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PETERSENITE-(Ce), A NEW MINERAL FROM MONT SAINT-HILAIRE, AND ITS STRUCTURAL RELATIONSHIP TO OTHER REE CARBONATES

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ABSTRACT

Petersenite-(Ce) is a newly discovered rare-earth-element carbonate from the Poudrette Quarry, Mont Saint-Hilaire, Quebec. It is grey, with a slight pink tint, forming striated prisms up to 7 mm in length. It has a vitreous luster and is translucent to transparent. Petersenite-(Ce) is relatively soft (Mohs hardness ~3) and very brittle, having a conchoidal fracture. Optical properties are biaxial, with moderate dispersion, no absorption, and indices of refraction α 1.623(1), β 1.636(1), γ 1.649(1), $2V_{\text{meas.}} = 89.7(5)^{\circ}$ and $2V_{\text{calc.}} = 89.8^{\circ}$, and an optical orientation of X = b, $Y \land c = 30^{\circ}$ and Z = a. Electron-microprobe analyses gave Na₂O 17.38, CaO 1.32, BaO 0.32, SrO 1.70, La₂O₃ 14.49, Ce₂O₃ 23.66, Pr₂O₃ 2.00, Nd₂O₃ 5.82, Sm₂O₃ 0.60 and CO₂ (from crystal-structure analysis) 32.92, total 100.21 wt.%, which yields the empirical formula (Na_{3.75}Ca_{0.16})(Ce_{0.96}La_{0.59}Nd_{0.23}Sr_{0.11}Pr_{0.08}Sm_{0.02}Ba_{0.01})(CO₃)₅ or, ideally, Na₄(REE)₂(CO₃)₅ for Z = 4, with $D_{\text{calc.}} = 3.67$ g/cm³ and $D_{\text{meas.}} = 3.69$ g/cm³. Crystal-structure analysis shows the mineral to be monoclinic, space group $P2_1$, with a = 20.872(4), b = 6.367(1), c = 10.601(2) Å and $b = 120.50(3)^{\circ}$. The strongest lines in the X-ray powder pattern [d in Å(I)(hkl)] are 9.13(3) (201,001), 5.22(5)(211,011,201,401), 4.13(3)(501,301), 3.70(4)(412,012), 2.607(10)(402,802,420), 2.148(3)(814,014) and 1.921(3)(432,032,430). The crystal structure has the subcell of burbankite and remondite-(Ce), but the supercell changes the ratio of [8]-coordinated to [10]-coordinated cations. The layered structure is compared to that of other REE carbonates.

Keywords: petersenite-(Ce), new mineral species, rare-earth element, carbonate, crystal structure, Mont Saint-Hilaire, Quebec.

SOMMAIRE

La petersenite-(Ce) est un carbonate à terres rares (TR) nouvellement découvert à la carrière Poudrette, au mont Saint-Hilaire, Québec. Il s'agit d'un minéral gris, avec une teinte rosâtre, se présentant en prismes striés translucides à transparents, à éclat vitreux et jusqu'à 7 mm en longueur. Sa dureté de Mohs est environ 3; elle est très cassante, et la fracture est concoïdale. C'est un minéral biaxe, ayant une dispersion moyenne, aucune absorption, et les indices de réfraction suivants: α 1.623(1), β 1.636(1), γ 1.649(1), $2V_{\text{meas.}} = 89.7(5)^{\circ}$ et $2V_{\text{calc.}} = 89.8^{\circ}$. L'orientation optique donne X = b, $Y \land c = 30^{\circ}$, et Z = a. Une analyse à la microsonde électronique démontre la présence de: Na₂O 17.38, CaO 1.32, BaO 0.32, SrO 1.70, La₂O₃ 14.49, Ce₂O₃ 23.66, Pr₂O₃ 2.00, Nd₂O₃ 5.82, Sm₂O₃ 0.60 et CO₂ (des résultats de l'analyse de la structure cristalline 32.92, pour un total de 100.21% (en poids), ce qui mène à la formula empirique (Na_{3.75}Ca_{0.16}) (Ce_{0.96}La_{0.59}Nd_{0.23}Sr_{0.11}Pr_{0.08}Sm_{0.02}Ba_{0.01})(CO₃)₅ ou, plus simplement, Na₄(TR)₂(CO₃)₅ pour Z = 4, avec $D_{\text{calc.}} = 3.67$ g/cm³ et $D_{\text{meas.}} = 3.69$ g/cm³. Une analyse de la structure cristalline montre qu'il s'agit d'un minéral monoclinique, groupe spatial $P2_1$, avec α 20.872(4), b 6.367(1), c 10.601(2) Å et β 120.50(3)°. Le cliché de diffraction X (méthode des poudres) possède les raies les plus intenses suivantes [d en Å(D(D(D(D)) = 9.13(3)(D(D), 5.22(5)(D(D)101, 201, D(D), 4.13(3)(D(D), 3.70(4)(D12, 2.607(10)(402, 802, 420), 2.148(3)(D(D14, 01) et 1.921(3)(D32, 2.430). La structure contient la sous-maille de burbankite et remondite-(Ce), mais la super-maille change le rapport des cations à coordinence [8] à [10]. La structure en feuillets est comparée à celle d'autres carbonates à terres rares.

