

Metabarcoding Reveals High Diversity of Benthic Foraminifera Driven by Atlantification of Coastal Svalbard

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1 Metabarcoding reveals high diversity of benthic foraminifera driven by atlantification of

2 coastal Svalbard

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12 Abstract (300 words)

Arctic marine biodiversity is undergoing rapid changes due to global warming and modifications of oceanic water masses circulation. These changes have been demonstrated in the case of mega- and macrofauna, but much less is known about their impact on the biodiversity of smaller size organisms, such as foraminifera that represents a main component of meiofauna in the Arctic. Several studies analysed the distribution and diversity of Arctic foraminifera. However, all these studies are based exclusively on the morphological identification of specimens sorted from sediment samples. Here, we present the first assessment of Arctic foraminifera diversity based on metabarcoding of sediment DNA samples collected in fjords and open sea areas in Svalbard Archipelago. We obtained a total of 5,968,786 reads that represented 1,384 ASVs. More than half of the ASVs (51.7%) could not be assigned to any group in the reference database suggesting a high genetic novelty of Svalbard foraminifera.

The sieved and unsieved samples resolved comparable communities, sharing 1023 ASVs, comprising over 97% of reads. Our analyses show that the foraminiferal assemblage differs between the localities, with communities distinctly separated between fjord and open sea stations. Each locality was characterized by a specific assemblage, with only a small overlap in the case of open sea areas. Our study demonstrates a clear pattern of the influence of water masses on the structure of foraminiferal communities. The stations situated on the western coast of Svalbard that is strongly influenced by warm and salty Atlantic Water (AW) are characterized by much higher diversity than stations in the northern and eastern part, where the impact of AW is less pronounced. This high diversity and specificity of Svalbard foraminifera associated with water mass distribution indicate that the foraminiferal metabarcoding data can be a very useful tool for inferring present and past environmental conditions in the Arctic.

Keywords (4 to 6 keywords)

37 Benthic foraminifera, sedimentary DNA, metabarcoding, Svalbard, atlantification

1. Introduction

The Arctic Ocean is strongly impacted by the increased influence of warm and saline Atlantic Water, so-called "atlantification", which causes sea-ice retreat and sea surface temperature increases ¹⁻³, directly affecting the entire ecosystem. Such changes in physical drivers lead to shifts of Atlantic species ranges towards the Arctic ⁴, increase in productivity ⁵, and changes in the timing of spring phytoplankton bloom ⁶. These changes may create shifts in food webs, affecting planktonic and bottom communities ^{6,7}. Particularly, Svalbard ecosystems are currently affected by increased heat transport from northward-flowing currents ^{3,8-10}. The changing environmental conditions in this region introduces a significant impact on shaping biodiversity and the biogeography of many taxonomic groups, such as birds and mammals ^{11,12}, fish ^{13,14}, zooplankton ¹⁵⁻¹⁷, phytoplankton ^{18,19}, and planktonic foraminifera ^{20,21}.

Foraminifera are a major component of benthic communities, from transitional and marine coastal areas to the deep-sea zones, where they have been demonstrated to be of significant importance in terms of both high abundance and diversity ²²⁻²⁴. Benthic foraminifera are recognized as important ecological indicators of environmental stress because they are particularly sensitive to abrupt climate change ²⁵⁻²⁷.

Traditionally, the morphological observations of foraminiferal mineral shells (so called tests), which belong to either the class Tubothalamea or Globothalamea, are the basic feature used to assess foraminiferal diversity. However, morphological identification is time-consuming and expertise-demanding, making it costly and unpractical, particularly for large-scale surveys. Recently, the metabarcoding of environmental DNA (eDNA) samples have provided new insights into biodiversity and ecological distribution of numerous taxonomic groups and offer an alternative to the traditional morphology-based approach ²⁸⁻³⁰. Metabarcoding consists in high-throughput sequencing of short DNA barcodes that include enough information for species identification to get a comprehensive inventory of all organisms

present in a given sample, e.g., foraminiferal sequences derived from the 37f hypervariable region of the 18S small subunit (SSU) rRNA gene ^{31,32}. To better understand large-scale patterns of biodiversity and distribution in various groups, this method is increasingly being employed, particularly in marine environments ^{22,29,33,34}. While numerous foraminiferal metabarcoding studies were conducted in various coastal areas ³⁵⁻³⁸, and the deep-sea ^{22,39-41}; the application of eDNA metabarcoding to monitor foraminiferal diversity in the Arctic was limited to a few paleogenomic studies using foraminifera as proxies in palaeoceanographic reconstructions ⁴²⁻⁴⁴.

In the conventional morphology-based foraminiferal studies the sediment samples are sieved before the specimens are sorted. In metabarcoding approaches, the DNA is also extracted from sieved sediment samples, which has several benefits: potentially reducing sample heterogeneity, detecting completely small and low abundant taxa, achieving a higher number of reads, or decreasing primer bias due to the reduction in the amount of DNA template produced by the large specimens ^{45,46}. Some DNA metabarcoding studies have shown that the preprocessing of samples does not significantly alter metazoan diversity patterns ^{47,48}. However, the effectiveness of sieving versus non-sieving in the case of foraminiferal metabarcoding has not been examined yet.

The two main goals of this study are to investigate whether metabarcoding of sieved sediment is effective for the assessment of foraminiferal biodiversity and how the foraminiferal communities respond to rapid environmental shifts of Arctic marine ecosystems. Taxonomic composition, diversity, and distribution of benthic foraminifera were analyzed in fjords and open water areas in Svalbard in order to 1) compare species composition and diversity patterns inferred from sieved and unsieved sediment samples, 2) describe the spatial diversity of Svalbard foraminiferal communities, and 3) identify new potential bioindicators of atlantification.

2. Study area

The oceanography of Svalbard region is shaped mainly by the interplay between warm and saline Atlantic Water (AW) and cold Arctic Water (ArW), as well as locally formed water masses ^{16,49}. AW is transported northward along the Spitsbergen shelf edge as the West Spitsbergen Current (WSC, Fig. 1) ^{50,51}. WSC is one of the major heat contributors to the Arctic Ocean⁵², transporting heat from low latitudes into the Arctic and transferring it to the atmosphere and adjacent water masses ⁵³. Between 78°N and 80°N, the WSC bifurcates into an eastern (Svalbard) branch and a western (Yermak) branch ⁵⁴. The Svalbard Branch flows northeasterly, staying close to the continental margin of Svalbard ⁵⁴. The Yermak Branch streams northwards and further recirculates southward as the Return Atlantic Current ⁵⁵. The Svalbard area is also under the influence of cold Arctic Water (ArW) that is transported from the north-eastern Barents Sea by the East Spitsbergen Current (ESC), also called Sørkapp Current or the Coastal Current ⁵⁶. Mixing of ArW and AW results in the formation of Transformed Atlantic Water (TAW) which expanded across the shelf and penetrated the fjords ^{49,57}.

Isfjorden (IS) and Wijdefjorden (WIJ) are located on the west coast Spitsbergen, along the main pathway of AW inflow. Both fjords are linked directly to shelf and slope area ^{10,58} and therefore, their oceanographic conditions are shaped mainly by the inflow of AW and TAW. Isfjorden is considered to be the most AW-impacted fjord of Spitsbergen ⁵⁷. Rijpfjorden (RIJ) is a north-facing fjord, located on the northern coast of Nordaustlandet. The oceanography of Rijpfjorden is dominated by cold ArW, with a less pronounced impact of AW. However, episodic inflows of AW may occur in ice-free periods. As such, it is considered to be a typical Arctic fjord. In most of the year, Rijpfjorden is covered by sea ice and/or drifting ice packs ⁵⁹.

The southeastern Nordaustlandet (NAL) and the eastern Edgeøya (EDG) are strongly impacted by the presence of large ice caps, making them one of the largest glacierized areas of Svalbard ⁶⁰. The tidewater cliffs supply the surrounding areas with large amounts of turbid meltwater ⁶¹. Water masses around Nordaustlandet and Edgeøya are dominated by ArW, carried by the ESC. However, in periods of strong WSC activity, the presence of AW is also pronounced ⁶².

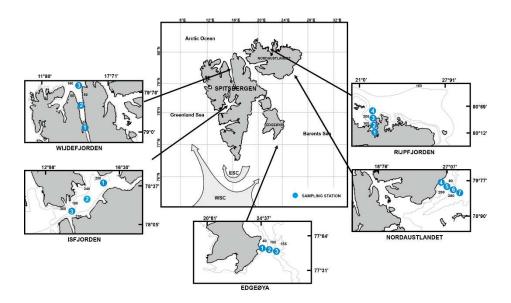


Figure 1. Map showing the location of sampling stations. Abbreviations: WSC - West Spitsbergen Current; ESC - East Spitsbergen Current.

3. Material and Methods

3.1. Sampling

The samples were collected at 15 sampling stations from five localities at Western, Northern and Eastern sides of the Svalbard Archipelago (Fig. 1), including three fjord sites (Isfjorden, Wijdefjorden, Rijpfjorden) and two open marine areas in front of tidewater glaciers (Edgeøya, Nordaustlandet). Sampling stations coordinates and sampling depths can be found in Table S1. Surface sediment samples were collected with the use of a box corer during the

cruise of R/V *Oceania* in August 2016. The upper 2 cm of sediment has been sampled from the surface of approximately 50 cm^2 . Samples for sedimentary eDNA analysis were split into two: one half remained unsieved and the other half has been wet sieved on $500 \mu m$, $100 \mu m$, and $63 \mu m$ sieves. A fraction smaller than $63 \mu m$ was retained. Samples were transferred to sterile containers and frozen at -20° C. In each sampling station, physical properties of the water column from a vertical conductivity-temperature-depth (CTD) profiler were obtained using a Mini CTD Sensordata SD202 at intervals of 1 second. Water temperature was reported in degrees Celsius (°C), turbidity was presented in Formazine Turbidity Units (FTU). Water masses were classified according to Cottier, et al. 49 . Supplementary Table S2 contains detailed information.

