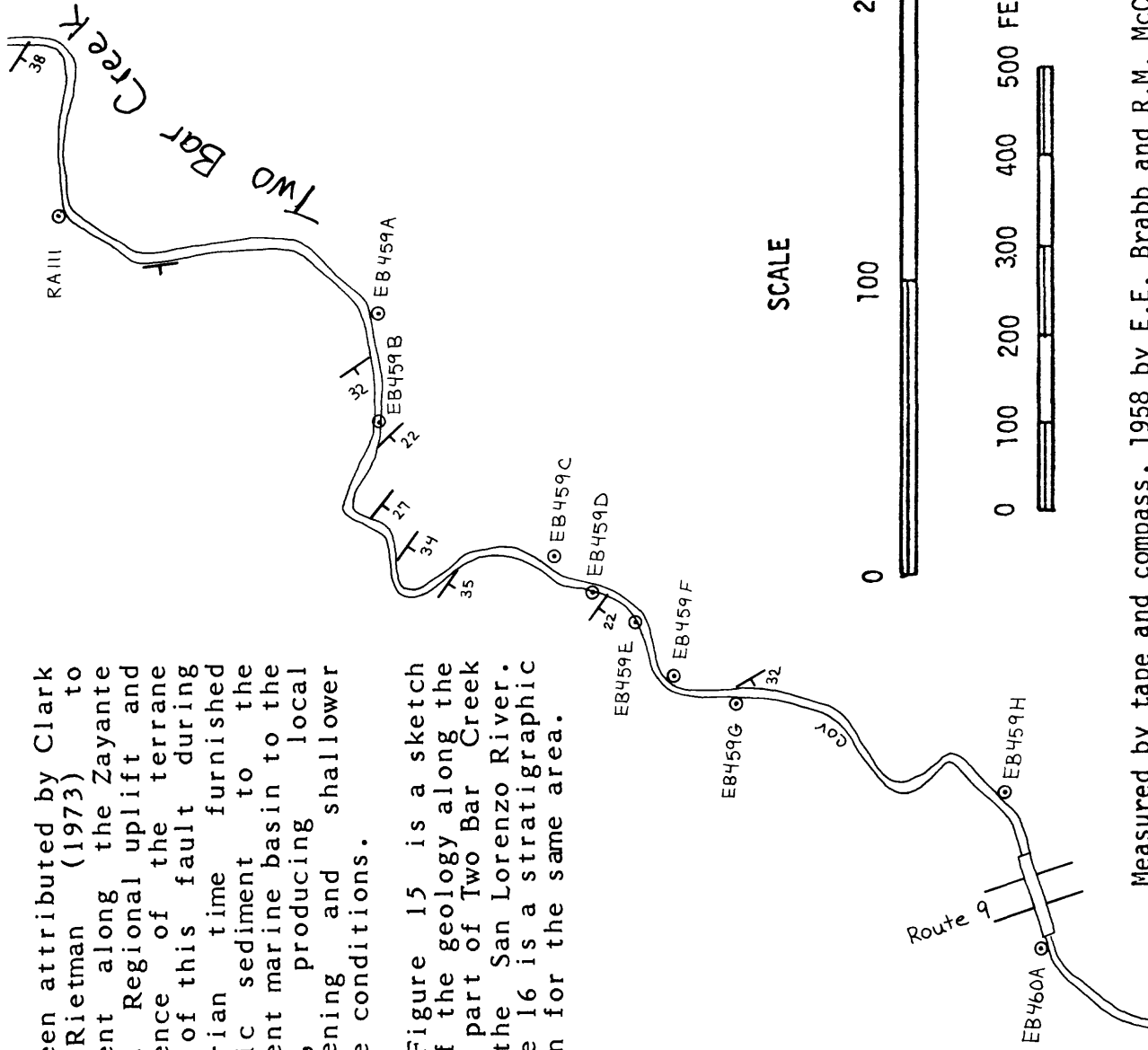


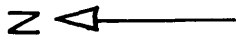
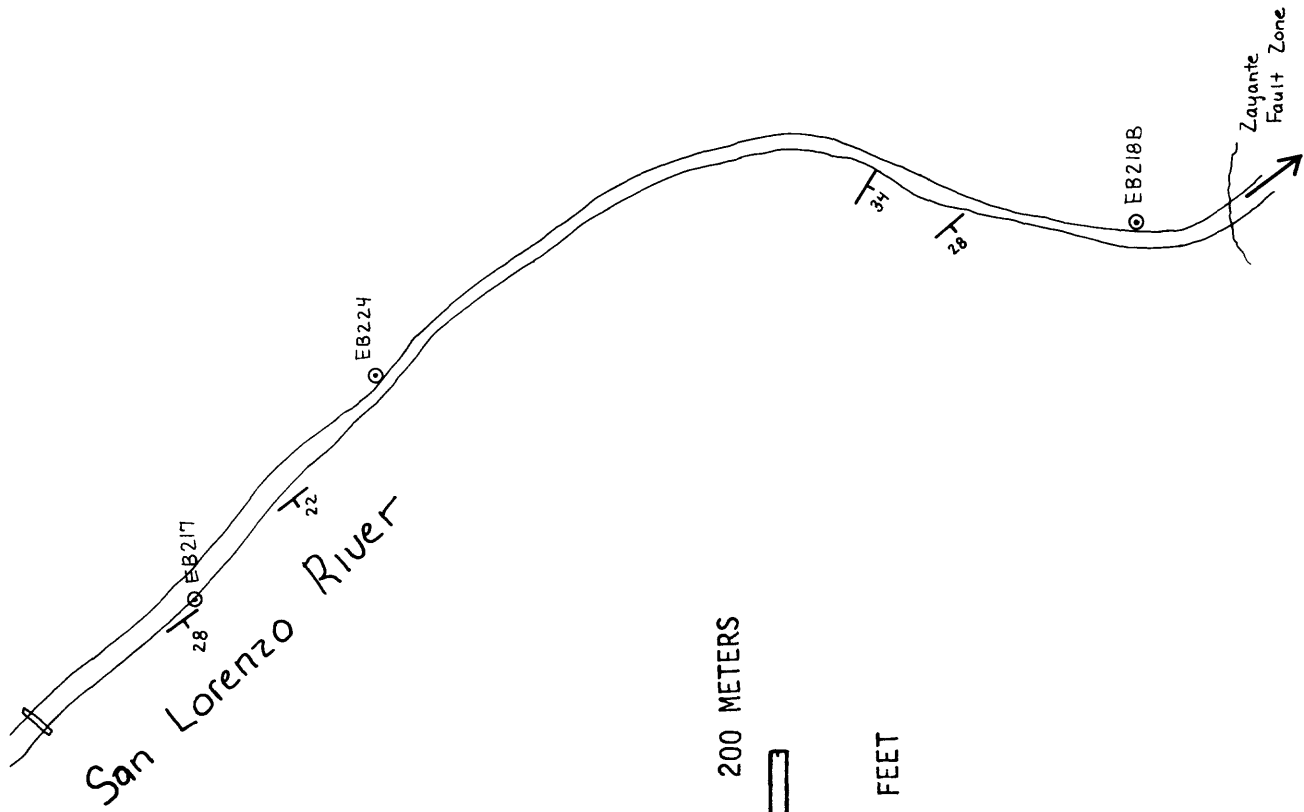
Figure 15.--Geologic sketch map of the lower Two Bar Creek and San Lorenzo River area showing formation contacts and localities where fossils and rocks were collected

has been attributed by Clark and Rietman (1973) to movement along the Zayante fault. Regional uplift and emergence of the terrane south of this fault during Zemorrian time furnished clastic sediment to the adjacent marine basin to the north, producing local coarsening and shallower marine conditions.

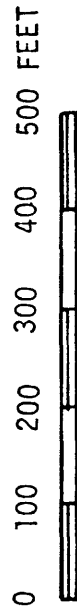
Figure 15 is a sketch map of the geology along the lower part of Two Bar Creek and the San Lorenzo River. Figure 16 is a stratigraphic column for the same area.



Measured by tape and compass, 1958 by E.E. Brabb and R.M. McCollom, Jr.



SCALE



continued

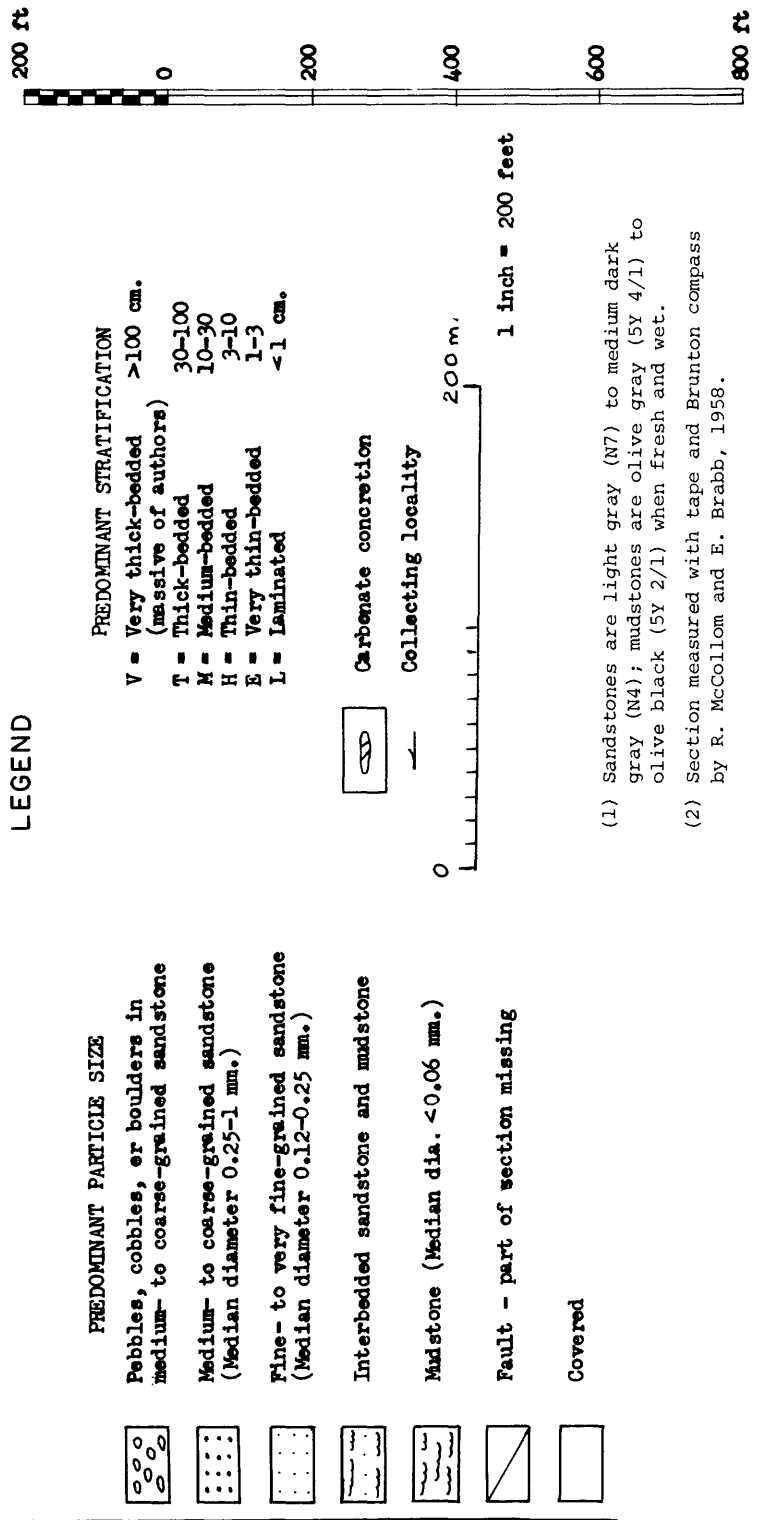
STRATIGRAPHIC COLUMN

OF THE VAQUEROS SANDSTONE AND UPPER PART OF THE
SAN LORENZO FORMATION, TWO BAR CREEK AND
SAN LORENZO RIVER SECTION

Secs. 18 and 19, T. 9 S., R. 2 W., Mount Diablo Base and Meridian

From report by Earl E. Brabb, 1960

"Geology of the Big Basin area, Santa Cruz Mountains, California"



BEAR CREEK SECTION

The Bear Creek area has been the training ground for a number of notable geology graduates of Stanford University who studied under the late Professor H. G. Schenck, such as R. M. Kleinpell, T. W. Dibblee, Jr., L. C. Forrest, H. D. Hobson, and L. J. Simon. All worked in the Bear Creek area, but no one has published previously a detailed locality map or stratigraphic column. Several paleontologic reports have been published, however. The earliest are by Arnold (1906 and 1908), who described some of the mollusks. Reports by Hobson (1932) and Cushman and Hobson (1935) mentioned collections from the Bear Creek area, including those that have holotypes for Robulus barbati, Cassidulina crassipunctata, Anomalina californiensis and Cibicides pseudoungerianus evolutus. Schenck (1936, p. 67-8) collected Acila dalli and other fossils from the Bear Creek area--his localities are shown on a map by Brabb (1960, fig. 14). Kleinpell (1938, table 10) lists benthonic foraminifers from the upper part of the San Lorenzo Formation along Bear Creek. Burchfiel (1964) has a generalized geologic map that shows a few of the important fossil localities. In particular, he discusses the Teilzone of Pecten sanctaecruzensis, which has been used to correlate rocks in the Santa Cruz Mountains with those in the Santa Lucia Range, 120 km to the south. Fairchild and others (1969, fig. 5) provided a geologic and locality map of the Upper Bear Creek area and a check list of foraminifers. Their discussion of the age and paleoecology of the faunas is very instructive, but their samples along the creek are difficult to locate precisely because of the paucity of cultural features. Their Bear Creek samples, together with their samples from the Brown School area nearby, provide the most extensive documentation yet published of the character of foraminifers from the middle part of the Butano Sandstone.

The geologic sketch map, figure 17, shows the upper part of the Bear Creek area. This section should be continued to the northeast to include the lower part of the Butano Sandstone. A section should also be measured along the lower part of Bear Creek where many of the fossil collections mentioned above were taken. Figure 18 is a stratigraphic column for the part of the section measured for the October conference.

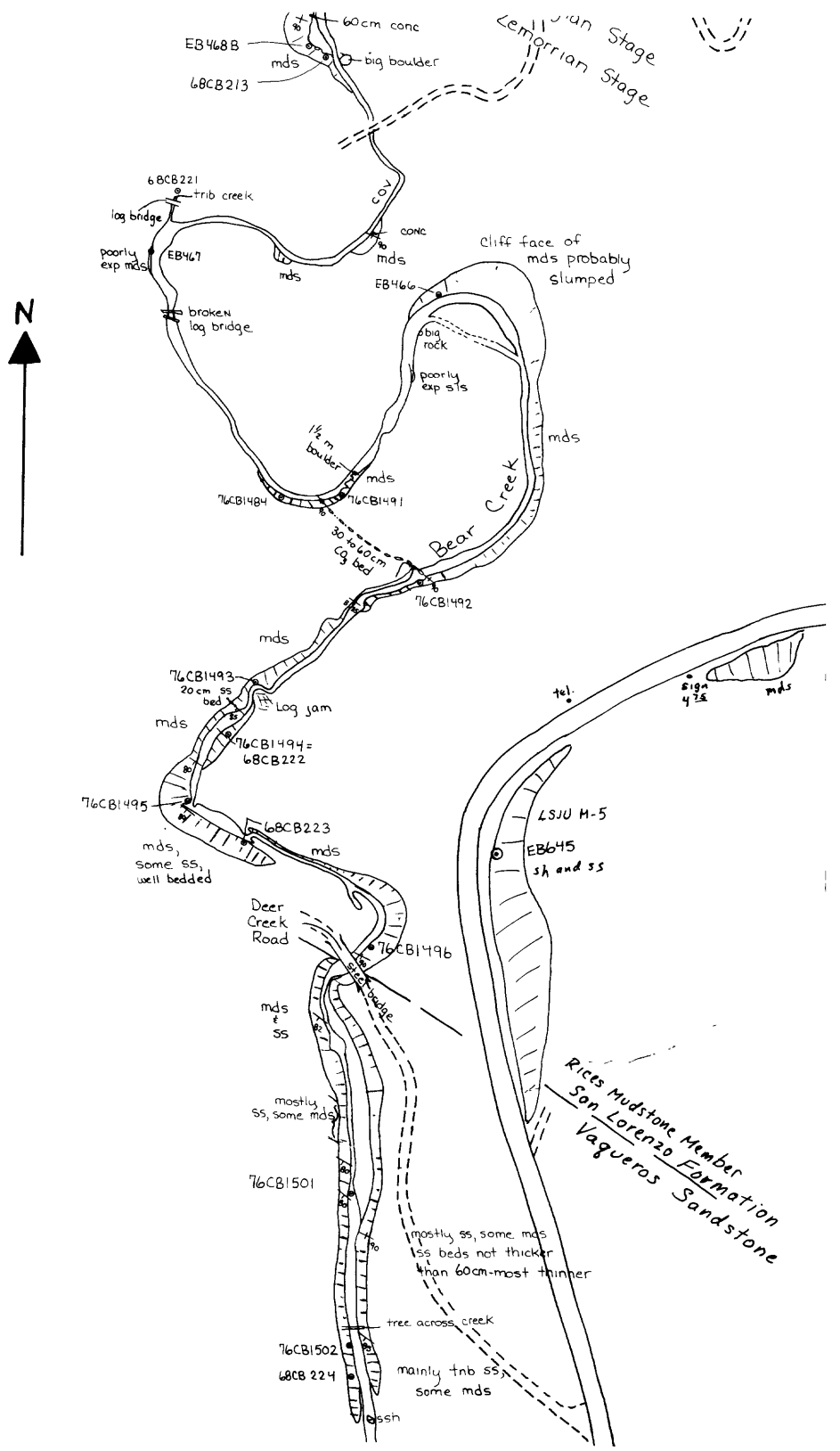
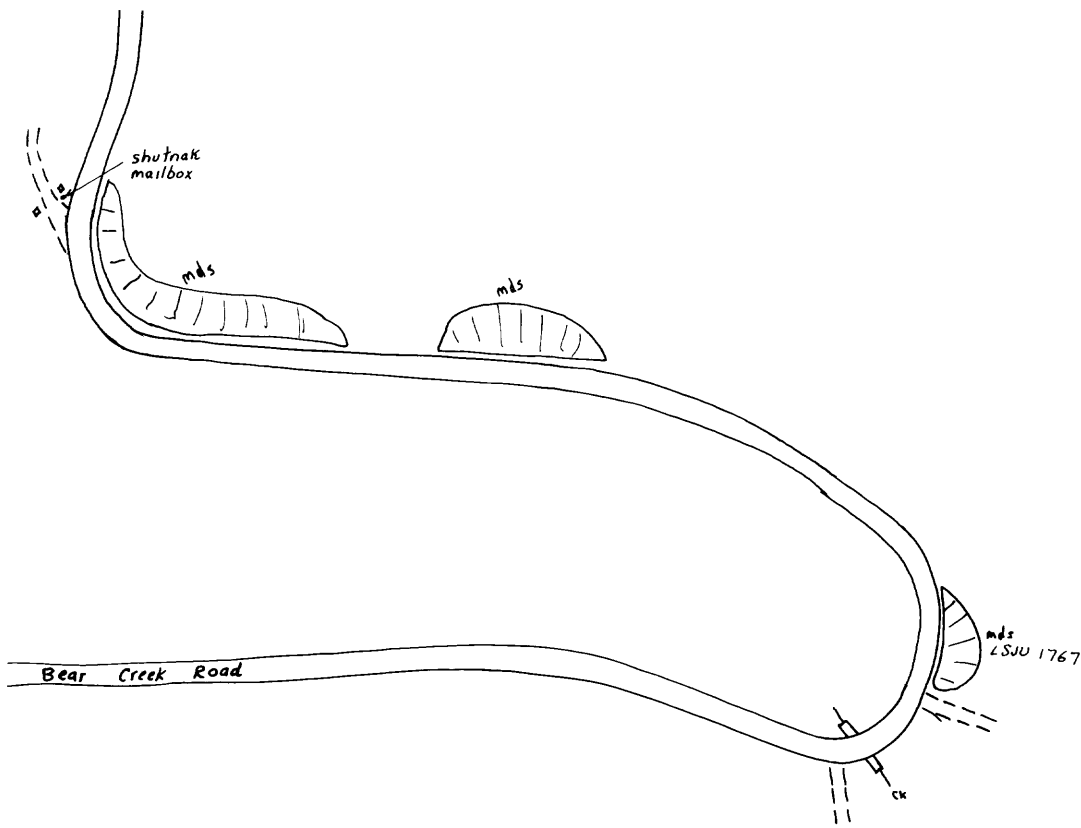


Figure 17.



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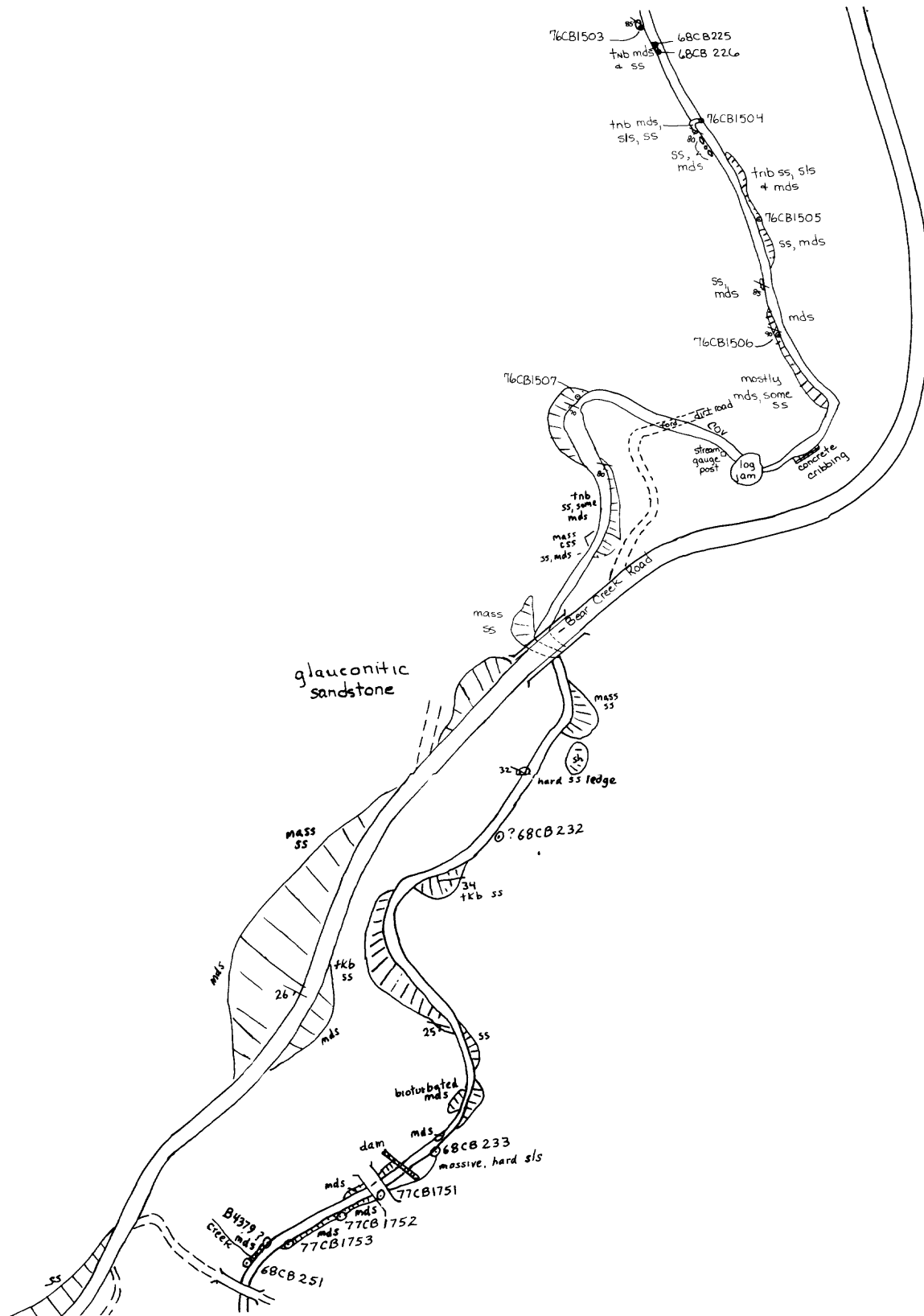
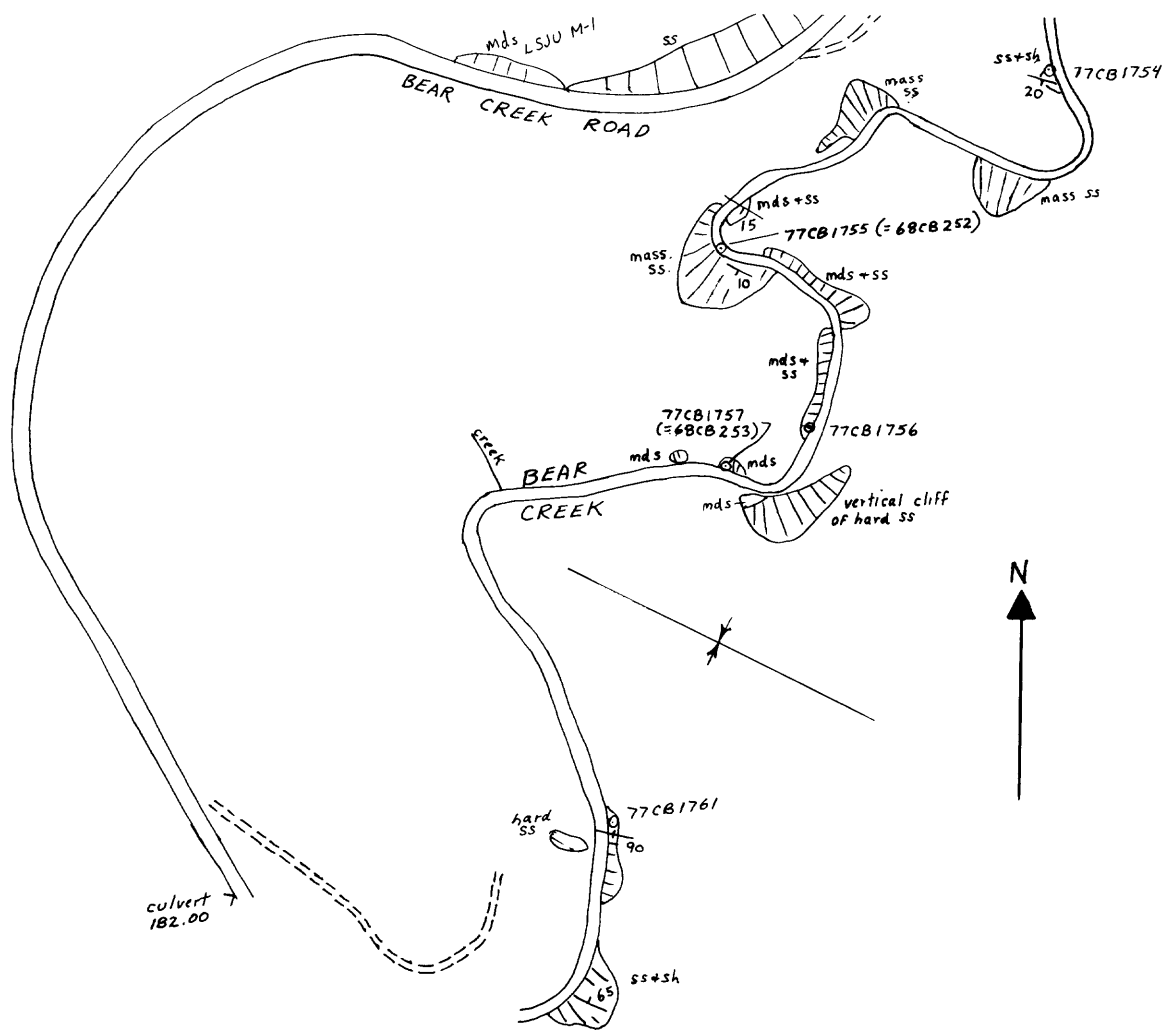