(Traduit par la Rédaction)

Mots-clés: petersenite-(Ce), nouvelle espèce minérale, terres rares, carbonate, structure cristalline, mont Saint-Hilaire, Québec.

INTRODUCTION

At Mont Saint-Hilaire, Quebec, 34 carbonate mineral species have been identified, as well as several unknowns (Horvath & Gault 1990). Of the identified carbonates, eleven contain essential amounts of rareearth elements (*REE*). At the Canadian Museum of Nature, we have been routinely studying this interesting crystal-chemical class.

The new mineral described here, petersenite-(Ce), becomes the fourth member of the burbankite group, which also includes remondite-(Ce) and khanneshite. Petersenite-(Ce), found originally in the Poudrette quarry, Mont Saint-Hilaire, Rouville County, Quebec, is named in honor of Dr. Ole V. Petersen, Geologisk Museum, Copenhagen, Denmark, in recognition of his significant contributions to the understanding of the mineralogy and genesis of alkaline rocks. The new mineral and the name were approved by the Commission on New Minerals and Mineral Names, IMA. Cotype material is housed in the collection of the Canadian Museum of Nature under catalog no. 81511.

OCCURRENCE

Mont Saint-Hilaire is one of the ten Monteregian Hills rising above the Saint Lawrence Lowlands. This almost linear arrangement of monadnocks stretches for almost 150 km eastward from Oka to Megantic. The Hills are alkaline intrusive complexes. Horvath & Gault (1990) briefly described the geology, gave detailed descriptions of the minerals, and have compiled a very complete list of references for the Mont Saint-Hilaire locality.

In the spring of 1991, J.V.V. collected samples that later became the cotype specimens for petersenite-(Ce), in the Poudrette quarry, from a late-stage pegmatitic phase of the nepheline-sodalite syenite, near the contact with the hornfels unit. Petersenite-(Ce) is associated with the following minerals: a member of the astrophyllite group, aegirine, albite, analcime, biotite, calcite, catapleiite, a member of the chlorite group, epididymite, eudialyte, fluorite, galena, microcline, polylithionite, rhodochrosite, serandite and sphalerite. At present, we must consider petersenite-(Ce) a rare mineral since only ten specimens are known, each with only a few milligrams of the new mineral.

Subsequent to the initial discovery of the cotype samples, petersenite-(Ce) has been identified from two other associations at Mont Saint-Hilaire. It also occurs as yellow, transparent, acicular crystals attaining several millimeters in length, embedded in trona, from a sodalite-rich inclusion in nepheline syenite, and as mauve, acicular crystals up to 1 mm in length in a pegmatite phase in nepheline syenite associated with microcline, albite, aegirine and shomyokite-(Y).

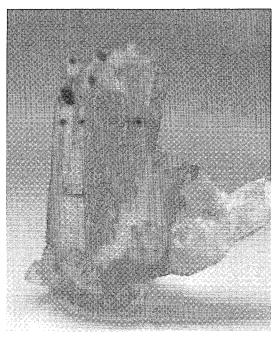


Fig. 1. Petersenite-(Ce) crystal with a dark coating, 7 mm in length, on albite.

PHYSICAL AND OPTICAL PROPERTIES

Petersenite-(Ce) forms striated prisms, owing to parallel growth, up to 7 mm in length (Fig. 1). The type material is grey with a slight pinkish tint and a vitreous luster. It has a white streak and is translucent to transparent in thin splinters. It shows no fluorescence with either long-wave or short-wave ultraviolet light. Petersenite-(Ce) is relatively soft (Mohs hardness about 3), very brittle, with no apparent cleavage, and a conchoidal fracture. The density, measured by suspension in Clerici solution, is 3.69 ± 0.03 g/cm³, which agrees well with the calculated density of 3.67 g/cm³.

Petersenite-(Ce) is biaxial, with indices of refraction α 1.623(1), β 1.636(1) and γ 1.649(1) (for λ = 590 mm); $2V_{\text{meas.}}$ = 89.7(5)° (extinction curves), $2V_{\text{obs.}}$ = 90 ± 5° (optic axis), and $2V_{\text{calc.}}$ = 89.8°; moderate dispersion; no pleochroism. As the 2V is close to 90°, neither the sign nor the dispersion formula could be determined. These optical properties are significantly different from those of remondite-(Ce), which is biaxial positive, with indices of refraction α = 1.632(2), β = 1.633(2), γ = 1.638(2) and $2V_{\text{meas.}}$ = 40(6)°. The optical orientation of petersenite-(Ce) is X = b, $Y \Lambda c = 30^\circ$, and Z = a.

CHEMICAL COMPOSITION

Chemical analysis was performed on a JEOL 733 Superprobe using Tracor—Northern 5600 automation. The wavelength-dispersion scan (WDS) mode was used; data reduction was done with the Tracor—Northern Task series of programs using a conventional ZAF correction routine. The operating voltage was 15 kV, and the beam current was 0.020 µA.

To prevent burn-up of the sample, the electron beam was defocused to 50 μm, and two different spots on the sample were used for each analysis after the area was checked for chemical homogeneity using the back-scattered electron detector. Each sample was analyzed for sodium first. Data for sodium were collected for 10 s to minimize burn-up. For all other elements in the samples, data were collected for 25 s or 0.50% precision, whichever was attained first. Data for standards were collected for 50 s or 0.25% precision, whichever was attained first.

