If $F \ge 2$ is the most common situation, then the phase rule may be adjusted accordingly:

 $F = C - \phi + 2 \ge 2$

$\phi \leq C (Eq \ 24.1)$

Goldschmidt's mineralogical phase rule, or simply the mineralogical phase rule

Suppose we have determined C for a rock Consider the following three scenarios:

a) $\phi = C$

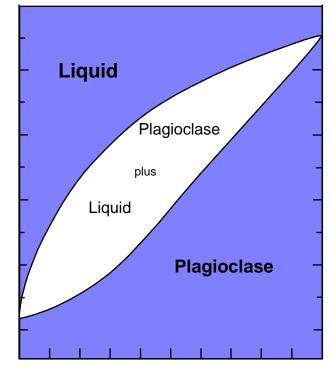
The standard divariant situation

The rock probably represents an equilibrium mineral assemblage from within a metamorphic zone

b) φ < C

Common with mineral systems that exhibit solid

solution



c) $\phi > C$

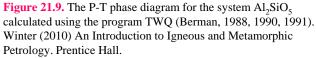
A more interesting situation, and at least one of three situations must be responsible:

1) F < 2

The sample is collected from a location right on a **univariant reaction curve (isograd) or invariant point**

Consider the following three scenarios:

C = 1common $\phi = 1$ 1.0 $\phi = 2$ rare 0.8 $\phi = 3$ only at the specific **Kyanite** Pressure (GPa) 9.0 P-T conditions of the 0.6 invariant point Sillimanite (~ 0.37 GPa and 500°C) 0.2 Andalusite 0 400 500 600 800 700 Temperature °C



2) Equilibrium has not been attained

The phase rule applies only to systems at equilibrium, and there could be any number of minerals coexisting if equilibrium is not attained

3) We didn't choose the # of components correctly

Some guidelines for an appropriate choice of C

- Begin with a 1-component system, such as CaAl₂Si₂O₈ (anorthite), there are 3 common types of major/minor components that we can add
 - a) Components that generate a new phase

Adding a component such as $CaMgSi_2O_6$ (diopside), results in an additional phase: in the binary Di-An system diopside coexists with anorthite below the solidus

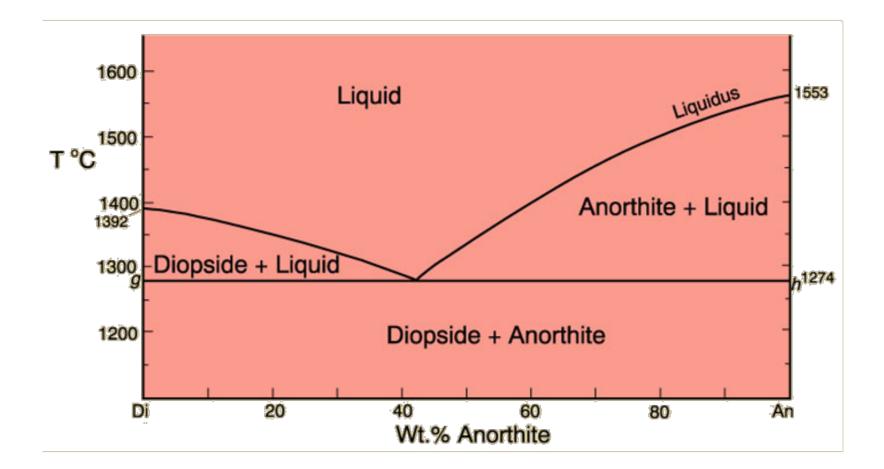


Fig. 6.11. Isobaric T-X phase diagram at atmospheric pressure. After Bowen (1915), Amer. J. Sci. 40, 161-185.

3) We didn't choose the # of components correctly

b) Components that substitute for other components

- Adding a component such as NaAlSi₃O₈ (albite) to the 1-C anorthite system would dissolve in the anorthite structure, resulting in a single solid-solution mineral (plagioclase) below the solidus
 - Fe and Mn commonly substitute for Mg
 - Al may substitute for Si
 - Na may substitute for K