Figure 17.--



continued

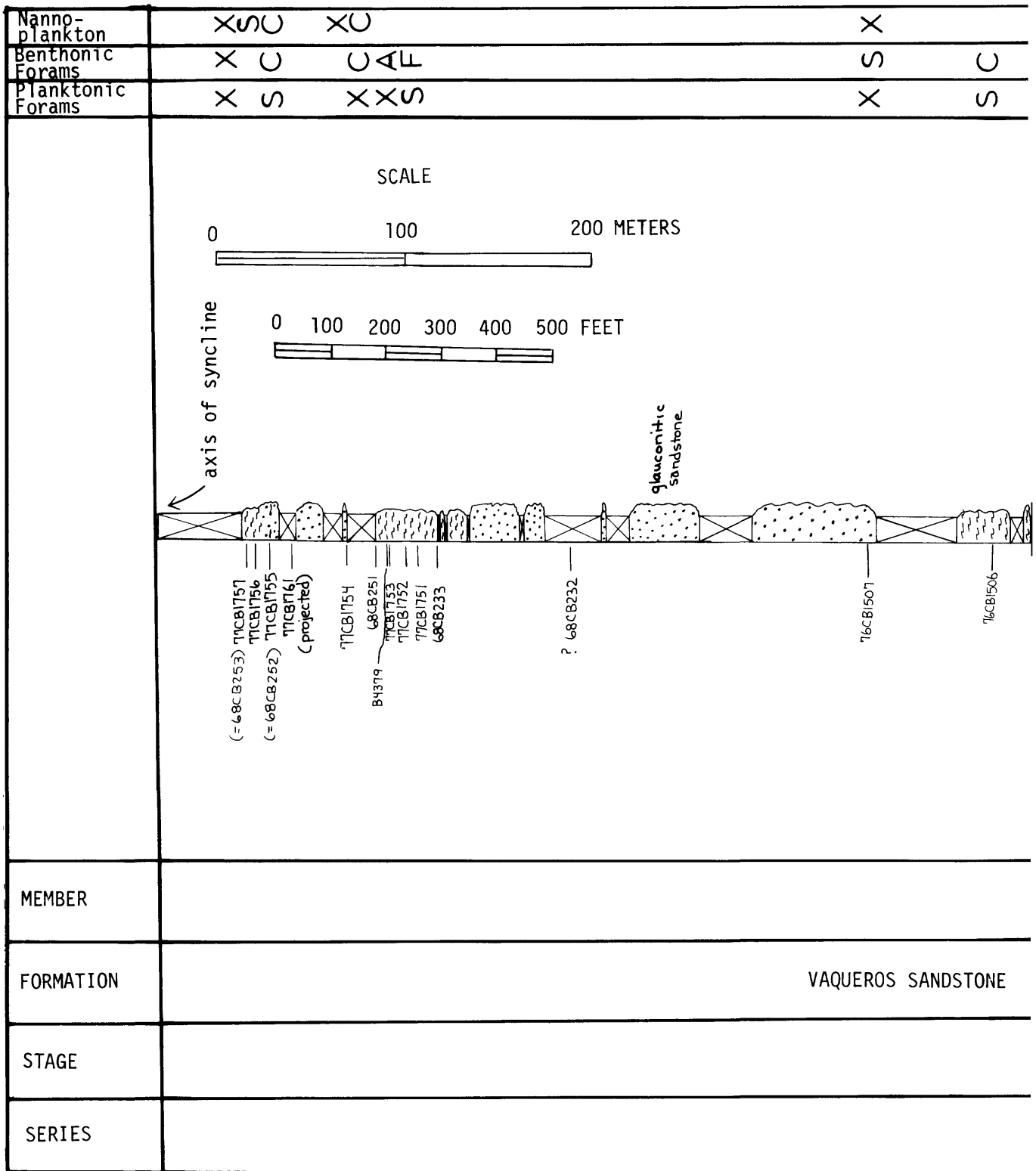
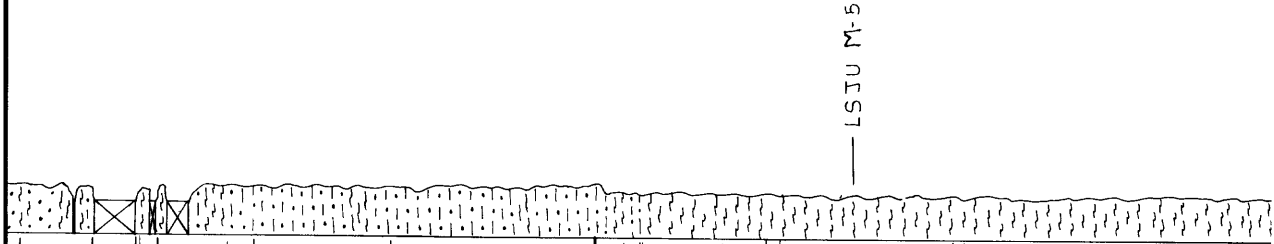


Figure 18.--Stratigraphic

K	S	C	C	S	S	S	C	C	S	S
D	A	C	C	A	C	A	A	C	A	A
K	S	X	X	S	S	S	S	X	S	S



76CB1505	CB1504	68CB225	68CB226	76CB1503	68CB224	76CB1502	76CB1501	76CB1496	68CB223	76CB1495	(= 68CB222) 76CB1494	76CB1493	76CB1484	76CB1492	76CB1491	EB467	68CB221	EB466
RICES MUDSTONE																		
SAN LORENZO																		

ZEMMORIAN

OLIGOCENE

column, Bear Creek area

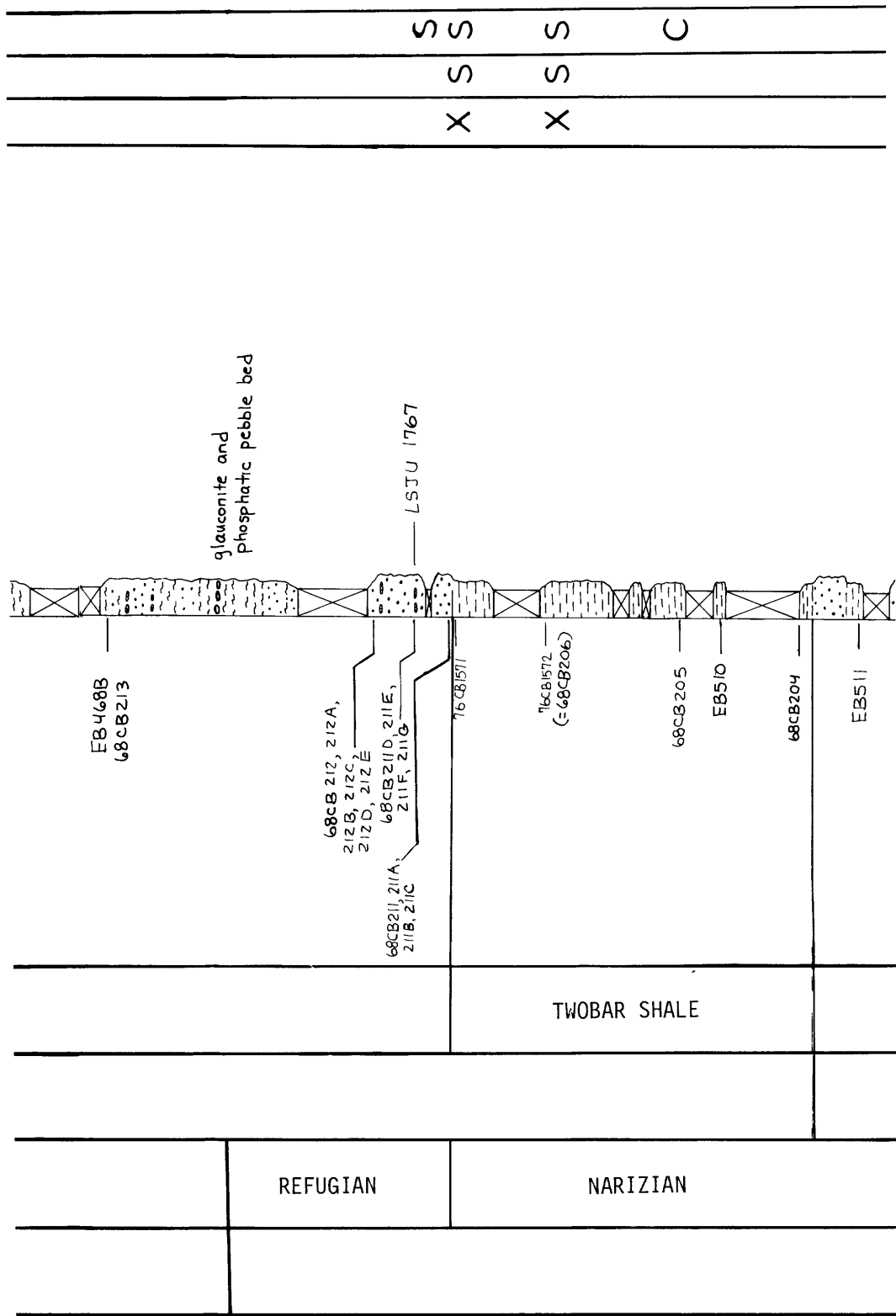
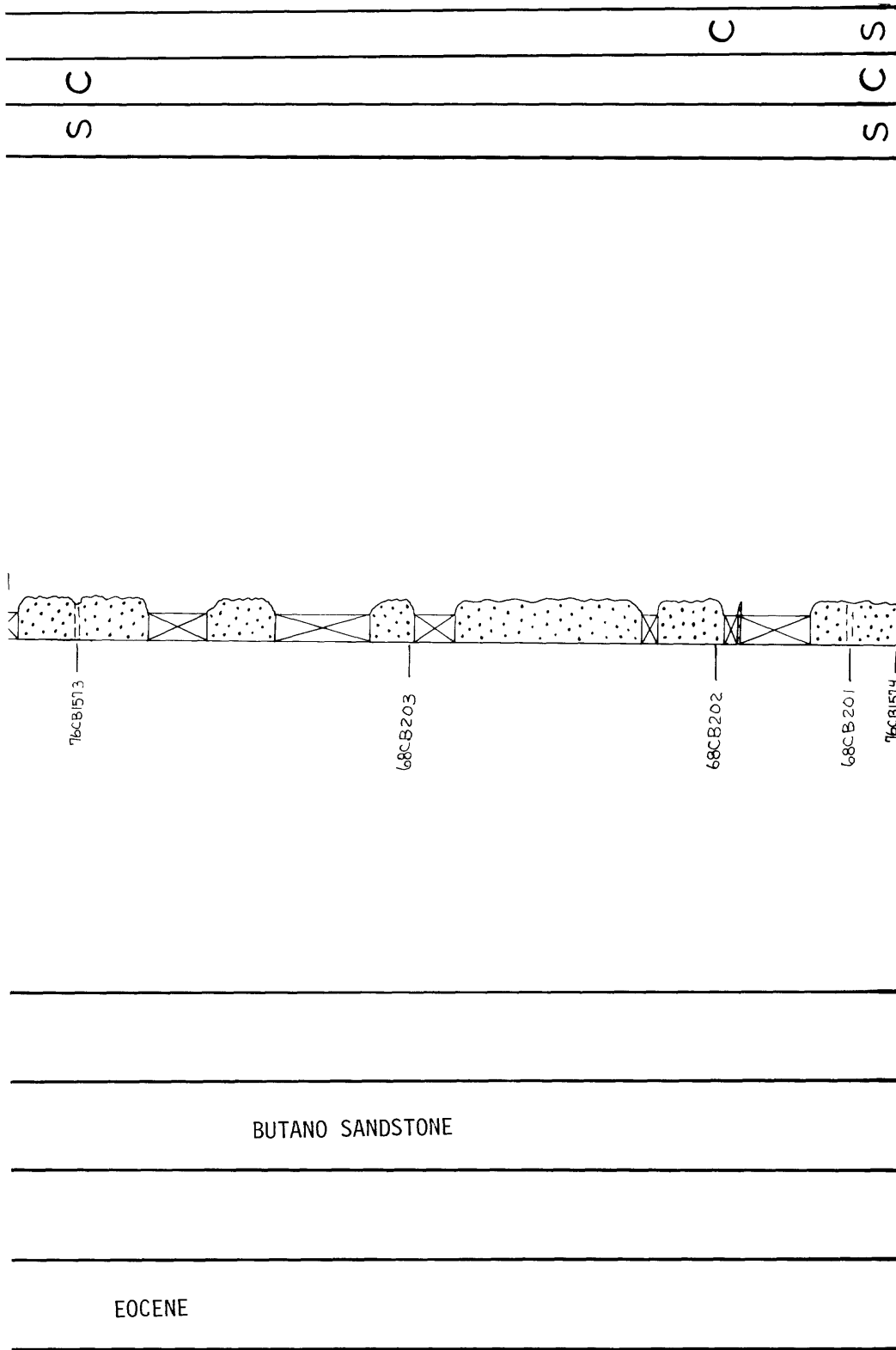


Figure 18.



Continued

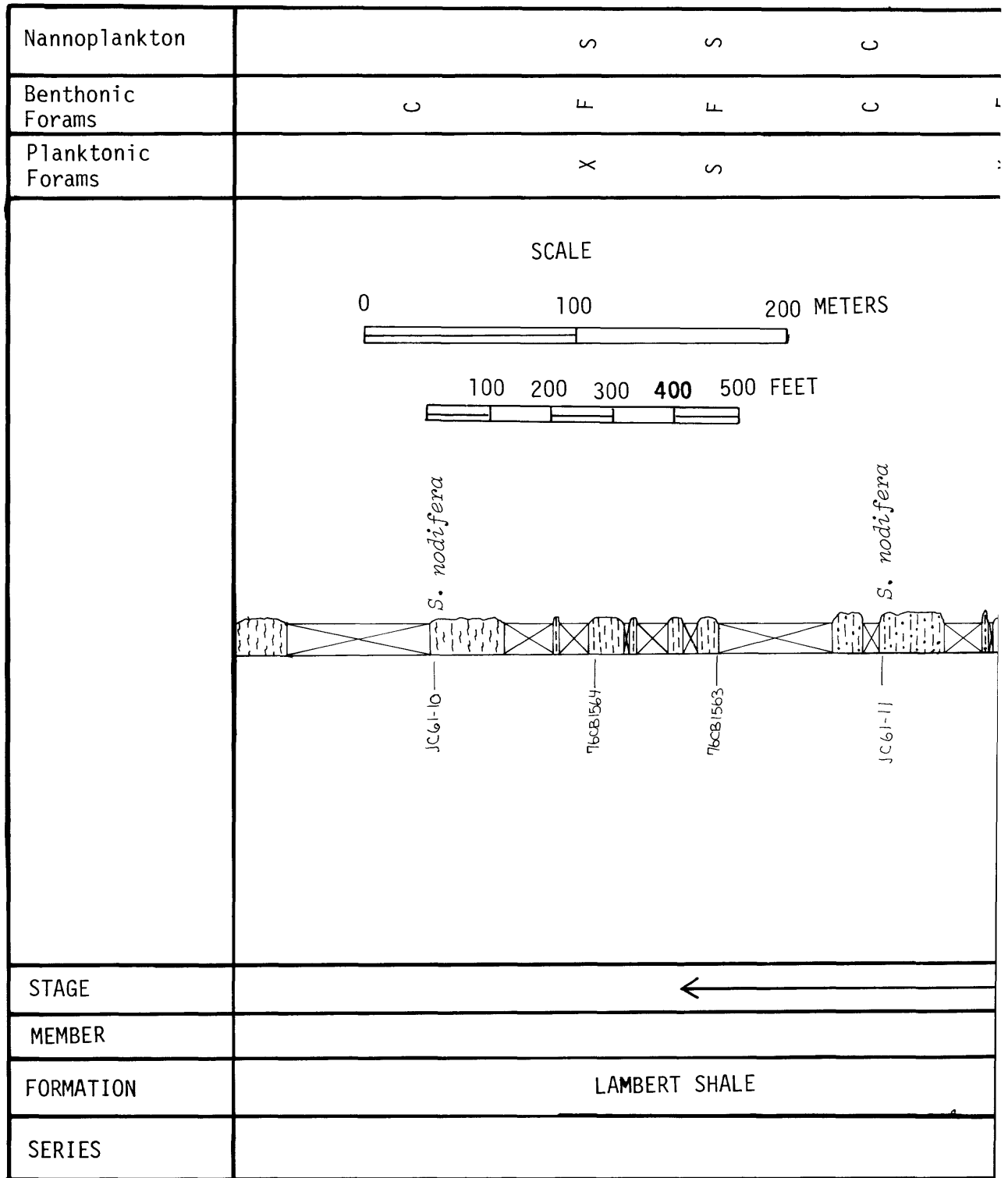
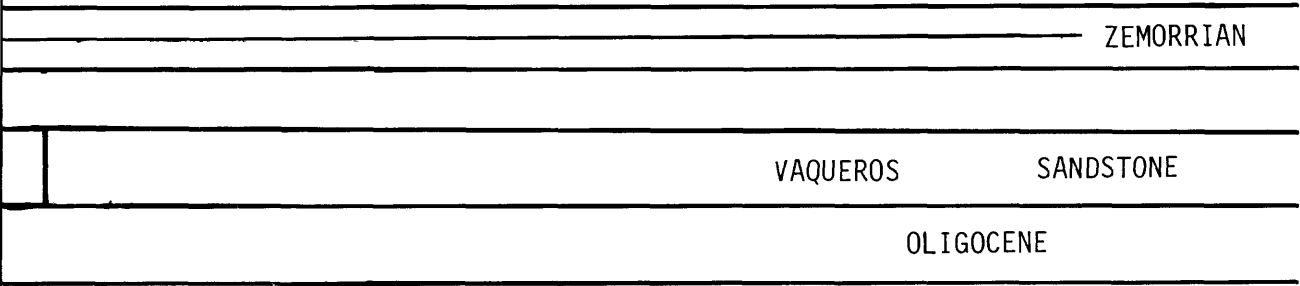
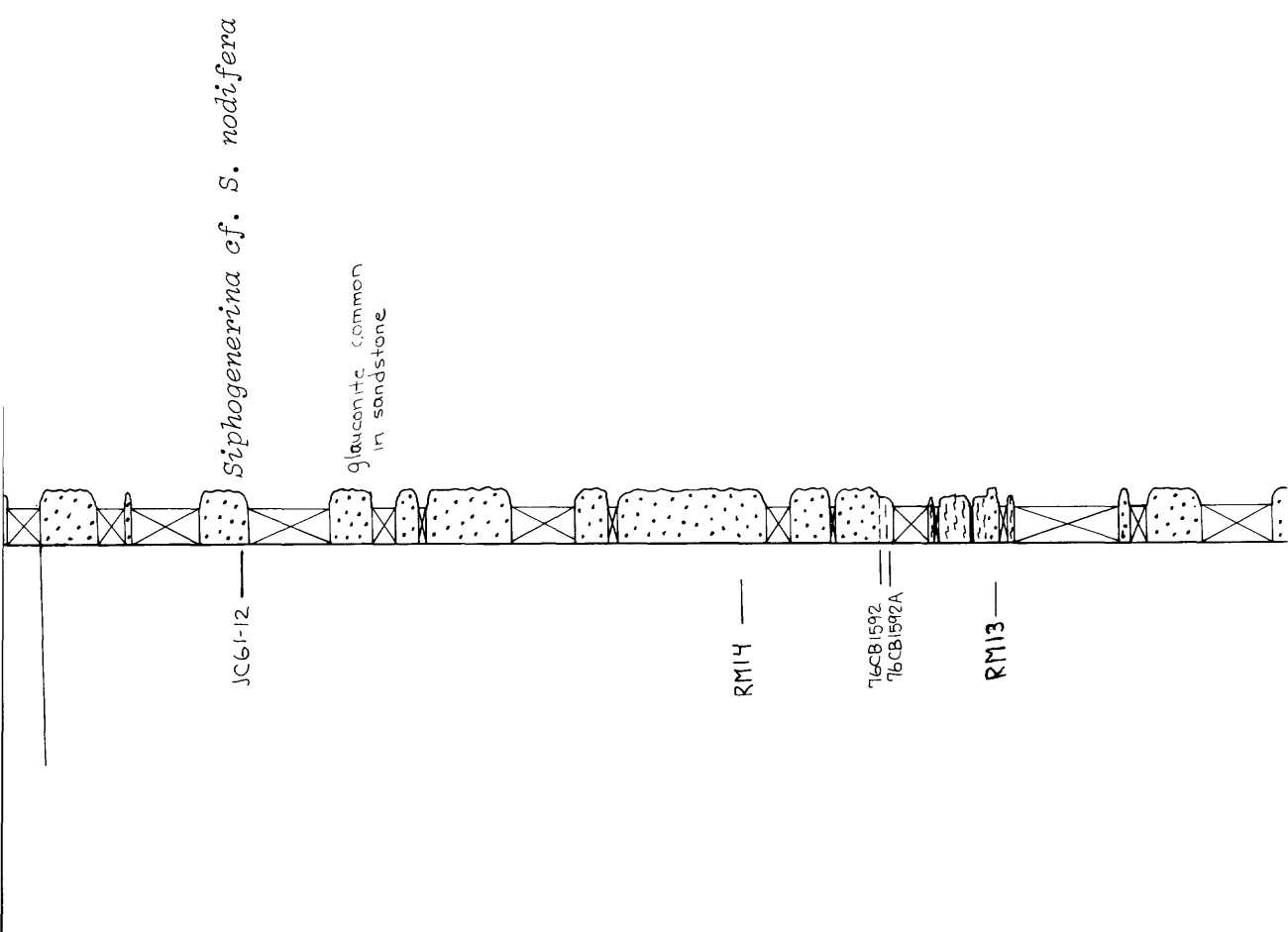
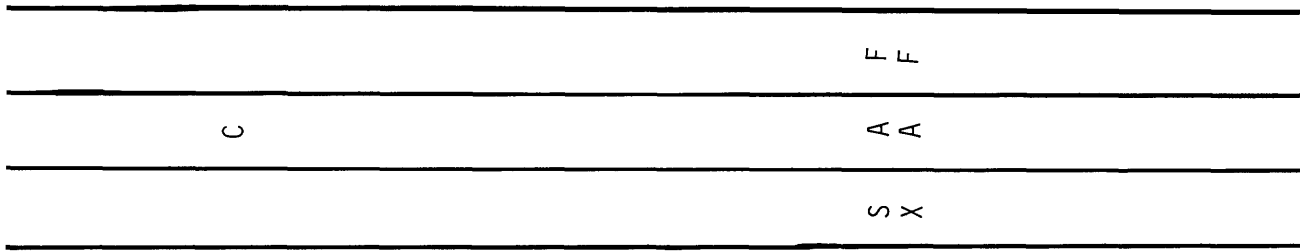


Figure 20.--Stratigraphic column, upper Zayan



reek area, north limb of San Lorenzo syncline.