Three analyses were performed on three grains. These analyses gave totals ranging from 99.87 to 100.63 wt. % (Table 1). A 100-s energy-dispersion spectrometer scan indicated no other elements with Z greater than 9 other than those reported here to be present. In addition, using WDS, the elements Y, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu and Th were sought but not detected. Concentrations of the REE were

TABLE 1. PETERSENITE-(Ce): CHEMICAL COMPOSITION

	1	2	3
Na ₂ O	17.38	16.81	17.16
CaD	1.32	1.31	10.54
BaO	0.32	0.22	
Sr0	1.70	1.66 1 4.96	3.98 11.60
La ₂ O ₃	14.49 23.66	24.40	14.99
Ce ₂ O ₃	2.00	2.16	1.49
Pr ₂ 0 ₃ Nd ₂ 0 ₃	5.82	5.89	3.34
Sm ₂ O ₃	0.60	0.51	0.50
COS	32.92*	32.08	35.24
	-		
Total	100.21	100.00	99.25
	Atomic propo	rtions based on 15	oxygens
Na	3.75	3.72	3.47
Ca	0.16	0.16	1.18
Ba	0.01	0.01	-
Sr	0.11	0.11	0.24
La	0.59	0.63	0.45
Ce	0.96	1.02	0.57
Pr	0.08	0.09	0.06
Nd	0.23	0.24	0.12
	0.02	0.02	0.02
Sm		T 00	
Sm C	5.00	5.00	5.00

Determined from crystal-structure analysis.

corrected for overlaps. The following standards were used: albite (Na), calcite (Ca), sanbornite (Ba), celestine (Sr), LaPO₄ (La), CePO₄ (Ce), PrPO₄ (Pr), NdPO₄ (Nd) and SmPO₄ (Sm). CO₂ was calculated by stoichiometry from the results of the crystal-structure analysis.

The empirical formula for petersenite-(Ce) based on 15 atoms of oxygen, as determined in the crystal-structure analysis, is $(Na_{3.75}Ca_{0.16})_{\Sigma 3.91}(Ce_{0.96}La_{0.59}Nd_{0.23}Sr_{0.11}Pr_{0.08}Sm_{0.02}Ba_{0.01})_{\Sigma 2.00}(CO_3)_5$ or, ideally, $Na_4(REE)_2(CO_3)_5$ for Z=4. The ideal formula of burbankite is $(Na_3Ca)_3(Sr_3Ba_3Ce)_3(CO_3)_5$ with Z=2, and that of remondite-(Ce) is $Na_3(Ca_3Ce_3Ca_3)_5$ with Z=2.

X-RAY CRYSTALLOGRAPHY AND CRYSTAL-STRUCTURE DETERMINATION

X-ray precession photographs show petersenite-(Ce) to be monoclinic with possible space-groups $P2_1$ and $P2_1/m$. Although the precession photographs show a strong hexagonal subcell related to burbankite and remondite-(Ce), the monoclinic supercell is so pronounced that a standard monoclinic setting, b axis unique, was chosen for the cell. X-ray powderdiffraction (XRPD) data obtained with a 114.6-mmdiameter Debye-Scherrer camera with CuKa (Nifiltered) radiation are given in Table 2, with refined unit-cell parameters and volume. For comparison, the X-ray powder-diffraction data for burbankite and remondite-(Ce) also are given in Table 2. The pronounced supercell of petersenite-(Ce) serves to easily differentiate it from either of the other two minerals and from the fourth member of the group, khanneshite, which is the Ba-analogue of burbankite and has virtually the same XRPD pattern. The strongest, differentiating lines in the petersenite-(Ce) pattern [d in Å] are: 6.84, 5.47, 4.13 and 3.18. In order to save space in Table 2, the d_{calc} for all overlapping hklreflections are averaged. Whether or not an hkl plane contributed to a reflection was determined from an XRPD pattern calculated from the crystal-structure refinement.

For the intensity-data measurements, a crystal fragment of petersenite-(Ce) from the cotype material (CMN # 81511) was ground to a sphere measuring 0.175 mm in diameter. Intensity data were collected on a fully automated Nicolet R3m four-circle diffractometer using the method of Grice & Ercit (1986). The data relevant to the structure determination are given in Table 3. There was no appreciable decrepitation of the crystal due to X-ray damage throughout the experiment. It is significant that most of the superlattice reflections are observed and quite intense.

The structure was solved using direct methods, and the refinement was done with the SHELXTL PC (Sheldrick 1990) package of programs. Scattering curves for neutral atoms and anomalous dispersion

Composition of petersenite-(Ce), results of three analyses averaged

^{2.} Calculated composition of patersenite-(Ce)
3. Composition of remondite-(Ce) (Cesbron et a7. 1988); total includes Eu₂O₂ 0.09, 6d₂O₃ 0.24, Dy₂O₃ 0.07, Ho₂O₃ 0.01, Er₂O₃ 0.03, Y₂O₃ 0.004, Y₂O₃ 0.024 wt. %.