3.2. Metabarcoding analyses

For the sieved samples, the genomic DNA from size fractions >500 μ m, 500-100 μ m, and 100-63 μ m were extracted from 0.25 g of sediment sample with DNeasy PowerSoil Kit (Qiagen, Hilden, Germany). The sediment fraction < 63 μ m and the unsieved part were extracted using DNeasy PowerMax Soil Kit (Qiagen, Hilden, Germany), which allows for the process of up to 10 g of sediment. In total, five amplicon libraries per station were prepared, corresponding to the fractions > 500 μ m, 500-100 μ m, 100-63 μ m and < 63 μ m, as well as unsieved sample.

The foraminifera-specific 37f hypervariable region of 18S rRNA gene was PCR amplified with the primers s14F1/s15 ^{31,44}, tagged with unique sequences of 5 nucleotides appended at 5' ends. Primer sequences and PCR conditions are detailed in Table S3. For each sample, 3 PCR replicates were obtained. PCR products were visualized by 1.5% agarose gel electrophoresis, quantified with Qubit 3.0 fluorometer (Thermo-Fisher Scientific Inc., Waltham, MA, USA), and the pool was purified with High Pure PCR Cleanup Micro Kit

(Roche Diagnostics GmbH, Mannheim, Germany). Library preparation was performed with TruSeq[®] DNA PCR-Free LT Library Prep Kit (Illumina Inc., San Diego, CA, USA) and was loaded onto a MiSeq instrument for a paired-end HTS run of 2 × 150 cycles using a v2 kit.

3.3. Data quality control and processing

Bioinformatics analyses were performed using the web application SLIM ⁶³. The reads were first demultiplexed using the double tag demultiplexing algorithm based on their unique barcode sequences. DADA2 ⁶⁴ was used for quality trimming and filtering sequences, dereplicating sequences, inferring amplicon sequence variants (ASVs), merging of forward and reverse sequences, and detection and removal of chimeras. A raw number of sequence reads were published by Pawłowska, et al. ⁶⁵. Subsequently, all the resulting ASVs tables were curated with the LULU algorithm ⁶⁶ to remove erroneous ASVs following the online tutorial (https://github.com/tobiasgf/lulu) with default parameters. Final quality filtering of ASVs involved the removal of unique (occurring in only one sample) and rare ASVs (having less than 10 reads).

The remaining ASVs were compared to the curated database of foraminiferal 18S rDNA sequences ^{67,68} and the PR2 database v4.11.1 ⁶⁹ using VSEARCH, implemented in SLIM, and BLASTN ⁷⁰ based on minimum similarity (-perc_identity 80%) and minimum coverage (-qcov_hsp 80%) for the taxonomic assignment to six taxonomic levels (phylum; class; order; family; genus; species). The representative sequences of ASVs that remained unclassified with the foraminiferal database, were aligned in a stand-alone BLAST using BLAST (v2.7.1) search against the NCBI's non-redundant nucleotide database. The sequences diverging by less than 1% were considered as belonging to the same species/genus. ASVs below 99% identity were classified at the family, order, or class or as unassigned foraminifera.

Finally, taxonomic compositions in terms of cluster abundance were compared among processing methods only using clusters reliably assigned at the species/genus level.

3.4. Statistical analysis

Before statistical analyses, the sample matrix was filtered to remove ASVs that were classified as planktic, or non-foraminifera. For each sample, datasets of 4 size fractions were combined as a sieved dataset and compared to an unsieved dataset in further analysis. All statistical analyses were performed in R, version 4.1.0 71 . All formal hypothesis tests were conducted on the 5% significance level ($\alpha = 0.05$).

To compare the community composition among methods and size fractions, Venn diagrams were constructed using *venn* package ⁷². The ASVs rarefaction curves were calculated to visualize whether or when a plateau was reached based on the number of eventually retained ASVs and reads using the iNEXT package ⁷³. We analyzed differences in the beta-diversity of the community composition by calculating a Non-Metric Multi-Dimensional Scaling (nMDS) on Bray-Curtis similarity coefficient using *vegan* package ⁷⁴ and a heatmap based on the Spearman's correlation with *pheatmap* package ⁷⁵. The influences of environmental factors were calculated with the *envfit* function. A global one-way analysis of similarities (ANOSIM) and permutational multivariate analysis of variance (PERMANOVA) were computed using the function *anosim*, *adonis* and *envfit* with 999 permutations and the Bray-Curtis distance matrix to test whether there were significant differences in community composition among methods and locations of sampling units using log10(1+x) transformed read abundance data.

Finally, Sparse Partial Least Squares (PLS) regression, available in the *mixOmics* package ^{76,77}, was used for the multivariate analysis of the combined foraminiferal datasets at ASVs level to identify ASVs that were more predictive of the observed environmental

response. Pairwise similarity matrices of an sPLS model with 2 components were computed and displayed by *cim* function. This approach enabled us to identify high correlations between certain ASVs and environmental parameters but without considering the structure of the foraminiferal community.

3.5. Data availability

All raw sequencing reads have been deposited in the NCBI Short Read Archive (SRA) database under Bioproject accession number PRJNA768352.

4. Results

4.1. CTD data

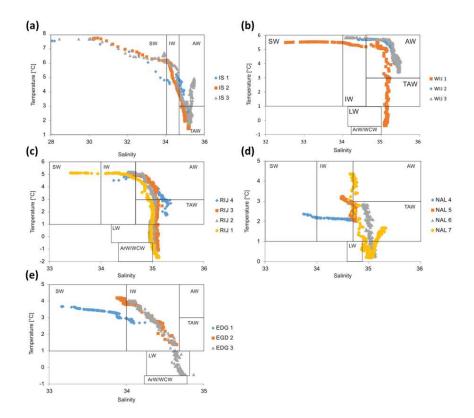


Figure 2. Temperature (°C) and salinity in Svalbard stations: (a) IS - Isfjorden, (b) WIJ - Wijdefjorden (c) RIJ - Rijpfjorden, (d) NAL-Nordaustlandet and (e) EDG - Edgeøya. Abbreviations: AW- Atlantic Water; TAW- Transformed Atlantic Water; ArW- Arctic Water;

WCW- Winter Cooled Water; SW- Surface Water, LW - Local Water, IW - Intermediate water.

Water masses are classified after Cottier et al. ⁴⁹.

Temperature, salinity, and turbidity for all sampling stations are presented in Figs. 2 and S1, respectively. AW and TAW dominated the water masses in Isfjorden (Fig. 2a) and Wijdefjorden (Fig. 2b), which have the highest temperatures and salinities of the investigated stations. Additionally, Surface Water (SW) and Intermediate Water (IW) were recorded at all stations in Isfjorden and station WIJ1. In Isfjorden, the highest temperature of 7.7°C was observed at the surface and progressively decreased towards the bottom to 1.4°C. Similarly, the temperature fluctuated from 5.8°C at the surface to - 0.4°C near the bottom of Wijdefjorden. Water temperatures above 0°C were noted up to 111 m depth at the station WIJ1 and in the whole water column at other stations. Salinity was the lowest at the surface, reaching 28.1 in Isfjorden and 32.5 in Wijdefjorden, respectively. The lowest salinity was recorded at inner Isfjorden and WiJdefjorden (IS1 and WIJ1) and the highest values near the mouth of these fjords (IS3 and WIJ3). The turbidity increased from the inner fjord (0.1 FTU) towards the fjord's mouth (particularly up to 12.5 FTU) in Isfjorden, reaching its maximum in the surface water at the station IS3. In contrast, turbidity was the highest in the surface layer and decreased from the inner fjord towards the fjord mouth, ranging from 5.2 to 0.1 FTU.

Rijpfjorden was characterized by the lowest near-bottom temperatures (as low as -1.7 °C) among the studied fjords. Three sites (RIJ1, RIJ2, and RIJ3) reported cold and saline Winter Cooled Water (WCW) in addition to other water masses (AW, TAW, SW, and IW). The temperature ranged from 5.2°C at the top to -1.7°C near the bottom, with the salinity varying from 33.4 to 35.4 throughout the water mass. The lowest near-bottom temperature was noted at the station RIJ1. The water temperature of the whole water column was observed to

be above 0°C at station RIJ4. Turbidity reached over 12.5 FTU at the station RIJ 3 in the near-bottom water layer and decreased to 0.1 FTU towards the mouth of the fjord.

In the region of eastern Svalbard, the Nordaustlandet stations were generally under influence of TAW, whereas AW was noted only at the glacier-distant station NAL7. NAL4 was the sole station where neither TAW nor AW was detected. SW and IW were other water masses recorded near the Nordaustlandet. Water temperature oscillated between 4.4°C at the surface to 0.2 °C near the bottom. Salinity at the surface ranged from 33.6 to 35.3 and increased towards the glacier-distant stations. The water column had relatively low turbidity (<1 FTU). The only exception was glacier-proximal station NAL 4, where turbidity reached 54.8 FTU, which was the highest value of all studied sites.

Edgeøya stations were the most distinct locations, with the absence of Atlantic-origin waters. The IW was detected at all stations, while Local Water (LW) occurred only at station EGD3. Towards the bottom of the stations, the temperature in the water column varied between 4.2 °C and -0.5 °C and salinity ranged from 33.2 to 34.9. The highest turbidity was recorded at EGD1, it increased with depth to reach 48.8 FTU near the bottom. At the other stations, turbidity values oscillated from 0.3 to 4.3 FTU.

