

## Tracking biases in the regular echinoid fossil record: The case of *Paracentrotus lividus* in recent and fossil shallow-water, high-energy environments

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### ABSTRACT

The poor fossil record of regular echinoids is commonly attributed to the fact that they preferentially live as epibenthic grazers on hard rocky substrates in nearshore environments with high energy conditions, where the preservation potential of their tests is generally low. In addition, they commonly occur in the fossil record as fragmented material that hampers identification to low taxonomic level. Herein, a comparative analysis of fossil and recent remains of the regular sea urchin *Paracentrotus lividus* has been conducted, in order to evaluate the source of taphonomic and taxonomic biases affecting regular echinoid remains.

Echinoid remnants are recognized in great abundance both in the Upper Pleistocene marine deposit of Is Mesas and recent littoral environments of Sa Mesa Longa, both in Sardinia (Italy). They were rigorously categorized and analysed with respect to relative abundance, biostratinomic signatures and constructional detail of the skeleton using scanning electron microscopy. *Paracentrotus* remnants mainly consist of test fragments, isolated plates and fragmented spines. Plates of the apical system and elements of the jaw apparatus were also found.

Echinoid remnants in the Pleistocene deposit reflect those in recent environment, with respect to relative abundance and biostratinomic features, and are useful tools for paleoenvironmental and paleoecological reconstruction, even in shallow-water, high-energy environments. The abundance of large test fragments of *Paracentrotus lividus* both in the Pleistocene record and recent environments is due to a relatively robust test and stereomic interlocking. Difference in relative abundance of echinoid remains and their distribution in sediment fractions (> 2 mm and 1-2 mm) are due to different sampling methods, but also indicate the different preservation potential among various test elements. Herein, the identification of echinoid remains at a low taxonomic level can be assessed even at a very-coarse sand-size fraction (1-2 mm). This study demonstrates that regular echinoid fragments can be identified to a low taxonomic level when very

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specific characters are preserved, and their analysis may help to minimize the taxonomic bias affecting the regular echinoid fossil record.

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## INTRODUCTION

### Skeletal Architecture, Ecology and Preservation

The evolutionary history of echinoids and their paleoecology as deduced from the fossil record is biased by several factors. Preservational differences among echinoid taxa depend on skeletal architecture and organic connective tissues, lifestyles and the complex interplay of taphonomic processes affecting echinoid tests and spines (e.g., Smith, 1984; Allison, 1990; Kidwell and Baumiller, 1990; Greenstein, 1991, 1993a, 1993b, 1995; Nebelsick and Kampfer, 1994; Nebelsick, 1995; Moffat and Bottjer, 1999; Banno, 2008; Smith and Rader, 2009).

All echinoids possess a multi-plated mesodermal skeleton constructed by conjoined high-Mg calcite plates. Each skeletal element consists of a three-dimensional meshwork of mineralized trabeculae, the stereom, and of an organic component, represented by the connective tissue filling the pores of the stereom, the stroma (Smith, 1980, 1984; Kroh and Nebelsick, 2010; Grun and Nebelsick, 2018b). Test stability and thus preservation potential is determined by the degree of interlocking by stereom trabeculae of coronal plates at sutures and the presence of numerous collagenous fibers, which cross the sutures and bind the adjacent plates together (Smith, 1984; Donovan, 1991; Ellers et al., 1998; Grun and Nebelsick, 2018a; Grun et al., 2018a, b).

There is a wide spectrum of test architectures among echinoids and the rigidity of the test differs highly among evolutionary lineages (Smith, 2005, 2007). Among regular echinoids, camarodonts possess tests relatively more resistant to disarticulation having some degree of stereomic interlocking at plate suture faces (Régis, 1977, 1979; Telford, 1985; Dafni, 1986; Kidwell and Baumiller, 1990; Greenstein, 1991, 1995; Ellers et al., 1998). By contrast, echinothurioids and diadematoids

have tests with no stereom connection between the plates, which are embedded in a meshwork of collagenous fibers and dissociate rapidly after soft tissue decay, especially when subjected to post-mortem transportation and reworking (Smith, 1984; Greenstein, 1989, 1991, 1993a, 1993b), leading to low preservation potentials and a poor fossil record (Smith, 1984; Greenstein, 1991, 1995; Ellers et al., 1998). Cedaroid echinoids fall between these extremes (Kidwell and Baumiller, 1990; Greenstein, 1991).

Among irregular echinoids, clypeasteroids have the most extensive stereomic interlocking across plate boundaries and also possess internal supports, represented by different structural solutions, such as pillars in *Clypeaster* and buttresses in *Echinocyamus*, which conjoin the oral and aboral sides, leading to a high preservation of tests even in high energy settings (Kier 1977; Seilacher, 1979; Nebelsick and Kroh, 2002; Mitchell et al., 2006; Belaústequi et al., 2012; Mancosu and Nebelsick, 2013, 2017; Grun and Nebelsick, 2018a; Grun et al., 2018a,b; Nebelsick, 2020).

Besides differences in constructional morphology, preservation potential is related to differences in life habit and ecology among echinoids and, in turn, taphonomic processes affecting the echinoid test (Ernst and Seibertz, 1977; Kier, 1977; Smith, 1984; Greenstein, 1993a; Nebelsick, 1996). Irregular echinoids, since their appearance in the Early Jurassic, diversified as deposit feeders (Kier, 1982; Barras, 2008) often buried within mobile substrates in areas of net sedimentation where the preservation potential is generally higher (Smith, 1984; Nebelsick, 1996).

In contrast to irregular forms, regular echinoids live as epibenthic grazers in environments where their preservation potential is generally low and are less likely to be represented in the fossil record (Ernst and Seibert, 1977; Kier, 1977; Smith, 1984; Kidwell and Baumiller, 1990).

Regular echinoids predominate in rocky intertidal shores and sheltered back-reef and peri-reefal shallow water environments. Rocky shores, and their associated faunas, are poorly preserved in the rock record due to the prevalence of high energy conditions and active erosion or non-deposition of sediments (see also Smith, 1984; Greenstein 1993a; Sørensen and Surlyk, 2015).

Regular echinoids living in these environments frequently suffer lethal and non-lethal damage and breakage that can be caused by impact from rolling cobbles and against rocks, where the echinoids can be dislodged, especially during storms (Ebert, 1968). Additionally, bioerosion and predation (Sievers and Nebelsick, 2018 and references cited therein) and post-mortem physical reworking and transportation leads to fragmentation of sea urchin tests resulting in isolated and broken plates and spines.

Actuopaleontological studies based on both experimental and field observations shed light on the taphonomic processes affecting regular echinoid tests. Experimental taphonomic investigations on regular echinoids conducted by Kidwell and Baumiller (1990) and Greenstein (1991, 1995) reveal that reworking can induce rapid disarticulation of spines and precludes the preservation of their fragile complete tests. The investigated echinoids, pertaining to cidaroids, diadematoids and camarodonts, show a similar sequence of skeletal disarticulation ranging from intact to completely disarticulated tests. There are, however, differences in the timing of disarticulation phases depending on skeletal architectures, namely the distribution of collagenous connective tissues, the degree of stereomic inter-digititation, and temperature and other ambient environmental conditions. Field investigations in recent environments are consistent with these results and document that the distribution of regular echinoid skeletal remains in shallow water environments accurately reflects living population when not only whole tests, but also isolated fragments are considered (Greenstein, 1989, 1992, 1993b; 1995; Nebelsick, 1992a, b, 1996; Lewis et al., 2000; Schein and Lewis, 2001; Dynowski, 2012).

The echinoid fossil record can furthermore be influenced by sampling methods and taxonomic methodology. Regular echinoids are often preserved as fragments meaning that taxonomists may have difficulty in identifying them at low taxonomic level. Taxonomic studies are largely restricted to well-preserved fully articulated specimens, due to the difficulty in the identification of

taxa based on fragmentary material (e.g., Gordon and Donovan, 1992; Greenstein, 1991, 1993a, b). Some studies, however, have demonstrated the importance of including echinoid fragments for taxonomy (e.g., Smiser, 1933; Thompson et al., 2019), paleoecology (e.g., Gordon and Donovan, 1992; Thompson et al., 2019) and actualistic investigations (e.g., Nebelsick, 1992a, b; Schein and Lewis, 2001; Dynowski, 2012).

Herein, within an ongoing evaluation of the preservation of echinoids and the nature of their fossil record, a comparative study of the remains of the regular sea urchin *Paracentrotus lividus* from both the Late Pleistocene (Tarantian, MIS 5e) marine rocks and modern shore deposits of Sa Mesa Longa Beach, both in Sardinia (Italy), has been carried out. This study aims to rigorously quantify these remains, show in detail how skeletal architectures and morphological features and the presence of collagenous fibers affect specific preservation patterns, evaluate the source of biostratigraphic and taxonomic biases affecting the remains, and reiterate the use of echinoid test fragments in paleoecological reconstructions.

### **The Camarodont Echinoid *Paracentrotus lividus***

The shallow water echinoid fauna in the Mediterranean is strictly controlled by substrate and other ecological factors (e.g., Koehler, 1927; Tortonese, 1965; Ernst et al., 1973; Riedl, 1983). Hard substrates in shallow water are characterized by the regular echinoids *Paracentrotus lividus*, *Arbacia lixula* and *Sphaerechinus granularis*, while *Psammechinus miliaris* is often associated with sea grass meadows. Shallow water soft substrates are characterized by *Spatangus purpureus*, *Echinocardium* (diverse species), *Ova canaliferous* and *Echinocyamus pusillus* which are controlled, among others, by the grain size of the sediment in which they live.

*Paracentrotus lividus* (Lamarck, 1816) is a camarodont parechinid regular echinoid, which has a general moderately low hemispherical test. Ambulacral plates are polygeminate with 5 to 6 pore-pairs to a plate arranged in arcs and one primary imperforate tubercle to each compound plates. Interambulacral plates have larger primary tubercles and one or more secondary tubercles on either side at the ambitus. Spines are relatively short, about half the test diameter in length, with the base bearing a crenulated milled ring and collar, shaft and tip with fine longitudinal striation (Smith and Kroh, 2011).

*Paracentrotus lividus* is one of the most common regular echinoids of the Mediterranean and eastern Atlantic coasts of Europe and north Africa (Fernandez et al., 2012; Deli et al., 2017 and references therein) and inhabits littoral and shallow sublittoral, temperate to subtropical marine environments (Ernst et al., 1973; Turon et al., 1995; Boudouresque and Verlaque, 2007, 2020; Girard et al., 2011). *P. lividus* has a broad ecological distribution (Tortonese, 1965; Régis, 1979; Jacinto and Cruz, 2012), ranging from intertidal rock pools, coastal lagoons, and most commonly, in the shallow marine subtidal nearshore environment down to 20 m depth (Mortensen, 1927, 1943; Tortonese, 1965; Ebbling et al., 1966; Verlaque, 1984, 1987; Fernandez et al., 2001, 2012).

### Pleistocene Marine Deposits of Sardinia

Along the coasts of Sardinia, Upper Pleistocene marine deposits, related to the sea level highstand during the last interglacial period, extensively crop out and are found at heights up to 10 m above the current sea level (Ulzega and Ozer, 1982; Ulzega and Hearty, 1986; Carboni and Lecca, 1985; Barca et al., 2005; Lecca and Carboni, 2007; Carboni et al., 2014; Barca et al., 2016, 2017). These marine deposits pertain to the Calamosca (PVM1) subsynthem of the Portovesme (PVM) synthem (Barca et al., 2005, 2016, 2017). One of the best-known Late Pleistocene (Tarantian, MIS

5e) marine deposit is documented in the Bay of Cagliari (southern Sardinia), along the coast from Calamosca to Is Mesas (Latitude 39° 11'00"N, Longitude 9° 08'55"E) (Figure 1A), that represents the type section of the Thyrrenian Stage (Issel, 1914). This deposit lies on an abrasion platform cut onto the Upper Miocene (Tortonian) Limestones of Cagliari and consist of conglomerates and sandstones (Barca et al., 2005, 2017) containing a rich fossil invertebrate fauna and flora consisting of coralline algae.