TABLE 2. PETERSENITE-(Ce): X-RAY POWDER-DIFFRACTION DATA

	Pet	ersenit	e-(Ce)		Remon	dite-((e) ¹		Bu	rbankft	9 ²
I obs	d,Å obs	d,Å calc	hk?	I obs	d,Å obs	d,Å calc	hk1	I ob:	d,Å s obs	d,Å calc	hkl
3	9.13 6.84	9.13 6.86	001 101	- 2	9.04	9.04	101	2	9.103	9.105	100
3 2 2 5	5.47 5.22	5.46 5.22	-111 011	3	5.21 5.17	5.23 5.18	101 011	5	5.276	5.301	101
>1 >1 3	4.62 4.54 4.13	4.66 4.57 4.14	-311,111 -402 -501	2	4.52	4.53	002	2	4.533	4.533	200
>1 4	3.96 3.70	3.97	-112 -412	2	3.681 3.673	3.680 3.675	012 212	4	3.732	3.713	201
>1	3.46	3.47	-403	10	3 416	3.418	203	2	3.444	3.441	210
>1	3.28	3.29	-503	10	3.410	3.410	2.03	3	3.258	3 260	002
2	3.18	3.18	020	3	3.157	3.156	020	•			
2	3.04	3.04	-413	4	3.010	3.011	112	8	3.041	3.043	211
>1 >1	2.99 2.889	3.00 2.887	411 121	7	3.006	3.008	113	4	0 754	0.750	301
1	2.715	2,718	221	2	2.724	2.724	013	4	2.754	2./52	301
>1	2.649	2.650	-404	2	2.703	2.703	121	8	2.651	2.651	202
		2.611	-422	5	2.617		202	10	2.631	2.628	220
10	2.607	2.609	402	7 10	2.509 2.589	2.611 2.590	204 022				
>l >l >l >l	2.525 2.343 2.283 2.198	2.523 2.343 2.287 2.200	321 -423,-223 -523,303 -623,023	2 1	2.330 2.26	2.333 2.265	411 004)1 >1 >1	2.354 2.278 2.220	2.355 2.276 2.221	311 400 302
•		2.200	020,023	2	2.180	2.182	023	5	2.150		401
3	2.148	2.148	-913,313	7	2.132	2.132 2.128	014				
1	2.116	2.119	810	5	2.127	2.128	414	>1	2.112	2.114	103
_		2.007	-10.02,-230 -424	3	2.049	2.050	031	3	2.046	2.046	222
>1 2	2.032 2.018	2.017	-822,422	2 3 1	2.015 2.012 1.977	2.012	222 224 104	2	1.989	1.987	410
1	1.965	1.965	-10.12,621 -430,-432	2 1	1.977 1.971 1.962	1.9/2	104 215 324	2		1.961	203
3	1.921	1.924		4	1.908	1.908	032	_			
>1 >1 >1	1.854 1.831 1.807	1.855 1.836 1.809	-923,024 -115,-915 -921,721	2	1.840	1.840	024	2 >1	1.867 1.837	1.866 1.838	402 213
į >į	1.755 1.733 1.694	1.755	-425,-625 -10,22,622					3	1.756	1.759	322
>1 1 !	1.694 1.672 1.653 1.644 1.619	1.673	920,-225 821,-10.21 -12.14,414					>1 >1	1.698 1.664	1.697 1.664	412 421
>1 1 2 2 2 2 3 2 3 3 3 3 3 3 3 3 3 3 3 3	1.555 1.545 1.500 1.484	1.645 1.617 1.592 1.555 1.543	-921,721 -425,-625 -10,22,622 920,-225 821,-10,21 -12,14,414 812,-12,12 631,-831 040,-13,04 333,-933 830					>1 >1 >1	1.586 1.570 1.537	1.586 1.572 1.535	511 403 204
>1 >1 >1 >1 >1 >1 >1 >1 >1 >1 >1 >1 >1 >	1.467 1.446 1.415 1.405 1.387 1.370 1.358 1.344 1.322							>1	1.400	1.459	431

114.6 mm Debye-Scherrer camera, Cuga radiation, visually estimated intensities. Indexed with the aid of intensity data collected with a single-crystal diffractometer, on a cell having a 20.48(2), b.6.374(4), c 10.578(4) Å, B 120.43(8). Data from the literature: 1. Cesbrom et al. (1988), 2. Chao & Chen (1974).

TABLE 3. PETERSENITE-(Ce): STRUCTURE-REFINEMENT DATA

Ideal formula: Space group:	Na ₄ REE ₂ (CO ₃) ₅ P2 ₁	a 20.872(4) Å b 6.367(1) Å
Z:	4	c 10.601(2) A
Crystal size:	Sphere d = 0.175 mm	β 120.50(1)° V 1213.9(4) Å ³
Rad./Mon:	Mo/graphite	Total no. $ F_0 $: 3845
μ :	6.32 mm ⁻¹	$No[F] > 4\delta$: 3506
Min transmission	0.441	Final R (obs): 3.5%
Max transmission	0.467	Final R, (obs): 3.0%
$R = \sum (F_o - F_c)/2$ $R_u = [\sum w(F_o - F_c)]$	·	-