4.2. Metabarcoding data

We obtained a total of 5,968,786 raw paired-end reads. After bioinformatic processing, the numbers of the raw reads were reduced to 5,579,202 with 4,836,419 in a sieved dataset, and 742,783 in an unsieved dataset. The number of reads per sample is indicated in Table S4. One sample, unsieved IS2, produced a low number of reads and was not included in the analysis of foraminiferal diversity. After LULU curation step and strict filtering of ASVs, 1384 ASVs (1354 ASVs of sieved and 1053 ASVs of unsieved samples) representing 5,483,500 reads

(98.28% of the total reads count) were retained for downstream analysis. The average number of sequences per station were 317,306 for the sieved and 51,976 for the unsieved datasets.

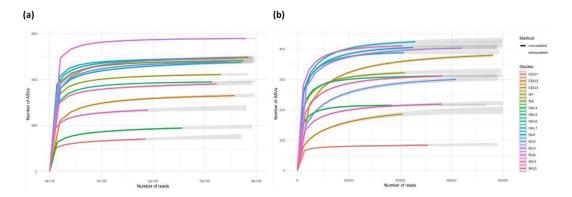


Figure 3. Rarefaction curves showing the relationship between sequencing depth and species richness in amplicon sequence variant (ASVs) of the 37f libraries from 15 stations of (a) unsieved (combined fractions) and (b) sieved samples.

The rarefaction curves were plotted at the sample level based on the number of retained ASVs and reads (Fig. 3). The rarefaction curves showed that the filtered ASV datasets reach saturation levels, indicating that most of the diversity had been captured and allowing for richness comparison among samples for all individual stations of each location and both methods. Considering the accumulation of ASVs richness across two datasets, sieved datasets exhibited the higher saturation degree of samples, respectively, and the diversity of NAL, IS, RIJ stations are higher than the individual station of EDG and WIJ.

4.3. Taxonomic composition of foraminiferal metabarcodes

Overall, the retained sequences were assigned to 1,384 foraminiferal ASVs. Among them, 758 ASVs were assigned to the class Monothalamea, 252 ASVs were assigned to the class Globothalamea and only 14 ASVs were assigned to the class Tubothalamea (Table S5). The 360 ASVs, classified as Foraminifera_X, had low similarity levels or were mainly assigned

to environmental sequences, that possibly represent non-described taxa. More than half of the ASVs (51.73%) were assigned with low similarity (< 0.9).

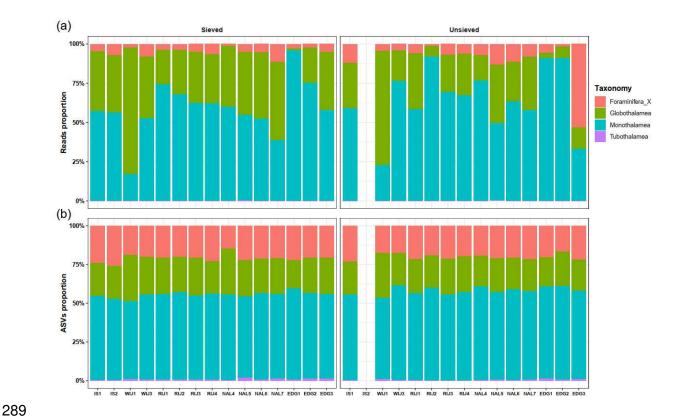


Figure 4. Proportions of reads (a) and ASVs (b) assigned to different foraminiferal classes detected in sieved and unsieved samples at different sites.

The sieved and unsieved sediment DNA samples resolved comparable communities at the class level (Fig. 4). Pairwise comparisons indicated no overall significant differences in community composition between sieved and unsieved datasets (ANOSIM statistic R <0, p > 0.05 and PERMANOVA, Table S6). In Fig. 5a, the Venn diagram showed that 1023 ASVs (corresponding to 97.23% of the reads) were shared among sieved and unsieved samples. The sieved dataset had 331 unique ASVs, while unsieved dataset comprised only 30 unique ASVs (corresponding to 0.1 and 2.67 % of the reads, respectively).

Taxonomic composition in sieved samples noticeably changed between size fractions (Figs. 5b,c). The < 63 μ m fraction comprises 1151 ASVs, corresponding to 83% of ASVs (Fig. 5b). It also recovered more unique ASVs (177, corresponding to 13.07 % of ASVs) than any other fractions. Shared foraminiferal ASVs among fractions including 424 ASVs (corresponding to 82.39% of the reads), mostly belonged to Monothalamea (59.72%), and Globothalamea (Rotaliida 25.17%, Textulariida 10.60%) as shown in Fig. S2. In all fractions (Fig. 5c), the monothalamous taxa made up from 58% to 80% of reads. Non-described monothalamiids dominated in the 500-100 μ m and < 63 μ m fractions (30.48 % and 33.85% of reads, respectively). For multichambered globothalamids, order Textulariida accounted for 15% to 20% of the reads in 500-63 μ m fractions, while Rotaliida represented more than 20% of reads in > 63 μ m fractions. Interestingly, most reads of unique ASVs were assigned to specific foraminiferal groups in each fraction, e.g., Foraminifera_XX (70.69%) in > 500 μ m fraction, Rotaliida (66.21%) in 500 - 100 μ m fraction, Textulariida (39.28%) in < 63 μ m fraction, and Clade Y of Monothalamea (70.24%) in 100 - 63 μ m fraction, see Fig. S2.

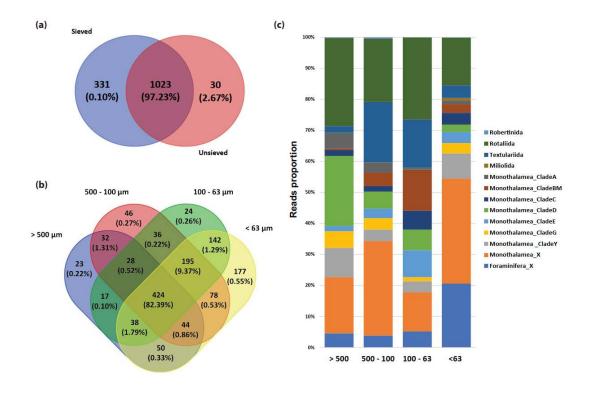


Figure 5. Venn diagrams showing the shared and unique numbers ASVs and proportion of reads for the sieved (combined fractions) and unsieved samples (a), and between different size fractions (b). Bar plots represent the proportion of reads assigned to the taxonomic composition of each fraction by order/clade (c).

The taxonomic composition of benthic foraminifera also changed between the locations. At the class level (Fig. 4), the monothalamous taxa were the dominant group which accounted for an average of 56.06 % and 61.77% of total ASVs and reads in both datasets, respectively. The highest proportion of monothalamiids (95.80%) was observed at the station EDG1 in the sieved dataset. The contribution of monothalamiids decreased in the deeper EDG stations in favour of the class Globothalamea. Comparatively, the average proportions of ASVs and reads assigned to Globothalamea were 22.60% of the ASVs and 30.75% of reads, respectively. The highest relative abundance of Globothalamea (80.80%) occurred in the WIJ3_Sieved sample. The Tubothalamea represented only a minor part of the total community (0.06% of ASVs and 7.42% of reads on average).

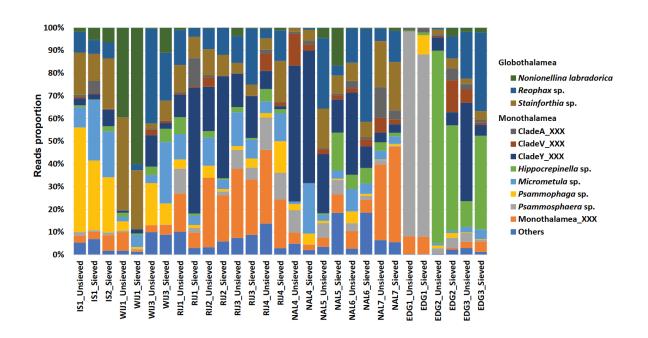


Figure 6. Patterns of relative abundance of dominant genera and species in the sieved (combined fractions) and unsieved sediment samples for each location.

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The variations of foraminiferal assemblage between different sampling localities were also reflected in the taxonomic composition of foraminiferal assemblages at the lower taxonomic level. To compare species composition in each station, all ASVs which had an identity percentage with the reference database of more than 99% and no less than 10 reads were picked and those attributed to the same taxa were merged (Fig. 6).

In the stations located at the western coast of Spitsbergen (IS, WIJ), foraminiferal communities were dominated by genera Psammophaga and Micrometula, which together made up to 57.35% of reads. The WIJ1 was the only station where the majority of sequences belonged to a textularid Reophax sp. and a rotalid Stainforthia sp. Another rotalid species Nonionellina labradorica was present at all IS and WIJ stations, mainly distributed in WIJ1 (sieved: 59.99%, unsieved: 39.46%), WIJ3_Sieved (11%) and IS2_Sieved (6.5%). The northern stations (RIJ) were dominated by monothalamiids assigned to Clade Y and Monothalamea XXX. At the outermost station RIJ4, also higher percentages of *Psammophaga* sp. and Psammosphaera sp. sequences occurred. At the stations located at eastern Svalbard (NAL), the number of Clade Y sequences decreased towards the glacier-distant stations, while the percentage of Monothalamea_XXX increased. Also, glacier-distal stations were characterized by a higher proportion of Globothalamea, mainly Reophax sp. and Stainforthia sp. The foraminiferal community in EDG1 was dominated by a monothalamid *Psammosphaera* sp. (up to 90%), while *Hippocrepinella* sp. dominated in EDG2 and EDG3 (from 11% to 84%). Also, the percentage of *Reophax* sp. sequences increased towards the glacier-distal stations, reaching up to 34.68% at the station EDG3_Sieved.

