

## THE FORMATION OF A CARLOSTURANITE-LIKE PHASE DURING THERMAL DECOMPOSITION OF SERPENTINE IN C-COMPLEX ASTEROIDS.

L. E. Jenkins<sup>1</sup>, A. J. King<sup>2</sup>, M. R. Lee<sup>1</sup>, L. Daly<sup>1,3,4</sup>, K. Ignatyev<sup>5</sup>, C. J. Floyd<sup>1</sup>, and P-E.M.C. Martin<sup>1</sup>, <sup>1</sup>School of Geographical & Earth Sciences, University of Glasgow, Glasgow, UK ([L.jenkins.1@research.gla.ac.uk](mailto:L.jenkins.1@research.gla.ac.uk)). <sup>2</sup>Planetary Materials Group, Natural History Museum, London, UK. <sup>3</sup>Australian Centre for Microscopy & Microanalysis, The University of Sydney, Australia. <sup>4</sup>Department of Materials, University of Oxford, Oxford, UK. <sup>5</sup>Diamond Light Source, Didcot, UK.

**Introduction:** Some Mighei-like carbonaceous chondrites (CMs) underwent post-hydration heating on their parent C-complex asteroid(s). This can result in the decomposition of serpentine, and the recrystallization of olivine and enstatite from its remnants [1]. In heated CMs, a transitional structure has been identified; it forms between the breakdown of serpentine and recrystallization of olivine and enstatite [2]. This phase's structure is unknown. In a prior heating experiment with X-ray diffraction (XRD) analysis of a bulk powdered CM, a peak at  $\sim 3.56\text{\AA}$  associated with this phase was observed at  $525^\circ\text{C}$ ; however due to peak overlaps with other phases, no other peaks of the transitional phase were identified [3]. To further characterize this phase, we heated two polished rock slices of the CM2 Murchison at  $400\text{--}550^\circ\text{C}$  and collected *in situ* micro XRD ( $\mu\text{XRD}$ ) data to minimize peak overlap during heating.

**Methods:** The rock slices of Murchison ( $100\ \mu\text{m}$  thick,  $\sim 1.8\ \text{mm}^2$  in area each) were heated in an inert  $\text{N}_2$  atmosphere using a THMS600 Linkam cell connected to a TMS 94 temperature controller and an ECP water circulation pump. The samples were kept in place by a pair of Kapton disks and Linkam's vertical sample holder. The I18 beamline at the Diamond Light Source (DLS) was used to collect  $\mu\text{XRD}$  data for an array of points for a select area in each sample using a spot size of  $2\ \mu\text{m}$  and a wavelength of  $0.826\ \text{\AA}$ .  $\mu\text{XRD}$  patterns were taken at room temperature prior to heating each sample  $400\text{--}550^\circ\text{C}$  in  $25^\circ\text{C}$  steps. The samples were held for two hours at each step and  $\mu\text{XRD}$  patterns were collected after an hour. Heating stopped at  $550^\circ\text{C}$  due to the disintegration of the securing Kapton.  $\mu\text{XRD}$  patterns of a set of standards were taken for comparison. The  $\mu\text{XRD}$  data was processed with Dawn-2.27 [4,5] and Rigaku SmartLab II software.

**Results:** At  $525^\circ\text{C}$ , targets yielding a broad peak at  $\sim 3.51\text{\AA}$  associated with the transitional phase were identified. The transitional structure was found to have additional peaks at  $2.28\text{\AA}$ ,  $2.31\text{\AA}$ ,  $2.35\text{\AA}$ ,  $2.49\text{\AA}$ ,  $2.54\text{\AA}$ ,  $3.04\text{\AA}$ , and  $3.51\text{\AA}$  (Fig. 1). The peak shapes were most similar to that of the antigorite standard (Fig. 1), however their d-spacings best matched carlosturanite ( $(\text{Mg}, \text{Fe})_{21}\text{Si}_{12}\text{O}_{28}(\text{OH})_{34}\cdot\text{H}_2\text{O}$ ), a phase related to antigorite [6].

**Discussion:** Because carlosturanite decomposes at  $400^\circ\text{C}$  and lacks a  $3.04\text{\AA}$  peak [6], it is not the transitional structure. However, the transitional structure likely shares structural similarities with carlosturanite, accounting for the similarities in their XRD patterns. It likely formed in the process of serpentine decomposition outlined by [7], wherein acceptor regions receive  $\text{Mg}^{2+}$  and  $\text{Si}^{4+}$  from donor regions, which in turn receive  $\text{H}^+$  from the acceptor regions to form  $\text{H}_2\text{O}$ . The  $\text{H}_2\text{O}$  is expelled while the acceptor regions form a partially disordered transitional structure [7]. This likely results in serpentine's tetrahedral sheet breaking up into strips similar to what is found within carlosturanite [8]. Other aspects of the crystal structure are not likely to be maintained due to the lack of  $\text{OH}^-$ , resulting in a partially disordered structure, accounting for differences in the XRD patterns between it and carlosturanite.

**Conclusions:** When CMs are heated to  $525^\circ\text{C}$ , their serpentine decomposes into a partially disordered transitional structure with structural similarities to carlosturanite. This phase may be diagnostic of moderately heated CMs.

**Acknowledgements:** We thank DLS (beamtime award SP31815). We thank Connor Brolly for assisting with equipment assembly and Holly Yu for PAT testing. We thank Christopher Syme and Ben Russell for advice regarding the Linkam cell use. We thank Finlay Walton and Zhiyu Liao for lending equipment. We thank Charlotte Slaymark for assisting with sample storage. We thank Claire Wilson for providing the Rigaku SmartLab II software used.

**References:** [1] Nakamura T. (2005) *Journal of Mineralogical and Petrological Sciences* 100: 260–270. [2] Akai J. (1988) *Geochimica et Cosmochimica Acta* 52: 1593–1599. [3] Jenkins L.E. et al. (2022) *85<sup>th</sup> MetSoc*, Abstract #6122. [4] Basham et al. (2015) *Journal of Synchrotron Radiation* 22 [5] Filik J. et al. (2017) *Journal of Applied Crystallography* 50 [6] Compagnoni R. et al. (1985) *American Mineralogist* 70: 767–772. [7] Ball M.C. & Taylor H.F.W. (1962) *Mineralogical Magazine* 33: 467–482. [8] Mellini et al. (1985) *American Mineralogist* 70: 773–781.

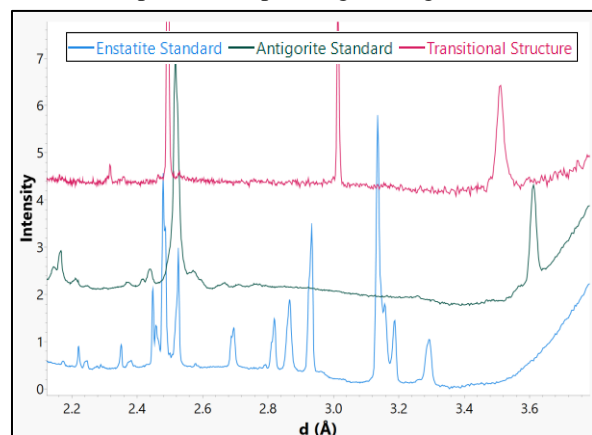


Fig. 1. Comparison of peak shapes for the transitional structure, and enstatite and antigorite standards.