The fossil content was first sampled by Lovisato and subsequently studied by Issel (1914) who established the “Piano Tirreno” (Tyrrhenian Stage) for the fossil-bearing deposit, which includes the characteristic and emblematic gastropod *Persististrombus latus*, previously known as *Strombus bubonius*. Subsequent studies improved the knowledge of the fossil fauna with respect to the taxonomy (Fontana Zanco, 1933; Comaschi Caria and Pastore, 1959; Comaschi Caria, 1968; Spano, 1982) and its paleoecological and paleoclimatic significance (Spano, 1993a, b). The fossil fauna is dominated by mollusks and the colonial scleractinian coral *Cladocora caespitosa* (Spano, 1993a, b). Together with a number of species commonly living in the Mediterranean today, such as *Arca noae*, and *Patella ferruginea*, these belong to a group of taxa traditionally called “Senegalese guests” or “warm visitors”, including *Persististrombus latus*,

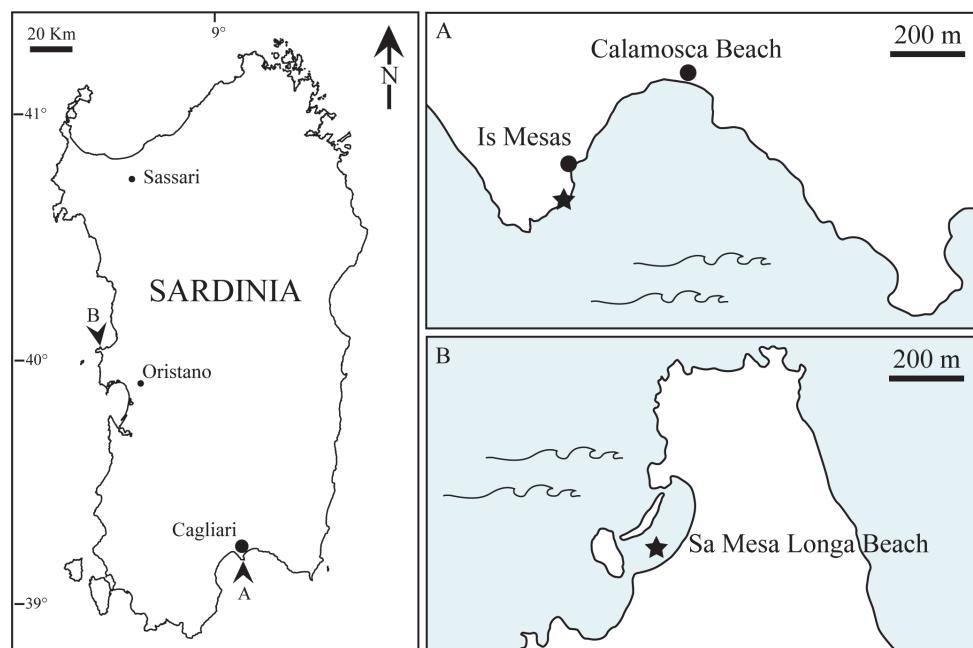


FIGURE 1. Location of A, Late Pleistocene deposit of Is Mesas and B, Sa Mesa Longa Beach, Sardinia, Italy.

*Conus ermineus* (syn. *C. testudinarius*) and *Gemophos viverratus*. These taxa still thrive in the tropics, especially in the warm coastal regions of West Africa, but are currently extinct in the Mediterranean (see Sessa et al., 2013 and references cited therein; Brunetti and Della Bella, 2014).

The age of these deposits has been debated (see Carboni et al., 2014 and references cited therein). Radiometric data using U/Th ratio on *Cladocora caespitosa* from Calamosca provided an age of 138±8 ka (Ulzega and Hearthy, 1986). More recently, U/Th dating on *Cladocora caespitosa* from the same locality provided an age of 122.3±5 ka for these deposits (Barca et al., 2005) which therefore pertain to MIS 5e (Marine Isotopic Stage 5e, ~130 ka to ~116 ka), the warmest interval of the Interglacial period (Rohling et al., 2007). This age is consistent with U/Th dating on *Caldochora caespitosa* colonies from the Upper Pleistocene deposits of San Giovanni di Sinis (Central-western Sardinia), the age of which is 128±4 ka (Carboni et al., 2014).

### Recent Collections from Sa Mesa Longa Beach

The Sa Mesa Longa Beach (Latitude 40° 02'47"N, Longitude 8°23'57"E) on the central west coast of Sardinia is located on the Capo Mannu Cape (Oristano Province) (Figure 1B). Sa Mesa Longa ("the long table" in Sardinian language) is a pocket beach extending for approximately 1200 m between two headlands and is characterized in its central part by a salient shape due to a rocky reef located approximately 100 m from the baseline. Two main wave directions are consistently predominant in magnitude and frequency: 310–340°N and 230–260°N, corresponding to the Maestrale (from northwest) and Libeccio (from southwest) winds in the Mediterranean Sea, respectively (Sulis and Annis, 2014). A variety of marine substrates occurs, including rocks and mobile sediments, which mainly consist of very coarse to coarse-grained bioclastic sands with pebbles and cobbles. Patches of the seagrass *Posidonia oceanica* also occur. Coralline algae rhodoliths are also present.

### MATERIAL AND METHODS

Field and laboratory investigations at Is Mesas were conducted with respect to paleontology, taphonomy and sedimentology. The frequency of echinoid remains was estimated per rock surface. The degree of close-packing determined as dense, loose or dispersed follows Kidwell and Holland (1991). The density of echinoid remains on the bedding plane was estimated using a 50 cm x

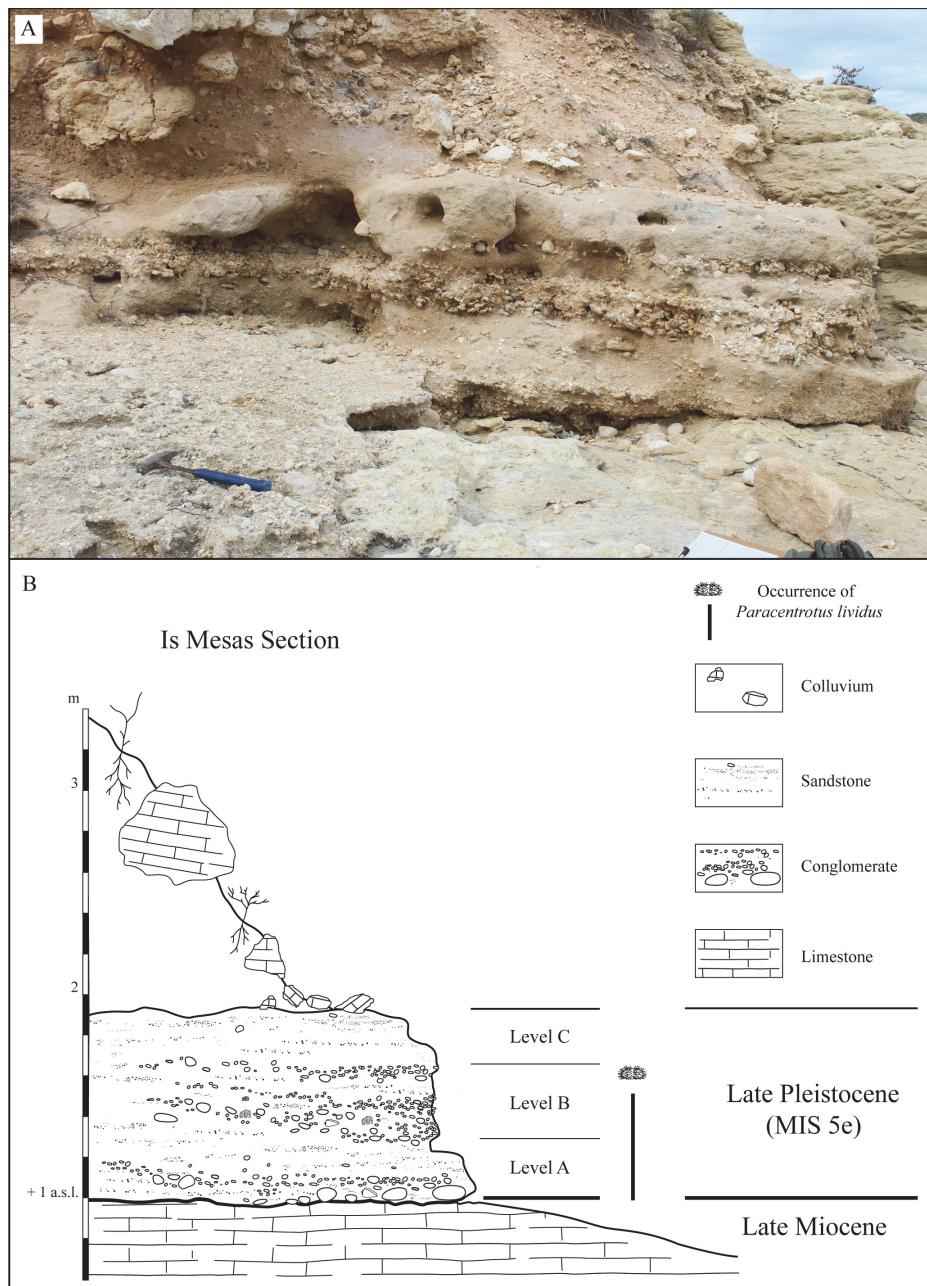
50 cm quadrat. Rock samples of poorly consolidated coarse-grained sandstones were collected from Level A of the Pleistocene sedimentary deposit (Figure 2) in order to recover fossil remains. Approximately 8 kg of the soft sedimentary rocks was sampled and wet-sifted using sieves with 2 mm and 1 mm meshes. The resulting dried residue was rich in echinoid fragments in both sediment fractions and was inspected using a stereo microscope. Spine remains were picked up directly from rock samples under the microscope to avoid breakage during the sieving process.

Remains of the recent *Paracentrotus lividus* were collected at Sa Mesa Longa Beach (Figure 1B), Central-western Sardinia, on multiple occasions between 2018 and 2020. Echinoid remains were exhaustively searched and collected by manually picking them from the sediment surface by snorkelling in shallow water near-shore environment from the shoreline down to 2 m depth. The remains were recovered on the sediment surface near patches of seagrass (*Posidonia oceanica*) as well as on and around rocky substrates. Additionally, bulk sediment samples were collected by using a hand corer at a depth from 0.5 to 1 m for a total of 2 kg of sediments and processed by the same method adopted for the fossil remains.

Echinoid remains, both fossil and recent, including test and spine remains, were examined with respect to taphonomic signatures, notably fragmentation, abrasion and encrustation. The preservation of test features, such as tubercles on the test surfaces and the crenulation of spine milled ring along with striation of spine shaft, was used to evaluate abrasion.

In this study, the taphonomic bias affecting the echinoid skeletal elements is directly evaluated by comparing the observed and expected abundance of skeletal elements. This was done using the following calculation: (OA/EA) x 100, with OA representing the observed skeletal element abundance in the sample and EA representing the expected skeletal element abundance present within a complete skeleton.

Photographs of the retrieved fossil and recent echinoid remains were taken with a digital camera and a Leica DMS1000 digital microscope at the Department of Chemical and Geological Sciences, Cagliari, Italy. Fossil specimens were cleaned with a hypersonic bath (3 minutes). Scanning electron microscopy (SEM) was performed at the State Museum of Natural History, Stuttgart, Germany. Gold-palladium coating with a thickness of 8 nm was accomplished using a Leica EM ACE200.



**FIGURE 2.** A, Field view of the Upper Pleistocene deposit of Is Mesas. B, Stratigraphic section of the deposit.

Scanning micrographs used a Zeiss EVO LS 15. Photos were enhanced (contrast, brightness), assembled and cut out using Adobe Photoshop (2019) and Adobe Illustrator (2018).

Echinoid classification at and above species level follows Kroh and Smith (2010) and Smith and Kroh (2011). Sutures are categorized according to Smith and Kroh (2011). Mollusk taxonomy follows WoRMS (World Register of Marine Species available at <http://www.marinespecies.org/index.php>). Samples are housed in the Museo di Geologia e

Paleontologia Domenico Lovisato, Università di Cagliari.

## RESULTS

### Sedimentology and Paleontology

The Upper Pleistocene marine deposit studied herein, which crops out in the locality Is Mesas, is approximately 1 m thick and overlies Upper Miocene, Tortonian pale yellow limestones, which pertain to the Calcare di Cagliari (Cagliari Limestones).

The studied deposit can be subdivided into three different levels from the base to top (Figure 2). The first level (Level A in Figure 2) is ca. 40 cm thick and consists of a conglomerate, containing sub-angular to rounded clasts along with pebbles, cobbles and boulders with a maximum size of 20 cm, gradually passing upward into a poorly sorted coarse to very coarse sandstone with abundant bioclasts. Clasts and pebbles consist predominantly of Miocene limestones and show bioerosion traces produced by endolithic bivalves (*Gastrochaenolites*) and clionid sponges (*Entobia*) and encrustation by coralline algae. Macrofossils are very abundant and are represented by the echinoid *Paracentrotus lividus* (Figure 3) and associated malacofauna (Table 1 and Figure 4). The remains of the coral *Cladocora caespitosa* and coralline algal rhodoliths also commonly occur. A single test fragment of the clypeasteroid echinoid *Echinocyamus* was found.

Level B consists of a matrix to clast-supported conglomerate with rounded pebbles and granules with an abundant fossil content similar to that of Level A. This conglomerate is surmounted by a medium to coarse sandstone (Level C) with floating pebbles and rare mollusk remains. The fossil remains in levels A and B are chaotically oriented in plan view and cross section and are not homogeneously distributed within the deposit, ranging from loosely packed and dispersed to very densely packed (sensu Kidwell and Holland, 1991).

Echinoid remains are very abundant (more than 500/m<sup>2</sup> in Level A) and occur in a wide spectrum of preservation (Figures 3, 5 and Table 2-Part 1). They are represented by variously sized test pieces composed of several articulated plates to isolated ambulacrals and interambulacrals plates showing fragmentation within plates (interplate fragmentation) as well as along plate boundaries (disarticulation or intraplate fragmentation). Test fragments composed of articulated plates are often confined to single rows or in the case of larger fragments to the offset paired rows of ambulacra or interambulacra. Genital and ocular plates of the apical system also occur sporadically.