4.4. Alpha and beta diversity patterns

4.4.1. Alpha diversity

The four alpha-diversity indices (Observed ASVs, Chao1, Simpson and Shannon) were measured separately for sieved and unsieved datasets (Fig. 7) and showed clear variation between different locations. On the one hand, the number of ASVs collected varied substantially depending on sample treatments and locations. In terms of sample treatment, sieved samples recovered higher Observed ASVs and Chao1 indices (Figs. 7a,b), but not Simpson or Shannon (Figs. 7c,d). On the other hand, the measured alpha-diversity indices tended to increase with increasing distance from the glacier (except station RIJ2). In general, the alpha-diversity indices of the fjords (IS, WIJ, RIJ) were higher than those of open sea areas (NAL, EDG), indicating higher foraminiferal diversity in fjords.

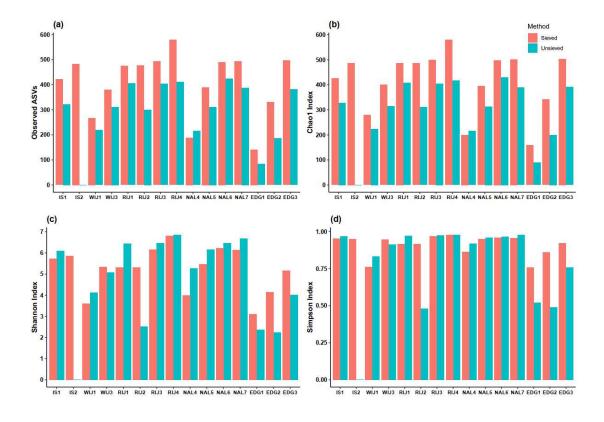


Figure 7. Bar plots of alpha diversity indices, including community richness (a: Observed ASVs, b: Chao1) and diversity (c: Shannon, d: Simpson) for the 15 sieved samples and 14 unsieved samples using retained ASV read abundances.

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4.4.2. Beta diversity

The non-metric multidimensional scaling plots (nMDS) and heatmap of all stations (Fig. 8) supported the findings of the ordinances, showing that foraminiferal communities detected in the sieved samples differed from those detected in the unsieved samples, but not significantly (Table S7). The nMDS and heatmap patterns also revealed spatial distribution of the foraminiferal communities among the different localities. In Fig. 8a, the nMDS analyses produced a similar pattern with sieved and unsieved datasets, although community segregation was observed in ordinations of EDG and NAL sites. The communities were distinctly separated between fjord stations and opened sea stations, as shown by low-stress values. Although the fjord samples formed tight clusters, the samples from each fjord were not overlapping with the samples from other fjord locations. On the contrary, the communities obtained from EDG and NAL sites formed clusters with much larger internal compositional differences and has an overlap between the two sites. Heatmap further clarified the community structuration with the stations and datasets (Fig. 8b), which were not visible on the nMDS (except for NAL5). The sampling sites were grouped in two main clusters: cluster 1 aggregating 3 stations (EDG1, EDG2, NAL4) and cluster 2 comprising the 12 remaining stations that were grouped into three subclusters. The stations of two fjords (IS and WIJ) had homogeneous communities and formed one separate subcluster. Two other subclusters are formed by i) NAL6, NAL5, EDG3, and ii) all RIJ stations and NAL7.

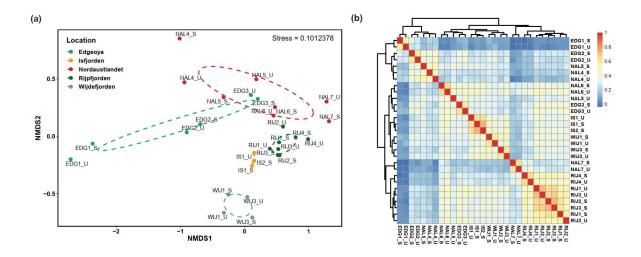


Figure 8. Community structuring of benthic foraminifera using nonlinear multidimensional scaling based on Bray-Curtis distance similarity coefficient (a) and heatmap based on Spearman's correlation coefficient for fractions and unsieved samples (b). Stress value is displayed on the plot.

4.4. sPLS Prediction Analysis

In the results of the sPLS regression, we detected several foraminiferal ASVs lineages for which relative sequence abundance correlated with environmental parameters (Fig. 9 and Table S8). The sPLS regression and subsequent hierarchical clustering suggested that the data was separated into three clusters (Fig. 9). These include lineages identified as potential indicators of water mass characteristics.

In cluster I.A, the ASVs exhibited a positive correlation with turbidity and negative correlation with factors such as depth, the salinity of bottom water and temperature of the surface water, with the ASVs predominantly affiliated as members of monothalamiids: *Hippocrepinella* sp. (ASV3, ASV10, ASV21), *Psammosphaera* sp. (ASV6), *Saccamminidae* sp. (ASV12), CladeY_spallogJAP (ASV22), STICKY_ICE (ASV39), *Pelosinella fusiformis* (ASV82), and globothalamids: *Buliminella* sp. (ASV14), *Nonionellina labradorica* (ASV20), *Cibicides* sp. (ASV91). The ASVs within cluster II revealed a strong and positive correlation

with temperature as well as a negative correlation with salinity in the surface water masses. This cluster included globothalamids: *Stainforthia* sp. (ASV1), *Virgulinella fragilis* (ASV79), *Reophax sp.* (ASV92), *Cibicidoides fletcheri* (ASV23), and monothalamiids: *Psammophaga* sp. (ASV25, ASV34, ASV42, ASV53, ASV64), *Micrometula* sp. (ASV4), ENFOR2_EnvHabIC19 (ASV45), CladeA (ASV85, ASV32). Some ASVs belonging to cluster II also had a strong positive correlation with the depth and bottom water salinity (Table S8).



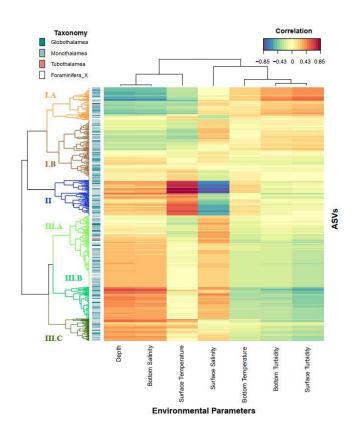


Figure 9. Clustered image map (CIM) of the first two sPLS dimensions, displaying pairwise correlations between foraminiferal ASVs of combined unsieved and sieved samples associated with environmental parameters. Correlations between ASVs and environmental parameters are depicted as a clustered heat map (detailed results in Table S8). Red and blue indicate positive and negative correlations, respectively.

Additionally, we observed positive correlations with depth and salinity in clusters III.A and III.B. Most of ASVs belonging to these clusters were classified as undetermined Monothalamea: Monothalamea_XX (39 ASVs), ENFOR XX (20 ASVs), CladeG (19 ASVs), CladeTIN (16 ASVs). In terms of ASV abundance, the dominant ASVs included environmental monothalamiids (ASV11, ASV44, ASV81, ASV66, ASV93, ASV88, ASV69, ASV78), CladeC_spsaccam (ASV59), *Gloiogullmia* sp. (ASV63), *Cibicides* sp. (ASV38), *Nonionella auris* (ASV31), rotalid (ASV96), *Reophax* sp. (ASV41, ASV13, ASV18), *Stainforthia* sp. (ASV77). Cluster III.C had positive associations with the depth, the bottom water salinity, and the surface temperature. ENFOR2_XXX (ASV408, ASV136, ASV125) exhibited their highest abundances of this cluster.

5. Discussion

5.1. Is pre-sieving useful for foraminiferal metabarcoding?

The methodological aim of this study was to compare the results of metabarcoding analyses based on sieved and unsieved sediment samples. Sieving is a common procedure in the conventional microscopic study of foraminiferal assemblage analyzing mainly hard-shelled, multi-chambered taxa preserved in fixed and dried sediment samples ⁷⁸. In contrast, metabarcoding studies of unsieved sediment samples usually provide a foraminiferal assemblage dominated by poorly known, soft-shelled or naked monothalamous taxa ^{37,41}. Because of this, it is difficult to compare the results of traditional morphology-based studies with those of metabarcoding analysis, which provide very different types of data ³⁵.

As shown by our study, the taxonomic composition of different size fractions is not the same. For example, the order Rotaliida was the most abundant in 500-100 μ m and 100-63 μ m size fractions. Also, another hard-shelled order Textulariida, which is microscopically studied in the 500-100 μ m fraction, in metabarcoding data is present mainly in fractions 500-100 μ m

and 100-63 μ m (Fig. 5c) This is congruent with the rotaliids and textulariids dominating microscopic assemblage found in >63 μ m sieved fraction. On the other hand, the smallest fraction (<63 μ m) was dominated by monothalamiids and undetermined Foraminifera (Fig. 5C), which may suggest the presence of some genetically unknown, tiny monothalamous species.

We also observed some differences between sieved and unsieved samples regarding the alpha diversity. The total number of recovered ASVs was clearly higher in sieved than in unsieved samples (approximately 30% ASVs). However, this could be explained by the difference in the number of DNA extractions, PCR amplifications, and sequencing depth. In the case of sieved samples, the datasets included four DNA extractions, one for each size fraction, while only one DNA extraction was performed for non-sieved sediment samples. In total, the number of sequences obtained for sieved fractions was several times higher, allowing to detection of more diversity in sieved compared to unsieved samples. However, no significant difference between the sieved/unsieved samples was observed in alpha-diversity measures such as Shannon's and Simpson's that take abundance and evenness of the sample into consideration as shown in Fig. 7. Although sieving might have been predicted to lead to a reduction in alpha diversity due to the loss of microfauna and extracellular debris, this has not been observed in previous studies ^{79,80}. In addition, NMDS of the beta-diversity matrices and correlation test showed sieved and unsieved samples clustered together (Fig. 8, Table S7), indicating that there is no significant difference in community composition inferred by the two methods.