TABLE 4. PETERSENITE-(Ce): ATOMIC COORDINATES AND ISOTROPIC DISPLACEMENT COEFFICIENTS ($\mathbf{x}10^3$, \mathbf{A}^3)

Atom	х	у	Z	U(eq)
Na1 Na2 Na3 Na4 Na5 Na6 Na7 Na8	0.5812(2) 0.9187(2) 0.2332(3) 0.7562(3) 0.9776(2) 0.5225(2) 0.2667(3) 0.2562(2)	0.4996(16) 0.9940(16) 0.8037(8) 0.3159(8) 0.3071(8) 0.8085(8) 0.3030(8) 0.3164(8)	0.3247(3) 0.1620(4) 0.0356(4) 0.5245(4) 0.9731(4) 0.0187(4) 0.5694(4) 0.9880(4)	15(1) 15(1) 15(2) 15(2) 16(2) 15(2) 14(2) 13(2)
Ce1 Ce2 Ce3 Ce4	0.6583(1) 0.9206(1) 0.8417(1) 0.5794(1)	0.0000 0.9968(3) 0.4975(3) 0.4979(3)	0.6419(1) 0.6556(1) 0.3254(1) 0.8144(1)	8(1) 7(1) 9(1) 7(1)
C1 C2 C3 C4 C5 C6 C7 C8 C9	0.3293(5) 0.8291(5) 0.6966(5)0979(5) 0.1002(5) 0.8027(5) 0.5975(5) 0.0010(7) 0.5999(5) 0.5029(7)	0197(22) 0208(22) 0.5334(16) 0.5358(15) 0.0344(16) 0.0299(18) 0.0313(16) 0.8277(18) 0.0323(17) 0.8275(17)	1470(8) 0.8080(9) 0.6840(8) 0.0942(8) 0.3141(8) 0.2900(9) 0.2907(8) 0.5080(12) 0.8865(9) 0.4978(11)	9(2) 10(2) 6(2) 5(2) 7(2) 9(2) 6(2) 12(2) 8(2) 9(2)
01 02 03 04 05 06 07 08 09 010 011 012 013 015 016 017 018 019 020	0.2945(3) 0.4005(4) 0.7086(4) 0.7086(4) 0.7949(3) 0.7914(4) 0.6841(4) 0.6843(4) 0.106(5) 0.9562(4) 0.8870(4) 0.8156(4) 0.8156(4) 0.8156(4) 0.3555(4) 0.3555(4) 0.9586(4)	0192(18)0275(14)5257(15)0258(15)0155(20)0260(15) 0.3480(13)3656(13) 0.3252(12) 0.1279(13) 0.8518(15)3650(13) 0.3451(12)3742(12) 0.1360(13)1499(13) 0.1242(12) 0.3502(14) 0.1291(13) 0.6248(12) 0.3084(14)	0760(6) 0767(6) 0.2899(6) 0.8795(6) 0.6663(6) 0.8716(6) 0.7046(7) 0.5699(7) 0.8708(7) 0.8708(7) 0.9613(8) 0.7764(7) 0.7105(7) 0.2020(7) 0.2634(7) 0.4111(7) 0.7401(8) 0.2011(8) 0.5889(7) 0.5517(8)	13(1) 11(1) 13(2) 12(2) 12(1) 13(2) 10(1) 13(1) 13(1) 13(1) 13(2) 15(2) 11(1) 12(1) 12(1) 12(1) 12(1) 12(1) 12(1) 17(2)
023 024 025 026 027 028 029 030	0.0343(4) 0.0677(4) 0.6479(4) 0.3879(4) 0.4560(4) 0.5678(4) 0.5346(4) 0.4589(4)	1241(13) 0.3068(14) 0.1255(12) 0.3458(12) 3637(13) 0.3075(14) 0.8760(13) 0.3100(13)	0.5459(7) 0.5597(7) 0.8597(7) 0.0646(7) 0.1361(7) 0.5742(8) 0.4250(8) 0.3667(7)	17(2) 11(1) 9(1) 16(2) 17(2) 17(2) 14(2)

corrections were taken from Cromer & Mann (1968) and Cromer & Liberman (1970), respectively. From the E-map, the positions of the nine atomic sites with the highest scattering power were chosen and refined to R=17%. The difference-Fourier maps of subsequent refinements showed additional atomic sites.

Table 4 contains the final positional and isotropic displacement parameters for the petersenite-(Ce)

TABLE 5. PETERSENITE-(Ce): SELECTED BOND-LENGTHS (Å)