3) We didn't choose the # of components correctly

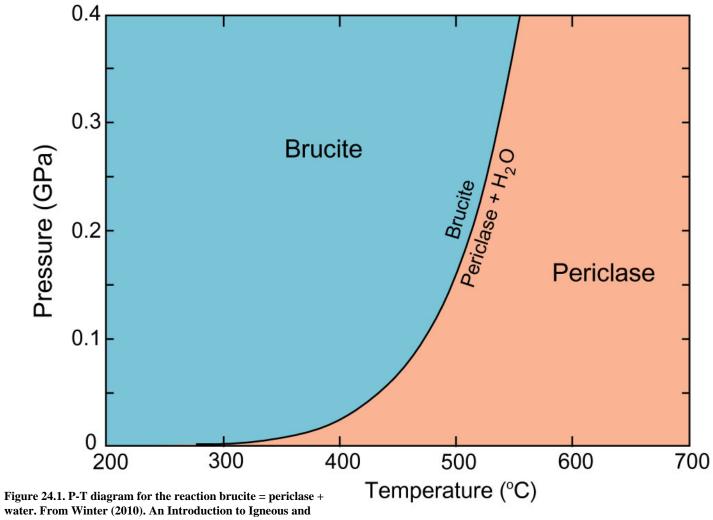
c) "Perfectly mobile" components

- Mobile components are either a freely mobile fluid component or a component that dissolves readily in a fluid phase and can be transported easily
- The chemical activity of such components is commonly controlled by factors external to the local rock system
- They are commonly ignored in deriving C for metamorphic systems

Consider the very simple metamorphic system, MgO-H₂O

- Possible natural phases in this system are periclase (MgO), aqueous fluid (H₂O), and brucite (Mg(OH)₂)
- How we deal with H₂O depends upon whether water is perfectly mobile or not
- A reaction can occur between the potential phases in this system:

 $MgO + H_2O \rightarrow Mg(OH)_2$ Per + Fluid = Bru



Metamorphic Petrology. Prentice Hall.

How do you know which way is correct?

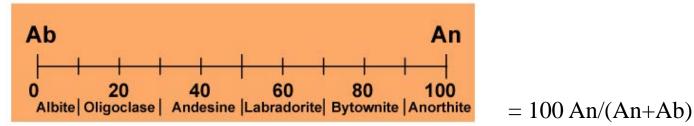
The rocks should tell you

- Phase rule = interpretive tool, not predictive
- If only see low-φ assemblages (e.g. Per or Bru in the MgO-H₂O system) → some components may be mobile
- If many phases in an area it is unlikely that all is right on univariant curve, and may require the number of components to include otherwise mobile phases, such as H₂O or CO₂, in order to apply the phase rule correctly

Chemographic Diagrams

Chemographics refers to the graphical representation of the chemistry of mineral assemblages

A simple example: the plagioclase system as a linear C = 2 components plot:

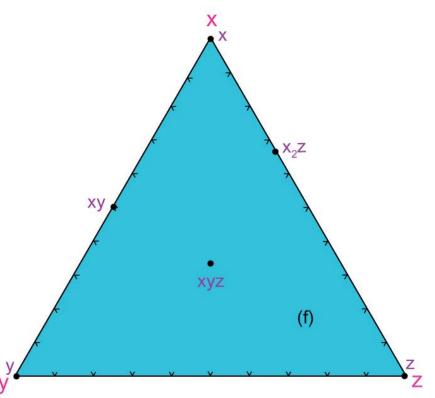


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Chemographic Diagrams

3-C mineral compositions are plotted on a triangular chemographic diagram as shown in Fig. 24.2

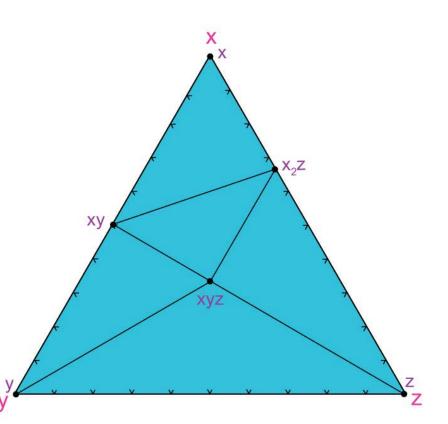
x, y, z, x₂z, xyz, and xy



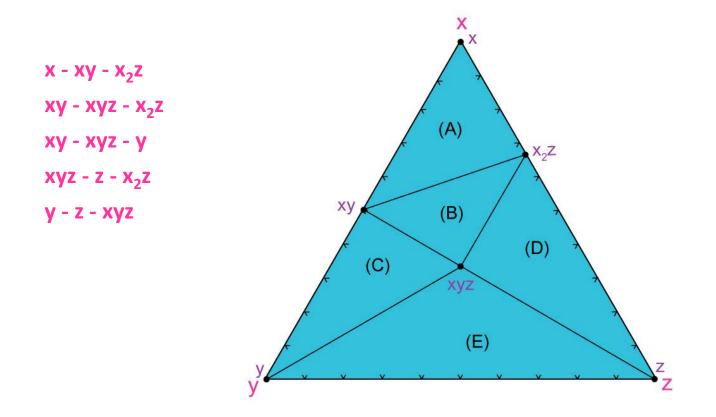
Suppose that the rocks in our area have the following 5 assemblages:

$$x - xy - x_2z$$
$$xy - xyz - x_2z$$
$$xy - xyz - y$$
$$xyz - z - x_2z$$
$$y - z - xyz$$

Figure 24.2. Hypothetical three-component chemographic compatibility diagram illustrating the positions of various stable minerals. Minerals that coexist compatibly under the range of P-T conditions specific to the diagram are connected by tie-lines. After Best (1982) *Igneous and Metamorphic Petrology*. W. H. Freeman.