Test fragments of *Paracentrotus lividus* show differential preservation (Figures 3, 5 and Table 2-Part 1). In better-preserved specimens, general morphological features of the surface such as primary and secondary tuberculation as well as paired ambulacrals pores can be readily recognized. Generally, the edges of the plates are abraded and rounded. The stereom microstructure is largely obscured by recrystallization, but can be recog-

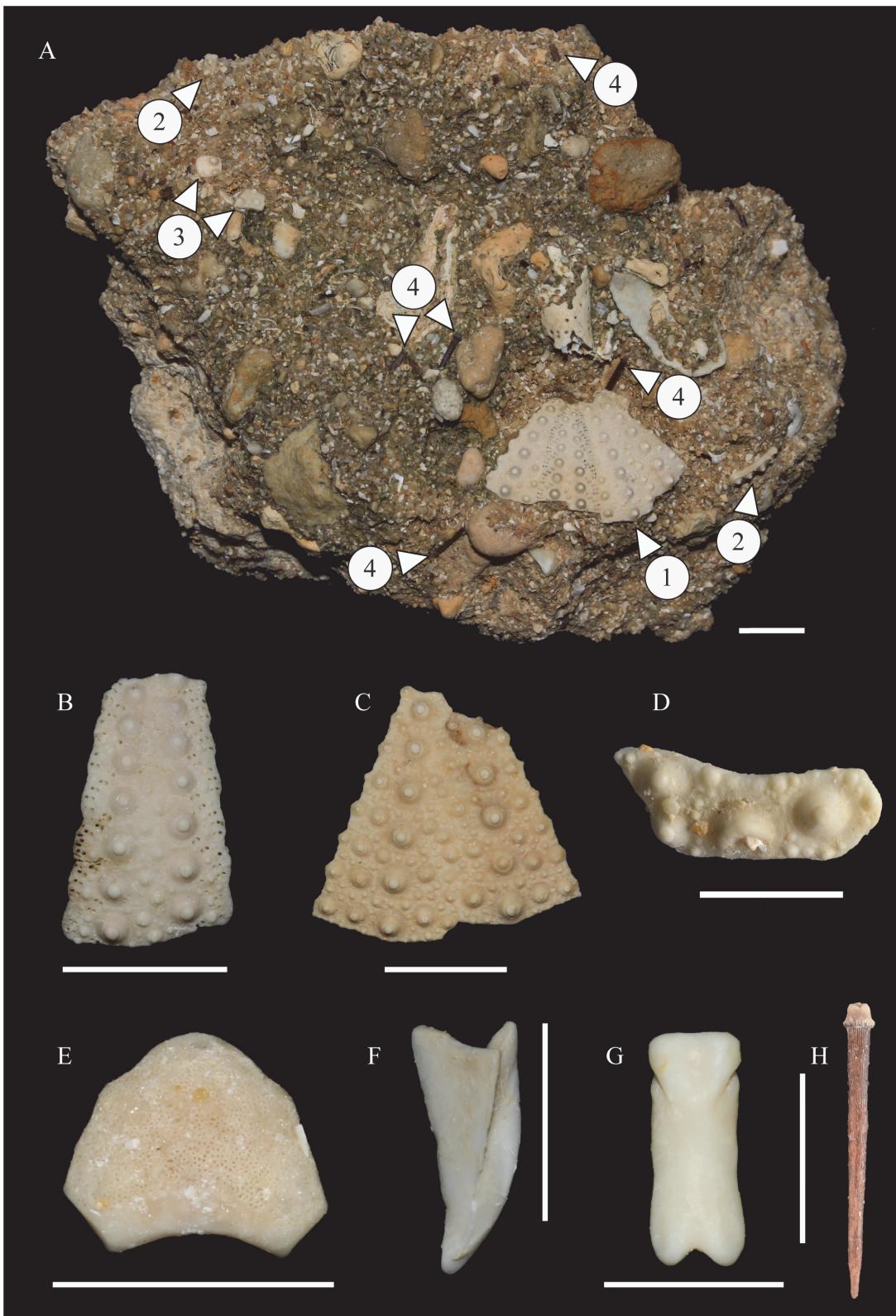
nized on some surfaces. The interdigitating knobs on the plate edges, which can be easily seen on recent plates boundaries (see below), cannot be recognized. The surface morphology shows some pitting and abrasion, and some tubercles have been fragmented. In one case, a cavity of the plate interior perhaps attributable to bioerosion is present.

Poorly preserved fragments show highly abraded surfaces on which the stumps of larger tubercles can be recognized. The smaller secondary tubercles have been abraded. Plate boundaries exposed by disarticulation as well as surfaces produced by fragmentation can be highly abraded.

Isolated Aristotle's lantern elements are common and mostly consist of hemipyramids and sub-ordinately rotulae and teeth. Epiphyses and compasses also occur sporadically. Fragmentation and abrasion are common (Table 2-Part 1). Spines are mostly fragmented although complete spines do occur, some of which reveal regenerated tips. Well-preserved spines show the base, milled ring collar and striations, which can be partially or totally masked in poorly preserved specimens. Encrustation on echinoid remains, tentatively assigned to coralline algae, is generally low and mostly restricted to spine remains. Isolated plates and spines often show grain indentations, which preserve the outlines of the terrigenous grains often with a smooth cement layer on the surfaces. In some cases, a secondary, calcitic overgrowth within the plates has occurred filling in the pores of the stereom.

Different degrees of fragmentation and abrasion were also observed on mollusk remains. Gastropods consist of complete shells with well-preserved surface details, some of which retain trace of their original color pattern, and variously sized shell fragments which can be reduced to the columellar pillar. Bivalves show a similar taphonomic gradient from well-preserved and articulated valves to disarticulated and fragmented, highly abraded remains. Mollusk shells show the presence of circular boreholes (*Sedilichnus*) and bioerosion (*Entobia*) attributed to clionid sponges.

The remains of the scleractinian coral *C. caespitosa* are preserved as portions of the entire colony to fragments of single branches, which show signs of abrasion, bioerosion and encrustation. Other major biotic constituents are common coralline algae rhodoliths, which range from 2 cm to 10 cm maximum length and are dominated by subspherical to ellipsoidal shape, although more flattened examples are also present. Growth forms



**FIGURE 3.** Remains of *Paracentrotus lividus* from Level A of the Is Mesas deposit (Late Pleistocene). A, 1) Fragment of test consisting of ambulacral and interambulacral plates still sutured together and showing interplate fracturing; 2) Isolated plates; 3) Fragments; 4) Spines. B, Large portion of ambulacral column showing intraplate fragmentation. C, Large portion of interambulacral column. D, Fragmented interambulacral plate. E, Madreporite. F, Hemipyramid. G, Rotula. H, Complete spine. A-C, H Scale bars equal 1 cm. E-G Scale bars equal 0.5 cm.

**TABLE 1.** Fauna of the Upper Pleistocene deposit of Is Mesas recognized in this study.

TAXON	HABITAT-SUBSTRATE-BIOCOENOSIS	REFERENCES	ABBREVIATIONS
Phylum Echinodermata			SupraL = Supralittoral
Class Echinoidea			L = Littoral
Order Camarodontia			IL = Infralittoral
Family Parechinidae			ML = Mesolittoral
	<i>Paracentrotus lividus</i>	IL, H (PA), P 1, 2, 17, references cited herein	CL = Circalittoral
			SL = Sublittoral
Phylum Mollusca			IN = Intertidal
Class Gastropoda			H = Hard substrates
Order ?			S = Soft sediments
Family Patellidae			P = <i>Posidonia</i> meadows
	<i>Patella caerulea</i>	IL, H (PA) 31, 33, 34	PA = Photophilic algae
	<i>Patella rustica</i>	ML, IL, H (PA) 2, 31, 33, 34	C = Coralligenous
Order Lepetellida			SABEL = <i>Sabellaria alveolata</i> reef
Family Fissurellidae			SA = Sciaphilic algae
	<i>Diodora gibberula</i>	IL to CL H (PA), P, C 17, 18, 31, 33, 34	Rh = Rhodolith beds
	<i>Diodora graeca</i>		
Order Neogastropoda			Literature list
Family Conidae			1) Mortensen, 1943
	<i>Conus ermineus</i>	L to SL, H, S 19, 28	2) Pérès and Picard, 1964
	<i>Conus ventricosus</i>	IL to CL, H (beneath boulders, PA), S 5, 11, 31, 33, 34	3) Lipkin and Safriel, 1971
Family Columbellidae			4) Conti and Rossini, 1985
	<i>Columbella rustica</i>	IL, H (beneath boulders, PA), P 2, 5, 6, 17, 27, 29, 31, 33, 34	5) Taylor, 1987
Family Muricidae			6) Russo et al., 1991
	<i>Hexaplex trunculus</i>	IN, IL to CL, H (PA), S 5, 16, 17, 33, 34	7) Hall-Spencer, 1998
Family Pisaniidae			8) Peharda et al., 2002
	<i>Gemophos viverratus</i>	IL, H, S 12, 28, 30	9) Rubio and Rolan, 2002
Family Nassariidae			10) Russo et al., 2002
	<i>Tritia incrassata</i>	IL, H (PA, SABEL), S, P 6, 11, 27, 29, 40	11) Antoniadou et al., 2005
Order Trochida			12) Rolán, 2005
Family Turbinidae			13) Sivan et al., 2006
	<i>Bolma rugosa</i>	IL to CL, H (PA, SA), P, C 11, 17, 18, 25, 31, 33, 34, 40	14) Spadini, 2006
Family Trochidae			15) Trigui El-Menif et al., 2007
	<i>Jujubinus exasperatus</i>	IL to CL, H (PA), P, C 4, 6, 11, 17, 18, 22, 25, 27, 31, 33, 34, 38, 40	16) Vasconcelos et al., 2007
	<i>Clanculus jussieui</i>	IL to CL, P, C 9, 14	17) Beqiraj et al., 2008
	<i>Clanculus corallinus</i>	IL to CL, H (PA), P, C 4, 11, 17, 18, 31, 33, 34	18) Casellato and Stefanon, 2008
	<i>Phorcus turbinatus</i>	ML, IL, H (PA) 2, 31, 33, 34	19) Rosenberg et al., 2009

**TABLE 1** (continued).

	TAXON	HABITAT-SUBSTRATE-BIOCOENOSIS	REFERENCES	ABBREVIATIONS
Order	Caenogastropoda ?			20) Sciberras et al., 2009
Family	Cerithiidae			21) Rufino et al., 2010
	<i>Cerithium lividulum</i>	IL, H (PA), S	31, 34	22) Albano and Sabelli, 2011
	<i>Cerithium vulgatum</i>	IL to CL, H, S, P	2, 11, 17, 25, 31, 33, 34	23) Gutierrez-Mas, 2011
	<i>Bittium latreillii</i>	IL to CL, H (PA, Rh), P, C	6, 10, 11, 20, 25, 38	24) Öztürk, 2011
Family	Epitoniidae			25) Albano and Sabelli, 2012
	<i>Epitonium clathrus</i>	IL to CL, H (PA), S	31, 34	26) Župan et al., 2012
Family	Triphoridae			27) Belgacem et al., 2013
	Triphoridae ind.			28) Sessa et al., 2013
Order	Littorinimorpha			29) Urra et al., 2013
Family	Triviidae			30) Brunetti and Della Bella, 2014
	<i>Trivia arctica</i>	IL to CL	31, 33	31) Öztürk et al., 2014
Family	Littorinidae			32) Pitacco et al., 2014
	<i>Melarhaphe neritoides</i>	SupraL to IL, H (PA)	2, 3, 5, 31, 33, 34	33) Alf and Haszprunar, 2015
Class	Bivalvia			34) Poursanidis et al., 2016
Order	Arcida			35) Chimienti et al., 2017
Family	Arcidae			36) Kersting et al., 2017
	<i>Arca noae</i>	IL to CL, H (PA, SABEL, SA), P	8, 17, 18, 26, 27, 31, 33, 34, 40	37) Morton and Puljas, 2017
Family	Glycymerididae			38) Donnarumma et al., 2018
	<i>Glycymeris nummaria</i>	IL to CL, S	2, 13, 23	39) Giacobbe and Renda, 2018
Family	Noetiidae			40) Casoli et al., 2019
	<i>Striarca lactea</i>	IL, H, P, C	15, 22, 25	
Order	Cardiida			
Family	Cardiidae			
	<i>Acanthocardia tuberculata</i>	IL to CL, S (SFBC), P	2, 8, 11, 17, 31, 33, 34	
Order	Carditida			
Family	Carditidae			
	<i>Cardita calyculata</i>	ML, IL, H (PA), S, P	2, 27, 29, 31, 33, 34	
Order	Venerida			
Family	Veneridae			
	<i>Irus irus</i>	ML, IL to CL, H (PA, SABEL), C, S, P	2, 11, 27, 31, 33, 40	
	<i>Venerupis geographica</i>	IL, P	29	
	<i>Clausinella fasciata</i>	IL, S	7, 21, 33	
Order	Lucinida			
Family	Lucinidae			
	<i>Lucinella divaricata</i>	IL to CL, S, P	11, 21, 35, 38	
	<i>Ctena decussata</i>	IL, P, S	35, 38, 39	
Order	Ostreida			
Family	Ostreidae			
	<i>Ostrea</i> sp.			