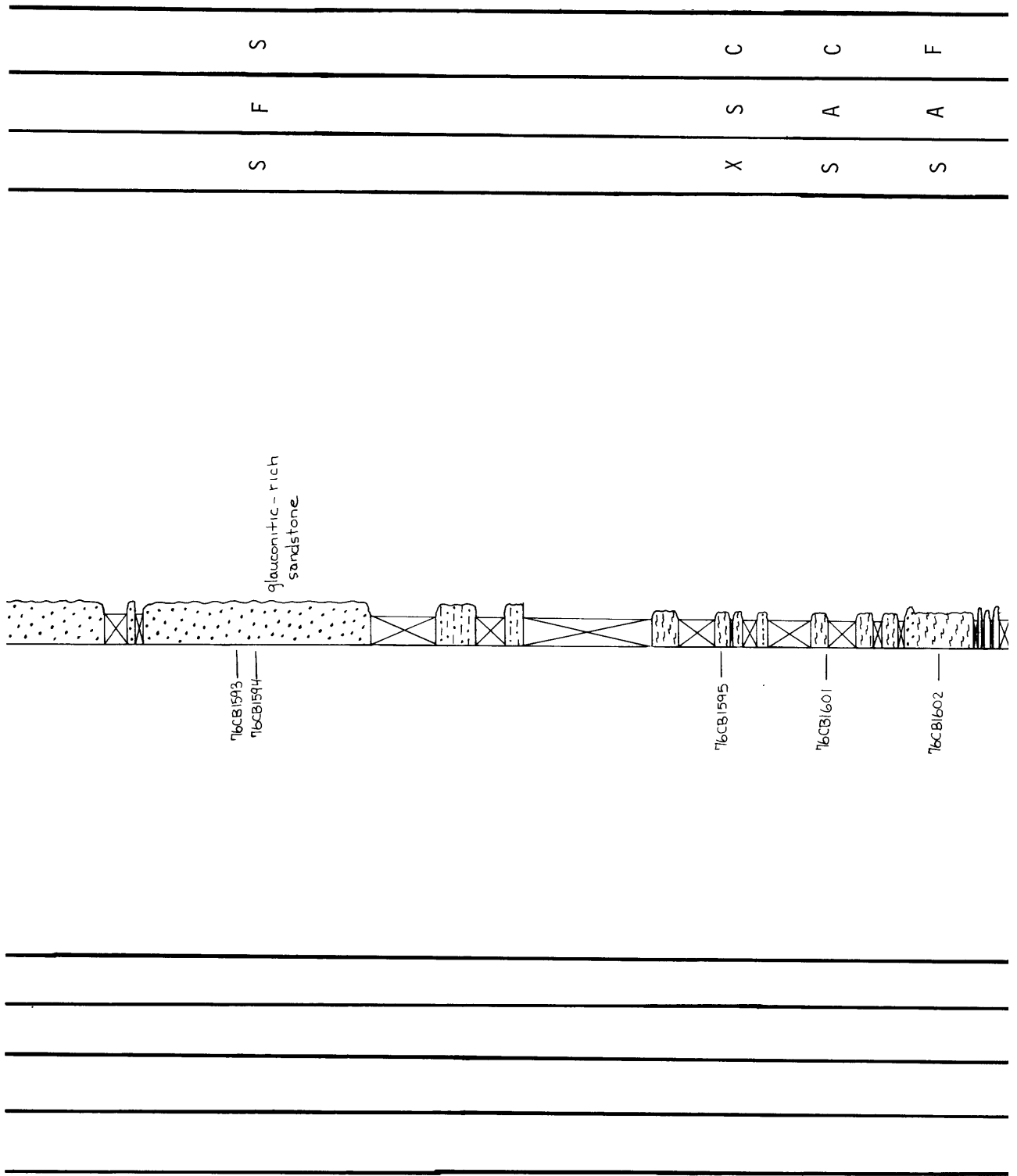
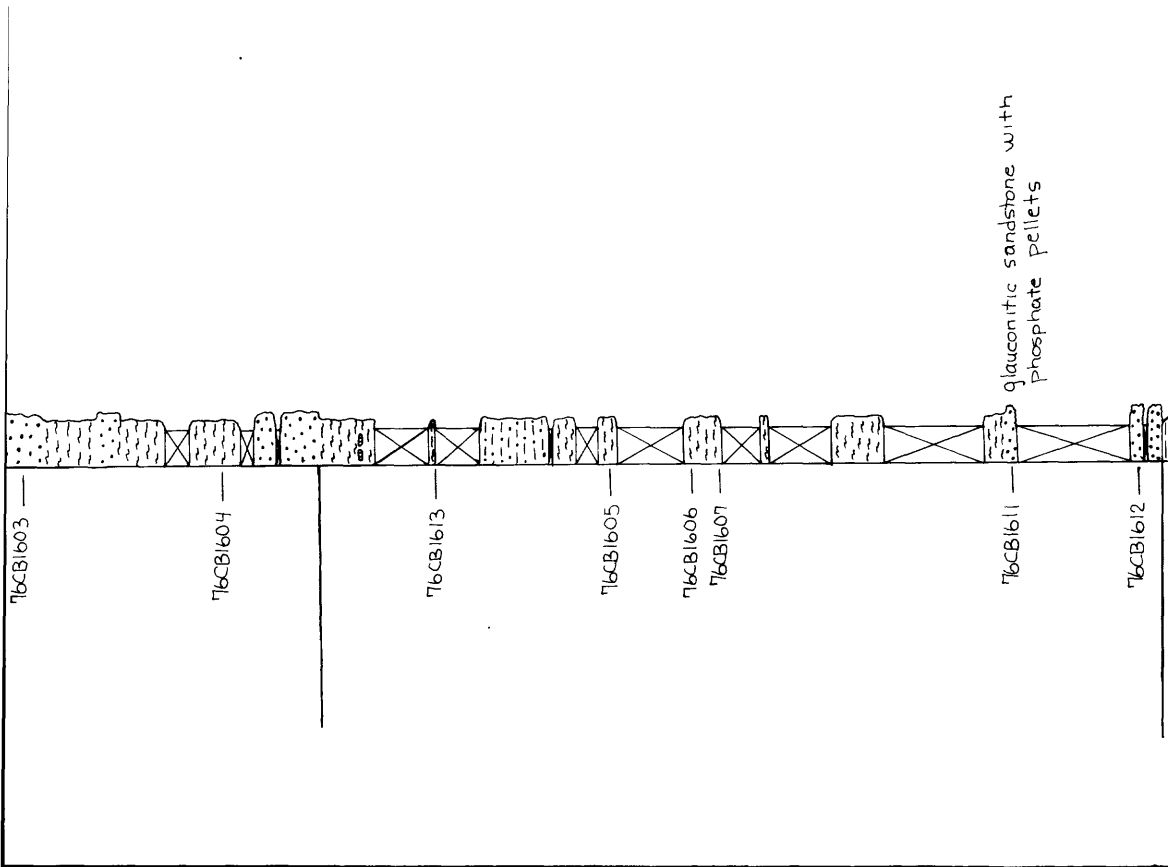


Figure 20.-

F	S		S S
U	A	F	C C
S	S	X	X S



		REFUGIAN ?
	RICES MUDSTONE	TWOBAR SHALE
	SAN LORENZO	
		EOCENE ?

continued

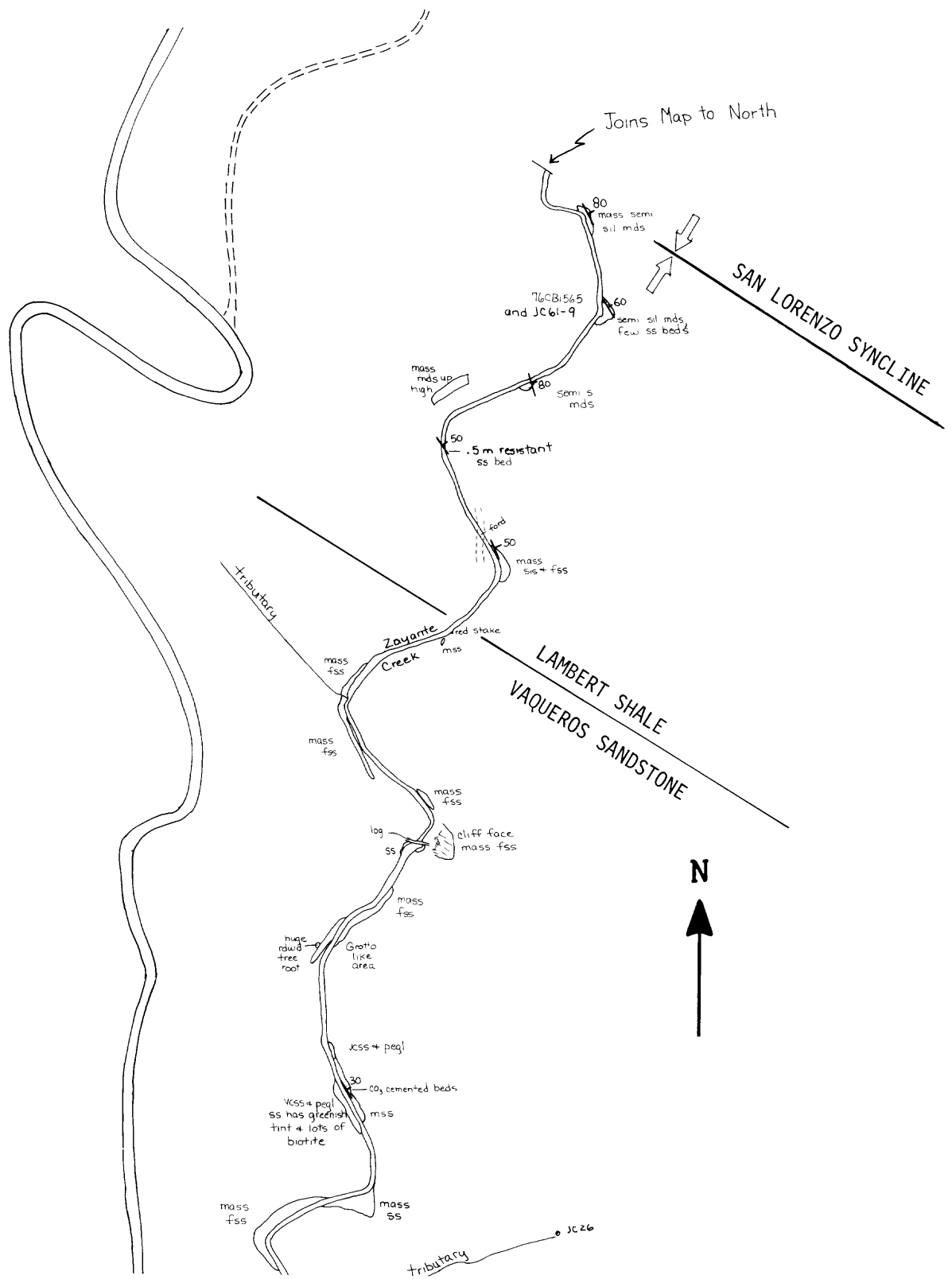


Figure 21.--Geologic sketch map of upper Zayante Creek area, south limb
 rocks and fossils were collected.

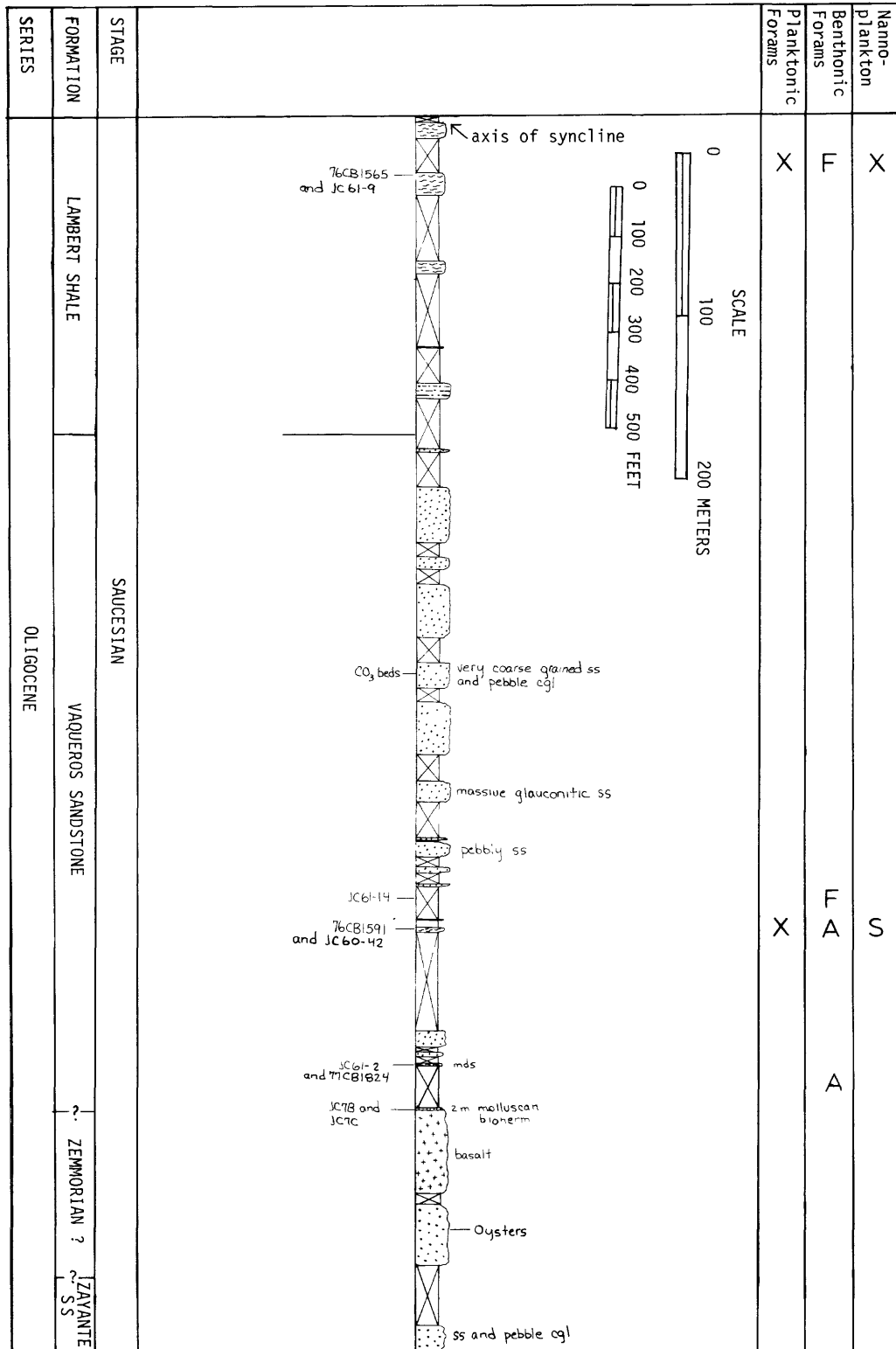


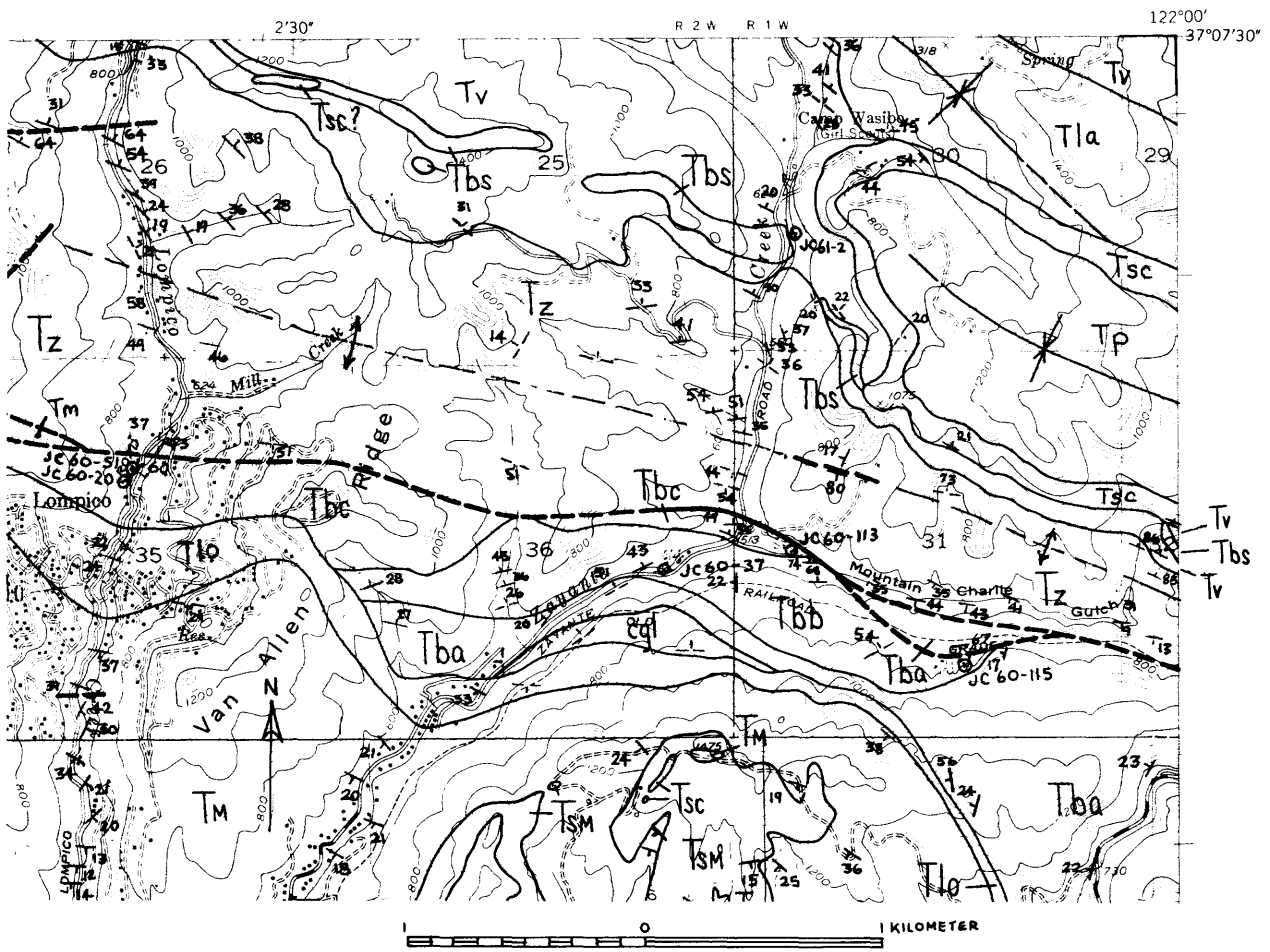
Figure 22.--Stratigraphic column, upper Zayante Creek area, south limb of San Lorenzo syncline.

Table 3.--Fossils from the Vaqueros Sandstone near Zayante Creek

	<u>USGS Cenozoic locality</u>	
	M5049 (JC7B)	M5050 (JC7C)
Gastropods		
Astraea aff. A. morani Loel & Corey.....	X	X
Bruclarkia santacruzana (Arnold).....		
Calyptraea cf. C. inornata (Gabb).....		X
Conus owenianus Anderson.....		X
Cymatium n. sp.....		X
Oliva sp.....		x
Olivella cf. O. pedroana var. subpedroana Loel & Corey.....		X
Scaphander(?) sp.....	X	
Sealesia sp.....		X
Turritella inezana Conrad.....	X	
Pelecypods		
Anadara(?) sp.....	X	
Chione latilaminosa Anderson & Martin.....	X	X
Chlamys sespeensis (Arnold).....		X
Crassatella granti (Wiedey).....	X	X
Dosinia margaritana Wiedey.....	X	X
Dosinia margaritana var. projecta Loel & Corey.....	X	X
Glycymeris sp.....	X	X
Mytilus cf. M. expansus Arnold.....		X
Trachycardium vaquerosense (Arnold).....	X	X
Vertipecten cf. V. perrini (Arnold).....	X	
Barnacle		
Balanus sp.....	X	
Echinoids.....	X	

LOMPICO SECTION

The area around Lompico (fig. 23) has the only exposures of Butano Sandstone with foraminifer and nannoplankton assemblages of early and middle Eocene age. The formation is truncated to the north by the Zayante fault and is overlapped to the south by the Lompico Sandstone of middle Miocene age, so that no creek or road has a complete section. The oldest beds, mapped informally by Clark (1966) as lower sandstone member (see fig. 24), crop out from Glenwood Highway, 4 km southeast of Lompico, to the vicinity of Zayante Creek. The middle siltstone member is generally poorly exposed and extensively landslid. An obscure locality (JC60-115) near an abandoned railroad grade has yielded a fairly rich foraminifer fauna (table 4). The youngest faunas that were recovered in stratigraphic sequence occur in the upper sandstone member near Lompico--one of them (JC60-51, fig. 24) contain nannoplankton that were correlated by F. R. Sullivan (written commun., 1964) with the Ulatisian Stage. Still younger faunas were collected from an unknown stratigraphic position within the Butano Sandstone south of the Zayante fault in the vicinity of the town of Boulder Creek. One of these (EB 444) has foraminifers of Narizian age, according to Brabb (1960, p. 35). Thus, the Zayante fault truncates progressively older beds from west to east.



EXPLANATION		
Pliocene	Tp	Purisima Formation
	Tsc	Santa Cruz Mudstone
	Tsm	Santa Margarita Sandstone
Miocene	Tm	Monterey Formation
	Tlo	Lompico Sandstone
Oligocene	Tz	Zayante Sandstone
	Tvc	Vaqueros Sandstone
	Tbs	Basalt
	Tbc	Butano Sandstone
	Tbb	Butano Sandstone, U. ss member
Eocene	Tba	Butano Sandstone, L. ss member
	Tla	Lambert Shale

Figure 23.--Geologic map of the Lompico area showing localities where fossils were collected. Base map from Felton 7.5' quadrangle. Geology by J. C. Clark.

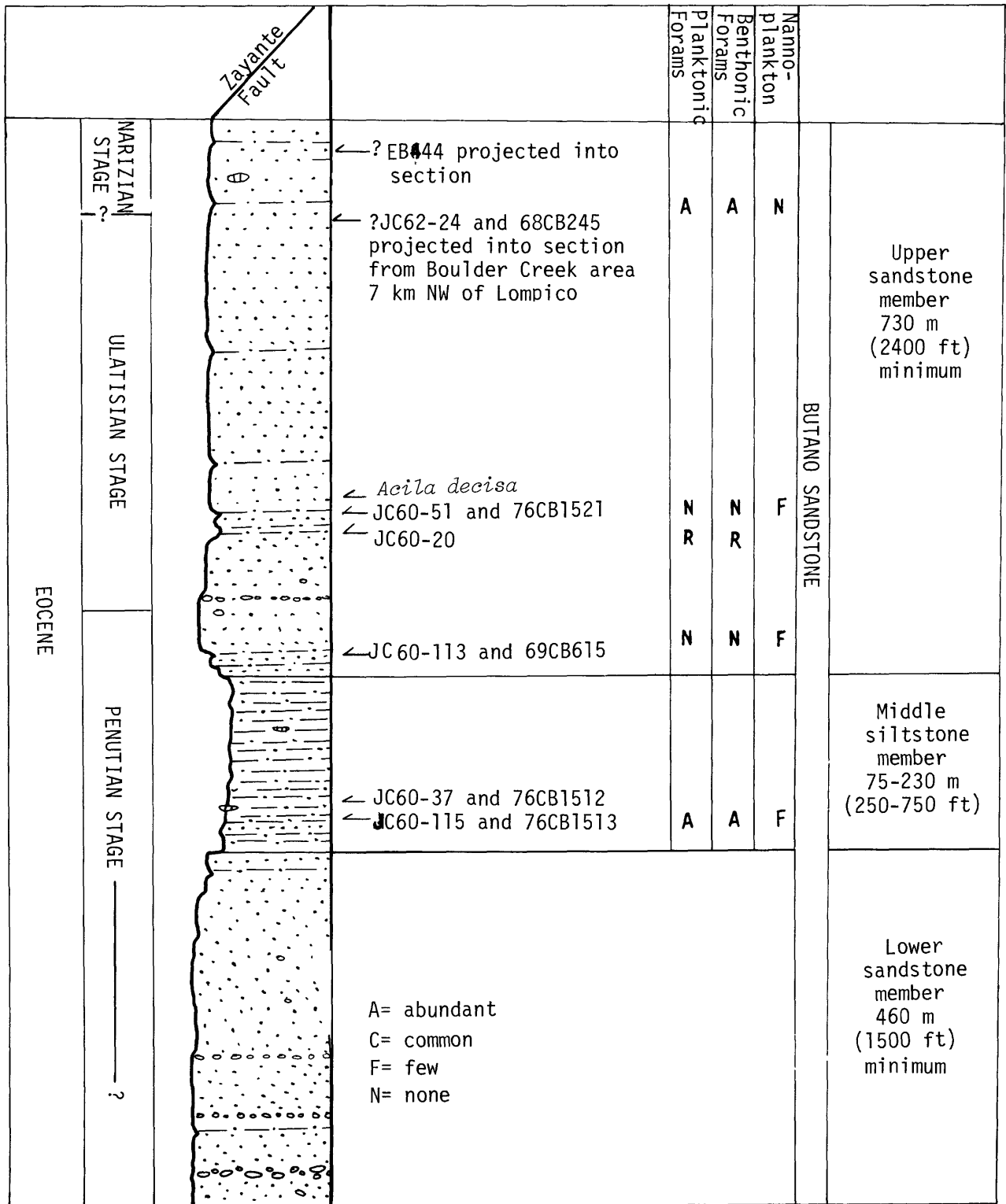


Figure 24.--Composite stratigraphic column of the Butano Sandstone in the vicinity of Lompico, California. Section estimated from geologic map by Clark (1966).