To conclude, the decision of whether the sediment samples should be sieved or not shall be based on the type of questions one wants to answer with metabarcoding data as well as the composition and characteristics of initial samples. Sieving procedure combined with metabarcoding analysis might be useful if targeted at particular groups of foraminifera (e.g., Rotaliida, Textulariida), for example, to compare with microscopic studies or to identify some

tiny species present in fine size fractions. In general, the unsieved samples provide a more complete overview of the taxonomic composition of the foraminiferal community. However, as shown by our study, both metabarcoding datasets reveal similar trends in foraminiferal diversity. Size-sieving might have some advantages, however, it also has some drawbacks, as (i) it is time-consuming, (ii) requires higher volume samples, and (iii) there is a possibility of cross-contamination between samples. Therefore, either extracting DNA directly from sediment or after sieving should be carefully considered when evaluating foraminiferal communities across metabarcoding studies.

5.2. Distribution patterns of foraminifera in Svalbard

The most striking result of this study is the variations of foraminiferal assemblage between different sampling localities. Although, there are some similar trends documented by diversity indices at different locations, such as the proportion of monothalamiids in near-glacier settings (Fig. 6), or the increase of alpha diversity from glacier proximal/inner to glacier-distant/outer stations (Fig. 7), which are in agreement with the previous morphology-based studies ⁸¹⁻⁸³, the composition of foraminiferal communities is generally specific to each location. Each fjord forms a separate cluster and only some stations at open-water areas overlap with each other (Fig. 8a).

The high-Arctic settings are usually considered as a cold system influenced at different levels by ArW during summer to late autumn ⁸⁴, and covered by sea ice in winter ^{9,59,85}. However, the increased influence of AW and winter sea-ice loss is observed in recent years ^{10,86,87}. We speculate that hydrographic conditions would lead to isolating populations from different settings and creating unique structures of the foraminiferal community.

It is well known that the unique habitats of fjords can support a high diversity and distinct biological communities 81-83,88-90. Fjords create a variety of habitats suitable for specific

species, where many species can converge and reach high population densities. Western Spitsbergen fjords are among the most AW-impacted areas. Both Isfjorden and Wijdefjorden are directly connected to the slope and shelf areas, which enables AW penetration into the fjords. Moreover, Isfjorden stations are located in the central basin of the fjord, which resulted in limited glacial influence. This led to the formation of foraminifera communities characterized by a relatively high proportion of globothalamids and the presence of monothalamiids community dominated by the genera *Psammophaga* and *Micrometula* (Fig. 6). Similar distribution patterns were previously observed in west Spitsbergen fjords ^{81,82}. Only station WIJ1 displayed a unique structure with a clear dominance of Rotaliida (Figs. 4 and 6). Station WIJ1 is located in the inner fjord, close to the glacier termini and is influenced by turbid meltwater runoffs. The dominance of Rotaliida contradicts previous studies, indicating that glacier proximal settings are dominated by monothalamous foraminifera ^{81,91}. However, this distribution pattern could be also explained by the natural patchiness of foraminiferal distribution.

In the northern site (RIJ), the dominant component of foraminiferal assemblages were undetermined monothalamiids (Figs 4 and 6). The dominance of these undetermined, small, soft-walled species was previously observed in areas characterized by a high level of environmental disturbance ^{42,81}. Northern Svalbard in general and Rijpfjorden in particular, are considered to be a typical Arctic setting, where sea ice forms in autumn and lasts until summer. Also, the drifting ice pack is often transported to the fjord during the summer ⁵⁹. Forming of sea ice is associated with brine rejection. Cold and dense brines sink to the bottom, which leads to the formation of cold and saline WCW. This process may create a more severe environment where monothalamiids thrive.

Foraminiferal communities of open water areas (eastern Svalbard) have generally lower diversity and form different groups compared to those from western Svalbard fjords. Stations

from the regions of Nordaustlandet and Edgeøya are located in front of large tidewater glaciers, releasing large amounts of turbid meltwater (Fig. S1). However, only Nordaustlandet was influenced by AW and TAW, while the Edgeøya oceanographic conditions were shaped mainly by local water masses. These led to the creation of distinctly different foraminiferal communities. NAL stations were characterized by a wide range of undetermined monothalamiids, while the EDG stations were dominated by a few monothalamous species representing genera *Hipocrepinella* and *Psammosphaera* (Fig. 6). In particular, station EDG1 exhibited a unique foraminiferal community, composed almost exclusively of *Psammosphaera* sp. Psammosphaerids were previously recorded in Spitsbergen fjords ^{82,90} and also in deep-sea areas affected by bottom currents ⁹². Environmental conditions at station EDG1 are probably the most dynamic and severe among the studied stations, as the large sub-bottom meltwater outflow was recorded (Fig. S1).

5.3. Influence of atlantification on foraminifera community

The response of benthic foraminifera to alterations in temperature and salinity in the water column are common and include expansions or retractions of distribution ranges or changes in assemblage compositions ⁹³⁻⁹⁵. We hypothesize that the composition of foraminiferal communities in our data resulted from water mass conditions. This hypothesis is strengthened by the clear clustering of the community in groups corresponding to different oceanographic regimes, in which stations from regions impacted by AW and/or sea-ice clustered separately.

As shown by the nMDS plot and heatmap (Fig. 8), the separation between two main clusters has a strong relationship with characteristics of water masses. Cluster 1 comprises exclusively the stations of the eastern part of the archipelago (EDG1, EGD2, NAL4, NAL5), characterized by colder, and less salty water, associated with turbid glacial meltwater ⁹⁶. On

the contrary, Cluster 2 includes mainly stations from the western and northern part of Svalbard where the impact of warmer and more saline Atlantic Water (AW) was much pronounced, confirmed by our CTD profile (Fig. 2). Also, sub-clusters that formed within cluster 2 reflected different impacts of AW. The first sub-cluster comprises stations located in the glacial-distant regions of Nordaustlandet and Edgeøya, influenced mainly by TAW. The second sub-cluster includes the most AW-impacted stations located on the western coast of the archipelago, while the third sub-cluster is composed of stations located in north-eastern Svalbard, influenced both by the inflow of AW and sea-ice.

The increased AW inflow, higher light availability, and the decline of sea-ice around Svalbard affect the primary productivity, changing both the timing of phytoplankton bloom and phytoplankton community structure. This may have significant effects on food web dynamics, affecting higher trophic levels, including benthic communities ⁹⁷. On the other hand, recent model projections indicated low mean habitat loss of benthic macrofauna under recent climate changes, which questions the vulnerability of Arctic benthos to atlantification ⁹⁸. This stands in clear opposition to the morphological observations that testate foraminifera communities from Svalbard fjords revealed significant changes, both in terms of abundance and species composition, related to atlantification ⁹⁹. Our study confirms the impact of AW on foraminiferal communities, suggesting that AW is one of the primary factors shaping the benthic foraminifera assemblages and thus foraminifera may be potential indicators of atlantification.

Through sPLS analysis of combining datasets, we identified foraminiferal taxa that could become potential bioindicators of "atlantification". This group of species includes some monothalamiids belonging to genera *Psammophaga* and *Micrometula* as well as some undetermined monothalamous species belonging to environmental lineage ENFOR2 and Clade A. The genera *Psammophaga* and *Micrometula* are widespread in many coastal areas including

polar regions ^{100,101} and are considered as bioindicator candidates in several studies ^{37,102}. However, the limited knowledge about the ecology of those taxa, as well as lack of reference database of their distribution in the North Atlantic region precludes making any general conclusions. Among potential bioindicators, there are also some globothalamids, such as *Stainforthia* sp., *Virgulinella fragilis*, *Reophax* sp., and *Cibicidoides fletcheri*. *Cibicidoides* species is common in the North Atlantic ¹⁰³, but to the best of our knowledge, it was not recorded in Svalbard before. One of the major signs of atlantification is the northward shift of boreal species, the trend observed in the case of zooplankton, fish, and benthic organisms ⁹⁷. A recent morphological study of Svalbard foraminifera revealed the presence of boreal species *Melonis affinis* in the northern part of the archipelago ⁹⁹. *Reophax* sp. and, *Stainforthia* sp. are commonly found in Svalbard ⁸³. Also, species belonging to the genus *Reophax* are considered as indicators of AW ¹⁰⁴. The agreement of our results with previous morphology-based studies further proves the potential of foraminifera in biomonitoring studies.

Apart from indicators of atlantification, we identified monothalamiids that show a strong correlation with turbidity, but not with depth or salinity. These species included *Hippocrepinella* sp., *Psammosphaera* sp., *Saccamminidae* sp., CladeY_spallogJAP, and several unassigned STICKY_ICE ASVs. This correlation is particularly strong in stations closer to the coast, which is probably caused by enhanced turbidity due to sediment-laden meltwater plumes. *Hippocrepinella* sp., *Psammosphaera* sp., *Saccamminidae* sp., as well as monothalamiids belonging to Clade Y were previously recorded in Svalbard ^{82,90,91}. Clade Y comprises mainly undetermined monothalamiids, known from environmental sequencing. They were abundantly sequenced in the settings characterized by a high level of environmental disturbance, suggesting that they are highly resistant to environmental disturbance ³⁷. Moreover, morphological studies reported *Hippocrepinella* sp., *Psammosphaera* sp., *Saccamminidae* sp., in the shallow-water parts of the fjords, located close to meltwater

outflows ^{81,82,90}. These findings underline how important can be to include soft-shelled monothalamous foraminifera in metabarcoding studies to enhance limited knowledge about their ecology for potential use in biomonitoring.