**TABLE 1** (continued).

TAXON	HABITAT-SUBSTRATE-BIOCOENOSIS	REFERENCES	ABBREVIATIONS
Order Mytilida			
Family Mytilidae			
	<i>Mytilaster minimus</i> ML, IL, H, (PA), C	2, 18, 31, 33, 34, 37, 40	
Order Pectinida			
Family Anomiidae			
	<i>Anomia ephippium</i> IL, P	29	
Class Scaphopoda			
Order Dentaliida			
Family Dentaliidae			
	<i>Antalis vulgaris</i> IL to CL, S, P, C	2, 17, 24, 25, 31, 33, 34	
	<i>Antalis inaequicosta</i> IL to CL, H (C), S, P	17, 24, 31	
Phylum Cnidaria			
Class Anthozoa			
Order Scleractinia			
Family Scleractinia incertae sedis			
	<i>Cladocora caespitosa</i> L to shallow SL, H, S, P	17, 32, 36	

are dominated by the presence of encrusting thalli and low protuberances.

#### Recent *Paracentrotus lividus* Remains

In Sa Mesa Longa Beach (Figure 6A) remains of *Paracentrotus* occur abundantly on and within the sediments, around rocky bottoms on which living specimens commonly occur (Figure 6B-H). Other echinoids are represented by very rare remains of the regular echinoid *Arbacia lixula* (Figure 6I) along with highly abundant tests of the minute clypeasteroid echinoid *Echinocyamus pusillus* (Figure 6J).

The remains of *Paracentrotus* from Sa Mesa Longa Beach occur in a broad spectrum of preservation (Table 2-Part 2 and 3; Figure 6B-H) similar to that of the studied fossil counterparts. They mainly consist of isolated coronal elements, most of which belonging to ambulacrals and interambulacrals plates (see Table 2-Part 2 and 3). Fragments consisting of several plates still sutured together are also frequent. Genital and ocular plates of the apical system were also collected, but are relatively rare. Additionally, complete specimens lack spines, the apical system and the Aristotle's lantern. Fragmented remains representing large parts of the test are present, though rare. Intraplate fragmentation along with surface abrasion occur very commonly (see Table 2-Part 2 and 3). Both complete denuded specimens and test fragments show

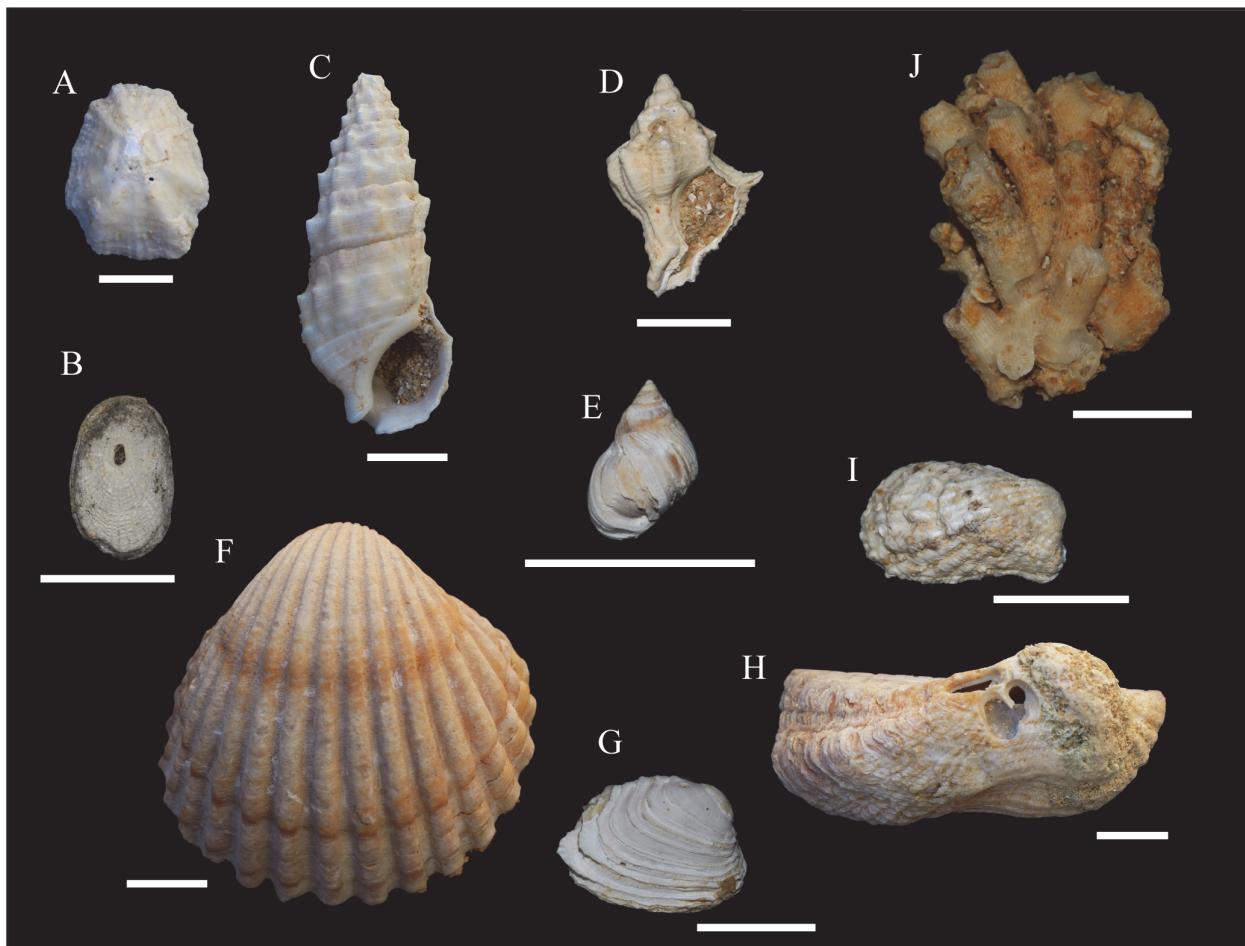
encrustation by polychaete serpulids, including the genus *Spirorbis* and coralline algae (Figure 6B-E).

Numerous disarticulated elements of the Aristotle's lantern were recovered and include mostly hemipyramids and subordinately rotulae and teeth. Epiphyses and compasses also occur sporadically. Lantern elements are commonly affected by fragmentation and abrasion while encrustation is low and restricted to hemipyramids and teeth. Spines also abundantly occur. Their state of preservation is also similar to those in the fossil deposits ranging from pristine and complete, which were rarely found, to fragmentary and abraded. Some specimens show regenerated tips. Spine remains also show encrustation by coralline algae.

The very rarely encountered remains of *Arbacia lixula* consist exclusively of test fragments and isolated plates. The preservation of clypeasteroid *Echinocyamus pusillus* ranges from complete pristine specimens to fragmented and abraded ones.

#### Morphology, Test Thickness and Plate Sutures of *Paracentrotus lividus*

The test of *Paracentrotus* is moderately thick with thickness decreasing slightly from near the apex to the ambitus. The ambulacrals plates are usually thicker than adjoining interambulacrals plates at the ambitus. Both ambulacrals and interambulacrals plates are thicker at their margins. The extent of plate thickening is greater along the radial



**FIGURE 4.** Associated fauna from levels A and B of the Is Mesas deposit (Late Pleistocene). A, *Patella caerulea*. B, *Diodora gibberula*. C, *Cerithium vulgatum*. D, *Hexaplex trunculus*. E, *Melarhaphe neritoides*. F, *Acanthocardia tuberculata*. G, *Irus irus*. H, *Arca noae*. I, *Cardita calyculata*. J, *Cladocora caespitosa*. Scale bars equal 1 cm.

sutures, particularly along the perradial sutures between ambulacral plates, which form a radial ridge when viewed from inside the test. The peristomial margin is thick with the perignathic girdle forming a reinforcing palisade just inside the peristomial edge. Plates are interconnected along sutures by collagenous fibers, crossing plate boundaries. Galleried stereom (sensu Smith, 1980) occurs at the sutures between interambulacral plates (Figure 7). This stereom, which possesses long parallel galleries running in one direction only and shows no pore alignment perpendicular to this direction, is functionally associated with long bundles of collagen fibers at the sutures (Smith, 1980).

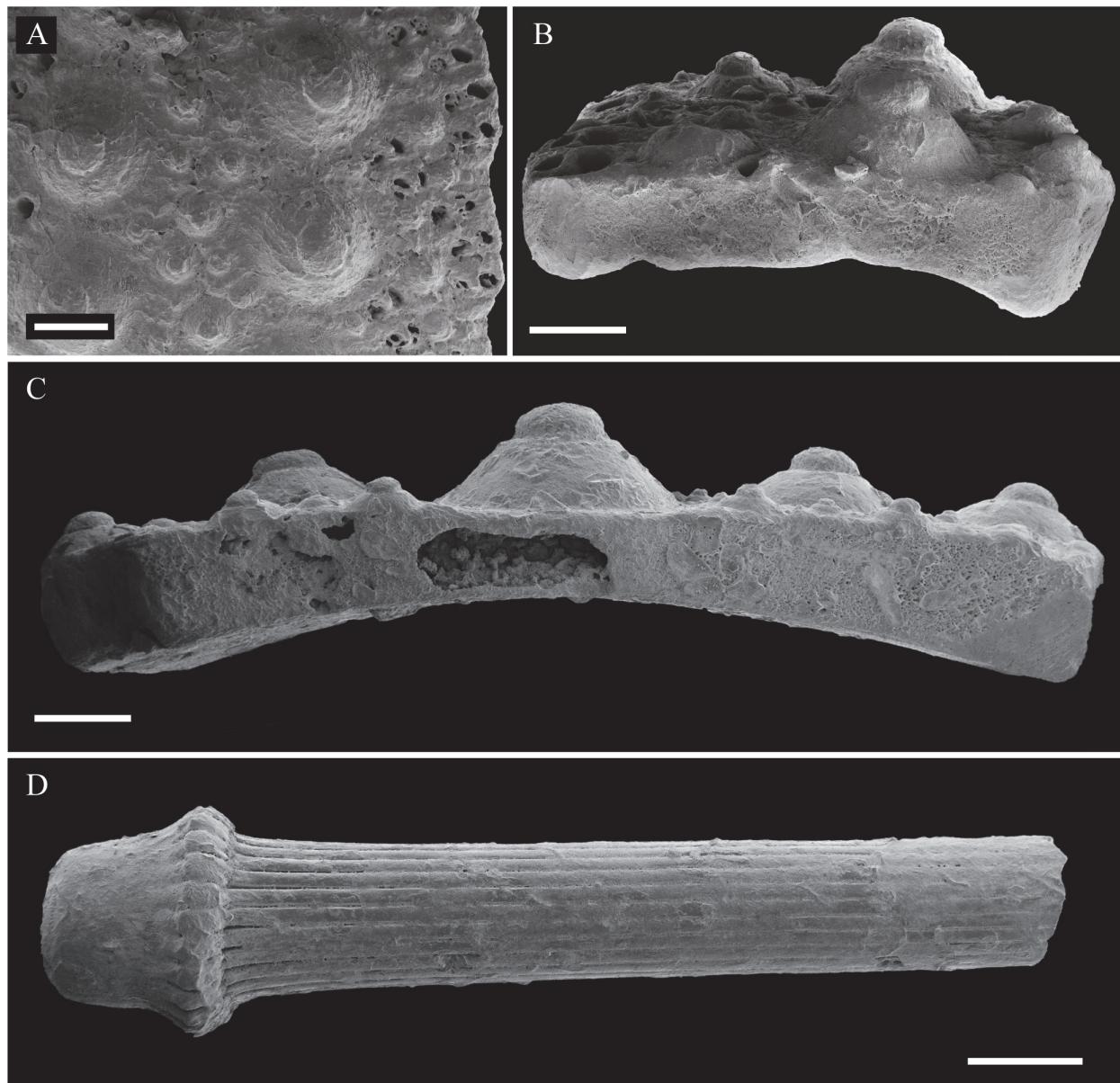
In *Paracentrotus*, parallel galleries can be observed along interradial and adradial sutures between interambulacral plates (aboral to oral direction) (Figure 7A-B) whereas no pore alignment can be observed along latitudinal sutures in interambulacral plates (Figure 7C). In ambulacral

plates, galleried stereom can be observed along the adoral sutures joining ambulacral and interambulacral plates, along the adradial sutures, as well as along the perradial suture joining ambulacral plates together (Figure 7D). Additionally, knob-like trabecular protrusions interlock with cavities formed by the opposing plate's stereom thus interlocking the adjacent plates (Figure 7C). This occurs for both radial (longitudinal) sutures (adradial, interradial and perradial) and circumferential (latitudinal) sutures (adapical and adoral).

## DISCUSSION

### Paleoenvironmental and Paleoecological Reconstruction

Geological context, fossil content, including fauna and flora, and sedimentological and taphonomic features indicate a littoral environment with a bathymetric range compatible with the intertidal



**FIGURE 5.** Scanning electron micrographs of *Paracentrotus* remains from Is Mesas. A, Details of an ambulacral column affected by abrasion. B, Two ambulacral plates still sutured together. C, Interambulacral plate showing a cavity possibly related to bioerosion. D, Spine fragment showing fine surface details such as the crenulated milled ring and shaft striation. Scale bars equal 1 mm.

to upper infralittoral zones. This reconstruction is consistent with the paleoecological interpretation of Spano (1993a, b), which was largely based on the study of the mollusk fauna.