Table 4.--Foraminifers from the middle siltstone member of the Butano Sandstone near Lompico.

LOCALITY	JC60-115
<i>Allomorpha trigona</i> Reuss	R
<i>Ammodiscus glabratus</i> Cushman & Jarvis	VR
<i>Amphimorphina</i> cf. <i>A. ignota</i> Cushman & Siegfus	VR
<i>Anomalina garzaensis</i> Cushman & Siegfus	VR
<i>Anomalina</i> aff. <i>A. regina</i> Martin	R
<i>Asterigerina crassaformis</i> Cushman & Siegfus	C
<i>Bathysiphon eocenicus</i> Cushman & G. D. Hanna	A
<i>Bathysiphon</i> spp.	R
<i>Bifarina nuttalli</i> Cushman & Siegfus (?)	VR
<i>Bolivina crenulata</i> Cushman	VR
<i>Bolivina</i> (?) sp.	VR
<i>Bulimina</i> aff. <i>B. alazanensis</i> Cushman	R
<i>Bulimina macilenta</i> Cushman & Parker	R
<i>Bulimina ovata</i> d'Orbigny	VR
<i>Bulimina pupoides</i> d'Orbigny	VR
<i>Bulimina</i> aff. <i>B. pyrula</i> d'Orbigny, Mallory	A
<i>Bulimina semicostata</i> Nuttall	F
<i>Bulimina</i> cf. <i>B. tarda</i> Parker & Bermudez	VR
<i>Bulimina</i> aff. <i>B. whitei</i> Martin	VR
<i>Bulimina</i> spp.	VR
<i>Cassidulina</i> cf. <i>C. globosa</i> Hantken	VR
<i>Chilostamella cylindroides</i> Reuss	R
<i>Cibicides martinicensis</i> Cushman & Barksdale	F
<i>Dentalina</i> cf. <i>D. coleii</i> Cushman & Dusenbury	VR
<i>Dentalina consobrina</i> d'Orbigny	C
<i>Dentalina</i> cf. <i>D. jacksonesis</i> (Cushman & Applin)	R
<i>Dentalina</i> sp.	VR
<i>Dentalina</i> (?) sp.	VR
<i>Discorbis</i> sp.	VR
<i>Dorothyia</i> sp.	R
<i>Eponides</i> aff. <i>E. umbonata</i> (Reuss)	VR
<i>Globigerina nitida</i> Martin	A
<i>Globigerina triloculinoides</i> Plummer	A
<i>Globigerina yeguaensis</i> Weinzierl & Applin	R
<i>Globigerina</i> spp.	VR
<i>Globorotalia acarinata</i> (Subbotina)	A
<i>Globorotalia imitata</i> Subbotina	R
<i>Globorotalia pseudomenardii</i> Bolli	F
<i>Globorotalia pseudotopilensis</i> (Subbotina)	A
<i>Globorotalia reissi</i> Loeblich & Tappan	VR
<i>Globorotalia rex</i> Martin	F
<i>Gyroldina orbicularis</i> var. <i>obliquata</i> Cushman & McMasters	VR
<i>Gyroldina</i> (?) sp.	VR
<i>Heterohelix</i> (?) sp.	VR
<i>Hoeglundina eocenica</i> (Cushman & M. A. Hanna)	VR
<i>Lagena acuticosta</i> Reuss	VR
<i>Lenticulina</i> (?) sp.	VR
" <i>Nadogenerina</i> " cf. " <i>N. adolphina</i> " (d'Orbigny)	C
" <i>Nadogenerina</i> " <i>kressenbergensis</i> (Gumbel)	R
" <i>Nadogenerina</i> " spp.	VR
<i>Nodosaria ewaldi</i> Reuss, Mallory	VR
<i>Nodosaria latejugata</i> Gumbel	R
<i>Nodosaria</i> sp.	VR
<i>Pleurostomella</i> cf. <i>P. acuta</i> Hantken	R
<i>Pleurostomella alternans</i> Schwager	VR
" <i>Pseudoglandulina</i> " (?) sp.	VR
" <i>Pseudohastigerina</i> " <i>micra</i> (Cole)	VR
<i>Robulus inornatus</i> d'Orbigny	A
<i>Robulus pseudovortex</i> Cole	R
<i>Robulus</i> spp.	F
<i>Siphonina wilcoxensis</i> Cushman	VR
<i>Virgulina</i> cf. <i>V. danvillensis</i> Howe & Wallace	VR
<i>Virgulina yazooensis</i> Cushman & Todd	VR
<i>Virgulina</i> sp.	VR
diatoms	A
ostracods	VR
nannoplankton	F

VR=very rare; R=rare; F=few; C=common; A=abundant.

Benthonic identifications by J. C. Clark.

Planktonic identifications by J. C. Clark and Yokichi Takayanagi.

From Clark (1960, table 3).

SMITH GRADE SECTION

Mollusks from the Smith Grade area were important in documenting the presence of rocks of Paleocene age in the Ben Lomond Mountain area (Brabb, 1959 and Cummings and others, 1964, p.184). The rocks containing the fossils were mapped subsequently by Clark (1966), who also collected foraminifers from the Smith Grade section. Figure 25 is a sketch of the geology in the area, figure 26 is a stratigraphic column for the rocks of Paleocene age, and tables 5 and 6 list foraminifers and larger invertebrates from these rocks.

The basal sandstone beds commonly yield a shallow-water molluscan fauna associated with diverse calcareous benthonic foraminifers, whereas the overlying siltstone beds yield abundant large arenaceous foraminifers. At two localities (JC60-126=M4668 and JC60-120=M4669), mollusks characteristic of the "Martinez Stage" (Paleocene) of the California molluscan chronology are associated with benthonic foraminifers of the Ynezian Stage of Mallory (1959).

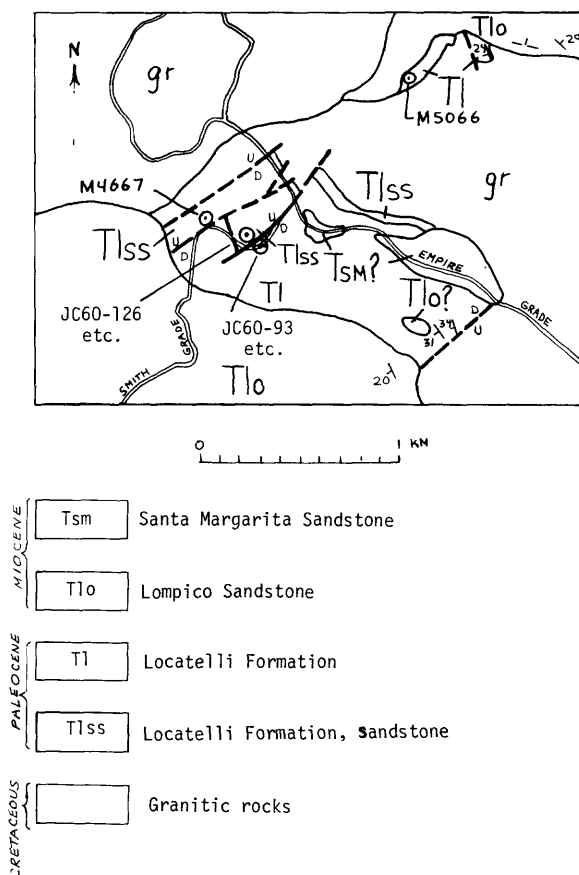


Figure 25.--Geologic map of the Smith Grade-Empire Grade area showing localities where fossils were collected. Base map from Felton 7.5' quadrangle. Geology from Clark (1966).

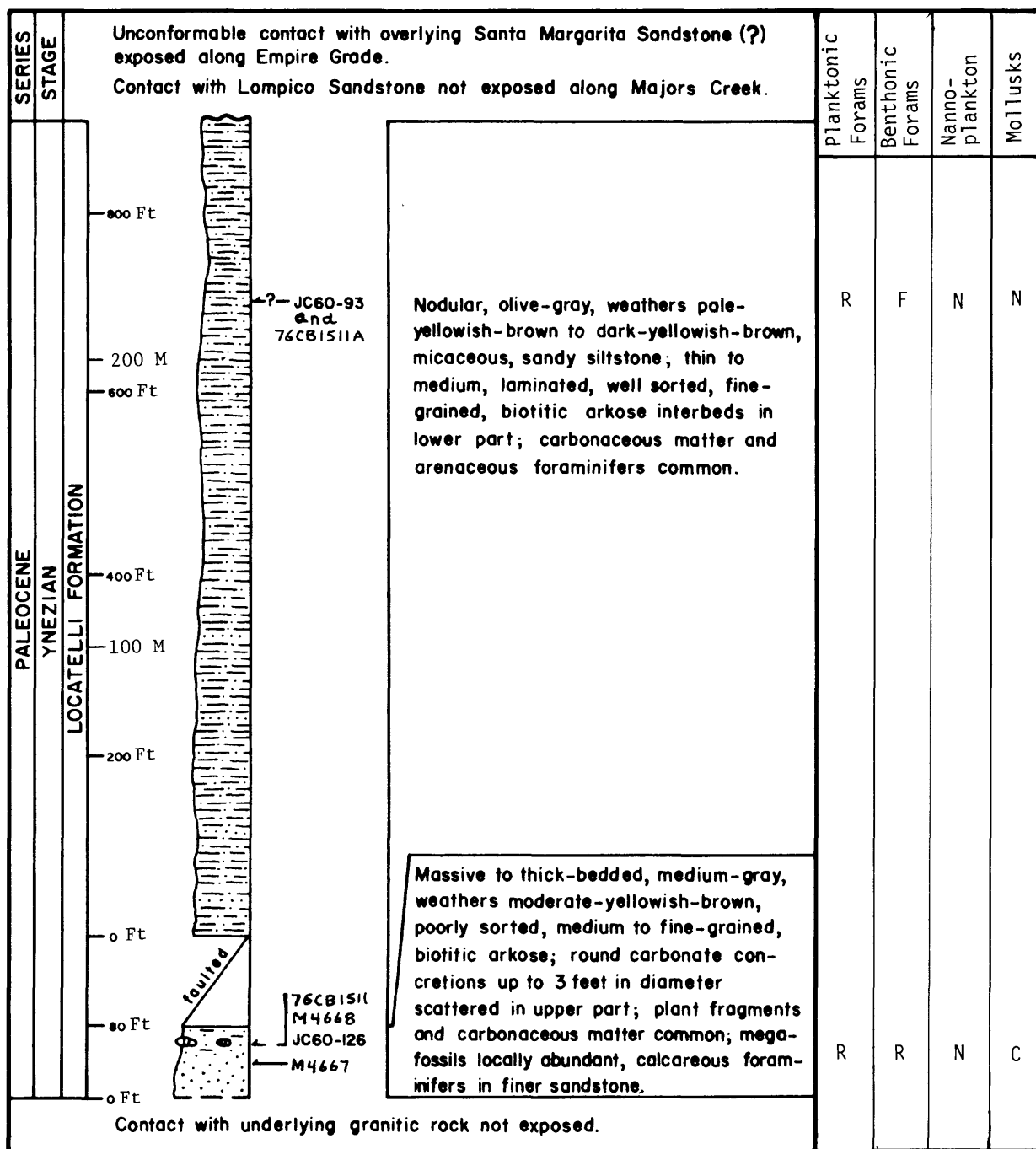


Figure 26.--Stratigraphic column, Smith Grade section. From Clark (1966, fig. 3).

Table 5.--Foraminifers from the Locatelli Formation.
From Clark (1960, table 2).

LOCALITY	JCG6-126 (JC14)	JCG6-120 (JC23)	JCG6-93	Vine Hill Ss.
<i>Ammobaculites</i> sp.			VR	
<i>Ammodiscoides turbinatus</i> Cushman			VR	
<i>Ammodiscus glabratus</i> Cushman & Jarvis			VR	
<i>Ammodiscus</i> sp.			VR	
<i>Anomalina dorri</i> var <i>aragonensis</i> Nuttall	VR	C		X
<i>Anomalina regina</i> Martin	VR			
<i>Bathysiphon eocenicus</i> Cushman & G D Hanna			VR	X
<i>Bathysiphon</i> spp.			F	
<i>Bathysiphon</i> sp.		R		
<i>Bulimina</i> cf <i>B. pachecoensis</i> Smith		VR		X
<i>Cassidulina globosa</i> Hanke		VR		
<i>Cibicides martinezensis</i> Cushman & Barksdale	C	C		
<i>Cibicides whitei</i> Martin	F			
<i>Cibicides</i> spp.		F		
<i>Cibicides</i> sp.	F			
<i>Dentalina alternata</i> (Jones)	VR			X
<i>Dentalina</i> cf <i>D. approximata</i> Reuss	R			X
<i>Dentalina</i> cf <i>D. communis</i> (d'Orbigny)	VR			X
<i>Dentalina</i> cf <i>D. consobrina</i> (d'Orbigny)	F			X
<i>Dentalina</i> sp.	VR			
<i>Dentalina</i> (?) spp.	VR			
<i>Dorothia</i> sp.		R	VR	
<i>Dorothia</i> (?) sp.			VR	
<i>Eggerella</i> (?) sp.			VR	
<i>Entosolenia hexagona</i> Williamson		VR		?
<i>Eponides mexicana</i> (Cushman)	F			
<i>Eponides</i> cf <i>E. mexicana</i> (Cushman)		VR		
<i>Gaudryina</i> sp.	VR	C		
<i>Globigerina aequiensis</i> Loeblich & Tappan		R		
<i>Globigerina triloculinoides</i> Plummer	R	VR	VR	X
<i>Globigerina</i> sp.	VR		VR	
<i>Globocanusa</i> (?) sp.		VR		
<i>Globorotalia convexa</i> Subbotina			VR	
<i>Glomospira charoides</i> (Jones & Parker)			VR	
<i>Glomospira charoides</i> var <i>corona</i> Cushman & Jarvis			VR	
<i>Guttulina</i> sp.	VR	R		
<i>Gyrodina soldani</i> var <i>octocamerata</i> Cushman & G D Hanna	F	C		X
<i>Haplophragmoides caronata</i> (Brady)			F	
<i>Haplophragmoides excavata</i> Cushman & Waters			R	
<i>Haplophragmoides</i> cf <i>H. excavata</i> Cushman & Waters		C		
<i>Haplophragmoides kirki</i> Wickenden			VR	
<i>Haplophragmoides</i> cf <i>H. longifissus</i> Israelsky		C		
<i>Haplophragmoides protullisatus</i> Israelsky			C	
<i>Haplophragmoides</i> sp.		VR		
<i>Hyperammina</i> cf <i>H. elongata</i> (Brady)			C	
<i>Hyperammina</i> (?) sp.			VR	
<i>Lagena acuticasta</i> Reuss		VR		
<i>Lenticulina</i> sp.	F			
<i>Lituituba</i> cf <i>L. lituiformis</i> (Brady)			VR	
" <i>Nodogenerina</i> " <i>adolphina</i> (d'Orbigny)	C	VR		
" <i>Nodogenerina</i> " (?) sp.		C		
<i>Nodosaria arundinea</i> Schwoeger		VR		X
<i>Nodosaria deliciae</i> Martin	VR	F		
<i>Nodosaria latejugata</i> Gumbel	R			X
<i>Nodosaria</i> cf <i>N. latejugata</i> Gumbel		VR		
<i>Nodosaria</i> spp.	VR	VR		
<i>Nodosaria</i> sp.			VR	
<i>Palosina complanata</i> Franke			R	
<i>Planularia</i> sp.	VR			
" <i>Pseudoglobulina</i> " <i>canica</i> (Neugeboren)	F			X
<i>Pseudoglobulina nahealensis</i> Cushman & Todd		F		
<i>Quinqueloculina</i> sp.	VR	VR		
<i>Reophax</i> sp.		R		
<i>Rabulus inornatus</i> d'Orbigny	C	VR		X
<i>Rabulus</i> aff. <i>R. midwayensis</i> (Plummer)	VR	VR		
<i>Rabulus</i> spp.	R			
<i>Rzehakina epigona</i> var <i>lata</i> Cushman & Jarvis			F	
<i>Silicosigammina californica</i> Cushman & Church		R	C	X
<i>Spiraplectammina perplexa</i> Israelsky			C	
<i>Spiraplectammina richardi</i> Martin	C			
<i>Spiraplectammina</i> cf <i>S. tejonensis</i> Mallory		VR		
<i>Spiraplectammina</i> (?) sp.		VR	VR	
<i>Trifarina</i> sp.		R		
<i>Trachammina</i> cf <i>T. globigeriniformis</i> (Parker & Jones)			R	
<i>Trachammina</i> spp.			F	
<i>Trachammina</i> sp.		C		
<i>Trachamminoides</i> (?) sp.			VR	
<i>Vaginulina suturalis</i> Cushman	VR			X
<i>Vaginulinopsis saundersi</i> (Hanna & Hanna)	A	R		X
<i>Valvulinera martinezensis</i> Smith	R	C		
<i>Verneulina</i> sp.	VR	VR		
ostracods	F	VR		
radiolaria			VR	

VR=very rare; R=rare; F=few; C=common; A=abundant.
X=recorded from Vine Hill Sandstone by Smith (1957, Fig. 4).
Benthonic identifications by J C Clark.
Planktonic identifications by J C Clark and Yokichi Takayanagi.

Table 6.--Paleocene larger invertebrates from the Locatelli Formation.

		<u>USGS Cenozoic locality</u>				
		M4667	M4668	M4669	M5065	M5066
Gastropods						
	<i>Perissolax tricarnatus</i> WEaver.....	X	X			
	<i>Surculites</i> (?) sp.....			X		
	<i>Turritella</i> cf. <i>T. infragranulata</i> Gabb.....	X	X			
	<i>Turritella pachecoensis</i> Stanton.....	X	X		X	
	<i>Turritella</i> cf. <i>T. pachecoensis</i> Stanton.....					X
Pelecypods						
	<i>Acila</i> sp.....	X		X		X
	<i>Cucullaea mathewsonii</i> * Gabb.....			X		X
	<i>Cucullaea mathewsonii</i> Gabb(?).....	X				
	<i>Glycymeris</i> sp.....	X				
	<i>Lucinoma</i> sp.....	X	X			
	<i>Nuculana</i> sp.....	X	X			X
	<i>Pholadomya nasuta</i> Gabb.....	X				
	<i>Pinna</i> n. sp.....	X				
Scaphopod						
	<i>Dentalium</i> sp.....	X	X			
Echinoids						
	<i>Cidaris martinezensis</i> Kew.....	X				
	<i>Periaster</i> n. sp.....	X			X	
	<i>Pericosmus</i> sp.....	X				
	Echinoid spines.....				X	
Arthropod						
	<i>Raninoides</i> sp.....			X		
Locality	Description of locality, collector, and date collected					
M4667	Cut on northeast side of abandoned section of Smith Grade, 10m (30 ft) northwest of new road and 30m (100 ft) east of fork of Majors Creek, S-1/2, sec. 29, T. 10S., R.2W., Felton quadrangle. Lower part of Locatelli Formation, Paleocene. Same locality as LSJU 3401A. Collected by J. C. Clark, 1960.					
M4668	North side of deep cut along Smith Grade, approximately 300 m (1000 ft) southwest of where Smith Grade joins Empire Grade, Felton quadrangle. Lower part of Locatelli Formation, Paleocene. Collected by J. C. Clark, 1960.					
M4669	Bed in unnamed, westward-flowing tributary to San Lorenzo River, about 530m (1730 ft) east of where tributary joins river, in Henry Cowell Redwoods State Park, Felton quadrangle. Lower part of Locatelli Formation, Paleocene. Collected by J. C. Clark, 1960.					
M5065	North bank of Gold Gulch, about 90m (300 ft) upstream from where stream passes under Highway 9, Felton quadrangle. Lower part of Locatelli Formation, Paleocene. Collected by J. C. Clark, 1962.					
M5066	Bed in Gold Gulch, approximate elevation 160m (520 ft), Felton quadrangle. Lower part of Locatelli Formation, Paleocene. Collected by J. C. Clark, 1962.					

MOUNTAIN CHARLIE GULCH SECTION

The Mountain Charlie Gulch area has a fairly complete section from Butano Sandstone to Lambert Shale, but most of the units are not well exposed. The reason for including the section in this report is that it yields faunas of both early and late Saucesian age that by the correlation of Tipton and others (1973) straddle the Oligocene-Miocene boundary. A sketch map of the geology and a stratigraphic column are shown on figure 27. Foraminifers from the section are listed on table 7.

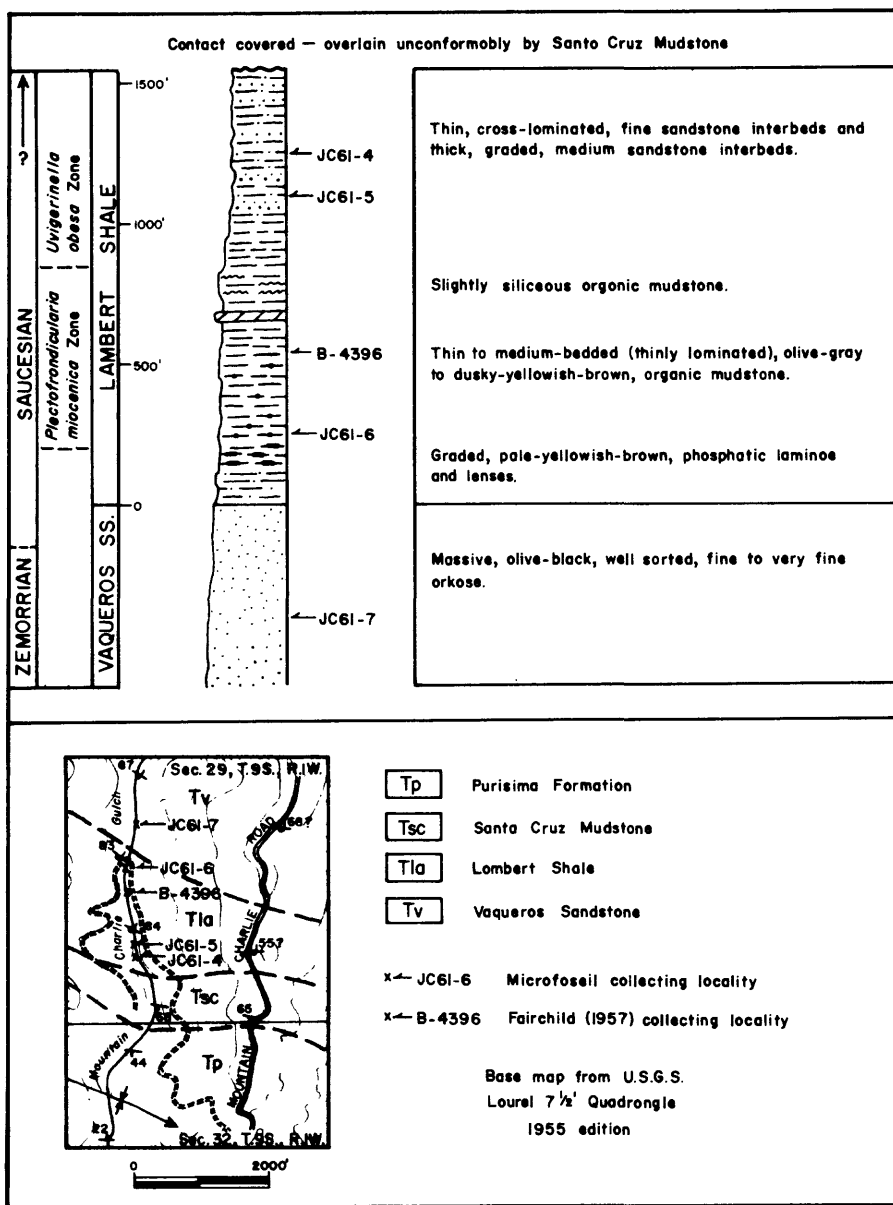


Figure 27.--Geologic sketch map and stratigraphic column for the Mountain Charlie Gulch area. From Clark (1966, fig. 6).

Table 7.--Foraminifers from the Lambert Shale in the Mountain Charlie Gulch area. From Clark (1960, table 5).