6. Conclusions

This study is the first to use high-throughput sequencing to comprehensively analyze the foraminiferal communities within marine sediments from Svalbard, which helps to provide more knowledge of foraminiferal diversity and distribution patterns in the Arctic's fjords. The DNA sequencing results from sieved and unsieved sediment revealed a high diversity of the Svalbard foraminifera compared to traditional morphology-based studies and variation in the taxonomic composition of foraminiferal communities from five sampling areas. Foraminiferal diversity and species richness increased from glacier proximal/inner to glacier-distant/outer stations and were higher in the fjords than in the open water. Moreover, the structure of foraminiferal community is clearly influenced by different water masses, with a particular impact of Atlantic Water in the Svalbard region. Numerous potential molecular foraminiferal bioindicators for atlantification were identified. This should be confirmed by analysing many more samples from reference areas in North Atlantic. With the increasing numbers of metabarcoding studies, the impact of atlantification on Arctic benthic communities identified in this study could be better assessed and expanded to those organisms that are not covered by the conventional morphological approach.

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Authors contributions

N-LN, Joanna P, Jan P and MZ conceived the idea for the study and developed the research methodology; Joanna P and MZ collected the samples and performed laboratory analysis; N-LN produced the dataset; N-LN and Joanna. P performed the bioinformatic analyses and wrote the manuscript with contributions and comments from all co-authors. All authors reviewed and approved the final manuscript.

Declaration of Competing Interest

The authors declare no competing interests.

References

Beszczynska-Möller, A., Fahrbach, E., Schauer, U. & Hansen, E. Variability in Atlantic water temperature and transport at the entrance to the Arctic Ocean, 1997– 2010. ICES Journal of Marine Science **69**, 852-863, doi:10.1093/icesjms/fss056 (2012).Polyakov, I. V. et al. Greater role for Atlantic inflows on sea-ice loss in the Eurasian Basin of the Arctic Ocean. Science 356, 285-291, doi:10.1126/science.aai8204 (2017).

- Onarheim, I. H., Smedsrud, L. H., Ingvaldsen, R. B. & Nilsen, F. Loss of sea ice
- during winter north of Svalbard. Tellus A: Dynamic Meteorology and Oceanography
- **66**, 23933, doi:10.3402/tellusa.v66.23933 (2014).
- Berge, J., Johnsen, G., Nilsen, F., Gulliksen, B. & Slagstad, D. Ocean temperature
- oscillations enable reappearance of blue mussels Mytilus edulis in Svalbard after a
- 650 1000 year absence. *Marine Ecology Progress Series* **303**, 167-175 (2005).
- 651 5 Slagstad, D., Ellingsen, I. H. & Wassmann, P. Evaluating primary and secondary
- production in an Arctic Ocean void of summer sea ice: An experimental simulation
- 653 approach. *Prog Oceanogr* **90**, 117-131,
- doi:https://doi.org/10.1016/j.pocean.2011.02.009 (2011).
- 655 6 Zajączkowski, M., Nygård, H., Hegseth, E. N. & Berge, J. Vertical flux of particulate
- matter in an Arctic fjord: the case of lack of the sea-ice cover in Adventfjorden 2006–
- 657 2007. *Polar Biology* **33**, 223-239, doi:10.1007/s00300-009-0699-x (2010).
- Kortsch, S., Primicerio, R., Fossheim, M., Dolgov, A. V. & Aschan, M. Climate
- change alters the structure of arctic marine food webs due to poleward shifts of boreal
- generalists. *Proceedings of the Royal Society B: Biological Sciences* **282**, 20151546,
- doi:doi:10.1098/rspb.2015.1546 (2015).
- 662 8 Serreze, M. C. & Barry, R. G. Processes and impacts of Arctic amplification: A
- research synthesis. Global and Planetary Change 77, 85-96,
- doi:https://doi.org/10.1016/j.gloplacha.2011.03.004 (2011).
- Dai, A., Luo, D., Song, M. & Liu, J. Arctic amplification is caused by sea-ice loss
- under increasing CO2. *Nat Commun* **10**, 121, doi:10.1038/s41467-018-07954-9
- 667 (2019).
- Nilsen, F., Cottier, F., Skogseth, R. & Mattsson, S. Fjord–shelf exchanges controlled
- by ice and brine production: The interannual variation of Atlantic Water in Isfjorden,

670 Svalbard. Continental Shelf Research 28, 1838-1853, 671 doi:https://doi.org/10.1016/j.csr.2008.04.015 (2008). 672 11 Vihtakari, M. et al. Black-legged kittiwakes as messengers of Atlantification in the 673 Arctic. Sci Rep 8, 1178, doi:10.1038/s41598-017-19118-8 (2018). 674 12 Descamps, S. et al. Climate change impacts on wildlife in a High Arctic archipelago -675 Svalbard, Norway. Glob Chang Biol 23, 490-502, doi:10.1111/gcb.13381 (2017). 676 13 Frainer, A. et al. Climate-driven changes in functional biogeography of Arctic marine 677 fish communities. Proc Natl Acad Sci U S A 114, 12202-12207, 678 doi:10.1073/pnas.1706080114 (2017). 679 14 Fossheim, M. et al. Recent warming leads to a rapid borealization of fish 680 communities in the Arctic. *Nature Climate Change* **5**, 673-+ (2015). 681 Weydmann-Zwolicka, A. et al. Zooplankton and sediment fluxes in two contrasting 15 682 fjords reveal Atlantification of the Arctic. Sci Total Environ 773, 145599, 683 doi:10.1016/j.scitotenv.2021.145599 (2021). 684 16 Hop, H. et al. Pelagic Ecosystem Characteristics Across the Atlantic Water Boundary 685 Current From Rijpfjorden, Svalbard, to the Arctic Ocean During Summer (2010– 686 2014). Frontiers in Marine Science 6, doi:10.3389/fmars.2019.00181 (2019). 687 17 Grabowski, M. et al. Contrasting molecular diversity and demography patterns in two 688 intertidal amphipod crustaceans reflect Atlantification of High Arctic. Marine Biology 689 **166**, 155, doi:10.1007/s00227-019-3603-4 (2019). 690 18 Barton, A. D., Irwin, A. J., Finkel, Z. V. & Stock, C. A. Anthropogenic climate 691 change drives shift and shuffle in North Atlantic phytoplankton communities. Proc 692 Natl Acad Sci U S A 113, 2964-2969, doi:10.1073/pnas.1519080113 (2016).

Neukermans, G., Oziel, L. & Babin, M. Increased intrusion of warming Atlantic water leads to rapid expansion of temperate phytoplankton in the Arctic. Glob Chang Biol , 2545-2553, doi:10.1111/gcb.14075 (2018). Meilland, J. et al. Population dynamics of modern planktonic foraminifera in the western Barents Sea. *Biogeosciences* 17, 1437-1450, doi:10.5194/bg-17-1437-2020 (2020).Ofstad, S. et al. Shell density of planktonic foraminifera and pteropod species Limacina helicina in the Barents Sea: Relation to ontogeny and water chemistry. PLoS One 16, e0249178, doi:10.1371/journal.pone.0249178 (2021). Schoenle, A. et al. High and specific diversity of protists in the deep-sea basins dominated by diplonemids, kinetoplastids, ciliates and foraminiferans. Commun Biol , 501, doi:10.1038/s42003-021-02012-5 (2021). Gooday, A. J. & Jorissen, F. J. Benthic foraminiferal biogeography: controls on global distribution patterns in deep-water settings. Ann Rev Mar Sci 4, 237-262, doi:10.1146/annurev-marine-120709-142737 (2012). Murray, J. W. Ecology and Applications of Benthic Foraminifera. (Cambridge University Press, 2006). Kawahata, H. et al. Perspective on the response of marine calcifiers to global warming and ocean acidification—Behavior of corals and foraminifera in a high CO2 world "hot house". *Progress in Earth and Planetary Science* **6**, **5**, doi:10.1186/s40645-018-0239-9 (2019). Wittmann, A. C. & Pörtner, H.-O. Sensitivities of extant animal taxa to ocean

acidification. Nature Climate Change 3, 995-1001, doi:10.1038/nclimate1982 (2013).

716 27 Prazeres, M., Roberts, T. E. & Pandolfi, J. M. Variation in sensitivity of large benthic 717 Foraminifera to the combined effects of ocean warming and local impacts. *Scientific* 718 Reports 7, 45227, doi:10.1038/srep45227 (2017). 719 28 Bohmann, K. et al. Environmental DNA for wildlife biology and biodiversity 720 monitoring. Trends Ecol Evol 29, 358-367, doi:10.1016/j.tree.2014.04.003 (2014). 721 29 Holman, L. E. et al. Animals, protists and bacteria share marine biogeographic 722 patterns. Nature Ecology & Evolution 5, 738-746, doi:10.1038/s41559-021-01439-7 723 (2021).724 30 Pawlowski, J., Bonin, A., Boyer, F., Cordier, T. & Taberlet, P. Environmental DNA 725 for biomonitoring. *Molecular Ecology* **30**, 2931-2936, 726 doi:https://doi.org/10.1111/mec.16023 (2021). 727 Lejzerowicz, F., Esling, P. & Pawlowski, J. Patchiness of deep-sea benthic 31 728 Foraminifera across the Southern Ocean: Insights from high-throughput DNA 729 sequencing. Deep Sea Research Part II: Topical Studies in Oceanography 108, 17-26, 730 doi:https://doi.org/10.1016/j.dsr2.2014.07.018 (2014). 731 Pawlowski, J. & Lecrog, B. Short rDNA barcodes for species identification in 32 732 foraminifera. J Eukaryot Microbiol 57, 197-205, doi:10.1111/j.1550-733 7408.2009.00468.x (2010). 734 33 Armbrecht, L. H. et al. Ancient DNA from marine sediments: Precautions and 735 considerations for seafloor coring, sample handling and data generation. Earth-736 Science Reviews 196, 102887, doi:https://doi.org/10.1016/j.earscirev.2019.102887 737 (2019).738 34 Vargas, C. d. et al. Eukaryotic plankton diversity in the sunlit ocean. Science 348,

1261605, doi:doi:10.1126/science.1261605 (2015).