This interpretation is consistent with the recent distribution of *Paracentrotus lividus*, which is common on firm substrates, such as gently sloping rocky walls or loose boulders, and from areas of high hydrodynamic energy in the first meters of exposed shores to calmer waters in bays, lagoon or deeper zones (Régis, 1979; Lozano et al., 1995;

Fernandez and Boudouresque, 1997; Fernandez et al., 2003; Zavodnick, 2003; Bonaviri et al., 2005; Ceccherelli et al., 2007; Fernandez et al., 2012). *Paracentrotus lividus* also occurs in seagrass meadows such as *Posidonia oceanica* and *Zostera marina*, while it is uncommon in meadows of *Cymodocea nodosa* (Boudouresque and Verlaque, 2020 and references cited therein). It is found also on colonies of *Cladocora caespitosa* (see Koukouras et al., 1998; Pitacco et al., 2014), an endemic coral living throughout the Mediterranean

**TABLE 2.** Occurrence, relative abundance and taphonomy of *Paracentrotus lividus* remains from Late Pleistocene (Part 1) and recent environments (Part 2 and 3). Fragmentation refers to intraplate fracturing. B = spine fragments retaining the base, T = retaining the tip, S = missing both the base and tip.

**Part 1.**

Locality State of preservation <b>Test</b>	N (%)	Is Mesas (Pleistocene)		
		Fragmentation (%)	Abrasion (%)	Encrustation (%)
<b>Complete test (no apical system, no spines)</b>				
More than half test with spines attached				
More than half test without spines attached				
Single ambulacral plates	613 (21.81)	475 (77.49)	508 (82.87)	6 (0.98)
Two ambulacral plates	91 (3.24)	75 (82.42)	75 (82.42)	1 (1.09)
Three to more ambulacral plates	97 (3.45)	86 (88.66)	85 (87.63)	1 (1.03)
Auricles with ambulacral plates	18 (0.64)	17 (94.44)	18 (100)	
Single interambulacral plates	755 (26.87)	572 (75.76)	570 (75.49)	5 (0.66)
Two interambulacral plates	132 (4.70)	96 (72.73)	99 (75)	6 (4.55)
Three to more interambulacral plates	39 (1.39)	25 (64.10)	26 (66.66)	1 (2.56)
Ambulacral and Interambulacral plates	113 (4.02)	104 (92.04)	103 (91.15)	4 (3.54)
Indeterminable plates	684 (24.34)	684 (100)	666 (97.37)	2 (0.29)
Total	2542 (90.46)	2134 (83.95)	2150 (84.58)	26 (1.02)
<b>Plates of the apical system</b>				
Madreporite	2 (0.07)	1 (50)	1 (50)	
Genital plates	15 (0.53)	1 (6.66)	9 (60)	
Ocular plates	4 (0.14)	1 (25)	1 (25)	
Total	21 (0.74)	3 (14.29)	10 (47.62)	
<b>Aristotle's lantern elements</b>				
Hemipyramids	123 (4.38)	109 (88.62)	118 (95.93)	
Rotulae	54 (1.92)	12 (22.22)	24 (44.44)	
Epiphyses	18 (0.64)	10 (55.55)	14 (77.77)	
Compasses	4 (0.14)		1 (25)	
Teeth	48 (1.71)	47 (97.92)	32 (66.66)	
Total	247 (8.80)	178 (72.06)	189 (76.52)	
Total	2810 (100)	2315 (82.38)	2349 (83.59)	26 (0.93)
<b>Spines</b>				
Complete	64 (4.47)		15 (23.44)	5 (7.81)
Fragments	1368 (B=461; T=208; S=699)	1368	975 (71.27)	124 (9.06)
Total	1432	1368 (95.53)	990 (69.13)	129 (9.01)
Sampling method	Bulk sampling and sieving			

on rocky and sandy bottoms where it can form extensive populations (Kersting et al., 2017; Zunino et al., 2018). *Paracentrotus lividus* is uncommon on soft substrata, such as sandy and detritic bottoms (Zavodnik 1980, 2003), although in coastal lagoons it can live on coarse sands or even

on muddy sediments (Fernandez et al., 2003; Boudouresque and Verlaque, 2020).

Individuals in shallower water under very exposed conditions and/or in intertidal rock pools, resist dislodgment by waves by boring into the hard substratum (e.g., sandstone, limestone, granite), creating cup-shaped cavities where they live per-

**TABLE 2** (continued).

Part 2.

Locality	N (%)	Sa Mesa Longa (Recent)		
State of preservation		Fragmentation (%)	Abrasion (%)	Encrustation (%)
<b>Test</b>				
Complete test (no apical system, no spines)	2 (0.11)		1 (50%)	2 (100%)
More than half test with spines attached	1 (0.05)	1 (100%)		
More than half test without spines attached	2 (0.11)			
Single ambulacral plates	272 (14.98)	137 (50.37)	203 (74.63)	14 (5.15)
Two ambulacral plates	57 (3.14)	24 (42.11)	40 (70.18)	8 (14.04)
Three to more ambulacral plates	141 (7.76)	103 (73.05)	104 (73.76)	42 (29.79)
Auricles with ambulacral plates	11 (0.60)	10 (90.91)	9 (81.82)	
Single interambulacral plates	487 (26.82)	306 (62.83)	346 (71.05)	36 (7.39)
Two interambulacral plates	167 (9.20)	81 (48.50)	112 (67.07)	41 (24.55)
Three to more interambulacral plates	98 (5.40)	47 (47.96)	68 (69.39)	36 (36.73)
Ambulacral and Interambulacral plates	163 (8.98)	113 (69.33)	78 (47.85)	63 (38.65)
Indeterminable	192 (10.57)	191 (99.48)	175 (91.15)	2 (1.04)
Total	1593 (87.72)	1013 (63.59)	1136 (71.31)	244 (15.32)
Plates of the apical system				
Madreporite	5 (0.27)	2 (40)	3 (60)	
Genital plates	7 (0.39)	2 (28.57)	2 (28.57)	
Ocular plates	4 (0.22)		1 (25)	
Total	16 (0.88)	4 (25)	6 (37.5)	
Aristotle's lantern elements				
Hemipyramids	108 (5.95)	78 (72.22)	103 (95.37)	5 (4.63)
Rotulae	61 (3.36)	12 (19.67)	32 (52.46)	
Epiphyses	10 (0.55)	6 (60)	7 (70)	
Compasses	3 (0.16)		1 (33.33)	
Teeth	25 (1.37)	21 (84)	14 (56)	4 (16)
Total	207 (11.39)	117 (56.52)	157 (75.85)	9 (4.35)
Total	1816	1134 (62.44)	1299 (71.53)	253 (13.93)
Spines				
Complete	63 (4.17)		11 (17.46)	2 (3.17)
Fragments	1447 (B=769; T=119; S=559)	1447	1170 (80.86)	82 (5.66)
Total	1510	1447 (95.83)	1181 (78.21)	84 (5.56)
Sampling method	Surface collecting			

manently or temporarily (Mortensen, 1943; Jacinto and Cruz, 2012). High population densities, up to hundreds of specimens, usually occur in shallow habitats, on gently sloping rocks, pebbles or boulders, in intertidal rocky pools and lagoons (Bonaviri et al., 2005; Fernandez et al., 2012; Ouréns et al., 2014; Boudouresque and Verlaque, 2020).

The identified fauna (Table 1) allows a highly structured environment with rocky substrates and sandy substrates with loose boulders and pebbles and secondary hardgrounds, including coralline algal rhodoliths to be inferred. Patches of *Posidonia oceanica* possibly occurred as indicated by the presence of numerous gastropod species that

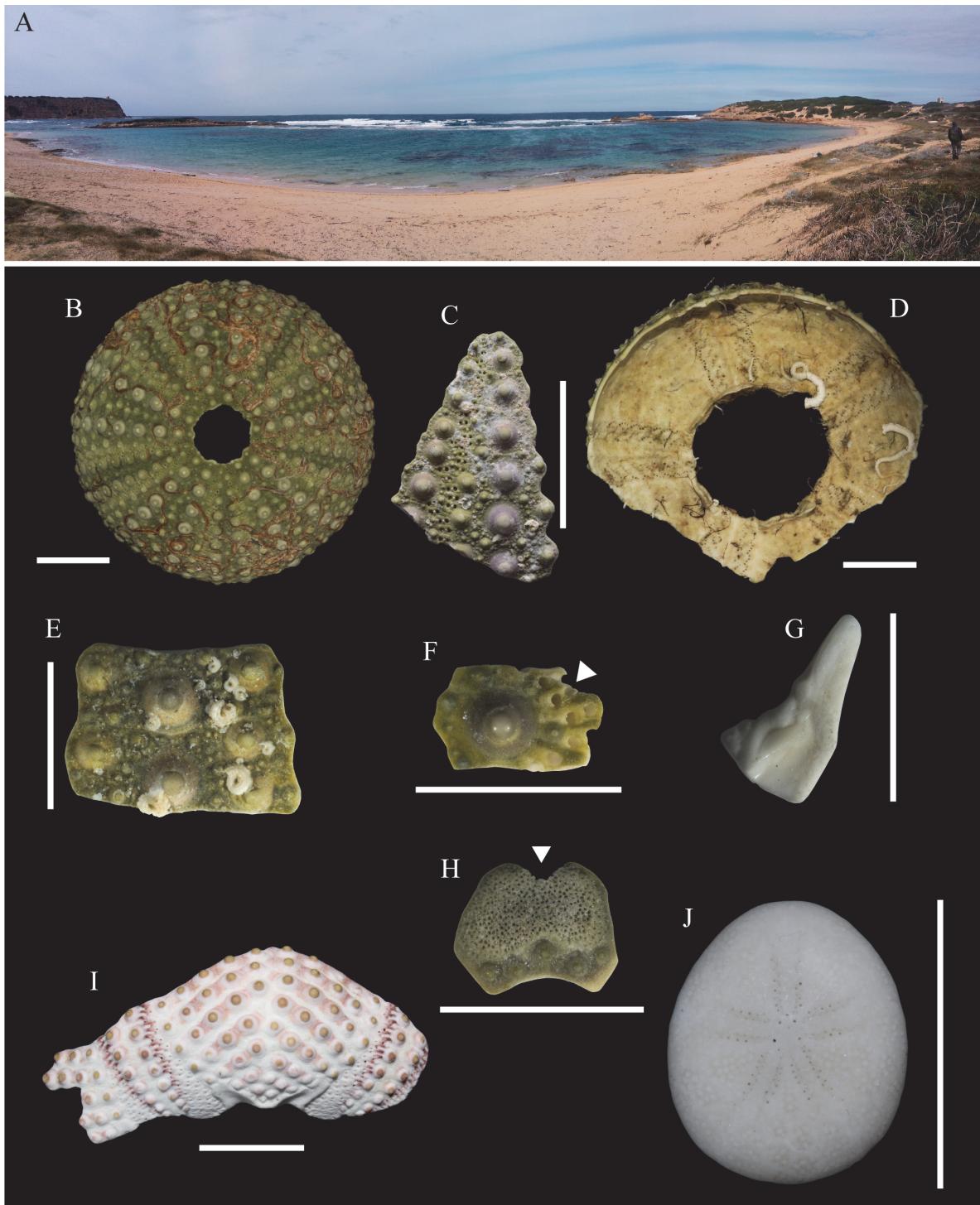
**TABLE 2** (continued).**Part 3**

Locality	N (%)	Sa Mesa Longa (Recent)		
State of preservation		Fragmentation (%)	Abrasion (%)	Encrustation (%)
<b>Test</b>				
Complete test (no apical system, no spines)				
More than half test with spines attached				
More than half test without spines attached				
Single ambulacral plates	603 (23.05)	457 (75.79)	585 (97.01)	6 (0.99)
Two ambulacral plates	38 (1.45)	30 (78.95)	35 (92.11)	
Three to more ambulacral plates	17 (0.65)	16 (94.12)	16 (94.12)	1 (5.88)
Auricles with ambulacral plates	3 (0.11)	3 (100)	3 (100)	
Single interambulacral plates	526 (20.11)	386 (73.38)	506 (96.18)	13 (2.47)
Two interambulacral plates	30 (1.15)	23 (76.67)	29 (96.67)	4 (13.33)
Three to more interambulacral plates	15 (0.57)	10 (66.67)	13 (86.67)	
Ambulacral and Interambulacral plates	11 (0.42)	9 (81.82)	11 (100)	
Indeterminable plates	1250 (47.79)	1250 (100)	1250 (100)	
Total	2493 (95.30)	2184 (87.61)	2448 (98.19)	24 (0.96)
Plates of the apical system				
Madreporite	2 (0.07)	1 (50)	2 (100)	
Genital plates	14 (0.54)	6 (42.86)	12 (85.71)	
Ocular plates	4 (0.15)		4 (100)	
Total	20 (0.76)	7 (35)	18 (90)	
Aristotle's lantern elements				
Hemipyramids	53 (2.03)	53 (100)	53 (100)	
Rotulae	17 (0.65)	3 (17.65)	12 (70.59)	
Epiphyses	10 (0.38)	7 (70)	9 (90)	
Compasses	3 (0.11)	3 (100)	1 (33.33)	
Teeth	20 (0.77)	20 (100)	20 (100)	
Total	103 (3.94)	86 (83.49)	95 (92.23)	
Total	2616	2277 (87.04)	2561 (97.90)	24 (0.91)
Spines				
Complete	48 (3.02)		26 (54.17)	
Fragments	1542 (B=508; T=76; S=958)	1542	1456 (94.42)	16 (1.03)
Total	1590	1542 (96.98)	1482 (93.21)	16 (1.01)
Sampling method	Bulk sampling and sieving			