TABLE 5. I	ETEROLITE (Ge): GEEEGTEE	DOID LLITE (11)
Nal-02B Nal-03A Nal-08A Nal-019B Nal-020A Nal-021B Nal-027A Nal-029A	2.847(9) 2.869(10) 2.544(7) 2.504(14) 2.631(12) 2.656(12) 2.506(8) 2.978(13)	Na5-04 2.534(10) Na5-010 2.313(11) Na5-014 2.458(7) Na5-04C 2.444(8) Na5-010C 2.738(10) Na5-012B 2.292(13) Na5-013A 2.822(10) Na5-013C 2.537(8)
Na2-04A Na2-06A Na2-010B Na2-011B Na2-012A Na2-013B Na2-016A Na2-023C	2.829(8) 2.872(6) 2.599(13) 2.510(13) 2.542(13) 2.514(10) 2.534(11) 3.010(12)	Na6-020 2.730(10) Na6-02A 2.446(9) Na6-02C 2.555(10) Na6-019C 2.284(7) Na6-020B 2.310(8) Na6-026B 2.451(11) Na6-027B 2.537(12) Na6-027C 2.828(11)
Na3-011 Na3-012 Na3-01A Na3-06C Na3-07B Na3-09B Na3-016C Na3-025B	2.705(9) 2.284(11) 2.418(11) 2.431(11) 2.431(8) 2.746(8) 2.429(8) 2.425(9)	Na7-03B 2.468(10) Na7-05B 2.443(9) Na7-08B 2.425(11) Na7-015B 2.423(9) Na7-017B 2.410(11) Na7-018C 2.746(10) Na7-019A 2.282(8) Na7-021A 2.698(10)
Na4-05 Na4-07 Na4-014 Na4-018 Na4-03A Na4-08A Na4-021B Ce1-028 Ce1-05A Ce1-06A Ce1-06A Ce1-09A Ce1-09A Ce1-025C Ce1-025C Ce1-025C	2.476(12) 2.373(11) 2.431(8) 2.446(11) 2.383(8) 2.724(11) 2.688(10) 2.349(9) 2.732(8) 2.610(5) 2.493(8) 2.573(9) 2.731(8) 2.41(8) 2.554(9) 2.554(9) 2.554(9) 2.555(7) 2.688(9)	Na8-01B 2.493(12) Na8-06C 2.390(10) Na8-09C 2.441(8) Na8-011A 2.359(9) Na8-016C 2.722(8) Na8-017C 2.392(9) Na8-025B 2.684(8) Na8-025B 2.684(8) Na8-026A 2.446(10) Cc3-01B 2.697(9) Cc3-01C 2.731(5) Cc3-03A 2.612(8) Cc3-015B 2.565(8) Cc3-015B 2.565(8) Cc3-015B 2.567(8) Cc3-017A 2.518(8) Cc3-022A 2.686(7) Cc3-023D 2.557(9) Cc3-024A 2.569(8)
Ce2-022 Ce2-048 Ce2-05A Ce2-010A Ce2-013A Ce2-014A Ce2-015A Ce2-023A Ce2-023A	2.581(10) 2.609(8) 2.683(8) 2.482(6) 2.556(8) 2.483(9) 2.674(8) 2.532(6) 2.558(10) 2.704(10)	Ce4-025 2.683(8) Ce4-01D 2.687(5) Ce4-02C 2.599(7) Ce4-09A 2.524(10) Ce4-020B 2.463(10) Ce4-026B 2.476(8) Ce4-027D 2.553(9) Ce4-028A 2.718(9) Ce4-029B 2.564(6) Ce4-030A 2.591(8)
C1-01 C1-02 C1-03C	1.285(14) 1.281(11) 1.306(9)	C6-016 1.302(15) C6-017 1.227(14) C6-018B 1.315(13)
C2-04 C2-05 C2-06	1.294(11) 1.296(10) 1.269(15)	C7-019D 1.266(14) C7-020A 1.312(11) C7-021B 1.294(10)
C3-07 C3-08A C3-09A	1.262(13) 1.284(12) 1.312(14)	C8-022B 1.292(20) C8-023B 1.279(13) C8-024B 1.243(15)
C4-010D C4-011C C4-012C	1.279(13) 1.271(15) 1.277(13)	C9-025A 1.313(15) C9-026B 1.269(13) C9-027C 1.254(14)
C5-013D C5-014B C5-015C	1.253(11) 1.289(13) 1.299(10)	C10-029 1.284(19) C10-028B 1.277(14) C10-030B 1.244(12)

structure, and Table 5 lists the interatomic distances. The final stages of the least-squares refinement involved a conversion to anisotropic displacement parameters for the four Ce (rare-earth element) sites

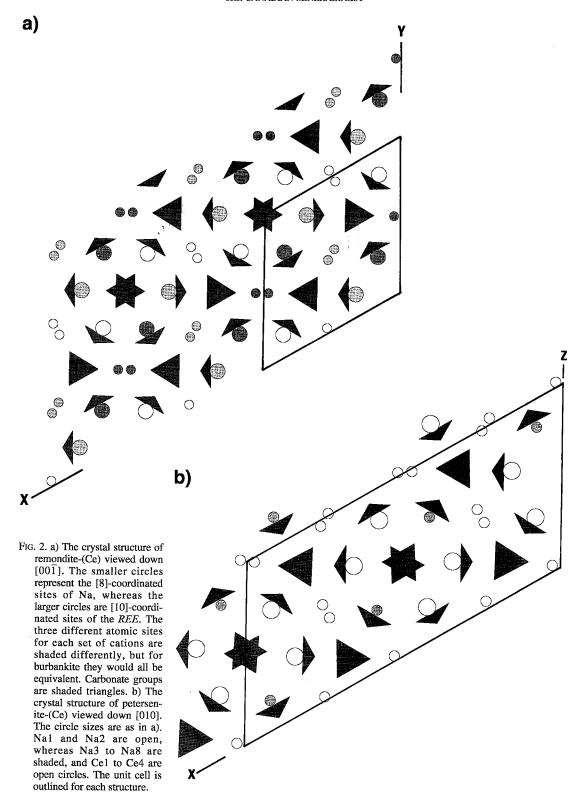
and the eight Na sites, and the addition of a weighting scheme involving an extinction correction. A refinement of the enantiomorph was tried but without significant improvement in the final R values. The observed and calculated structure-factors, as well as the anisotropic displacement factors, have been submitted to the Depository of Unpublished Data, CISTI, National Research Council of Canada, Ottawa, Canada K1A 0S2.