740 35 Frontalini, F. et al. Benthic foraminiferal metabarcoding and morphology-based 741 assessment around three offshore gas platforms: Congruence and complementarity. 742 Environ Int 144, 106049, doi:10.1016/j.envint.2020.106049 (2020). 743 36 Pawlowski, J. et al. Benthic monitoring of salmon farms in Norway using 744 foraminiferal metabarcoding. Aquacult Env Interac 8, 371-386 (2016). 745 37 Pawlowski, J., Esling, P., Lejzerowicz, F., Cedhagen, T. & Wilding, T. A. 746 Environmental monitoring through protist next-generation sequencing metabarcoding: 747 assessing the impact of fish farming on benthic foraminifera communities. *Molecular* 748 Ecology Resources 14, 1129-1140 (2014). 749 38 Chronopoulou, P. M., Salonen, I., Bird, C., Reichart, G. J. & Koho, K. A. 750 Metabarcoding Insights Into the Trophic Behavior and Identity of Intertidal Benthic 751 Foraminifera. Frontiers in microbiology 10, 1169, doi:10.3389/fmicb.2019.01169 752 (2019).753 39 Cordier, T., Barrenechea, I., Lejzerowicz, F., Reo, E. & Pawlowski, J. Benthic 754 foraminiferal DNA metabarcodes significantly vary along a gradient from abyssal to hadal depths and between each side of the Kuril-Kamchatka trench. Prog Oceanogr 755 756 **178** (2019). 757 40 Lejzerowicz, F. et al. Eukaryotic Biodiversity and Spatial Patterns in the Clarion-758 Clipperton Zone and Other Abyssal Regions: Insights From Sediment DNA and RNA 759 Metabarcoding. Frontiers in Marine Science 8, doi:10.3389/fmars.2021.671033 760 (2021).761 41 Lecroq, B. et al. Ultra-deep sequencing of foraminiferal microbarcodes unveils hidden richness of early monothalamous lineages in deep-sea sediments. Proc Natl 762

Acad Sci U S A 108, 13177-13182, doi:10.1073/pnas.1018426108 (2011).

764 Pawlowska, J. et al. Ancient DNA sheds new light on the Svalbard foraminiferal 42 765 fossil record of the last millennium. Geobiology 12, 277-288, doi:10.1111/gbi.12087 (2014).766 767 Pawlowska, J., Wollenburg, J. E., Zajaczkowski, M. & Pawlowski, J. Planktonic 43 768 foraminifera genomic variations reflect paleoceanographic changes in the Arctic: 769 evidence from sedimentary ancient DNA. Sci Rep 10, 15102, doi:10.1038/s41598-770 020-72146-9 (2020). 771 Barrenechea Angeles, I. et al. Planktonic foraminifera eDNA signature deposited on 44 772 the seafloor remains preserved after burial in marine sediments. Sci Rep 10, 20351, 773 doi:10.1038/s41598-020-77179-8 (2020). 774 45 Elbrecht, V., Peinert, B. & Leese, F. Sorting things out: Assessing effects of unequal 775 specimen biomass on DNA metabarcoding. Ecol Evol 7, 6918-6926, 776 doi:10.1002/ece3.3192 (2017). 777 46 Leray, M. & Knowlton, N. DNA barcoding and metabarcoding of standardized 778 samples reveal patterns of marine benthic diversity. P Natl Acad Sci USA 112, 2076-779 2081 (2015). 780 47 Brandt, M. I. et al. Evaluating sediment and water sampling methods for the 781 estimation of deep-sea biodiversity using environmental DNA. Sci Rep 11, 7856, 782 doi:10.1038/s41598-021-86396-8 (2021). 783 48 Sinniger, F. et al. Worldwide Analysis of Sedimentary DNA Reveals Major Gaps in 784 Taxonomic Knowledge of Deep-Sea Benthos. Frontiers in Marine Science 3, 785 doi:10.3389/fmars.2016.00092 (2016). 786 49 Cottier, F. et al. Water mass modification in an Arctic fjord through cross-shelf

exchange: The seasonal hydrography of Kongsfjorden, Svalbard. Journal of

788 Geophysical Research: Oceans 110, doi:https://doi.org/10.1029/2004JC002757 789 (2005).790 50 Blindheim, J. & Østerhus, S. in The Nordic Seas: An Integrated Perspective 11-37 791 (2005).792 51 Loeng, H. Features of the physical oceanographic conditions of the Barents Sea. 793 Polar Research 10, 5-18, doi:https://doi.org/10.1111/j.1751-8369.1991.tb00630.x 794 (1991).795 52 Spielhagen, R. F. et al. Enhanced Modern Heat Transfer to the Arctic by Warm 796 Atlantic Water. Science 331, 450-453, doi:doi:10.1126/science.1197397 (2011). 797 53 Saloranta, T. M. & Haugan, P. M. Northward cooling and freshening of the warm 798 core of the West Spitsbergen Current. Polar Research 23, 79-88, 799 doi:10.3402/polar.v23i1.6268 (2004). 800 54 Aagaard, K., Foldvik, A. & Hillman, S. R. The West Spitsbergen Current: Disposition 801 and water mass transformation. Journal of Geophysical Research: Oceans 92, 3778-802 3784, doi:https://doi.org/10.1029/JC092iC04p03778 (1987). 803 Bourke, R. H., Weigel, A. M. & Paquette, R. G. The westward turning branch of the 55 804 West Spitsbergen Current. Journal of Geophysical Research: Oceans 93, 14065-805 14077, doi:https://doi.org/10.1029/JC093iC11p14065 (1988). 806 56 Sternal, B. et al. Postglacial variability in near-bottom current speed on the 807 continental shelf off south-west Spitsbergen. Journal of Quaternary Science 29, 767– 808 777, doi:10.1002/jqs.2748 (2014). 809 57 Nilsen, F., Skogseth, R., Vaardal-Lunde, J. & Inall, M. A Simple Shelf Circulation 810 Model: Intrusion of Atlantic Water on the West Spitsbergen Shelf. Journal of 811 Physical Oceanography 46, 1209-1230, doi:10.1175/jpo-d-15-0058.1 (2016).

812 58 Kowalewski, W., Rudowski, S. & Zalewski, S. M. Seismoacoustic studies within 813 Wijdefjorden, Spitsbergen. Polish Polar Research vol. 11, 287-300-287-300 (1990). 814 59 Ambrose Jr., W. G., Carroll, M. L., Greenacre, M., Thorrold, S. R. & McMahon, K. 815 W. Variation in Serripes groenlandicus (Bivalvia) growth in a Norwegian high-Arctic 816 fjord: evidence for local- and large-scale climatic forcing. Global Change Biology 12, 817 1595-1607, doi:https://doi.org/10.1111/j.1365-2486.2006.01181.x (2006). 818 60 Dowdeswell, J. A. et al. Digital Mapping of the Nordaustlandet Ice Caps from 819 Airborne Geophysical Investigations. *Annals of Glaciology* **8**, 51-58, 820 doi:10.3189/S0260305500001130 (1986). 821 61 Dowdeswell, J. A. & Bamber, J. L. On the glaciology of Edgeøya and Barentsøya, Svalbard. Polar Research 14, 105-122, doi:10.3402/polar.v14i2.6658 (1995). 822 823 62 Knies, J., Brookes, S. & Schubert, C. J. Re-assessing the nitrogen signal in continental 824 margin sediments: New insights from the high northern latitudes. Earth Planet Sc Lett 825 **253**, 471-484 (2007). 826 63 Dufresne, Y., Lejzerowicz, F., Perret-Gentil, L. A., Pawlowski, J. & Cordier, T. 827 SLIM: a flexible web application for the reproducible processing of environmental 828 DNA metabarcoding data. BMC Bioinformatics 20, 88, doi:10.1186/s12859-019-829 2663-2 (2019). 830 Callahan, B. J. et al. DADA2: High-resolution sample inference from Illumina 64 amplicon data. Nat Methods 13, 581-583, doi:10.1038/nmeth.3869 (2016). 831 832 65 Pawłowska, J., Pawlowski, J. & Zajączkowski, M. Dataset of foraminiferal 833 sedimentary DNA (sedDNA) sequences from Svalbard. Data in Brief 30, 105553, 834 doi:https://doi.org/10.1016/j.dib.2020.105553 (2020).

- 835 Frosley, T. G. et al. Algorithm for post-clustering curation of DNA amplicon data 66 836 yields reliable biodiversity estimates. Nat Commun 8, 1188, doi:10.1038/s41467-017-837 01312-x (2017). 838 Holzmann, M. & Pawlowski, J. An updated classification of rotaliid foraminifera 67 839 based on ribosomal DNA phylogeny. Marine Micropaleontology 132, 18-34, 840 doi:https://doi.org/10.1016/j.marmicro.2017.04.002 (2017). 841 68 Pawlowski, J., Holzmann, M. & Tyszka, J. New supraordinal classification of 842 Foraminifera: Molecules meet morphology. Marine Micropaleontology 100, 1-10, 843 doi:https://doi.org/10.1016/j.marmicro.2013.04.002 (2013). 844 69 Guillou, L. et al. The Protist Ribosomal Reference database (PR2): a catalog of 845 unicellular eukaryote small sub-unit rRNA sequences with curated taxonomy. Nucleic 846 Acids Res 41, D597-604, doi:10.1093/nar/gks1160 (2013). 847 70 Altschul, S. F., Gish, W., Miller, W., Myers, E. W. & Lipman, D. J. Basic local 848 alignment search tool. Journal of Molecular Biology 215, 403-410, 849 doi:https://doi.org/10.1016/S0022-2836(05)80360-2 (1990). 850 71 R Core Team. R: A language and environment for statistical computing. R 851 Foundation for Statistical Computing, Vienna, Austria, Available online at 852 https://www.R-project.org/. 2018). 853 72 Dusa, A. venn: Draw Venn Diagrams, https://cran.r- 854 project.org/web/packages/venn/venn.pdf> (2018). 855 73 Hsieh, T. C., Ma, K. H. & Chao, A. iNEXT: an R package for rarefaction and 856 extrapolation of species diversity (Hill numbers). Methods in Ecology and Evolution 857 7, 1451-1456, doi:https://doi.org/10.1111/2041-210X.12613 (2016). 858 74 Oksanen, J. A. I.
 - 39

Kolde, R. Pretty Heatmaps. (2019).