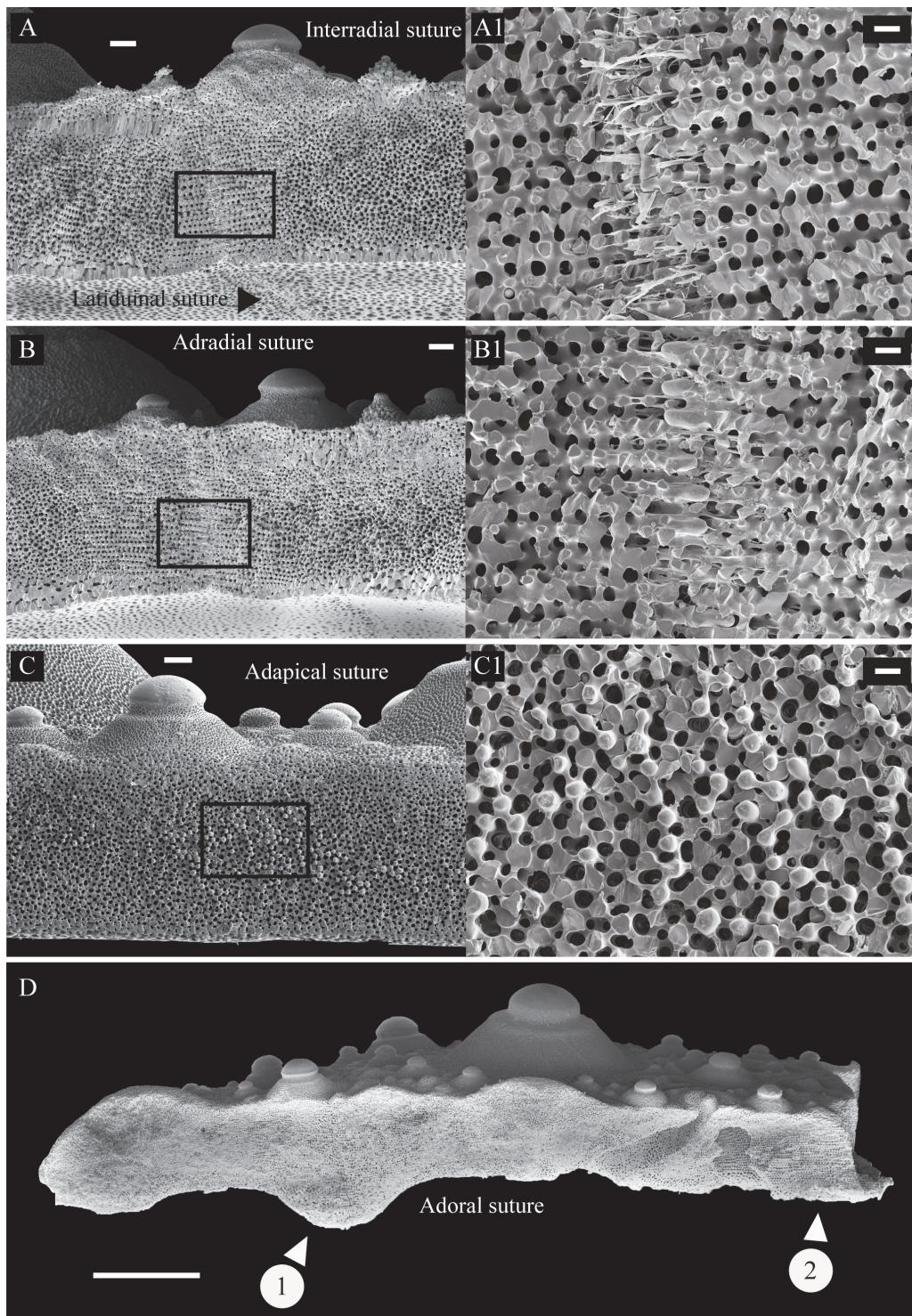
commonly live in seagrass meadows (Table 1). The fossil material studied herein, including echinoids, mollusks and coral remains, display a wide spectrum of biostratinomic signatures, with respect to fragmentation, encrustation and abrasion that suggests time-averaging in a nearshore environment. The high degree of fragmentation and abrasion of echinoid remains indicates in situ alteration

by waves and storms in a littoral environment with high hydrodynamic conditions.

The autochthony/parautochthony of echinoid remains within the Pleistocene deposit is suggested by the size-frequency distribution of some test elements in comparison with recent ones. The size-frequency distribution of the rotulae of the Aristotle's lantern within the fossil deposit is consis-



**FIGURE 6.** A, Panoramic view of Sa Mesa Longa Beach (Central-western Sardinia). B-H, Recent remains of *Paracentrotus lividus* from Sa Mesa Longa. B, Complete denuded test showing encrustation by serpulids. C, Test fragment made up of several ambulacral and interambulacral plates sutured together, showing inter- and intraplate fragmentation and encrustation by coralline algae and the polychaete *Spirorbis*. D, Internal view of a large test fragment affected by intraplate fragmentation and encrustation by serpulid worms. E, Two interambulacral plates showing abrasion and encrustation by *Spirorbis*. F, Single ambulacral plates showing fragmentation (white arrow). G, Epiphysis from the jaw apparatus showing fragmentation and abrasion. H, Madreporite from the apical system affected by fragmentation (white arrow). I, Test fragment of *Arbacia lixula* showing intraplate fracturing. J, Complete test of *Echinocyamus pusillus*. B, C, D, I, J Scale bar equals 1 cm. E-H scale bars equal 0.5 cm.



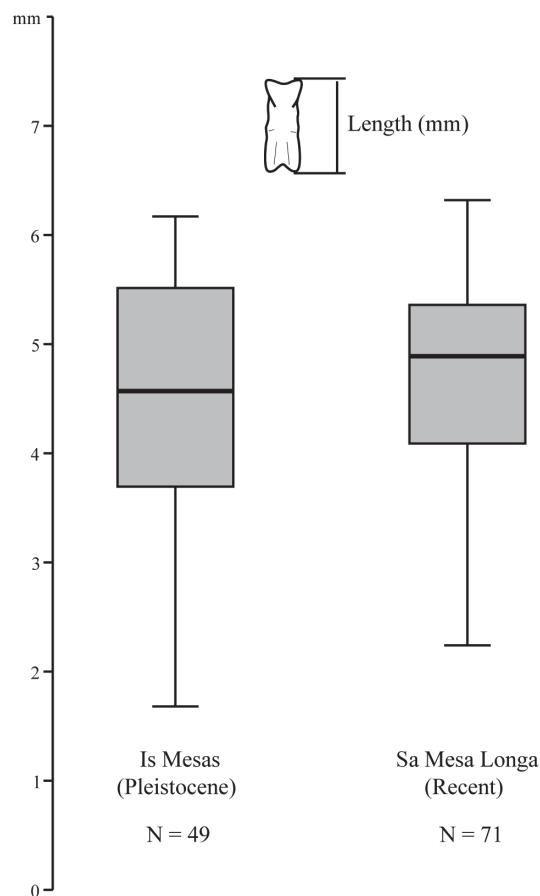
**FIGURE 7.** Scanning electron micrographs of suture faces in recent *Paracentrotus lividus*. A, Interradial suture between interambulacral plates showing galleried stereom with galleries running in aboral-oral direction. A1, Detail of the galleried stereom and collagen fibers. B, Adradial suture of interambulacral plates showing galleried stereom. B1, Detail of the parallel galleries and collagen fibers. C, Adapical suture of an interambulacral plates showing knob-like trabecular protrusions and cavities. C1, The close-up shows numerous knob-like protrusions some of which are interconnected to one another in twos, threes or more. D, Adoral suture of ambulacral plates; 1) Radial ridge at the boundary between ambulacral plates (running across perradial sutures) and 2) galleried stereom. A, B, C Scale bars equal 100 µm; A1, B1, C1 Scale bars equal 20 µm. D Scale bar equals 1 mm.

tent with that within recent setting (Figure 8) and documents the presence of small to large-sized individuals (ca. 13 to 55 mm in diameter on the basis of the ratio rotula length/test diameter in recent specimens), which typically co-occur in living population (e.g., Sala and Zabala, 1996; Gianguzza et al., 2006; Addis et al., 2012; Loi et al., 2017). This, along with the lack of sorting, as denoted for example by the presence of both primary and smaller miliary spines, excludes the transport of the echinoid remains over long distances.

#### Preservation of *Paracentrotus lividus*

Fossil remains of *Paracentrotus lividus* have been recognized before in Pleistocene deposits in the North Atlantic and Mediterranean coastlines (Borghi, 1995, 2003; Borghi et al., 2006; Ferreira Soares et al., 2007; Néraudeau and Masrour, 2008; Madeira et al., 2011). Borghi (1995, 2003) and Borghi et al. (2006) reported the occurrence of both the parechinid *Paracentrotus lividus* and the toxopneustid *Sphaerechinus granularis*, in Lower Pleistocene sandy deposits of Emilia and Sicily (Italy). These echinoids occur mainly as test fragments and disarticulated spines, although complete, denuded tests were sporadically found. Ferreira Soares et al. (2007) reported the occurrence of *Paracentrotus lividus* from the Lower/Middle Pleistocene deposits of Portugal. The echinoid remains, which are represented by two test fragments and several spines, occur in association with a mollusk fauna, including *Nucella lapillus*, *Patella vulgata*, *Littorina littorea* and *Mytilus edulis*, well adapted to living on intertidal and uppermost infralittoral hard substrates with high energy conditions. Néraudeau and Masrour (2008) also document the occurrence of *Paracentrotus lividus* both from the Pleistocene (Cap Dra) and the Holocene (Rabat) of Morocco.

*Paracentrotus lividus* also occurs in the Late Pleistocene (MIS 5e) deposits of Santa Maria Island (Azores), in association with *Sphaerechinus granularis* and the arbacioid *Arbacia lixula* (Madeira et al., 2011). *Paracentrotus* and *Sphaerechinus* were found as test fragments and disarticulated and fragmented spines only. *Arbacia lixula* occurs as complete tests, numerous test fragments and spines (Madeira et al., 2011). These three sea urchins commonly co-occur in littoral environments on a wide variety of substrates, including rocky bottoms (Sala et al., 1998; Bonaviri et al., 2005; Guidetti and Mori, 2005; Chiantore et al., 2008; Privitera et al., 2008; Agnetta et al., 2015), which is

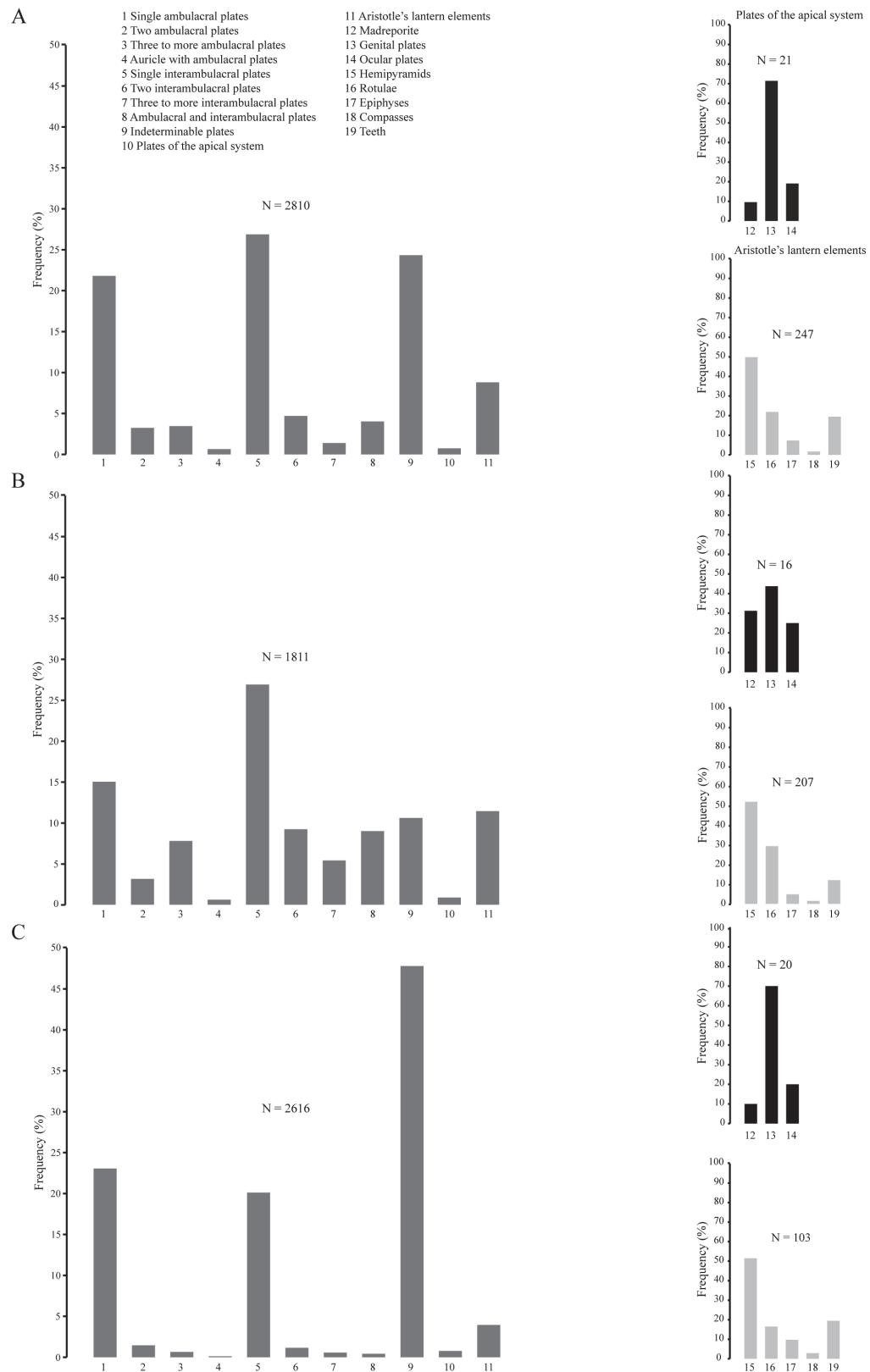


**FIGURE 8.** Size-frequency distribution of the rotulae of the Aristotle's lantern both in Pleistocene deposit and recent setting. Each box-plot represents 25 and 75 percent quartile of all values, Q1 and Q3 respectively. Black line inside the box represents the median. Whiskers are drawn from Q1 and Q3 to the largest values less than 1.5 times the Interquartile range (Q1-Q3). N = number of counted specimens.