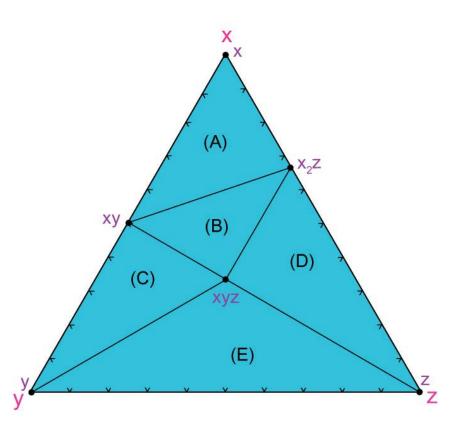


Note that this subdivides the chemographic diagram into 5 sub-triangles, labeled (A)-(E)



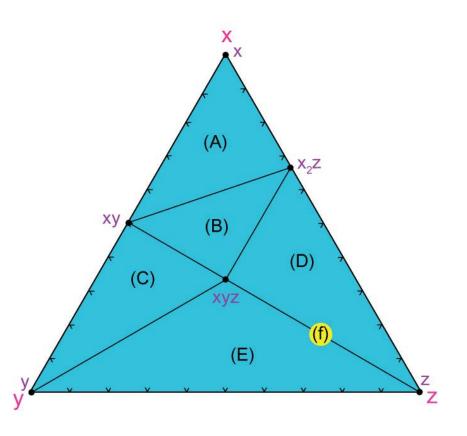
Common point corresponds to 3 phases, thus $\phi = C$

Figure 24.2. Hypothetical three-component chemographic compatibility diagram illustrating the positions of various stable minerals. Minerals that coexist compatibly under the range of P-T conditions specific to the diagram are connected by tie-lines. After Best (1982) *Igneous and Metamorphic Petrology*. W. H. Freeman.



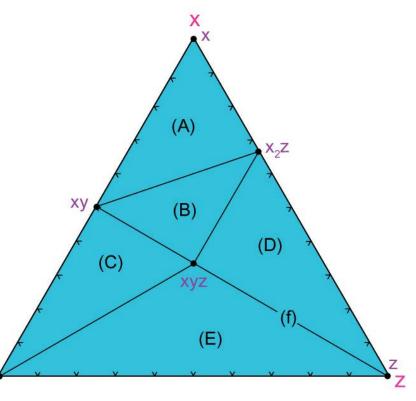
What happens if you pick a composition that falls directly on a tie-line, such as point (f)?

Figure 24.2. Hypothetical three-component chemographic compatibility diagram illustrating the positions of various stable minerals. Minerals that coexist compatibly under the range of P-T conditions specific to the diagram are connected by tie-lines. After Best (1982) *Igneous and Metamorphic Petrology*. W. H. Freeman.



In the unlikely event that the bulk composition equals that of a single mineral, such as xyz, then $\phi = 1$, but C = 1 as well

"compositionally degenerate"



Chemographic Diagrams

Valid compatibility diagram must be referenced to a specific range of P-T conditions, such as a zone in some metamorphic terrane, because the stability of the minerals and their groupings vary as P and T vary

- Previous diagram refers to a P-T range in which the fictitious minerals x, y, z, xy, xyz, and x₂z are all stable and occur in the groups shown
- At different grades the diagrams change
 - Other minerals become stable
 - Different arrangements of the same minerals (different tie-lines connect different coexisting phases)

A diagram in which some minerals exhibit solid solution

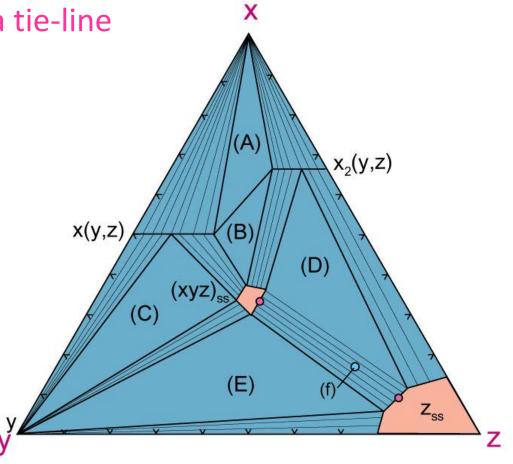
(A) x₂(y,z) **Figure 24.3.** Hypothetical three-component chemographic compatibility diagram illustrating the x(y,z)(B positions of various stable minerals, many of which (D) exhibit solid solution. After (C) (xyz)_{ss} Best (1982) Igneous and Metamorphic Petrology. W. H. (E) \mathbf{Z}_{SS} Ζ

Freeman.