FORMATION	VAQUEROS SS.		LAMBERT SH.	
	JC61-7	JC61-6	B-4396	JC61-5
LOCALITY (shown on Figure 6)				
<i>Baggina robusta</i> Kleinpell		F		
<i>Bathysiphon eocenicus</i> Cushman & GD Hanna			F	
<i>Bolivina marginata</i> Cushman		C	R	C
<i>Bolivina</i> cf. <i>B. perrini</i> Kleinpell				VR
<i>Bolivina</i> sp.				VR
<i>Bulimina rinconensis</i> Cushman & Laiming			R	
<i>Buliminella</i> cf. <i>B. brevior</i> Cushman		R		
<i>Buliminella curta</i> Cushman				R
<i>Buliminella subfusiformis</i> Cushman		VR		
<i>Cibicides americanus</i> (Cushman)		VR	C	
<i>Cibicides americanus</i> var. <i>crassiseptus</i> Cushman & Laiming	VR			
<i>Cibicides</i> cf. <i>C. americanus</i> var. <i>crassiseptus</i> Cushman & Laiming	VR			
<i>Cibicides floridanus</i> (Cushman)	F			
<i>Cibicides</i> aff. <i>C. floridanus</i> (Cushman)			F	
<i>Cibicides</i> cf. <i>C. floridanus</i> (Cushman)				VR
<i>Cyclammina cancellata</i> Brady	C			
<i>Dentalina</i> cf. <i>D. pauperata</i> d'Orbigny				VR
<i>Dentalina quadrulata</i> Cushman & Laiming			F	R
<i>Dentalina spinosa</i> d'Orbigny			R	
<i>Dentalina</i> sp.		VR		
<i>Eponides</i> cf. <i>E. keenani</i> Cushman & Kleinpell				VR
<i>Globigerina</i> sp.				VR
<i>Globigerina</i> spp.		F		
<i>Gyroldina condani</i> (Cushman & Schenck)			R	
<i>Gyroldina saldani</i> d'Orbigny			R	VR
<i>Gyroldina</i> sp.		F		
<i>Haplophragmoides</i> spp.				R
" <i>Nodogenerina</i> " <i>advena</i> Cushman & Laiming	F	?		F
<i>Nonion castiferum</i> (Cushman)				R
<i>Nonion</i> cf. <i>N. pompilioides</i> (Fichtel & Moll)	F			
<i>Plectofrondicularia</i> aff. <i>P. basi-spinata</i> Cushman			?	
<i>Plectofrondicularia miocenica</i> var. <i>directa</i> Cushman & Laiming				R
<i>Prateonina</i> (?) sp.			R	
<i>Pullenia salisburyi</i> R.E. & K.C. Stewart			R	
<i>Pullenia</i> cf. <i>P. salisburyi</i> R.E. & K.C. Stewart				VR
" <i>Pulvinulinella</i> " sp.				VR
<i>Robulus inornatus</i> (d'Orbigny)			R	
<i>Robulus simplex</i> (d'Orbigny)	C		R	
<i>Robulus</i> sp.				R
<i>Siphogenerina kleinpelli</i> Cushman		VR	R	VR
<i>Siphogenerina nodifera</i> Cushman & Kleinpell	VA			
<i>Siphogenerina transversa</i> Cushman		A	R	
<i>Spiraplectammina</i> (?) sp.				F
<i>Tritaxilina coleii</i> Cushman & Siegfus			R	
<i>Uvigerina auberiana</i> d'Orbigny			A	
<i>Uvigerina</i> cf. <i>U. beccarii</i> Fornasini			F	
<i>Uvigerina rustica</i> Cushman & Edwards			R	
<i>Uvigerinella obesa</i> Cushman				A
<i>Uvigerinella obesa</i> var. <i>impolita</i> Cushman & Laiming				VR

VR=very rare; R=rare; F=few; C=common; A=abundant; ?=questionable identification.
 JC61-7 identifications by micropaleontologists of Humble Oil & Refining Company.
 JC61-6, JC61-5 identifications by J.C. Clark.
 B-4396 identifications by Fairchild (1957, Fig. 6).

ANO NUEVO SECTION

Stratigraphic contrasts across the active San Gregorio fault in the vicinity of Año Nuevo have been reported in an abstract by Clark and Brabb (1977). A summary of the stratigraphic sequences is shown on figure 28. The geology of the area and fossil localities are shown on figure 29.

Even within the area west of the San Gregorio fault, the geologic section is complicated by folds and faults. Figure 30 is a stratigraphic column for the area, and figure 31 shows the section in the vicinity of Año Nuevo Point. Some of the section is exposed only at very low stages of spring tides. Some of the section exposed in September, 1976 was covered by sand in August, 1977. Table 8 lists foraminifers from the Año Nuevo area; table 9 provides a description of most of the localities.

The Vaqueros(?) and Monterey Formations near Año Nuevo Point yield calcareous benthonic foraminifers diagnostic of the Zemorrian, Saucesian, and Relizian Stages of Kleinpell (1938) and are characteristic of bathyal depths. Large, well-preserved, benthonic foraminifers are especially abundant in the siltstone beds of the Vaqueros(?) 1.6 km (1 mi) north of the point, where the common occurrence in the lower part of the section of Siphogenerina nodifera is diagnostic of Zemorrian age. The upper part of this synclinal section ranges into the early Saucesian.

The Vaqueros(?) section that is exposed near the axis of the anticline that is east of Año Nuevo Point is mostly younger, ranging from early Saucesian to Relizian. Sheared mudstone on the east flank of the anticline (USGS, loc. Mf2192 = JC68-3) locally yields a Relizian fauna that includes Cassidulina laevigata carinata, Cibicides of C. floridanus, "Nonion" pompilioides var., Pullenia bulloides, P. moorei, and Spheroidina variabilis. These species are diagnostic of the Pseudosaucesian faunal facies of Beck (1952). Although these deep-water species are generally considered typical of the Saucesian Stage, they are now known to recur in any stratigraphic position from the Upper Saucesian through the Lower Mohnian (upper Miocene). Planktonic foraminifers are sparse in the Vaqueros(?) beds north and east of Año Nuevo Point.

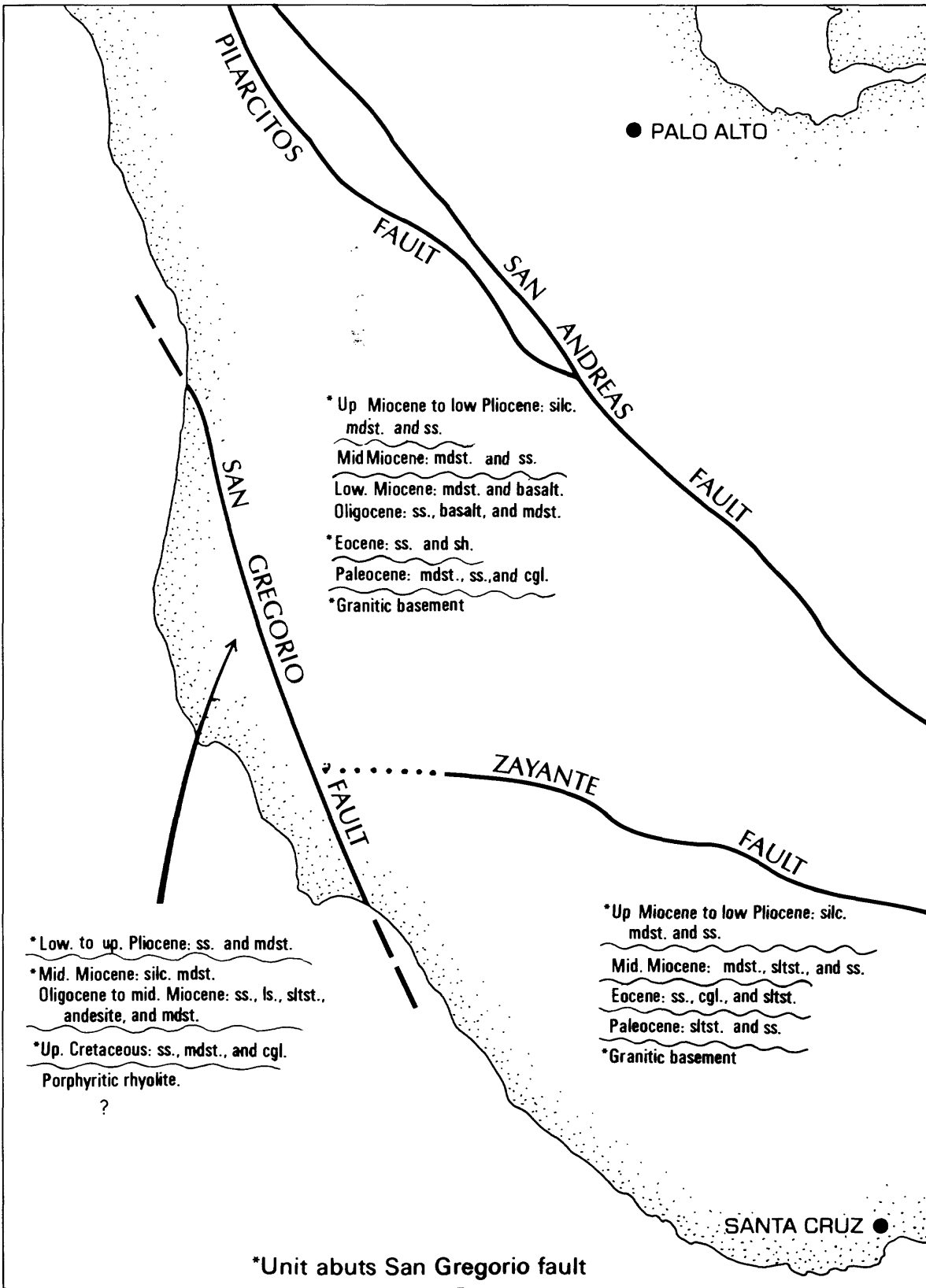
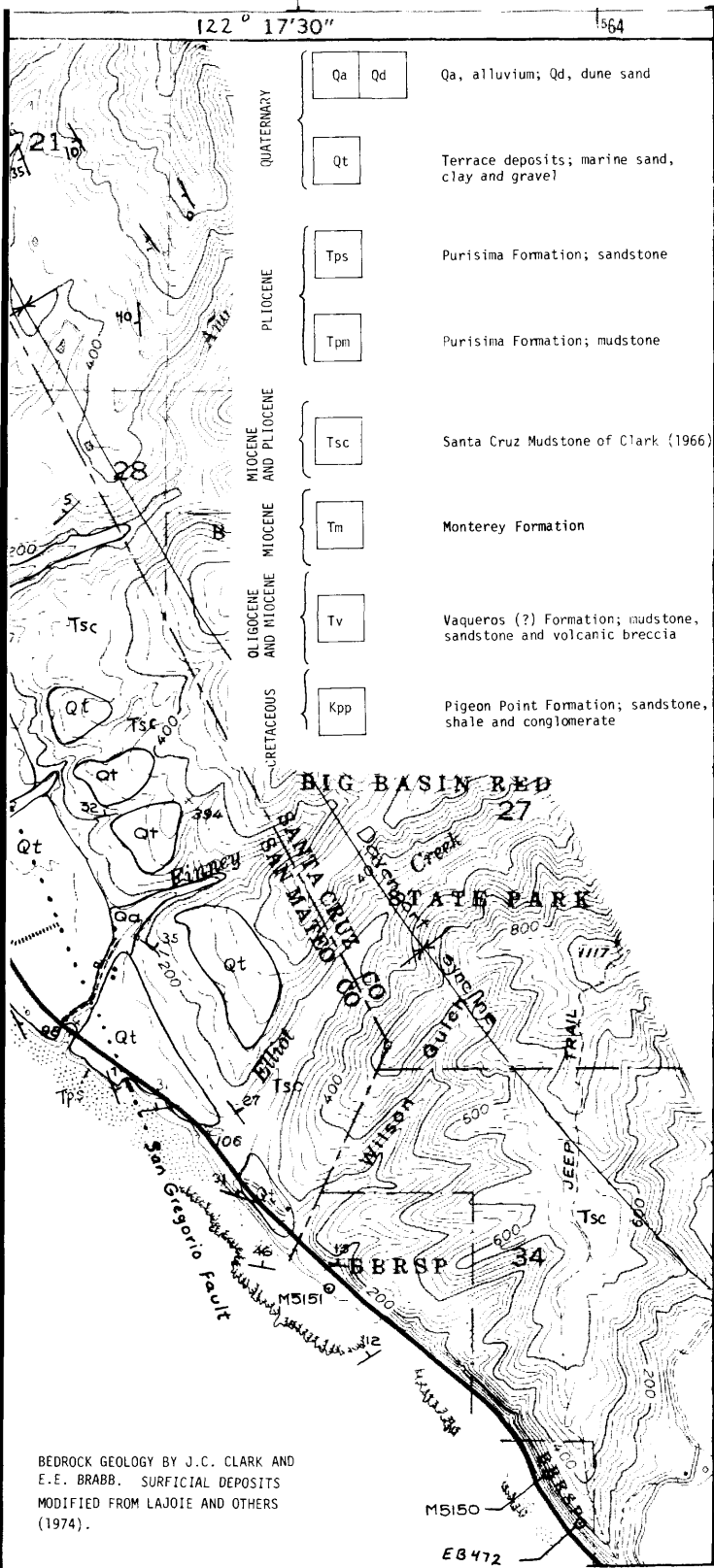


Figure 28.--Stratigraphic contrasts across the San Gregorio and Zayante faults



Benthonic foraminifers are not as abundant or well preserved in the overlying siliceous beds of the Monterey Formation, where they are diagnostic of Relizian age and characteristic of middle bathyal depths. Planktonic foraminifers are common in one sample from Año Nuevo Island (USGS loc. Mf4531 = JC67-5) but are difficult to process from this siliceous mudstone and when recovered are poorly preserved.

The Monterey beds in the Año Nuevo area are lithologically similar to the rhythmically bedded siliceous strata of the type Monterey Formation, 65 km (40 mi) to the south, but are older than those of the type section, which are mostly of Mohnian (upper Miocene) age.

(Cross section on next page.)

of the Año Nuevo area.

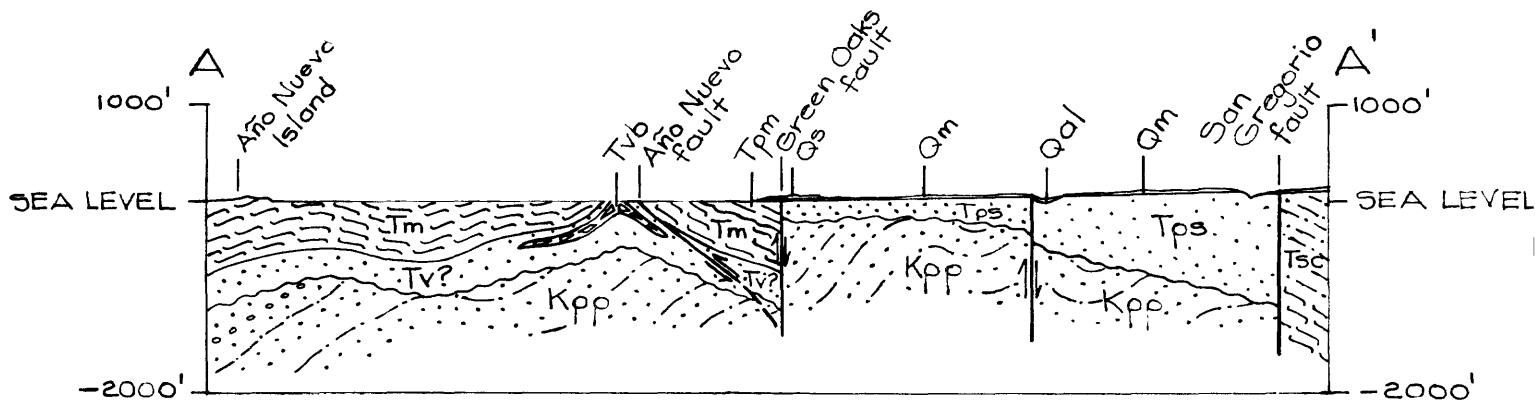


Figure 29--continued

SYSTEM	SERIES	STAGE	FORMATION	LITHOLOGY	THICKNESS Meters (feet)	DESCRIPTION
TERTIARY	PLIOCENE		Purisima Formation		360+ (1200+)	Thick to very thick bedded olive-gray, fine-grained lithic sandstone with thick to very thick interbeds of rhyolitic tuff and a few carbonate concretions
			Unconformity		76 (250)	Medium to thick bedded light-olive-gray, nodular mudstone
			Monterey Formation		215+ (700+)	Thin bedded and thinly laminated olive-gray to dusky-yellowish-brown, siliceous mudstone
TERTIARY	OLIGOCENE TO MIOCENE	Zemarran, Relizian	Vaqueras(?) Formation		230+ (770+)	Thick to very thick bedded arkasic sandstone and sandy limestone, olive-gray bioturbated siltstone, and dusky-yellowish-brown phosphatic mudstone. Upper part contains altered andesitic breccia
		Saucesan, Relizian	Unconformity			
CRETACEOUS	UPPER CRETACEOUS	Santonian(?), Campanian, and Maastrichtian	Pigeon Point Formation		2600+ (8500+)	Interbedded sequence of brownish-gray fine- to coarse-grained sandstone and dark-gray sandy siltstone in lower part, and sandy pebble to cobble conglomerate, pebbly mudstone, coarse-grained sandstone, and laminated and cross-bedded fine-grained sandstone in upper part
			Contact not exposed			
CRETACEOUS OR OLDER						Porphyrritic rhyolite, greenish, with phenocrysts of quartz, plagioclase, and orthoclase(?) in a silicic matrix

Figure 30.--Stratigraphic column for the area west of the San Gregorio fault

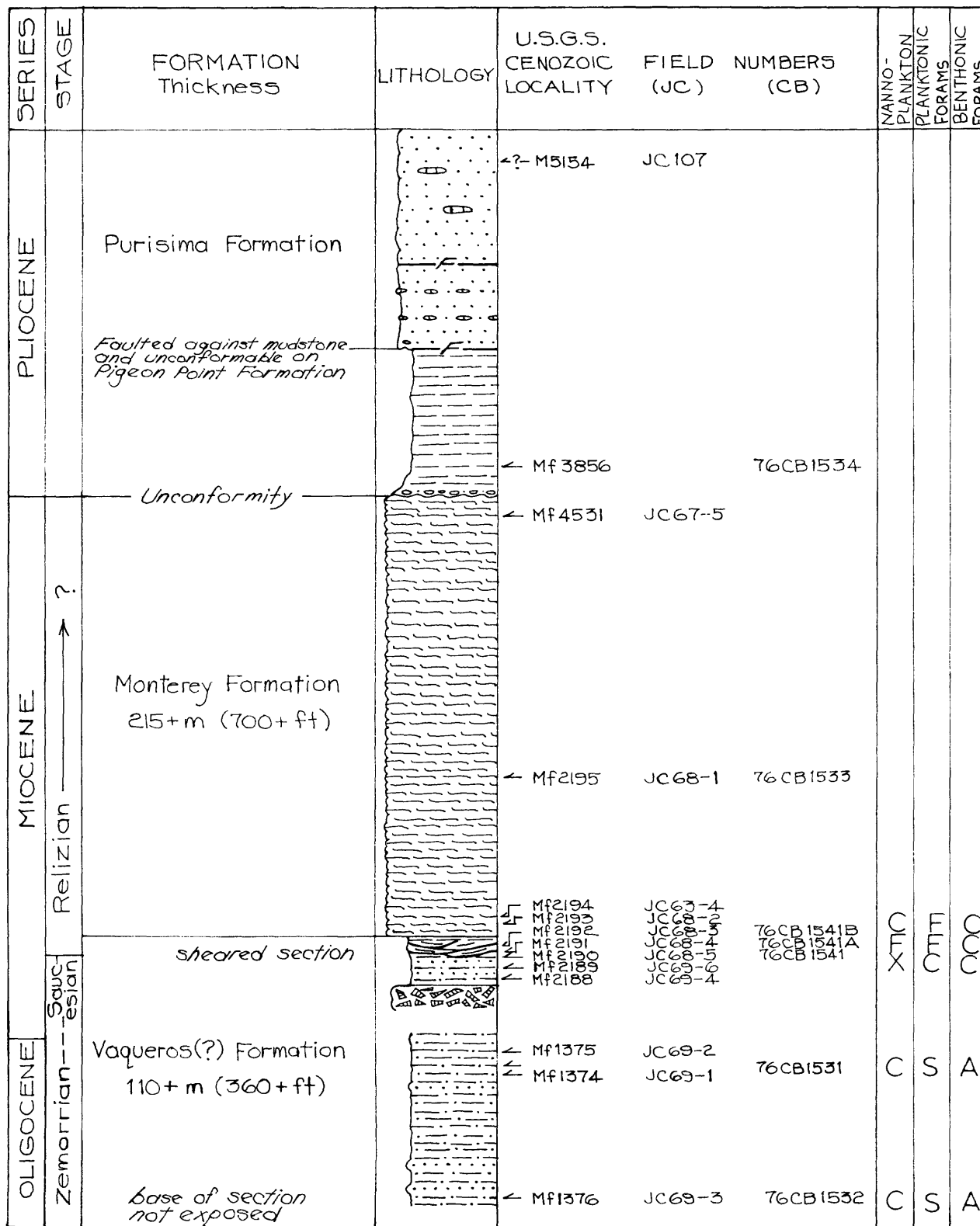


Figure 31.--Stratigraphic column for the Año Nuevo area.

Table 8.--Foraminifers from the Año Nuevo area

[Locality MF1374-MF1376 identifications by R. L. Pierce; other identifications by J. C. Clark and R. M. Kleinpell;

A=Abundant, C=Common, F=Few, R=Rare, VR=Very Rare]

	USGS Cenozoic locality and field number					
	MF1374 JOC69-1	MF1375 JOC69-2	MF1376 JOC69-3	MF2188 JOC69-4	MF2189 JOC69-6	MF2190 JOC69-5
Anomalina californiensis Cushman & Hobson			R			
Anomalina aff. A. salinasensis Kleinpell						
Baggina robusta Kleinpell						
Baggina(?) sp.						
Bolivina advena Cushman						
Bolivina advena var. striatella Cushman						
Bolivina cf. B. advena var. striatella Cushman						
Bolivina breviar Cushman						
Bolivina californica Cushman						
Bolivina cf. B. floridana Cushman						
Bolivina aff. B. hughesi Cushman						
Bolivina marginata Cushman						
Bolivina cf. B. marginata Cushman						
Bolivina aff. B. seminoda Cushman						
Bolivina aff. B. vaughani Nuttall						
Bulimina alligata Cushman & Laming						
Bulimina carnerosensis Cushman & Kleinpell						
Bulimina inflata Seguenza						
Bulimina aff. B. inflata Seguenza						
Bulimina montereyana Kleinpell						
Bulimina ovata d'Orbigny						
Bulimina rinconensis Cushman & Laming(?)						
"Bulimina(?) sp." Kleinpell						
Bulimina sp.						
Buliminella curta Cushman						
Buliminella subfusiformis Cushman						
Buliminella cf. B. subfusiformis Cushman						
Buliminella sp.						
Cassidulina crassa d'Orbigny of Kleinpell						
Cassidulina laevigata var. carinata Cushman						
Cassidulina pulchella d'Orbigny(?)						
Cassidulina subglobosa H. B. Brady						
Cassidulina sp.						
Chilostomella sp.						
Cibicides americanus var. crassiseptus Cushman & Laming						
Cibicides dohertryi Galloway & Morrey						
Cibicides floridanus (Cushman)						
Cibicides cf. C. floridanus (Cushman)						
Cibicides aff. C. lidgei Cushman & Schenck						
Cibicides pseudorangeanus var. evolutus Cushman & Hobson						
Dentalina pauperata d'Orbigny						
Dentalina quadrulata Cushman & Laming						
Dentalina sp.						
Dentalina(?) sp.						
Epistominella superuviana (Cushman)						
Epistominella cf. E. superuviana (Cushman)						
Lpistominella sp.						
Epistominella(?) sp.						
Eponides keenani Cushman & Kleinpell						
Eponides namus (Reuss)						
Eponides rostriformis Cushman & Kleinpell						
Eponides umbonatus (Reuss)						
Fioriulus incisus (Cushman)						
Fronducularia sp.						
Gaudryina(?) sp.						
Globigerina sp.						
Globigerina spp.						
Gyrodina soldanii d'Orbigny						
Gyrodina cf. G. soldanii d'Orbigny						
Lenticulina hughesi (Kleinpell)						
Lenticulina warmani (Barbat & von Estorff)						
Lenticulina sp.						
Lenticulina spp.						
Nodosaria longiscata d'Orbigny						
"Nodosaria" pomplioides (Fichtel & Moll) var.						
"Nodosaria" cf. "N." pomplioides (Fichtel & Moll) var.						
Plectofronducularia advena (Cushman)						
Plectofronducularia californica Cushman & Stewart						
Plectofronducularia miocenica Cushman						
Plectofronducularia vaughani Cushman						
Pullenia bulloides d'Orbigny						
Pullenia moorei Kleinpell						
Pullenia aff. P. quinqueloba (Reuss)						
Pullenia cf. P. bulloides d'Orbigny						
Pullenia salisburyi R. E. & K. C. Stewart						
Pullenia aff. P. salisburyi R. E. & K. C. Stewart						
Siphogenerina branneri (Bagg)						
Siphogenerina cf. S. collina Cushman						
Siphogenerina cf. S. hughesi Cushman						
Siphogenerina kleinpelli Cushman						
Siphogenerina masi Cushman & Parker						
Siphogenerina multicostata Cushman & Jarvis						
Siphogenerina multicostata Cushman & Jarvis or Kleinpell						
Siphogenerina nodifera Cushman & Kleinpell						
Siphogenerina cf. S. reedi Cushman						
Siphogenerina transversa Cushman						
Sphaeroidina bulloides d'Orbigny						
Sphaeroidina variabilis Reuss						
Stilostomella advena (Cushman & Laming)						
Suggrundia sp.						
Uvigerina cf. U. uberiana d'Orbigny						
Uvigerina cf. U. hesarum Fornasini						
Uvigerina cf. U. mexicana Nuttall						
Uvigerina sp.						
Uvigerinella californica Cushman						
Uvigerinella aff. U. californica Cushman						
Uvigerinella californica var. ornata Cushman						
Uvigerinella obesa Cushman						
Uvigerinella obesa var. impolita Cushman & Laming						
Uvigerinella cf. U. obesa var. impolita Cushman & Laming						
Valvulineria aff. V. californica var. obesa Cushman						
Valvulineria cf. V. californica var. obesa Cushman						
Valvulineria depressa Cushman						
Valvulineria cf. V. depressa Cushman						
Virgulina sp.						
Virgulina(?) sp.						