859

860 76 Rohart, F., Gautier, B., Singh, A. & Le Cao, K. A. mixOmics: An R package for 'omics feature selection and multiple data integration. PLoS Comput Biol 13, 861 862 e1005752, doi:10.1371/journal.pcbi.1005752 (2017). 863 77 Le Cao, K. A., Rossouw, D., Robert-Granie, C. & Besse, P. A sparse PLS for variable 864 selection when integrating omics data. Stat Appl Genet Mol Biol 7, Article 35, 865 doi:10.2202/1544-6115.1390 (2008). 866 78 Schönfeld, J. et al. The FOBIMO (FOraminiferal BIo-MOnitoring) initiative -867 Towards a standardised protocol for soft-bottom benthic foraminiferal monitoring 868 studies. Marine Micropaleontology 94-95, 1-13, 869 doi:https://doi.org/10.1016/j.marmicro.2012.06.001 (2012). 870 79 Brannock, P. M. & Halanych, K. M. Meiofaunal community analysis by high-871 throughput sequencing: Comparison of extraction, quality filtering, and clustering 872 methods. Marine Genomics 23, 67-75, doi:https://doi.org/10.1016/j.margen.2015.05.007 (2015). 873 874 80 He, X., Sutherland, T. F. & Abbott, C. L. Improved efficiency in eDNA metabarcoding of benthic metazoans by sieving sediments prior to DNA extraction. 875 876 Environmental DNA n/a, doi:https://doi.org/10.1002/edn3.172 (2020). 877 81 Sabbatini, A., Morigi, C., Negri, A. & Gooday, A. Distribution and biodiversity of 878 stained monothalamous foraminifera from Tempelfjord, Svalbard. Journal of Foraminiferal Research 37, 93-106, doi:10.2113/gsjfr.37.2.93 (2007). 879 880 82 Majewski, W., Pawłowski, J. & Zajączkowski, M. Monothalamous foraminifera from 881 West Spitsbergen fjords, Svalbard: a brief overview. Polish Polar Research vol. 26,

882

269-285-269-285 (2005).

883 83 Hald, M. & Korsun, S. Distribution of modern benthic foraminifera from fjords of 884 Svalbard, European Arctic. Journal of Foraminiferal Research - J FORAMIN RES 27, 885 101-122, doi:10.2113/gsjfr.27.2.101 (1997). 886 84 Wallace, M. I. et al. Comparison of zooplankton vertical migration in an ice-free and 887 a seasonally ice-covered Arctic fjord: An insight into the influence of sea ice cover on 888 zooplankton behavior. Limnology and Oceanography 55, 831-845, 889 doi:https://doi.org/10.4319/lo.2010.55.2.0831 (2010). 890 85 Leu, E., Søreide, J. E., Hessen, D. O., Falk-Petersen, S. & Berge, J. Consequences of 891 changing sea-ice cover for primary and secondary producers in the European Arctic 892 shelf seas: Timing, quantity, and quality. *Prog Oceanogr* **90**, 18-32, 893 doi:https://doi.org/10.1016/j.pocean.2011.02.004 (2011). 894 86 Dahlke, S. et al. The observed recent surface air temperature development across 895 Svalbard and concurring footprints in local sea ice cover. *International Journal of* 896 Climatology 40, 5246-5265, doi:https://doi.org/10.1002/joc.6517 (2020). 897 87 Pavlova, O., Gerland, S. & Hop, H. in The Ecosystem of Kongsfjorden, Svalbard 898 (eds Haakon Hop & Christian Wiencke) 105-136 (Springer International Publishing, 899 2019). 900 88 Walseng, B. et al. Freshwater diversity in Svalbard: providing baseline data for 901 ecosystems in change. *Polar Biology* **41**, 1995-2005, doi:10.1007/s00300-018-2340-3 902 (2018).903 Włodarska-Kowalczuk, M., Pawłowska, J. & Zajączkowski, M. Do foraminifera 89 904 mirror diversity and distribution patterns of macrobenthic fauna in an Arctic glacial 905 fjord? Marine Micropaleontology 103, 30-39, 906 doi:https://doi.org/10.1016/j.marmicro.2013.07.002 (2013).

907 90 Gooday, A. J. et al. Monothalamous foraminiferans and gromiids (Protista) from 908 western Svalbard: A preliminary survey Published in collaboration with the 909 University of Bergen and the Institute of Marine Research, Norway, and the Marine 910 Biological Laboratory, University of Copenhagen, Denmark. Marine Biology 911 Research 1, 290-312, doi:10.1080/17451000510019150 (2005). 912 91 Pawłowska, J. et al. Palaeoceanographic changes in Hornsund Fjord (Spitsbergen, 913 Svalbard) over the last millennium: new insights from ancient DNA. Clim. Past 12, 914 1459-1472, doi:10.5194/cp-12-1459-2016 (2016). 915 92 Kaminski, M. et al. Modern agglutinated Foraminifera from the Hovgård Ridge, Fram 916 Strait, west of Spitsbergen: evidence for a deep bottom current. Annales Societatis 917 *Geologorum Poloniae* **85**, 309-320, doi:10.14241/asgp.2015.006 (2015). 918 93 Langer, M. R., Weinmann, A. E., Lotters, S., Bernhard, J. M. & Rodder, D. Climate-919 Driven Range Extension of Amphistegina (Protista, Foraminiferida): Models of 920 Current and Predicted Future Ranges. *Plos One* **8** (2013). 921 94 Dong, S. S., Lei, Y. L., Li, T. G. & Jian, Z. M. Responses of benthic foraminifera to 922 changes of temperature and salinity: Results from a laboratory culture experiment. Sci 923 China Earth Sci 62, 459-472 (2019). 924 95 Weinmann, A. E. & Goldstein, S. T. Changing structure of benthic foraminiferal 925 communities: implications from experimentally grown assemblages from coastal 926 Georgia and Florida, USA. Mar Ecol-Evol Persp 37, 891-906 (2016). 927 96 Meslard, F., Bourrin, F., Many, G. & Kerhervé, P. Suspended particle dynamics and 928 fluxes in an Arctic fjord (Kongsfjorden, Svalbard). Estuarine, Coastal and Shelf 929 Science 204, 212-224, doi:https://doi.org/10.1016/j.ecss.2018.02.020 (2018).

930 97 Csapó, H. K., Grabowski, M. & Węsławski, J. M. Coming home - Boreal ecosystem 931 claims Atlantic sector of the Arctic. Science of The Total Environment 771, 144817, 932 doi:https://doi.org/10.1016/j.scitotenv.2020.144817 (2021). 933 98 Renaud, P. E. et al. Arctic Sensitivity? Suitable Habitat for Benthic Taxa Is 934 Surprisingly Robust to Climate Change. Frontiers in Marine Science 6, 935 doi:10.3389/fmars.2019.00538 (2019). 936 99 Kujawa, A., Łacka, M., Szymańska, N., Telesiński, M. & Zajączkowski, M. Could 937 Norwegian fjords serve as an analogue for the future of the Svalbard fjords? State and 938 fate of high latitude fjords in the face of progressive "atlantification". Polar Biology, 939 doi:10.1007/s00300-021-02951-z (2021). Gooday, A. J., Anikeeva, O. V. & Pawlowski, J. New genera and species of 940 100 941 monothalamous Foraminifera from Balaclava and Kazach'ya Bays (Crimean 942 Peninsula, Black Sea). Marine Biodiversity 41, 481-494, doi:10.1007/s12526-010-943 0075-7 (2011). 944 101 Altin-Ballero, D. Z., Habura, A. & Goldstein, S. T. Psammophaga sapela n. sp., a 945 new monothalamous foraminiferan from coastal Georgia, U.S.A.: Fine structure, 946 gametogenesis, and phylogenetic placement. Journal of Foraminiferal Research 43, 947 113-126, doi:10.2113/gsjfr.43.2.113 (2013). 948 102 Smith, C. W. & Goldstein, S. T. The Effects of Selected Heavy Metal Elements 949 (arsenic, Cadmium, Nickel, Zinc) On Experimentally Grown Foraminiferal 950 Assemblages from Sapelo Island, Georgia and Little Duck Key, Florida, U.S.A. 951 Journal of Foraminiferal Research 49, 303-317, doi:10.2113/gsjfr.49.3.303 (2019). 952 103 Dorst, S. & Schönfeld, J. Diversity Of Benthic Foraminifera on The Shelf and Slope 953 of The NE Atlantic: Analysis of Datasets. Journal of Foraminiferal Research 43, 238-954 254, doi:10.2113/gsjfr.43.3.238 (2013).

955	104	Majewski, W., Szczuciński, W. & Zajączkowski, M. Interactions of Arctic and
956		Atlantic water-masses and associated environmental changes during the last
957		millennium, Hornsund (SW Svalbard). Boreas 38, 529-544,
958		doi: https://doi.org/10.1111/j.1502-3885.2009.00091.x (2009).
959		

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