X

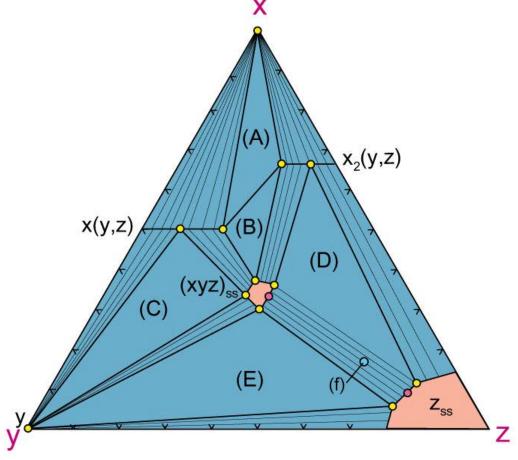
If X_{bulk} on a tie-line

Figure 24.3. Hypothetical three-component chemographic compatibility diagram illustrating the positions of various stable minerals, many of which exhibit solid solution. After Best (1982) *Igneous and Metamorphic Petrology*. W. H. Freeman.



X_{bulk} in 3-phase triangles F = 2 (P & T) so X_{min} fixed

Figure 24.3. Hypothetical three-component chemographic compatibility diagram illustrating the positions of various stable minerals, many of which exhibit solid solution. After Best (1982) *Igneous and Metamorphic Petrology*. W. H. Freeman.



Chemographic Diagrams for Metamorphic Rocks

- Most common natural rocks contain the major elements: SiO_2 , Al_2O_3 , K_2O , CaO, Na_2O , FeO, MgO, MnO and H_2O such that C = 9
- Three components is the maximum number that we can easily deal with in two dimensions
- What is the "right" choice of components?
- Some simplifying methods:

1) Simply "ignore" components

- Trace elements
- Elements that enter only a single phase (we can drop both the component and the phase without violating the phase rule)
- Perfectly mobile components

2) Combine components

Components that substitute for one another in a solid solution: (Fe + Mg)

3) Limit the types of rocks to be shown

Only deal with a sub-set of rock types for which a simplified system works

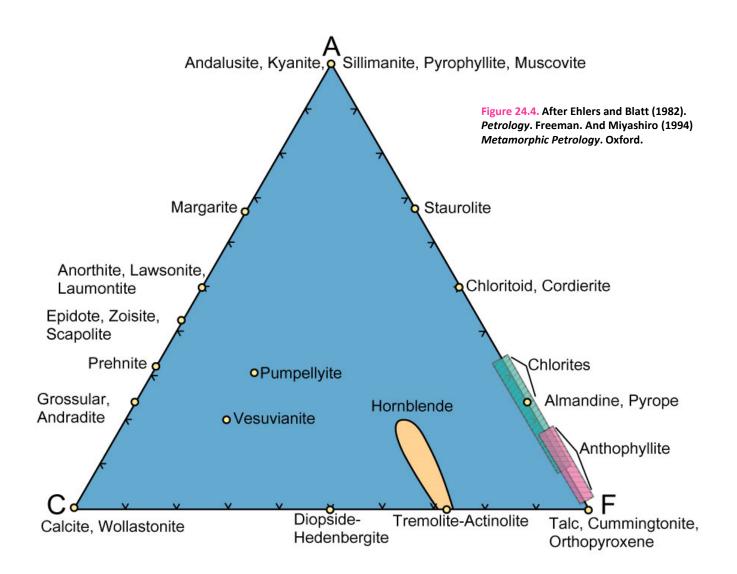
4) Use projections

I'll explain this shortly

The phase rule and compatibility diagrams are rigorously correct when applied to complete systems

- A triangular diagram thus applies rigorously only to *true* (but rare) 3-component systems
- If drop components and phases, combine components, or project from phases, we face some issues:
 - Gain by being able to graphically display the simplified system, and many aspects of the system's behavior become apparent
 - Lose a rigorous correlation between the behavior of the simplified system and reality

- Illustrate metamorphic mineral assemblages in mafic rocks on a simplified 3-C triangular diagram
- Concentrate only on the minerals that appeared or disappeared during metamorphism, thus acting as indicators of metamorphic grade



The three pseudo-components are all calculated on a molecular basis:

 $A = AI_2O_3 + Fe_2O_3 - Na_2O - K_2O$ $C = CaO - 3.3 P_2O_5$ F = FeO + MgO + MnO

The ACF Diagram $A = Al_2O_3 + Fe_2O_3 - Na_2O - K_2O$

Why the subtraction?