DESCRIPTION AND DISCUSSION OF THE STRUCTURE

The crystal structure of petersenite-(Ce) has many similarities to that of burbankite, which was solved by Voronkov *et al.* (1967) and later refined by Effenberger *et al.* (1985). The crystal structure of burbankite (Fig. 2a), in the noncentrosymmetric spacegroup $P6_3mc$, has two large cationic sites. The A site, with [8] coordination, is occupied by Na and Ca atoms, whereas the B site, with [10] coordination, is occupied by Sr and lesser amounts of Ba, Ca and rareearth elements.

The crystal structure of remondite-(Ce) (Fig. 2a) (Ginderow 1989), has the same unit-cell volume as burbankite, but the symmetry is lowered to that of the noncentrosymmetric space-group $P2_1$ owing to ordering involving the large cations. The A site of burbankite becomes three sites occupied entirely by Na atoms having [8] coordination in the remondite-(Ce) structure, whereas the B site of burbankite degenerates into three M sites, each with [10] coordination in remondite-(Ce). The order in remondite-(Ce) seems to primarily involve the M sites, as evidenced by their different refined occupancy-factors (Ginderow 1989). This subtle degree of order is sufficient to give biaxial optical properties and some doubling of lines in the XRPD pattern (Table 2).

In the petersenite-(Ce) structure, the order of cations is more pronounced than in remondite-(Ce). Although the distribution of the carbonate groups and the large cations in petersenite-(Ce) (Fig. 2b) is similar to that of burbankite and remondite-(Ce) (Fig. 2a), there is a change in the ratio of cations in [8] coordination (i.e., Na sites) to cations in [10] coordination (i.e., Ce sites). In burbankite and remondite-(Ce), this coordination ratio of [8] to [10] is 1:1, whereas in petersenite-(Ce) it is 2:1. This shift in the structure is sufficient to give the pronounced supercell, which is readily seen in both the XRPD pattern (Table 2) and the single-crystal X-ray-diffraction films. Refining the occupancy factor of each of the Na sites reveals that the Na1 and Na2 sites contain some atoms of higher atomic number (i.e., Ca or REE). In Figure 2a, it is evident that these two sites are [10]-coordination sites in remondite-(Ce) and subsequently in burbankite. Thus we may conclude that an increase in the proportion of Ca or REE (Table 1) causes a change from



the petersenite-(Ce) structure to the remondite-(Ce) structure while maintaining the same space-group but losing the supercell.

As (CO₃) groups do not polymerize owing to simple bond-valence considerations (Grice & Ercit 1986), the crystal structures of carbonate minerals tend to be layered, with (CO₃) slabs being interleaved with layers of large cations, predominantly alkaline, alkaline-earth and rare-earth cations. This rather simple arrangement gives rise to very few structuretypes but a large number of mineral species. The great diversity of species results from: 1) the numerous possible substitutions in the layer of cations, 2) variability of cation:(CO₂) ratio, 3) the variation in thickness and content of large cations of the carbonate layer, 4) addition of an (H₂O) layer, and 5) the addition of other anionic groups such as $(SiO_4)^{4-}$, $(PO_4)^{2-}$ and $(BO_3)^{3-}$ or $(BO_4)^{5-}$. The following brief discussion concerns itself with REE carbonate minerals only and their structural relationship to petersenite-(Ce). It will show how subtle changes in the crystal structure gives rise to so many species.

The thickness of the carbonate layer is determined by the orientation of the triangular polyhedron, which in its two extremes may be either "flat-lying" or

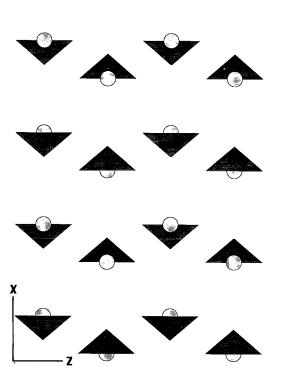


Fig. 3. The crystal structure of ancylite-(La) viewed down [010], showing "corrugated, flat-lying" carbonate layers and large REE atoms.

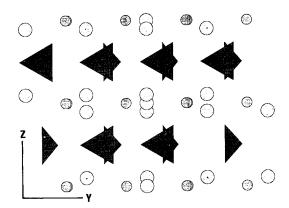


Fig. 4. The crystal structure of bastnäsite-(Ce) viewed down [100], showing "standing-on-edge" carbonate layers and the CeF layer, with Ce a small circle, and F, a large circle.