Table 9.--Locality descriptions for the Ano Nuevo area.

Ano Nuevo Mf (JC) localities

- Mf2188: California, San Mateo County, Ano Nuevo 7.5' quad.
From low intertidal exposure about 335m (1100 ft.) E. of Ano Nuevo Point. Unsectioned. Vaqueros(?) Formation, lower Miocene. From 2m (7 ft) stratigraphically above volcanic breccia exposed at low tide. Collected by J. C. Clark, 1969
(field number = JC69-4)
- Mf2189: California, San Mateo County, Ano Nuevo 7.5' quad.
From low intertidal exposure about 330 m (1080 ft) E. of Ano Nuevo Point. Unsectioned. Vaqueros(?) Formation, lower Miocene. From 8m (26 ft) stratigraphically above Mf2188. Collected by J. C. Clark, 1969
(field number = JC69-6)
- Mf2190: California, San Mateo County, Ano Nuevo 7.5' quad.
From low sea cliff at small beach about 400 m (1300 ft) E. of Ano Nuevo Point. Unsectioned. Vaqueros(?) Formation, lower Miocene. From 3.6m (12 ft) stratigraphically above volcanic breccia. Collected by J. C. Clark, 1968
(field number = JC68-5)
- Mf2191: California, San Mateo County, Ano Nuevo 7.5' quad.
Same locality as Mf2190; from 1.8 m (6 ft) stratigraphically above Mf2190. Unsectioned. Vaqueros(?) Formation, middle Miocene. Collected by J. C. Clark, 1968
(field number = JC68-4)
- Mf2192: California, San Mateo County, Ano Nuevo 7.5' quad.
Same locality as Mf2191, from about 1m (4 ft) stratigraphically above Mf2191. Unsectioned. Vaqueros(?) Formation, middle Miocene. Collected by J. C. Clark, 1968
(field number = JC68-3)
- Mf2193 California, San Mateo County, Ano Nuevo 7.5' quad.
From sea cliff on east side of small beach about 400m (1325 ft) E. of Ano Nuevo Point. Unsectioned. Monterey Formation, middle Miocene. Collected by J. C. Clark, 1968
(field number = JC68-2)
- Mf2194: California, San Mateo County, Ano Nuevo 7.5' quad.
From sea cliff about 2m (6 ft) above water level from E. side of small cove about 420m (1375 ft) E. of Ano Nuevo Point. Unsectioned Monterey Formation, middle Miocene. From 4m (13 ft) stratigraphically above Mf2193. Collected by J. C. Clark, 1963
(field number = JC63-4)
- Mf2195: California, San Mateo County, Ano Nuevo 7.5' quad.
From W. side of small headland about 640m (2100 ft) E. of Ano Nuevo Point. Unsectioned. Monterey Formation, middle Miocene. From 0.6m (2 ft) below unconformable contact with overlying Purisima Formation. Collected by J. C. Clark, 1968
(field number = JC68-1)
- Mf4531: California, San Mateo County, Ano Nuevo 7.5' quad.
From sea cliff 107m (350 ft) N.70°W. of lighthouse on Ano Nuevo Island. Unsectioned. Monterey Formation, middle Miocene. Collected by J. C. Clark, 1967
(field number = JC67-5)

LARGE FORAMINIFERS

Large foraminifers, such as Discocyclina sp., were collected from rocks of Eocene age and from reworked boulders in rocks of Miocene age at three areas in the San Francisco Bay region (fig. 32), and near Devils Den (fig. 2). In the San Francisco Bay region, the large foraminifers were collected near Mount Diablo (loc. EB 611), south of San Jose (locs. 74CB971, EB633, and 71CB983), and in the Sveadal area (Osburn locs. 1-1, 1-43, 1-49).

Mount Diablo Area

Many reports have been published on the Paleogene stratigraphy and paleontology of the Mount Diablo area. Colburn (1961) discusses most of this literature. The rocks are reasonably well exposed and fossiliferous but generally have shallower water assemblages than those in Santa Cruz Mountains.

Figure 33 shows just a few of the many fossil collections that have been made in a small part of the area. The map is from Watson (1941, pl. 4) and was part of a project of the late Hubert G. Schenck to study and correlate Eocene formations in California. Collection EB 611 was made at LSJU 986 from which Keenan (1932, pl. 3, figs. 1 and 2) described and illustrated two species of large foraminifers. The bed from which these fossils were obtained is just a few centimeters thick and was covered in 1976 when the section was revisited.

San Jose and Sveadal Areas

Large foraminifers from the San Jose area were first described and illustrated by Schenck (1929, p. 224-227). Bailey and Everhart (1964) have described the geology of the area. Locality 71CB983 (see fig. 34) is about one-half kilometer northwest of LSJU 309 from which the holotype of Discocyclina californica was obtained. Within a few centimeters of 71CB983C, which contains large foraminifers, another sample (71CB983A) was collected that contains the following smaller foraminifers, according to W. A. Berggren (written comm., 1977):

Acarinina coalingensis (Cushman and Hanna)
Subbotina patagonica (Todd and Kniker)
Acarinina acarinata Subbotina
A. soldadoensis (Bronnimann)
Morozovella subbotinae (Morozova)
M. gracilis (Bolli)
M. formosa (Bolli)
Several species of agglutinated foraminifers

Berggren believes that this assemblage is correlative with the Morozovella formosa zone (P7), early Eocene.

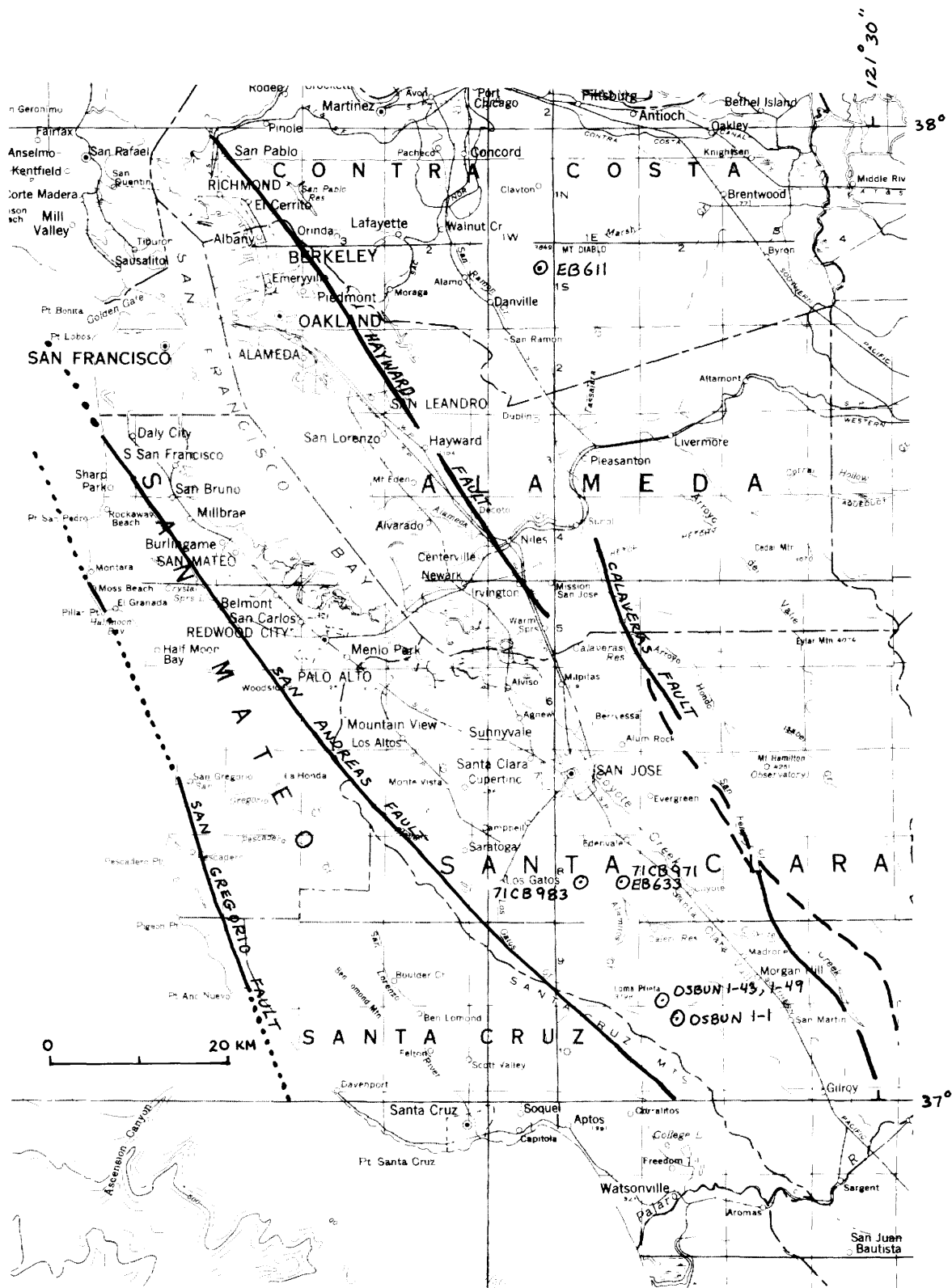


Figure 32.--Index map showing localities where large foraminifers, such as Discocyclina sp., were collected.

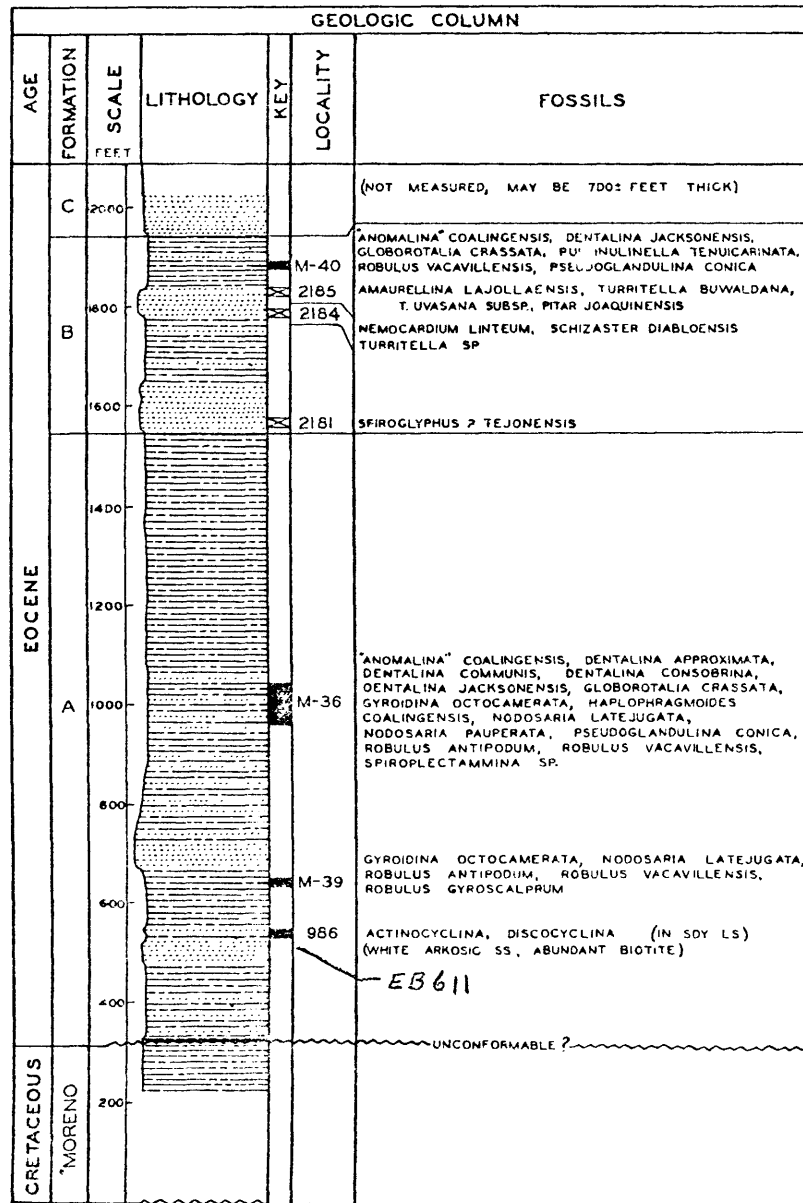
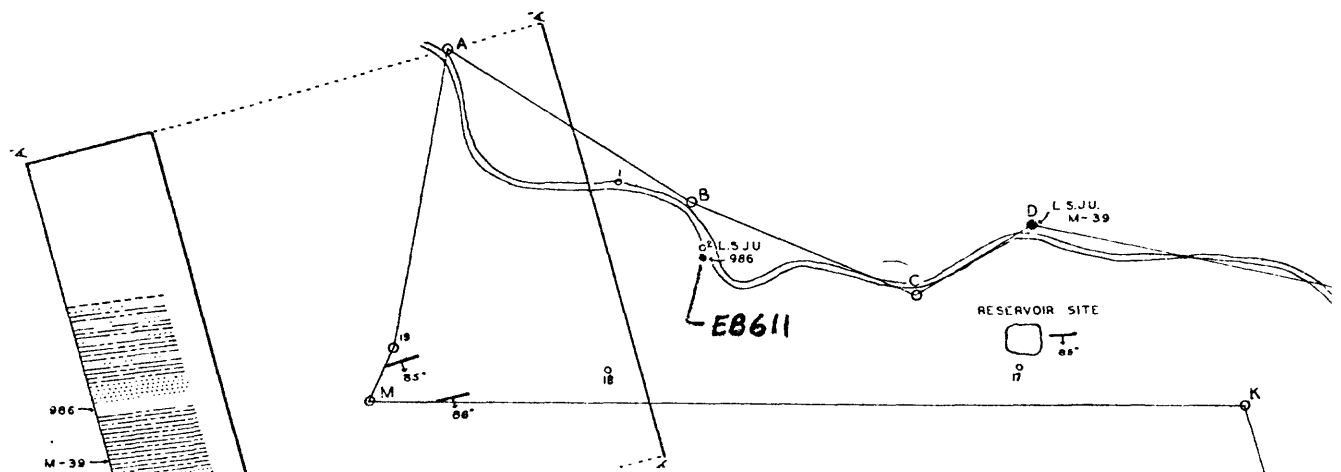
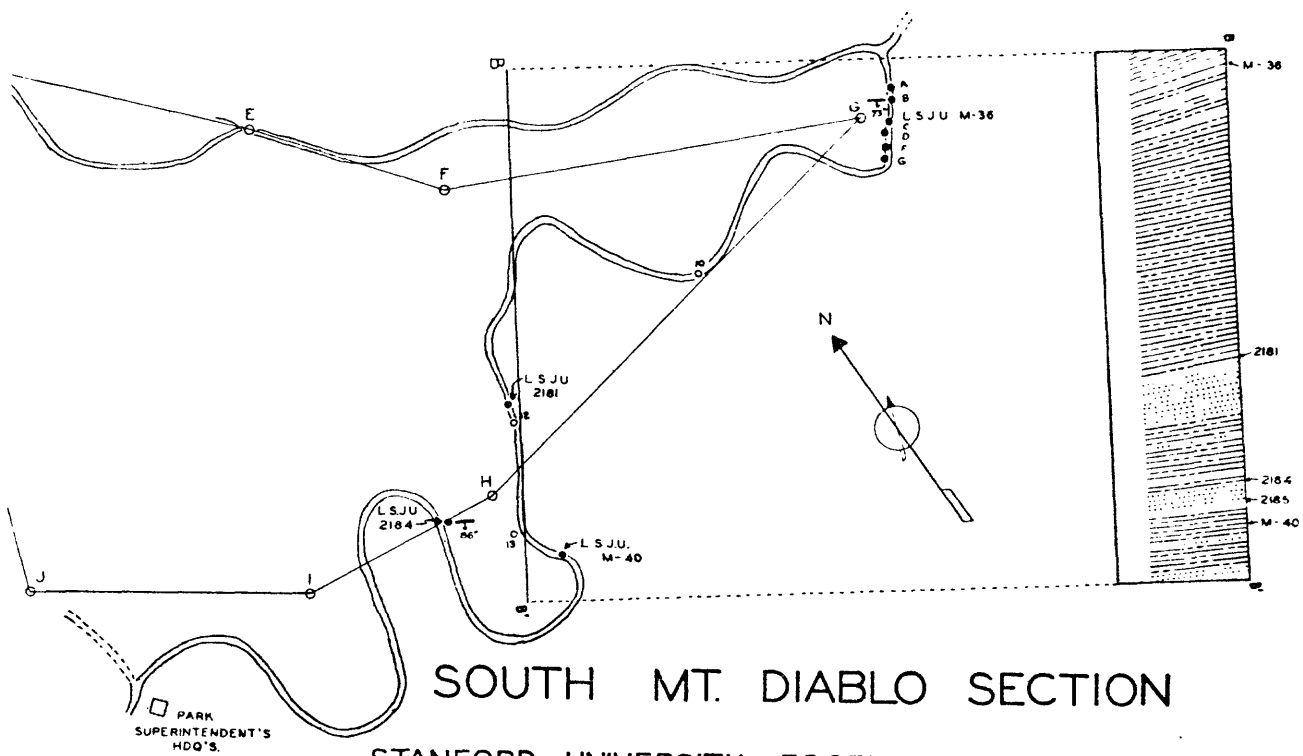


Figure 33.--
Locality map and
stratigraphic
column, Mount
Diablo area



SOUTH MT. DIABLO SECTION

STANFORD UNIVERSITY EOCENE PROJECT NO. 83
 MT. DIABLO QUAD., CONTRA COSTA CO., CALIF.

PLANE TABLE MEASUREMENT
 NOV. 1939 BY
 E. WATSON AND A. BRADBURY

L.S.J.U. LOC. 986; IN ROADCUT
 E. OF HILL 2018, U.S.G.S. TOPO-
 GRAPHIC MAP, 2.2 MI. S. 30° W.
 OF TOP OF MT. DIABLO

200 400 600 FEET

KEY

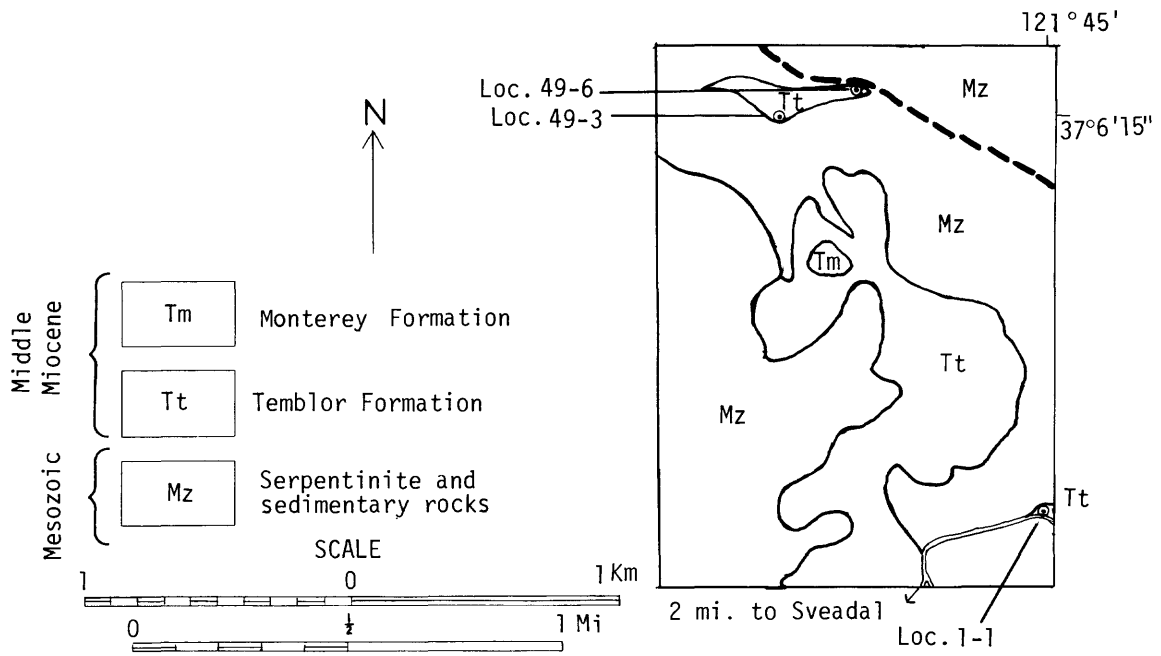
LOCALITY NUMBERS RECORDED IN
 S.J.U. REGISTER OF LOCALITIES

- MEGAFOSSILS
- MICROFOSSILS

Figure 33.--continued

Locality 71CB983 is now covered by a newly constructed house. Locality 71B971 is in an abandoned limestone quarry that was still accessible in 1971 when it was last visited.

In the Sveadal area near Loma Prieta (see fig. 32), Osburn (1975) discovered large foraminifers in limestone boulders within the basal part of the Temblor Formation of middle Miocene age. The localities and a stratigraphic column are shown on figure 35. The boulders were presumably derived from the erosion of the limestone in the vicinity of 71CB983 and 71CB971.



ERATHEM	SERIES	FORMATION	LITHOLOGY	THICKNESS	REMARKS
CENOZOIC	MIOCENE	Monterey Formation		18 m (60 ft)	
		Temblor Formation		600 m (2000 ft)	Osburn Collections 49-6, 49-3, 1-1 with large foraminifers
MESOZOIC	JURASSIC AND CRETACEOUS				unconformity serpentine

Figure 35.--Geologic map and stratigraphic column for the Sveadal area. Geology from Osburn (1975).

DEVILS DEN SECTION

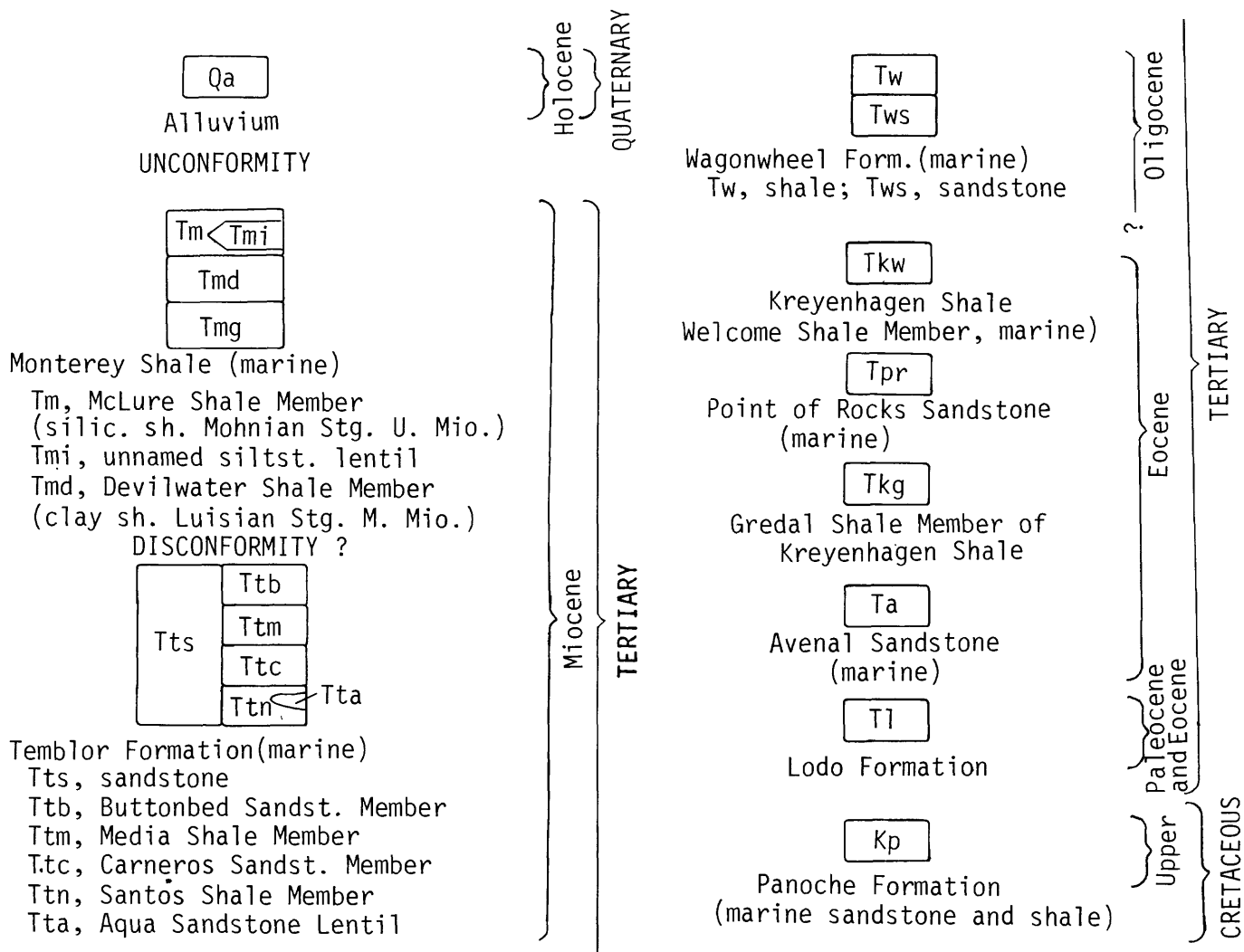
The Devils Den area (figs. 36 and 37) has a nearly complete section of Paleogene rocks that is easily accessible, moderately well exposed, and not too complicated structurally. Unfortunately, at most (but not all) localities the calcareous foraminifers have been leached out. Several trenches have been dug by oil company paleontologists in generally successful attempts to obtain fresh material--one of these may have provided the foraminifers that Mallory (1959) used to establish his Narizian Stage. The senior author found stakes and shallow pits that were probably left by Smith (1956) in his study of the Wagon Wheel Formation, but material dug from a depth of 1 m in these pits failed to yield any calcareous foraminifers. Auger holes or trenches to a depth of 3 m or more are probably required to recover calcareous material for most of the section.