- Na and K in the average mafic rock are typically combined with Al to produce Kfs and Albite
- In the ACF diagram, we are interested only in the other Al-bearing metamorphic minerals, and thus only in the amount of Al₂O₃ that occurs in excess of that combined with Na₂O and K₂O (in albite and K-feldspar)
- Because the ratio of Al₂O₃ to Na₂O or K₂O in feldspars is 1:1, we subtract from Al₂O₃ an amount equivalent to Na₂O and K₂O in the same 1:1 ratio

 $C = CaO - 3.3 P_2O_5$

F = FeO + MgO + MnO

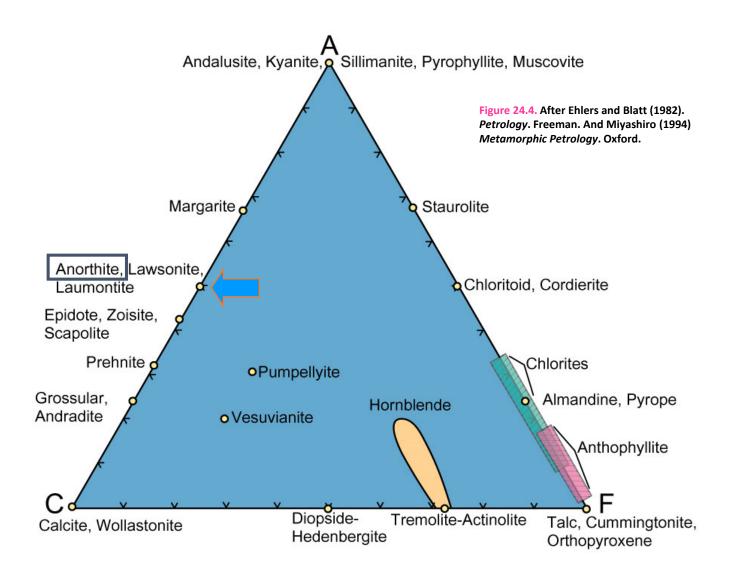
By creating these three pseudo-components, Eskola reduced the number of components in mafic rocks from 8 to 3

- Water is omitted under the assumption that it is perfectly mobile
- Note that SiO₂ is simply ignored
 - We shall see that this is equivalent to projecting from quartz
- In order for a projected phase diagram to be truly valid, the phase from which it is projected must be present in the mineral assemblages represented

An example:

Anorthite $CaAl_2Si_2O_8$ A = 1 + 0 - 0 - 0 = 1, C = 1 - 0 = 1, and F = 0Provisional values sum to 2, so we can normalize to 1.0 by multiplying each value by ½, resulting in A = 0.5 C = 0.5F = 0

Where does Ab plot? Plagioclase?



Plotting Rules

- 1. Obtain molecular proportions of oxides comprising the rock or mineral.
- Component A equals the molecular proportions of Al₂O₃ minus Na₂O minus K₂O. This formulation assumes that plagioclase is present and any K₂O and Na₂O present is in the form of K-feldspar (K₂O · Al₂O₃ · 6SiO₂) and albite (Na₂O · Al₂O₃ · 6SiO₂) dissolved in the plagioclase; note the 1:1 ratios of K₂O and Na₂O to Al₂O₃ in these alkali feldspars. If muscovite or biotite are present, this calculation for A is invalid
- 3. Component C equals the molecular proportions of CaO minus $\frac{10}{3}$ (P₂O₅) minus CO₂. These subtractions allow for the presence of ideal apatite (10CaO \cdot 3P₂O₅) and calcite (CaO \cdot CO₂).

- 4. Component F equals the molecular proportions of FeO plus MgO. In some cases molecular proportions equal to TiO_2 and Fe_2O_3 may be subtracted to allow for the presence of ideal ilmenite (FeO \cdot TiO₂) and magnetite (FeO \cdot Fe₂O₃).
- 5. The sum of A + C + F is found and the percentages of A, C, and F calculated for plotting

A typical ACF compatibility diagram, referring to a specific range of P and T (the kyanite zone in the Scottish Highlands)