"standing-on-edge". The "flat-lying" carbonate layers are of two types: a uniformly flat layer, as in the calcite and dolomite structures, for which no REE carbonates have been reported, and a "corrugated" layer, as in the aragonite structure, to which ancylite-(La) may be compared (Dal Negro et al. 1975). Figure 3 shows the "corrugated" layer of carbonate groups interleaved with the (REE, Ca) atoms in [10] coordination. REE minerals having "standing-on-edge" carbonate layers are much more common. The fluorcarbonate minerals, which are members of the bastnäsite parisite - röntgenite - synchysite group, owe their syntactic intergrowths to various stacking combinations of (CeF), (CO₃) and [Ca(CO₃)] layers (Ni et al. 1993). From the crystal-structure analysis of cebaite-(Ce) (Li & Hashimoto 1984), it is evident that the barium fluorcarbonates cebaite-(Ce), huanghoite-(Ce) and zhonghuacerite-(Ce) have syntactic intergrowths similar to those of the bastnäsite group. Figure 4 shows the bastnäsite-(Ce) structure (Ni et al. 1993) as an example of this type of carbonate "standing-onedge" layering.

In addition to the "flat" and "standing-on-edge" carbonate structures described above, a whole series of modifications exist with mixtures of the two types of layer. These crystal structures are invariably complex and found only in rare minerals. Figure 5 shows an integral carbonate layer in sahamalite-(Ce) (after Pertlik & Preisinger 1983) along (110), which is composed of "flat", "standing-on-edge" and "standing-on-base" slabs of carbonate. These are sandwiched by Mg and REE polyhedra. In the lanthanite structure (Dal Negro et al. 1977), there is a distinct "standing-on-edge" layer of carbonate on (100) between La polyhedra, with another (CO₃) polyhedron, "standing-on-base" sharing the water layer on (010) (Fig. 6).

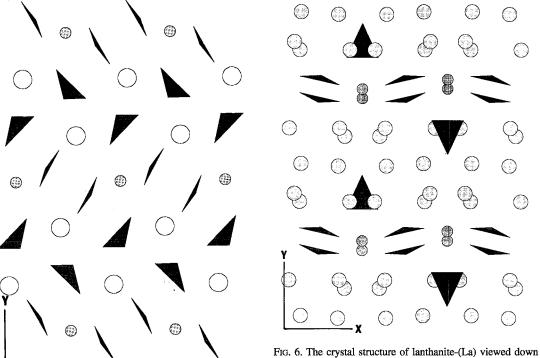


Fig. 5. The crystal structure of sahamalite-(Ce) viewed down [001], with the carbonate groups layered, but with varying orientations. The Mg is a small circle, and Ce, a large circle.

Fig. 6. The crystal structure of lanthanite-(La) viewed down [011], showing the two types carbonate layers, "standing-on-edge", and "standing-on-base", plus water interleaved with layer of La polyhedra.

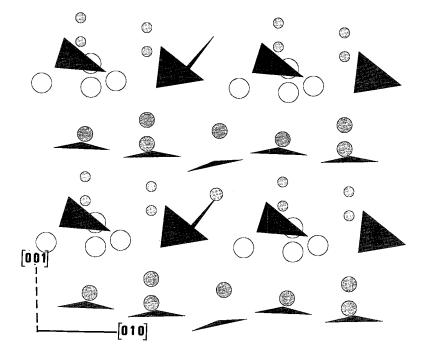


Fig. 7. The crystal structure of weloganite viewed down [100] (rotated 7° about [010]), showing the two types of carbonate layers, "flat-lying" and "standing-on-base", plus water (large, open circles) interlayered with layers of large cations. Na atoms are shown as small, lightly shaded circles, Zr, as circles of medium size and shading, and Sr atoms, as large, darker shaded circles.

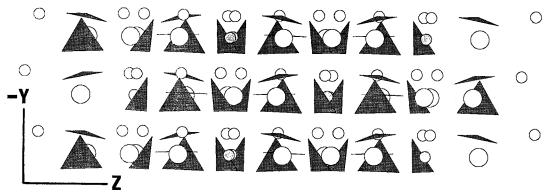


Fig. 8. The crystal structure of petersenite-(Ce) viewed down [100], showing the "flat-lying" [Na₃(CO₃)] layers and inclined [REE₂Na(CO₃)₄] layers. Atom designation is the same as in Figure 2, with smaller circles representing Na atoms, and larger circles, the REE atoms.

Donnayite-(Y) and mckelveyite-(Y) have the weloganite structure (Fig. 7) (Grice & Perrault 1975). In this structure, the layering is parallel to (001), with a carbonate - water layer of the type described for the lanthanite structure, as well as a "flat-lying" carbonate layer, but no longer consisting entirely of carbonate groups in the sense that it also contains Na atoms. It is this complexity of mixed layering that we have in petersenite-(Ce) and other members of the burbankite group. The layering on (010) (Fig. 8) is composed of four slabs; $2 \times [REE_2Na(CO_3)_4]$ and $2 \times [Na_3(CO_3)]$. In the first type of slab, the (CO₃) polyhedra are both "flat-lying" and inclined "standing-on-base", whereas in the second type of slab, the (CO₃) polyhedra are approximately "flat-lying". These mixed layered structures, although reducing the possibility of syntactic intergrowth, increase the structural variability and order within the REE carbonates.

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