Happily, the Barranda Mesa Water District constructed a water distribution system in the late 1960's that has opened new exposures of the lower part of the Paleogene sequence. In particular, the afterbay area of pump station "A" has a continuous section from the Avenal Sandstone through the Gredal Shale Member of the Kreyenhagen Shale to the lower part of the Point of Rocks Sandstone (see figs. 38 and 39). The lower part of this section is repeated by a fault, and the relation of the lowest beds to the underlying Panoche Formation of Late Cretaceous age could not be determined. A small amount of siltstone below the Avenal Sandstone has been provisionally mapped as Lodo Formation.

The Devils Den area has been remapped recently by Dibblee (1971, 1974). Dibblee (1973) also has discussed the stratigraphy and nomenclature of the various rock units in the Devils Den area.

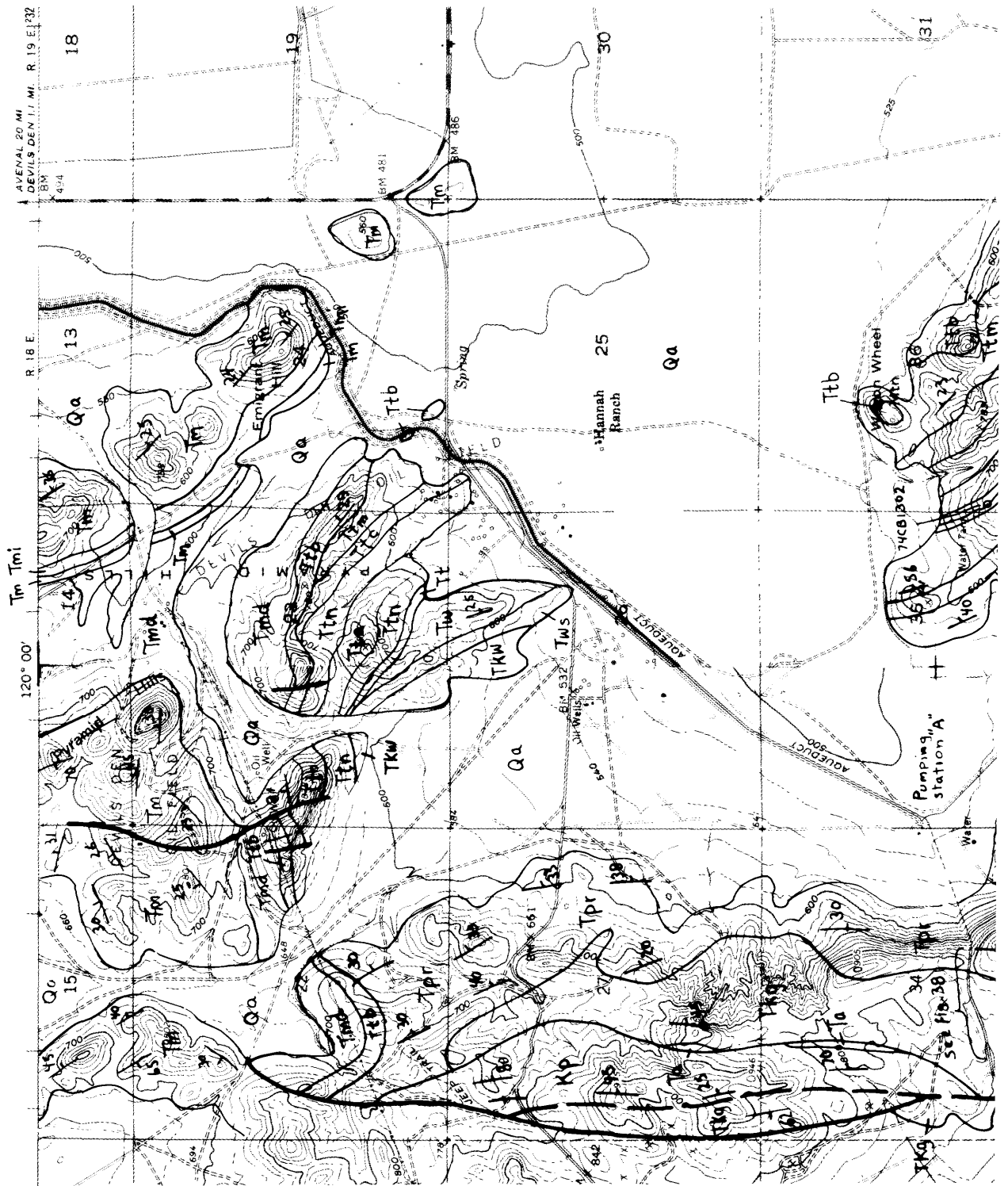
The Point of Rocks Sandstone and Kreyenhagen Shale of the Devils Den area are inferred by Clarke and Nilsen (1973) to have been contiguous with the Butano Sandstone and Twobar Shale Member of the San Lorenzo Formation in the Santa Cruz Mountain area (see figs. 3 and 5). They believe that these units have been displaced since Eocene time by about 300 km of right-lateral movement along the San Andreas fault.

EXPLANATION



(Map follows on next page.)

Figure 36.--Geologic map of the Devils Den area. Base maps from Sawtooth Ridge and Emigrant Hill 7.5' quadrangles. Geology from Dibblee (1971, 1973, and 1974).



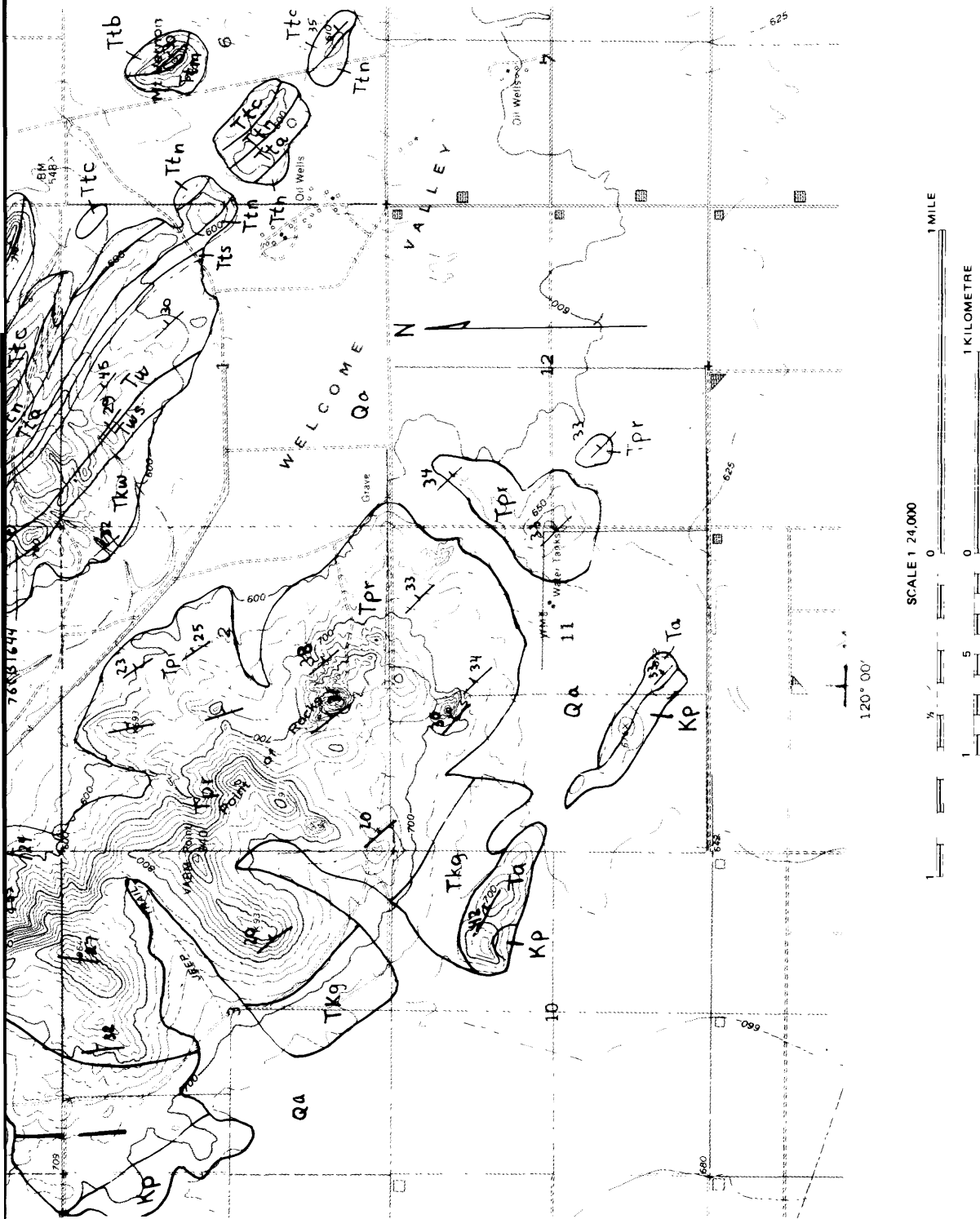
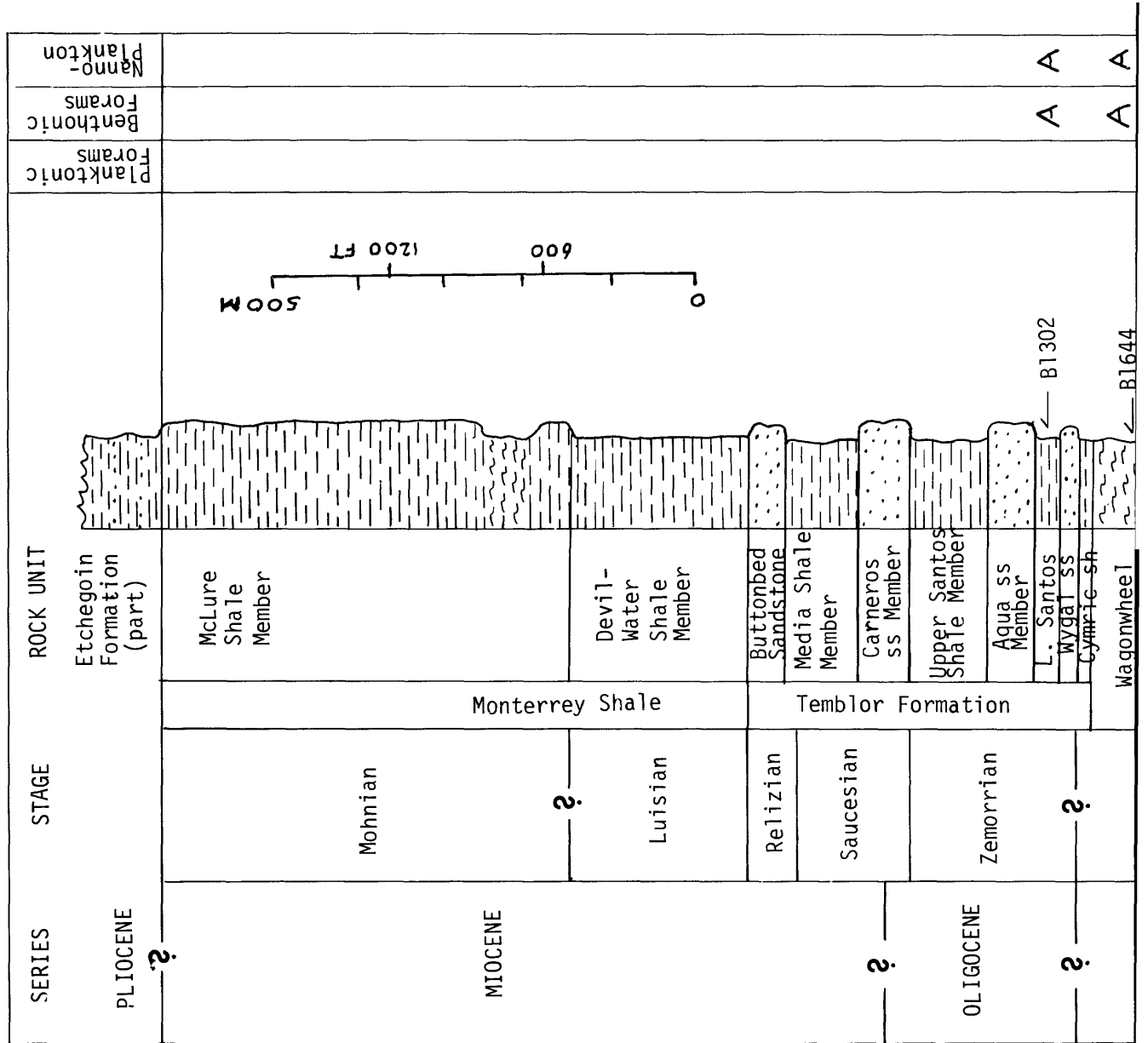


Figure 36.-- continued



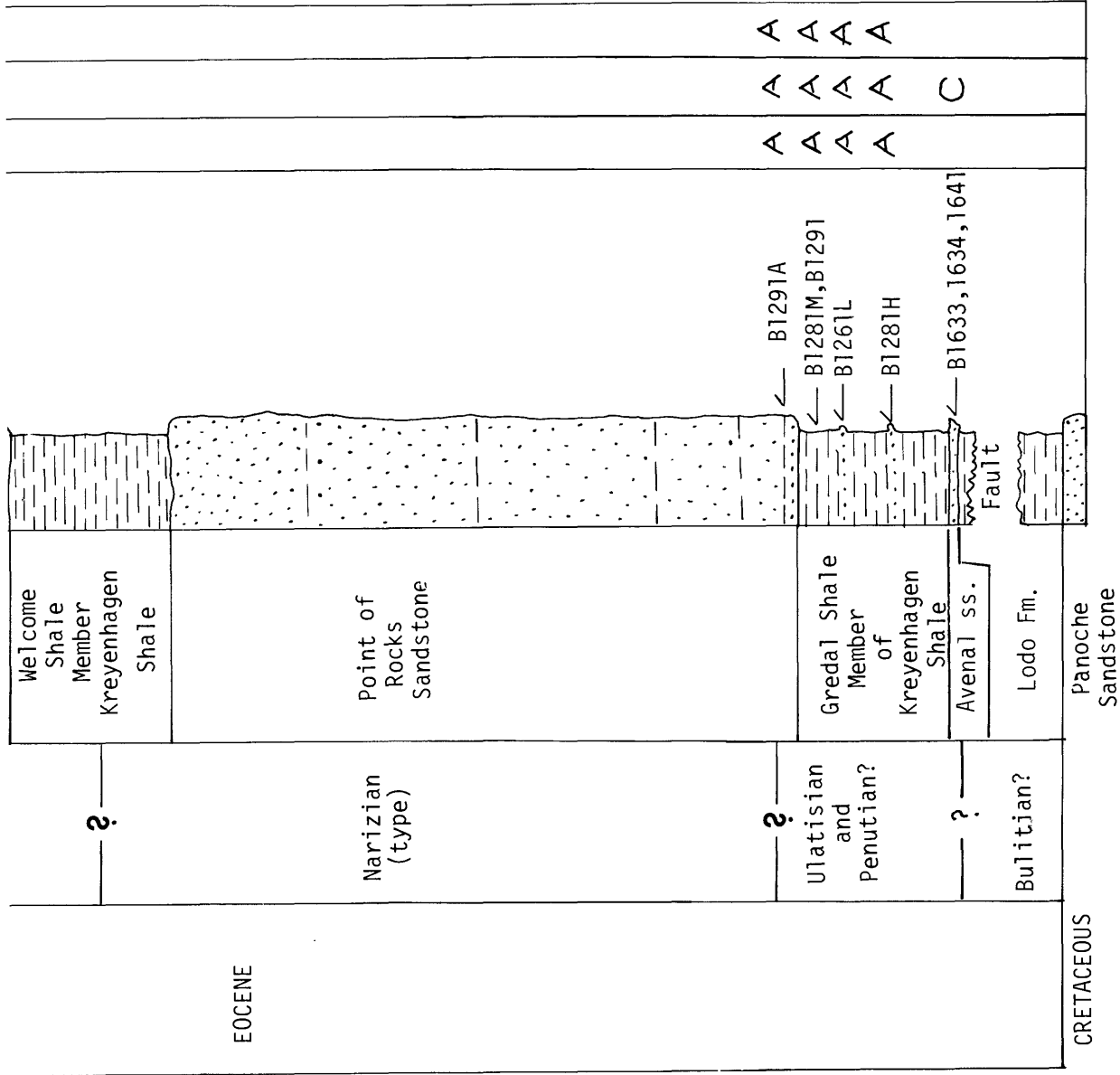


Figure 37.--Stratigraphic column for the Devils Den area. Data from Van Couvering and Allen (1943) and Dibblee (1973).

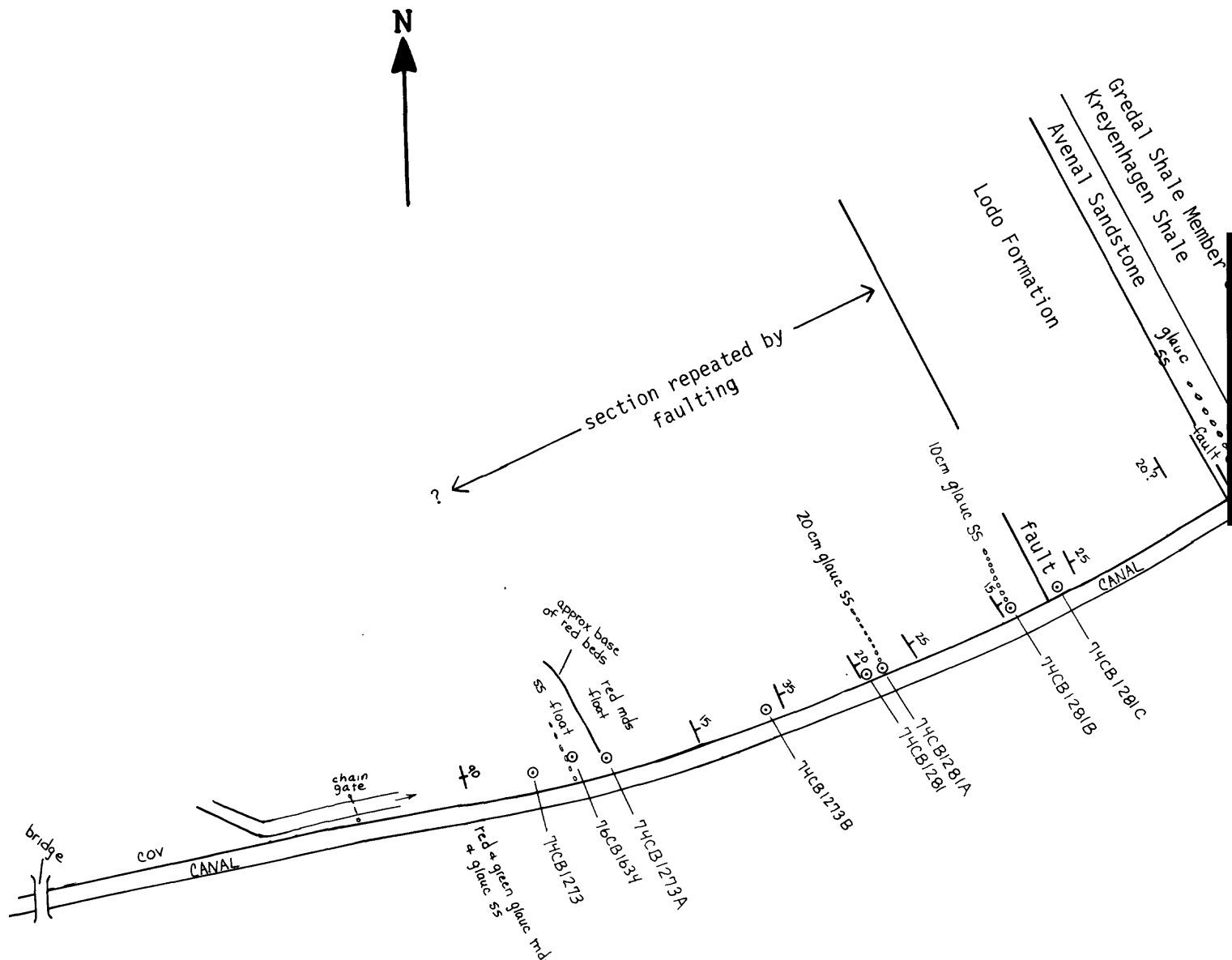
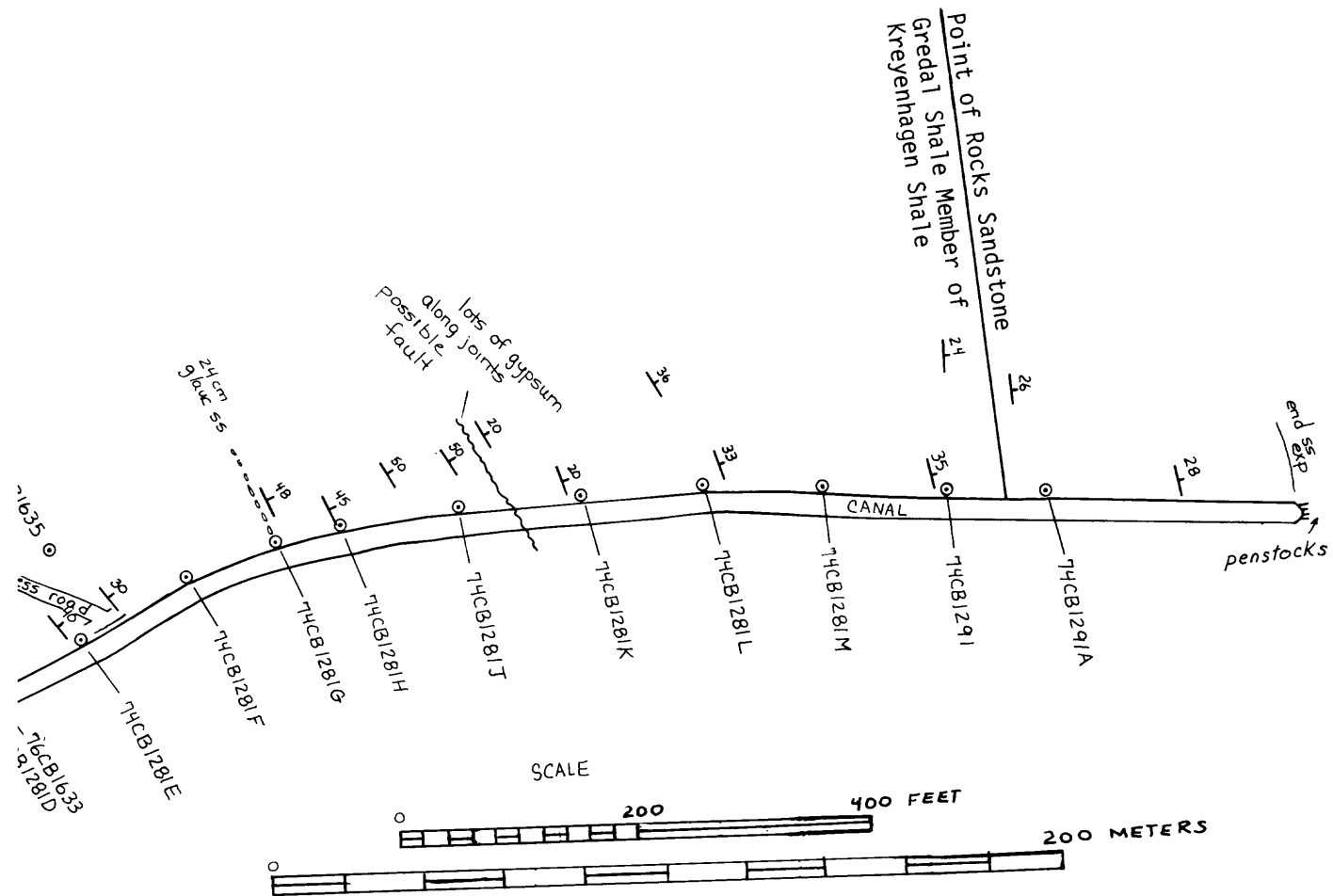