consistent with the paleoenvironmental setting of the echinoid-bearing deposits of the Late Pleistocene of Santa Maria Island (Ávila et al., 2009).

Observations reported herein document that skeletal remains of *Paracentrotus lividus* occur abundantly both in Pleistocene deposit and recent environments, displaying a generally high test stability in shallow water in a high energy setting. The relative abundance of these remains (Table 2 and Figure 9) is dependent on factors such as the number of skeletal elements per organism (Table 3), differences in preservation potential of the elements and sampling methods.

A taphonomic bias in the relative abundance of skeletal elements is recognized as complete tests and large fragments consisting of more than half the test are rare and only occur on the sedi-



**FIGURE 9.** Relative abundance of *Paracentrotus lividus* test remains. A, Is Mesas (Pleistocene-Bulk sample). B, Sa Mesa Longa Beach (Recent-Surface collection). C, Sa Mesa Longa Beach (Bulk sample). N = number of counted specimens.

**TABLE 3.** Relative abundance of test elements of *Paracentrotus lividus* with respect to element abundance in a recent specimen 45.02 mm in diameter. E = number of elements. EA = Expected element abundance (E/Total). OA = Observed element abundance (E/Total).

Test elements			Is Mesas			Sa Mesa Longa			Sa Mesa Longa		
	E	EA	E	OA	OA/EA %	E	OA	OA/EA %	E	OA	OA/EA%
Ambulacral plates	251	0.515	613	0.374	72.621	272	0.277	53.786	603	0.482	93.592
Interambulacral plates	191	0.392	755	0.461	117.602	487	0.496	126.530	526	0.420	107.142
Madreporite	1	0.002	2	0.001	50	5	0.005	250	2	0.001	50
Genital plates	4	0.008	15	0.009	112.5	7	0.007	87.5	14	0.011	137.5
Ocular plates	5	0.010	4	0.002	20	4	0.004	40	4	0.003	30
Hemipyramids	10	0.020	123	0.075	375	108	0.109	545	53	0.042	210
Rotulae	5	0.010	54	0.033	330	61	0.062	620	17	0.013	130
Epiphyses	10	0.020	18	0.011	55	10	0.010	50	10	0.008	40
Compasses	5	0.010	4	0.002	20	3	0.003	30	3	0.002	20
Teeth	5	0.010	48	0.029	290	25	0.025	250	20	0.016	160
Total	487		1636			982			1252		
Sampling method	Bulk sampling and sieving			Surface collecting			Bulk sampling and sieving				

ment surface in the recent setting (Table 2; Figure 9). Smaller test fragments, consisting of several plates still sutured together, are proportionally more abundant on the sediment surface in recent settings (see Table 2). This could be, at least in part, due to the shorter time these remains have undergone reworking, as suggested by the fact that both fragmentation and abrasion are slightly lower for recent remains collected on the surface than within the sediment.

Both in the Pleistocene deposit and in the recent environment, a number of large test fragments with ambulacral and interambulacral plates still sutured together, are frequently altered by abrasion and encrustation (Figure 10). This reveals that, not only the presence and distribution of bundles of collagenous fibers (which quickly decay after the death of the echinoids), but also plate thickness and interlocking across plates sutures by interdigitating knobs, are able to hold plates together long enough until burial. This leads to the high preservation potential of larger fragments even within high energy, shallow water environments.

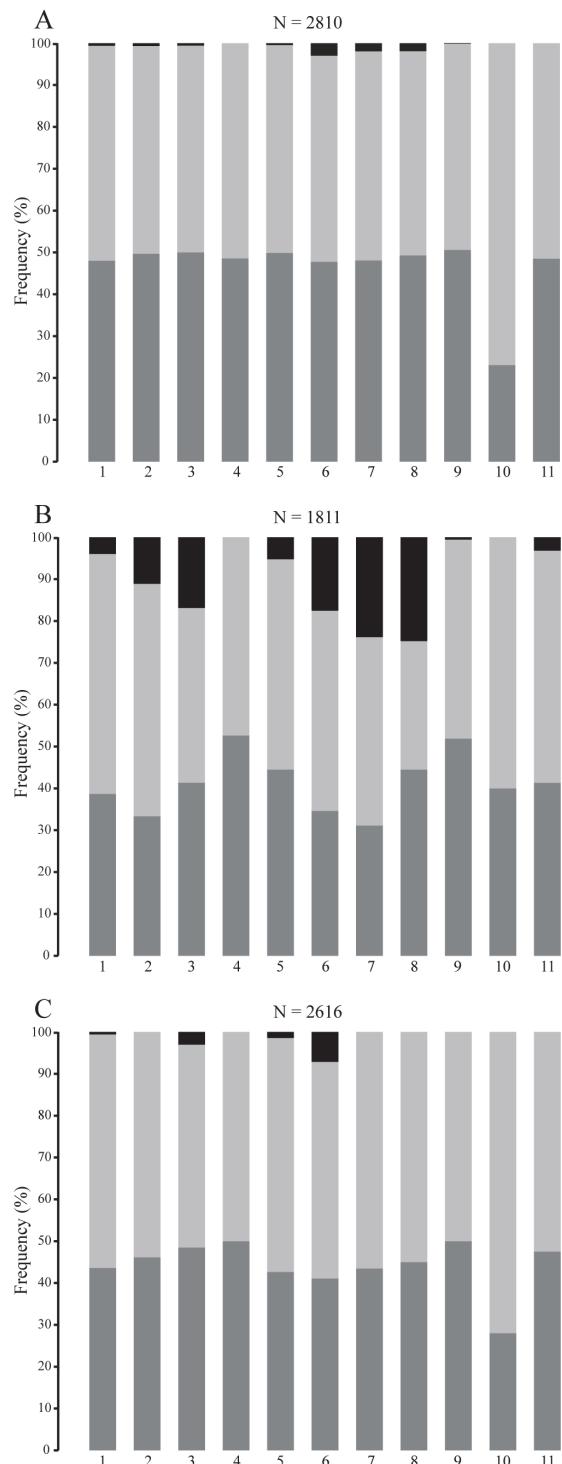
Although ambulacral plates are slightly more numerous than interambulacral plates in the test of *Paracentrotus*, single ambulacral plates are less represented than interambulacral plates both in fossil and recent settings (see Table 3), at least in the > 2 mm fraction (Figure 11A). Indeed, most ambulacral plates are preserved in the 1-2 mm fraction and occur mainly as fragmented remains.

By contrast, interambulacral plates mostly occur in the > 2 mm fraction (Figure 11B). These results suggest that, beyond discrepancies related to different sampling methods, the underrepresentation of ambulacral plates is related to a lower preservation potential with respect to that of interambulacral plates, likely due to the presence of pore-pairs leaving ambulacral plates more fragile (see Figure 6F).

This conclusion is corroborated by the analysis of bulk sediment samples from Sa Mesa Longa Beach (Figure 11). In shallow water, high-energy, environments which are conducive to constant post-mortem transportation and reworking, ambulacral plates are more rapidly reduced to sand-sized particles (1-2 mm fraction) than interambulacral plates.

Among the elements of the Aristotle's lantern, hemipyramids are the most abundant and represent more than half of the remains both among fossil and recent lantern elements (Table 3). This is interpreted as a result of their robustness that leads to a higher preservation potential than other elements such as epiphyses and compasses, which are fragile and are also more difficult to be recognized as fragments.

Spines remains are the most abundant skeletal elements both in fossil and recent settings. They consist mainly of fragmented remains (Figure 12), affected by abrasion and encrustation (Figure 13), although complete spines showing fine surface details do occur. Among fragmented spines, a



**FIGURE 10.** Biostratinomic signatures (fragmentation, abrasion and encrustation) of echinoid test remains. A, Is Mesas (Pleistocene-Bulk sample). B, Sa Mesa Longa Beach (Recent-Surface collection). C, Sa Mesa Longa Beach (Recent-Bulk sample). N = number of counted specimens. Numbers (1-11) are explained in Figure 9. Colors: dark grey, light grey and black represent fragmentation, abrasion and encrustation, respectively.

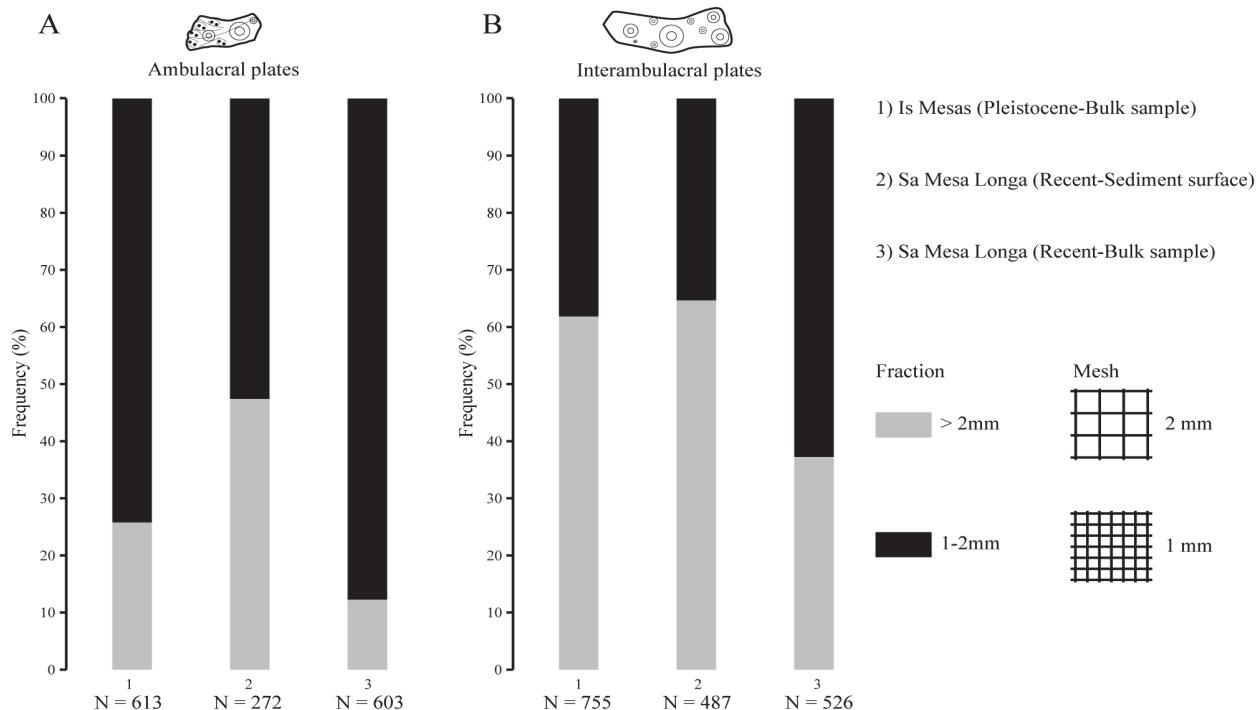
large proportion retain the base while tips are less abundant. This discrepancy is possibly due to the difference in preservation potential, which is higher for the proximal portion of the spines. Spine tips are more fragile and tend to be broken in shallow water environments where echinoid remains are continuously subject to reworking by waves even under fair-weather conditions. In addition, some spines show evidence of regrowth after breakage occurred during life, which often occurs in living specimens as documented by Ebert (1967, 1988) for *Strongylocentrotus purpuratus* and *Heterocentrotus mammillatus*.