Figure 38.--Geologic sketch map of the aqueduct near Devils Den showing formation contacts, structures, and localities where fossils were collected



Measured by tape and compass, 1974, by E.E. Brabb, J.R. LeCompte; and S.W. Moore

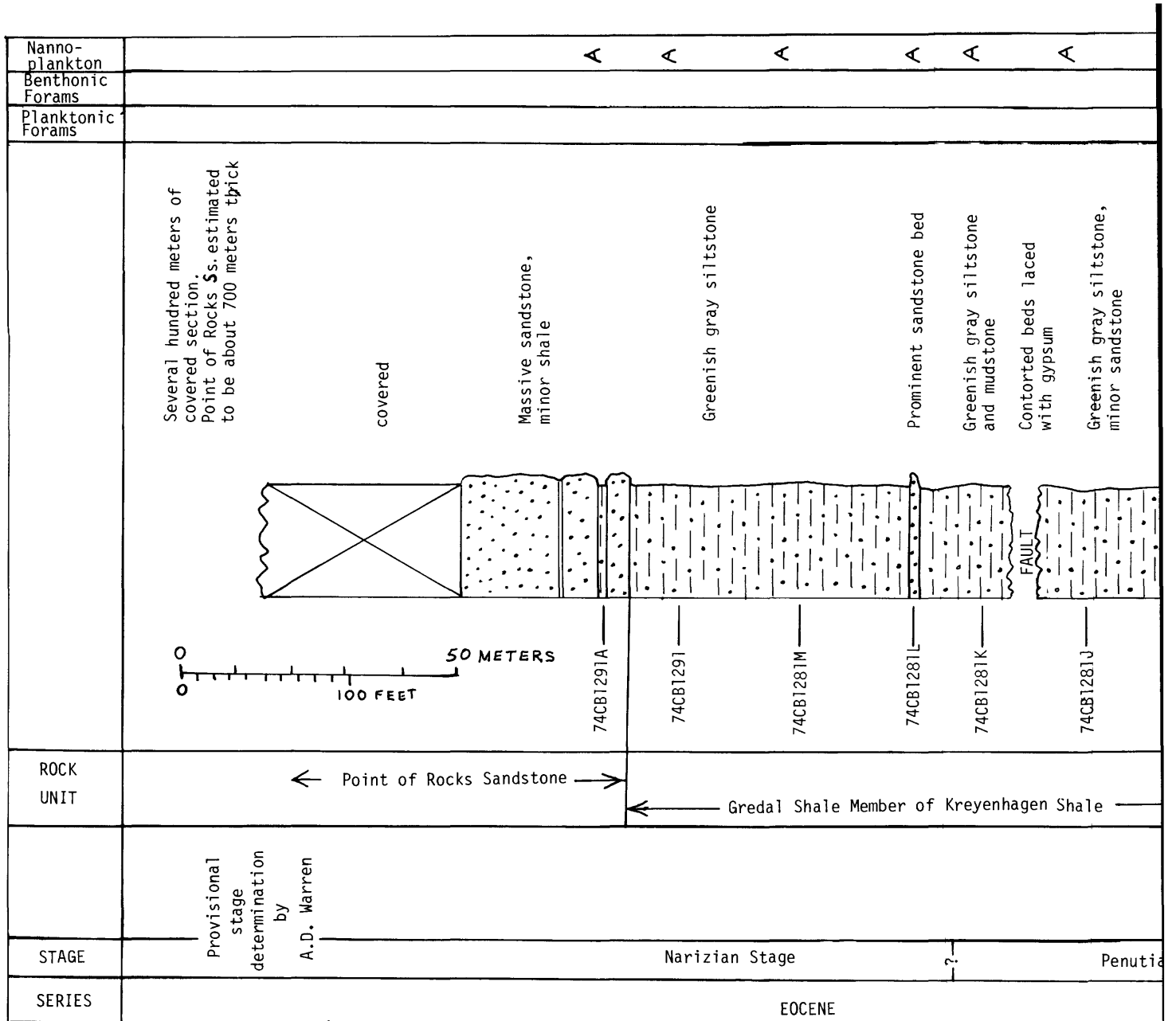
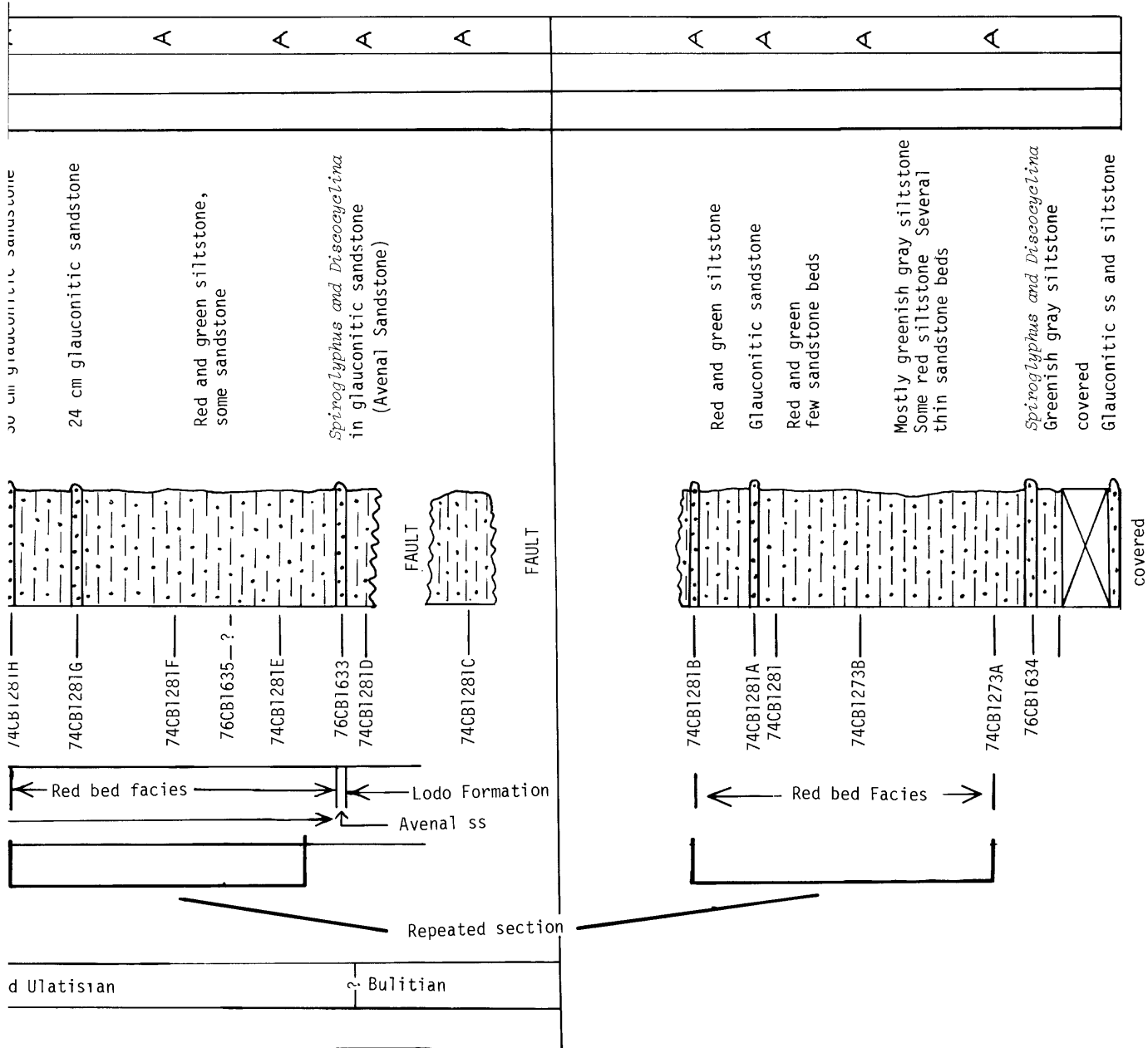


Figure 39.--Stratigraphic column, lower part of



section near Devils Den area, along aqueduct.

LODO GULCH SECTION

The Lodo Formation at its type locality (figs. 40 and 41) is one of the most important reference sections for the lower Paleogene in California. Foraminifers from the section have been described by Martin (1943), Israelsky (1951 and 1955), and Schmidt (1970). Nannoplankton have been described by Bramlette and Sullivan (1961). The integration of foraminifer and nannoplankton data and correlation with standard zonations has been done by Schmidt (1970, p. 50)--his time scale has been added to the stratigraphic column, figure 42.

The thickest section of the Lodo Formation (sec. A, fig. 40) is covered in its lower part. A supplementary section (B, fig. 40) was used by Israelsky (1951, pl. 1) to show the lower part of the formation and its contact with the underlying Moreno Shale of Late Cretaceous age. We have used still another section (C, fig. 40) to show this lower part. The section at C has locality LSJU 2073 shown on a map by Martin (1943, fig. 1) and on another map by Smith (1975, fig. 1). Smith lists more than 60 species of gastropods, pelecypods, brachiopods, and corals from LSJU 2073, and she compares them with similar species from the Bracheux sands, Rilly sands, and other units of Paleocene age in the Paris basin. The sandstone at LSJU 2073 is glauconitic and may have provided the material dated by Funnell (1964, p. 188, item 113) as 58.5 m.y. old.

The rocks in section A have slid extensively so that it is difficult in most places to obtain samples in correct stratigraphic position. Holes up to 1 m deep were dug with a shovel in an attempt to get below the landslide debris and into undisturbed rock, but some mixing of the samples is likely. A similar problem was noted by Israelsky (1951, p. 4).

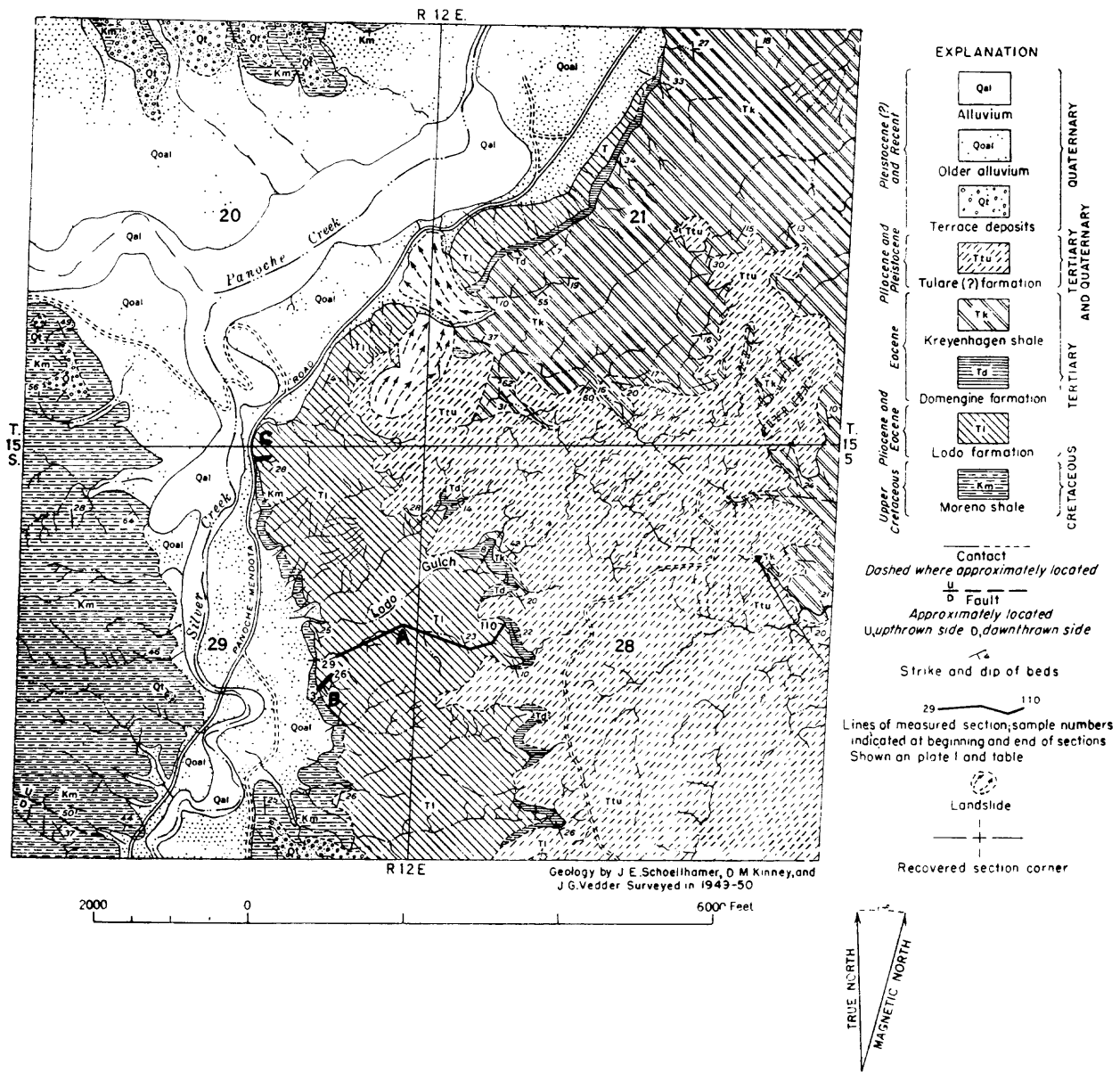


Figure 40.--Geologic map of the northwest Tumey Hills showing location of Lodo Gulch sections.

Approximate location of Panache Creek Road

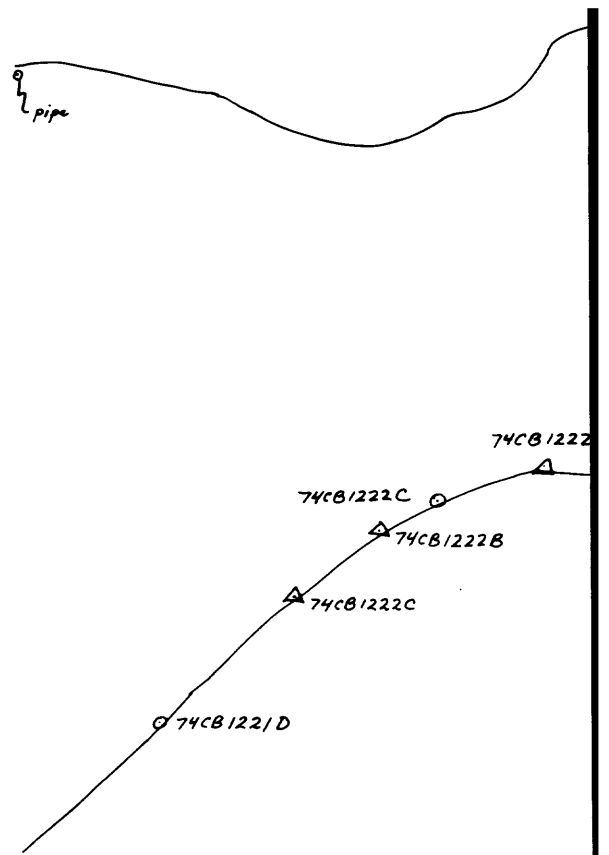
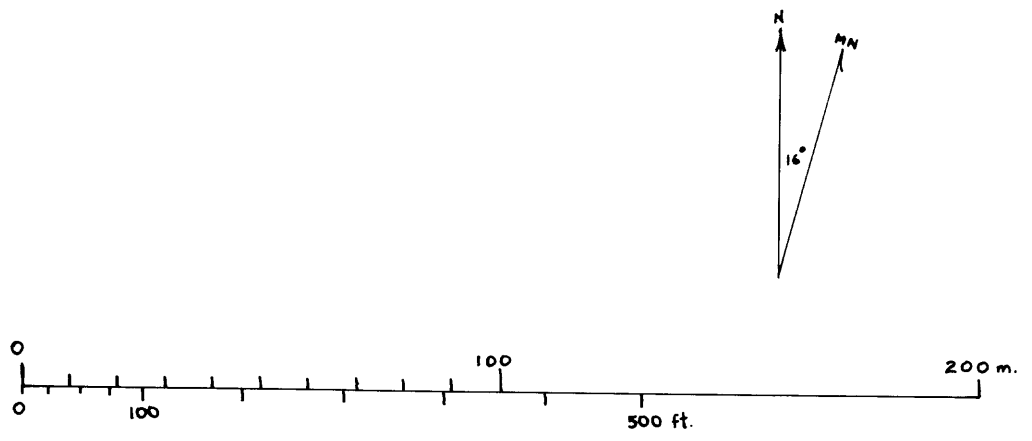
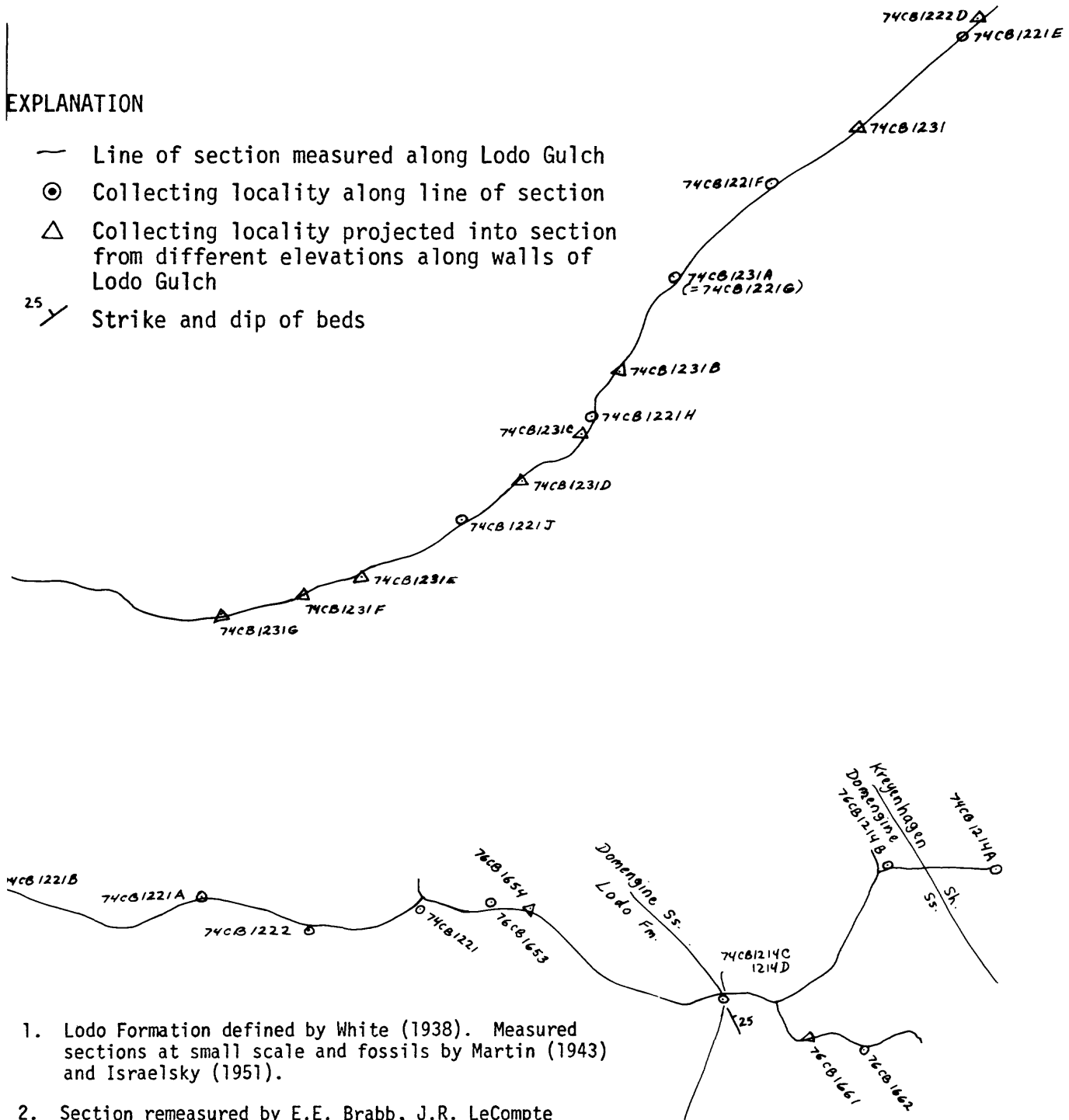


Figure 41.--Geologic sketch map of the Lod

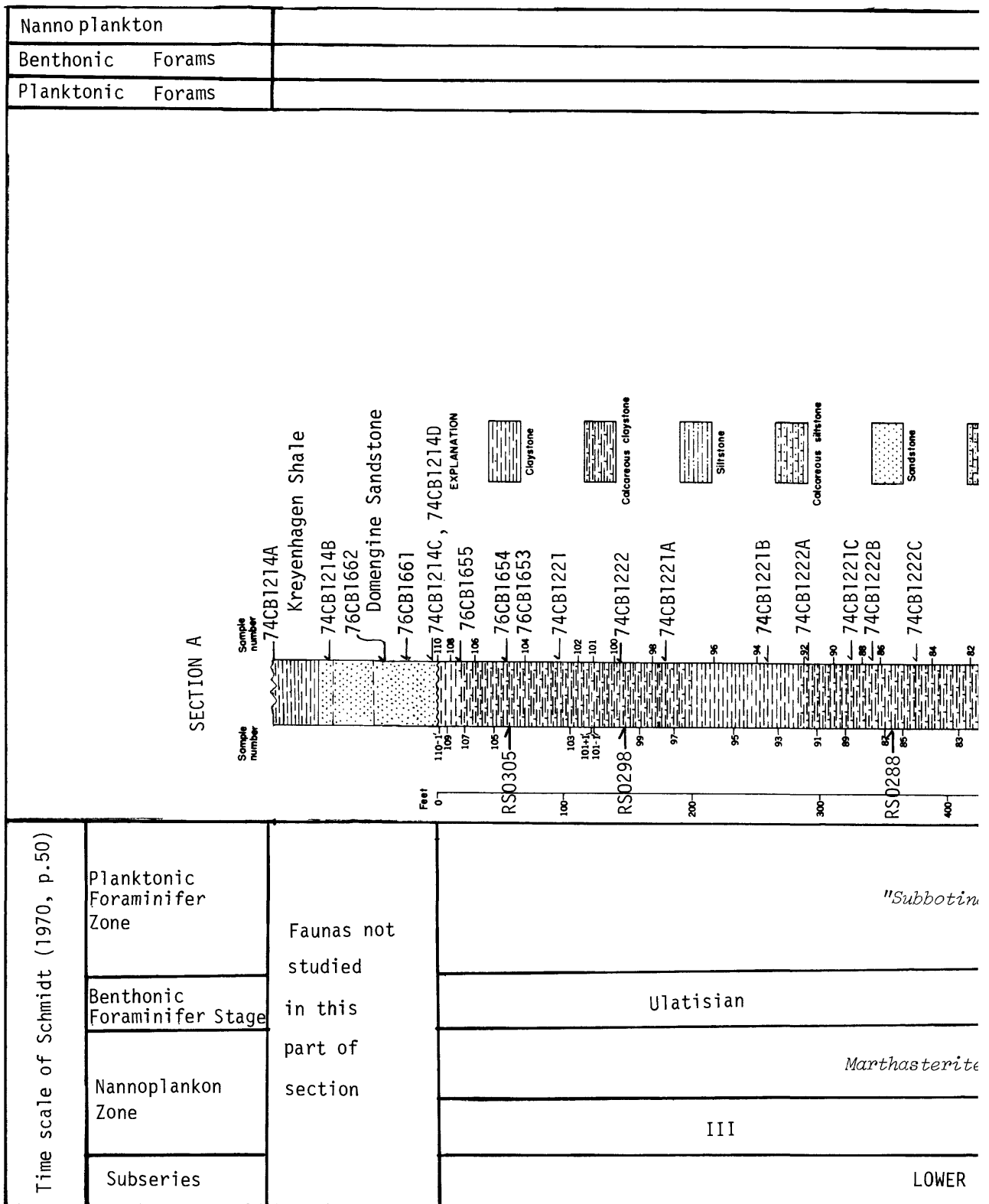
EXPLANATION

- Line of section measured along Lodo Gulch
- ⊙ Collecting locality along line of section
- △ Collecting locality projected into section from different elevations along walls of Lodo Gulch
- ²⁵ / Strike and dip of beds



1. Lodo Formation defined by White (1938). Measured sections at small scale and fossils by Martin (1943) and Israelsky (1951).
2. Section remeasured by E.E. Brabb, J.R. LeCompte and S.W. Moore, 10/15/74 to 10/17/74.
3. Most of the section is covered with landslide debris. Samples were collected by digging holes a few feet deep through the landslide deposits. Some mixing of the presumed bedrock with landslide debris is likely, a problem also noted by Israelsky (1951, p.4).

formation in its type section.



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