Along with fragmentation by water agitation, predation could be another cause of test breakage and disarticulation as documented by the finding of a recent *Paracentrotus* test still with spine attached and affected by a large opening (hole) on the aboral side, similar to that documented by Sievers and Nebelsick (2018) for *Sphaerechinus granularis*. *Paracentrotus* is commonly preyed by fish, asteroids, crustaceans, gastropods and birds, which can result in fragmentation (Ebling et al., 1966; Sala, 1997; Guidetti, 2004 and references therein; Bonaviri et al., 2010; Boudouresque and Verlaque, 2020).

Although *Paracentrotus* is common, there is a noticeable lack of *Arbacia lixula* remains in the Pleistocene deposit studied herein. *Arbacia* is one of the most common echinoids in shallow water high energy environments of the Mediterranean today and shares the same environment as *Paracentrotus*. *Arbacia lixula* also possesses a suite of surface characteristics which allows identification, and is also found, albeit rare, in the census of recent echinoid remains of the present study.

The lack of occurrences of *Arbacia* in Pleistocene deposits of the Mediterranean area, with the exception of a single well-preserved specimen from the Pleistocene of Livorno (Italy), under the name *A. pustulosa* (Stefanini, 1911), has led to it being regarded as a “post-glacial” invasive species into the Mediterranean Sea (Stefanini, 1911; Mortensen, 1935). More recently, Wangensteen et al. (2012) reported that *Arbacia lixula* is consistently absent from “Tyrrhenian” deposits and suggested that the expansion in the Mediterranean Sea possibly took place during the last interglacial period (MIS 5e). This is supported by the present study where the rigorous examination of echinoid fragments of Pleistocene sediments also did not reveal the presence of *Arbacia*.

The scarcity of *Echinocyamus* remains in the Pleistocene deposit of Is Mesas, is possibly a



**FIGURE 11.** Distribution of ambulacrals and interambulacrals in the two sediment fractions (> 2 mm and 1-2 mm) investigated herein. N = number of counted specimens.

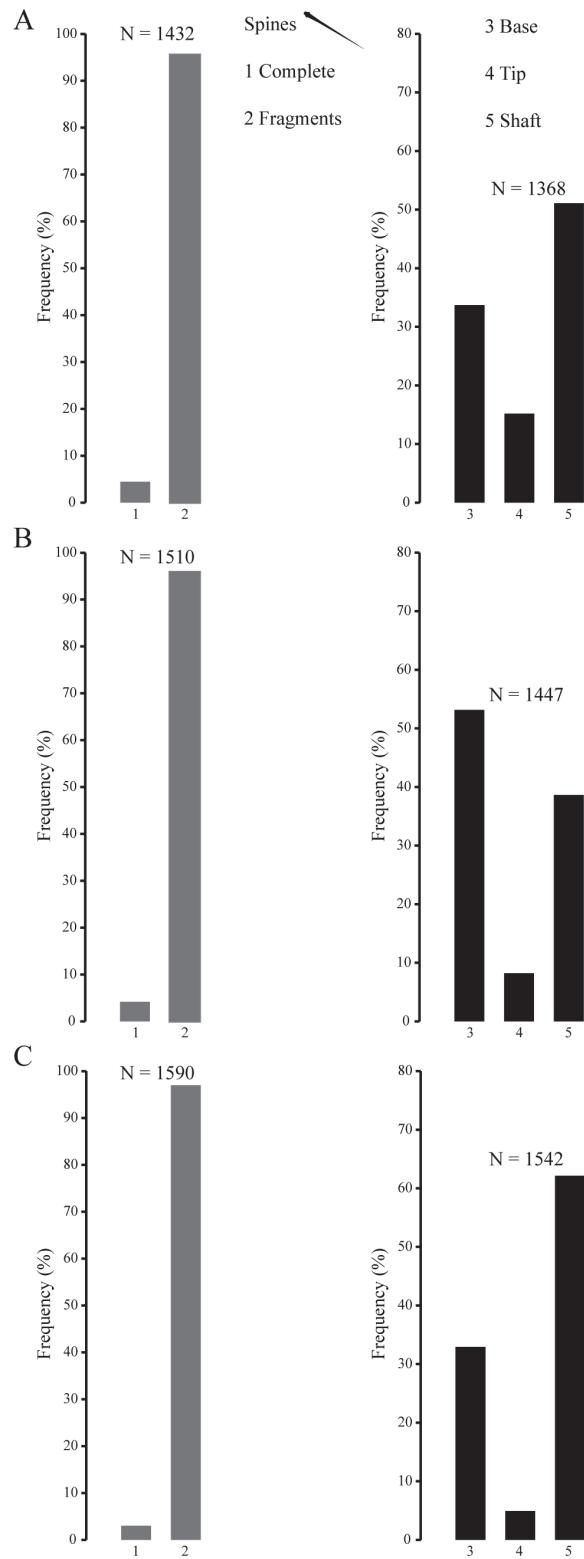
reflection of a low population density rather than a consequence of the high energy environment destroying all recognizable remnants of this echinoid. *Echinocyamus* test remains are common in both recent environments (e.g., Grun et al., 2014 and herein) and fossil deposits (e.g., Grun et al., 2017; Mancosu and Nebelsick, 2019), showing high preservation potential due to its skeletal architecture (Grun and Nebelsick, 2018a; Grun et al., 2018b).

#### Taxonomic Bias

The relatively poor fossil record of regular echinoids with respect to abundance and diversity over geological time, when compared to that of irregulars, is also the result of sampling methods and taxonomic methodology (Kier, 1977; Greenstein, 1993a). Regular echinoids in general are frequently preserved as fragmented or disarticulated test remains with some exception, generally related to rapid burial (see Mancosu et al., 2015 and references therein). The lack of complete tests generally hampers the identification of regular echinoids at genus and species level. Classification of some groups, such as cidaroids, diadematoids and camarodonts at sub-familial level mainly relies on features, such as pedicellariae, primary spines, apical disc and other test characteristics (Mortensen, 1928; Phelan, 1970; Smith, 1988;

Mooi et al., 2000; Coppard and Campbell, 2004, 2006a, b; Brosseau et al., 2012; Smith and Kroh, 2011), which are often missing in fossils or can only be observed in well-preserved complete tests. Therefore, the classification of fossil regular echinoids at low taxonomic levels based on fragmented material still remains problematic (Mortensen, 1940; Donovan et al., 2001; Kroh, 2005; Borghi and Stara, 2016). Greenstein (1991) states that, in terms of recognizing and identifying fossil echinoids, a size of 2 mm is a cut-off value since smaller test fragments would possibly not be detected by paleontologists in the field work or unlikely retain features that allow their identification at low taxonomic level.

The possibility of recognizing regular echinoid presence by analyzing fragmented material in fossil deposits has been explored by studies performed on the Pleistocene deposits of shallow water environments (Gordon, 1991; Gordon and Donovan, 1992; Donovan and Gordon, 1993; Donovan and Jones, 1994; Donovan et al., 1994; Borghi, 1995, 2003; Ragagni, 1997; Donovan, 2003, 2005; Borghi et al., 2006; Madeira et al., 2011). Identification of fossil material can be facilitated by direct comparisons with extant representatives (e.g., Gordon and Donovan, 1992; Donovan and Gordon, 1993; Ragagni, 1997).



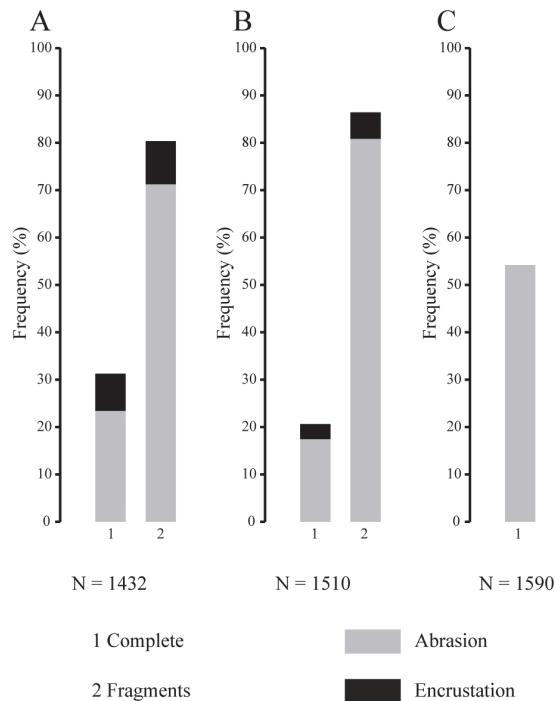
**FIGURE 12.** Relative abundance of *Paracentrotus lividus* spine remains. A, Is Mesas (Pleistocene-Bulk sample). B, Sa Mesa Longa Beach (Recent-Surface collection). C, Sa Mesa Longa Beach (Bulk sample). N = number of counted specimens.

Fossil remains of *Paracentrotus* have been sometimes misidentified as *Strongylocentrotus* Brandt, 1835 (Mortensen, 1943; Geys and Marquet, 1979) due to the fact that the naked tests of these sea urchins appear very similar in overall morphology. The globiferous pedicellariae and the morphology of the auricles of the perignathic girdle, however, allows these two genera to be distinguished (Mortensen, 1943; Smith and Kroh, 2011). Additionally, in *Paracentrotus*, the ocular plates in the apical disc are all exert whereas in *Strongylocentrotus* the posterior oculars (oculars I and V) are generally insert (in adults), in contact with the periproct (Mortensen, 1943; Smith and Kroh, 2011).

The genus *Paracentrotus* Mortensen, 1903, includes *P. lividus* (Lamarck, 1816) and *P. gaimardi* (Blainville, 1825), which is reported to occur along the coasts of Brazil and West Africa, and inhabits in large number in shallow rocky reefs (Mortensen, 1943; Cordeiro et al., 2014). They differ in the features of the apical plates, which show distinct radiating stripes in *P. gaimardi* and in their pedicellariae (Mortensen, 1943; Smith and Kroh, 2011). Although no pedicellariae have been found in the studied deposits, the occurrence of all test elements, including the auricles, the apical system plates and Aristotle's lantern elements, which themselves are rare finds in the fossil record, allows for the identification as *Paracentrotus lividus*. Herein, complete and well-preserved test elements occur abundantly both in > 2 mm and the 1–2 mm fractions allowing identification of fossil remains at low taxonomic level even in sand particle-size fraction.

## CONCLUSIONS

This study documents the presence of abundant remains of the regular echinoid *Paracentrotus lividus* in Late Pleistocene (Tarantian, MIS 5e) deposits of Is Mesas (southern Sardinia) and modern shore deposits of Sa Mesa Longa Beach (Central-Western Sardinia). In the Pleistocene deposit the occurrence of *Paracentrotus* and associated mollusk fauna, including both bivalves and gastropods, allows a shallow water environment with rocky substrate and adjacent coarse-grained soft sediments and possibly sea grass to be inferred. Additionally, patches of coralline-algae rhodoliths and colonies of *Cladocora caespitosa* occurred. Fauna diversity, taphonomy and sedimentological features point to an autochthonous to paraautochthonous assemblage (a within-habitat time averaged assemblage). Echinoid remains include



**FIGURE 13.** Biostratinomic signatures (abrasion and encrustation) of spine remains. A, Is Mesas (Pleistocene-Bulk sample). B, Sa Mesa Longa Beach (Recent-Surface collection). C, Sa Mesa Longa Beach (Recent-Bulk sample). N = number of counted specimens

variously sized test fragments and isolated ambulacrals and interambulacrals plates, plates of the apical system, elements of the Aristotle's lantern and complete and fragmented spines. Although no whole echinoid tests were found, *Paracentrotus lividus* occur as large test fragments depending to its relatively robust test and stereomic interlocking. This interpretation is consistent with observation in recent environments where *Paracentrotus* remains, which range from complete denuded tests to isolated elements, have been recognized in great abundance on and within the sediments around rocks where living populations occur.

Echinoid remains in the Pleistocene deposit reflect those in recent environment with respect to relative abundance and biostratinomic signatures, and are useful indicators of paleoenvironmental

and paleoecological conditions even in shallow water, high-energy environments. This study indicates that when rocky shore facies are preserved in the fossil record the remains of epibenthic regular echinoids may be locally abundant, mostly as disarticulated bioclasts.

Discrepancy in relative abundance among various test elements with respect to the expected skeletal element abundance in complete specimens and their different distribution in size fraction (1-2 mm and > 2 mm) are indicative of differences in preservation potential. Reworking of echinoid remains tends to increase the richness of taphonomically durable elements such as interambulacrals plates over ambulacrals ones and hemipyramids and rotulae over the other elements of the Aristotle's lantern. Differential resistance of test elements to taphonomic processes potentially led to taxonomic bias.

Besides preservation failure, the fossil record of regular echinoids may be biased by inadequate sampling that hampers identification at low taxonomic levels. Herein, the analysis of fragmented material, allows for the identification of echinoid remains at species level, even at a very coarse sand fraction (1-2 mm), when supported by comparison with extant species. This approach is difficult for other taxonomic groups such as diadematid echinoids, the systematics of which is based on features that are seldom preserved in the fossil record, and other now extinct echinoid taxa which lack comparable recent taxonomic representatives.

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